WAVES RECORDED AT ST. GOWAN LIGHT VESSEL

by

B C H FORTNUM

Data for August 1975 to July 1976
and January 1977 to December 1978
at position 51°31'N, 005°00'W

Summary Analysis and Interpretation Report

Report No 125
1981
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## CONTENTS

1. INTRODUCTION
   1.1 Site description
   1.2 Description of measuring and recording systems
   1.3 Details of calibration and maintenance
   1.4 Wave data coverage and return

2. WIND DATA – COMPARISON WITH THE LONG-TERM AVERAGE
   2.1 Monthly variation of wind speeds
   2.2 Yearly variation of wind speeds

3. WAVE DATA – DESCRIPTION AND DISCUSSION OF THE PRESENTATIONS
   3.1 Statistics of variations of wave heights
      3.1.1 Monthly variation of Hs
      3.1.2 Yearly variation of Hs
   3.2 Statistics of wave heights
      3.2.1 Occurrence of Hs
      3.2.2 Exceedance of Hs and Hmax(3hr)
   3.3 Design wave heights
      3.3.1 Weibull distribution of Hs
      3.3.2 Fisher-Tippett I distribution of Hs
      3.3.3 Fisher-Tippett III distribution of Hs
      3.3.4 Individual wave model
   3.4 Statistics of wave periods
      3.4.1 Occurrence of Tz
   3.5 Statistics of wave height and period combined
      3.5.1 Occurrences of Hs and Tz combined
   3.6 Statistics of persistence of wave conditions
      3.6.1 Persistence of calms of Hs
      3.6.2 Persistence of storms of Hs

4. ACKNOWLEDGMENTS

5. REFERENCE

APPENDIX I – Method of system calibration

APPENDIX II – Definitions of wave parameters and method of analysis of wave chart data

APPENDIX III – Details of methods used for calculating design wave heights

APPENDIX IV – The correction factor applied to wave heights
LIST OF FIGURES

1.1 Map showing location of site
1.4 Time series of Hs
2.1 Mean and standard deviation of the monthly means of wind speed
2.2 Mean of N largest values of wind speed
3.1.1 Monthly means of Hs for each year
3.1.2 Mean of N largest values of Hs
3.2.1.1 Percentage occurrence of Hs - spring
3.2.1.2 Percentage occurrence of Hs - summer
3.2.1.3 Percentage occurrence of Hs - autumn
3.2.1.4 Percentage occurrence of Hs - winter
3.2.1.5 Percentage occurrence of Hs - annual
3.2.2.1 Percentage exceedance of Hs and Hmax(3hr) - spring
3.2.2.2 Percentage exceedance of Hs and Hmax(3hr) - summer
3.2.2.3 Percentage exceedance of Hs and Hmax(3hr) - autumn
3.2.2.4 Percentage exceedance of Hs and Hmax(3hr) - winter
3.2.2.5 Percentage exceedance of Hs and Hmax(3hr) - annual
3.3.1 Cumulative distribution of Hs: Weibull scale
3.3.2 Cumulative distribution of Hs: Fisher-Tippett I scale
3.3.3 Cumulative distribution of Hs: Fisher-Tippett III scale
3.3.4 Return period v. wave height - individual wave model
3.4.1.1 Percentage occurrence of Tz - spring
3.4.1.2 Percentage occurrence of Tz - summer
3.4.1.3 Percentage occurrence of Tz - autumn
3.4.1.4 Percentage occurrence of Tz - winter
3.4.1.5 Percentage occurrence of Tz - annual
3.5.1.1 Occurrences (parts per thousand) of Hs and Tz combined - spring
3.5.1.2 Occurrences (parts per thousand) of Hs and Tz combined - summer
3.5.1.3 Occurrences (parts per thousand) of Hs and Tz combined - autumn
3.5.1.4 Occurrences (parts per thousand) of Hs and Tz combined - winter
3.5.1.5 Occurrences (parts per thousand) of Hs and Tz combined - annual
3.6.1.1 Mean durations and numbers of calms v. Hs - spring
3.6.1.2 Mean durations and numbers of calms v. Hs - summer
3.6.1.3 Mean durations and numbers of calms v. Hs - autumn
3.6.1.4 Mean durations and numbers of calms v. Hs - winter
3.6.1.5 Mean durations and numbers of calms v. Hs - annual
3.6.2.1 Mean durations and numbers of storms v. Hs - spring
3.6.2.2 Mean durations and numbers of storms v. Hs - summer
3.6.2.3 Mean durations and numbers of storms v. Hs - autumn
3.6.2.4 Mean durations and numbers of storms v. Hs - winter
3.6.2.5 Mean durations and numbers of storms v. Hs - annual
1. INTRODUCTION

1.1 Site description
The site at which the wave measurements were taken is shown on the map in figure 1.1. It is approximately 10 kilometres south of the south-western tip of Wales, at position 51°30.5"N, 004°59.8"W. Although it is moored in water of depth about 49 metres, there are several shoals in its vicinity; in particular a shoal to the north-east has a water depth above it of about 7 metres. The tidal currents in the area are quite strong, the maximum current being over 2 knots (1 metre/sec), with directions of approximately 100° and 280°. Since the vessel is stationed close to several shoals, it is probable that it experiences even stronger currents, which may occasionally cause an increase in the steepnesses of the waves to an unusually high level. This effect would be most pronounced for short, low-period waves.

1.2 Description of measuring and recording systems
The vessel is fitted with a shipborne wave recorder as described in TUCKER (1956), which provides information about the water surface elevation. This information is recorded for a 12-minute period (usually) once every 3 hours by a pen on paper chart rolls. The expression for the correction factor which is applied to the wave height data is given in Appendix IV, together with a table of its values for various values of Tz, using 1.62 metres as the depth of the pressure transducer on the vessel. The correction factor is greater than 2 for Tz of 4.3 seconds and less. Since the expression for the correction factor is rather inaccurate for these low values of Tz, data with Tz values of 4.5 seconds and less are not considered to be reliable. However all the data have been included in the presentations in this report, with a warning shown on those figures which, if used without regard for this limitation on the data, might yield misleading information.

1.3 Details of calibration and maintenance
During the period covered by this report an N10 4753 (valve) shipborne wave recorder was deployed on the light vessel.
(i) 5 August 1975 - Instrument installed on LV no. 10 following laboratory calibration in June 1975.
(ii) 19 August 1976 - LV no. 10 docked at Holyhead for re-fit. SBWR removed and calibration checked. Result of check was an overall change in sensitivity of −1.7% since June 1975.
(iii) 20 October 1976 - SBWR re-installed in LV no. 10. Vessel returned to St. Gowan station in December 1976.
(iv) 19 October 1979 - LV no. 10 docked at Swansea for re-fit. SBWR removed and calibration checked. Result of check was an overall change in sensitivity of −0.5% since August 1976.

1.4 Wave data coverage and return (figures 1.4(A) - 1.4(I))
The periods covered by the data are 1 August 1975 to 31 July 1976, and 1 January 1977 to 31 December 1978. For this period 718 of the 8768 possible chart records were either missing or classified as invalid, resulting in a data return of 91.8%; and 1.34% of the valid records were calms (see Appendix II). No attempt has been made to correct any bias which may have resulted from missing/invalid records, because of the uncertain reliability of available techniques. (Simple gap-filling by
linear interpolation, up to a maximum of 7 consecutive records, has been carried out for the purpose of persistence calculations only: see section 3.6.) The approximate times when missing/invalid records occurred may be derived from the plots in figure 1.4 which show $H_s$ as a time series. On these plots each vertical line represents a valid record, and the height of the line is proportional to the value of $H_s$ for that record: therefore these plots also indicate the variation of $H_s$ with time.

2. WIND DATA - COMPARISON WITH THE LONG-TERM AVERAGE

The meteorological station nearest to the wave measurement site is Milford Haven (51°43'N, 005°02'W) from which wind data have been analysed from the period August 1964 to December 1978. Winds approaching from directions which have very limited fetches associated with them have not been considered, so that only winds in the sector from 200° to 280° have been considered in this report (including a proportion of calms and variables). The data used are three-hourly synoptic wind speeds.

2.1 Monthly variation of wind speeds (figure 2.1)

For each month, the mean of the monthly means of wind speed is plotted. Comparing monthly means for the 'long-term' wind data and for the wind data from the years when waves were recorded, only March appears to show a significant and consistent difference (individual years about 10%-20% higher than the 'long-term' mean) although for two of the years November has a higher wind speed than the 'long-term' mean (by about 23% and 18%).

2.2 Yearly variation of wind speeds (figures 2.2(A) and 2.2(B))

The year-to-year variability of wind conditions is illustrated in these figures. It shows, for each year, the maximum value of wind speed, and also the means of the next $N$ highest wind speeds, where $N = 5, 10, 20, 50, 100$ (thus the highest 186 wind speeds are represented). It can be seen that for 1975/6, 1977 and 1978 the maximum wind speeds are higher than the maximum wind speeds for most of the other years, whilst the mean wind speeds for these same years lie in the middle or even the lower part of the range of mean wind speeds. This suggests that the worst storms during the period of wave recording were amongst the most severe of those for the 14 years considered.
3. WAVE DATA - DESCRIPTION AND DISCUSSION OF THE PRESENTATIONS

Where figures show seasonal data, the seasons are defined as follows:
- spring - March, April, May
- summer - June, July, August
- autumn - September, October, November
- winter - December, January, February

The maximum value of Hs in these three years of data is 9.1 metres; the associated value of Tz is 9.8 seconds, and of Hmax(3hr) is 17.1 metres. However, the largest value of Tz recorded is 13.7 seconds, with an associated Hs of 4.2 metres.

3.1 Statistics of variations of wave heights

3.1.1 Monthly variation of Hs (figure 3.1.1)
For each month, the mean of the significant wave height is calculated and plotted separately for each year. The summer months (June, July and August) show little variation in the mean Hs between years. For December, the yearly variation is very large (the 1978 mean is 2 times greater than the 1975/6 mean); and October, also, has a large yearly variation (the 1977 mean is more than twice the 1978 mean). The winter month of January has the smallest variation when expressed as the ratio of range to mean.

3.1.2 Yearly variation of Hs (figure 3.1.2)
The year-to-year variability of wave conditions is illustrated in this figure. It shows, for each year, the maximum value of Hs, and also the means of the next N highest values of Hs, where N = 5, 10, 20, 50, 100 (thus the highest 186 values of Hs are represented). From this figure there appears to be very little yearly variation in the highest values of Hs, although the first of the three years shows slightly lower values of the highest wave conditions than the latter two years.

3.2 Statistics of wave heights

3.2.1 Occurrence of Hs (figures 3.2.1.1-3.2.1.5)
The percentage occurrence of Hs is shown on histograms. The most frequently occurring values of Hs may be seen, from figure 3.2.1.5, to lie between 0.5 and 1.5 metres, accounting for 37% of the total.

3.2.2 Exceedance of Hs and Hmax(3hr) (figures 3.2.2.1-3.2.2.5)
These graphs may be used to estimate the fraction of the time during which Hs was greater than, or less than, a given height. For instance, from figure 3.2.2.4 it may be seen that during winter the significant wave height exceeded 4 metres for approximately 16 per cent of the time.

3.3 Design wave heights
The methods used to calculate the design wave height (the most probable height of the highest wave with a return period of 50 years) are described in Appendix III. The results obtained by the different methods are given below.

-3-  St. Gowan 75/78
3.3.1 Weibull distribution of Hs (figure 3.3.1)
The parameters of the Weibull distribution which most closely fits the
data are \( A = 0.24 \) metre, \( B = 1.91 \) metres and \( C = 1.39 \), and this
distribution is represented by the straight line in figure 3.3.1. Extrapolation
of this line to a return period of 50 years gives a value of Hs of 11.6 metres. The value of \( T_z \) associated with this Hs is approximately 10.5 seconds, resulting in a value for the design wave height of 21.7 metres.

3.3.2 Fisher-Tippett I distribution of Hs (figure 3.3.2)
The parameters of the Fisher-Tippett I distribution which most closely fits the data are \( a = 1.28 \) metre \(^7\) and \( b = 3.28 \), and this distribution is represented by the straight line in figure 3.3.2. (The parameters have been chosen for a best fit to the top 7 data points.) Extrapolation of this line to a return period of 50 years gives a value of Hs of 11.9 metres. The value of \( T_z \) associated with this Hs is approximately 10.5 seconds, resulting in a value for the design wave height of 22.3 metres.

3.3.3 Fisher-Tippett III distribution of Hs (figure 3.3.3)
The parameters of the Fisher-Tippett III distribution which most closely fits the data are \( A = 34.0 \) metres, \( B = 32.6 \) metres and \( C = 30.9 \), and this distribution is represented by the straight line in figure 3.3.3. Extrapolation of this line to a return period of 50 years gives a value of Hs of 11.7 metres. The value of \( T_z \) associated with this Hs is approximately 10.5 seconds, resulting in a value for the design wave height of 21.9 metres.

3.3.4 Individual wave model (figure 3.3.4)
The value of steepness used in the wave-by-wave method of determining design wave heights (as described in Appendix III) is 1:18, and the inverse mean period is 0.15 Hz. Using these values and the Fisher-Tippett III parameters given in section 3.3.3, the data which appear in figure 3.3.4 are obtained; by interpolation the design wave height is found to be 23.6 metres. A higher value of design wave height is expected from this method than from the methods described above, for the reasons stated in Appendix III.

3.4 Statistics of wave periods
The percentage occurrence of \( T_z \) is shown on a histogram.

3.4.1 Occurrence of \( T_z \) (figures 3.4.1.1-3.4.1.5)
The most frequently occurring values of \( T_z \) in the data set lie between
6.0 and 7.0 seconds (29% of the total), and all values of \( T_z \) lie between 2.5 and 14.0 seconds (figure 3.4.1.5).

3.5 Statistics of wave height and period combined
These figures (sometimes called "scatter" plots) show the numbers of wave records having particular combinations of values of Hs and \( T_z \). The numbers of wave records are presented as parts per thousand (the total number of valid observations being shown on each figure), except for those which would be less than one part per thousand; these are shown instead as single occurrences and are distinguished by being underlined.
3.5.1 Occurrences of Hs and Tz combined (figures 3.5.1.1-3.5.1.5)
On these figures points of equal occurrences are joined by contour lines to give an indication of the bivariate probability distribution of Hs and Tz, and to illustrate the correlation between them. A wave "steepness" (as defined in Appendix III) can be calculated for each (Hs,Tz) pair. A line is drawn on figure 3.5.1.5 showing a "steepness" of 1:12, which is the limiting "steepness" for the main body of the data. (Wave "steepnesses" as shown in this figure are less than the maximum of 1:7 for an individual wave, since Hs and Tz are parameters averaged over a number of waves most of which have steepnesses less than this maximum.)

3.6 Statistics of persistence of wave conditions
These figures show the means and standard deviations of the durations of storms and calms against each threshold value of Hs, and also the percentage of the total duration occupied by each event. Gaps in the data series of 7 or less records are filled (for the purpose of persistence calculations only) by linear interpolation; larger gaps are not filled, effectively reducing the series to a number of smaller sub-series, each with a correspondingly smaller total duration. 'Split' seasons (those in which the months are not consecutive) are not used in the persistence calculations. (For storms, the curves showing percentage of time occupied by the events are, for all practical purposes, the same as those showing percentage exceedance of Hs as described in section 3.2.2.)

3.6.1 Persistence of calms of Hs (figures 3.6.1.1-3.6.1.5)
Information about, for example, calms of Hs less than 0.7 metres at the St. Gowan station during winter can be derived from figure 3.6.1.4. The mean duration of such calms was approximately 15 hours (with a standard deviation of 21 hours); they occupied about 7% of the total duration of 4173 hours, i.e. about 290 hours; and therefore there were 19 or 20 such calm events during this period. Since this represents two seasons, there were on average about 10 such events each winter.

3.6.2 Persistence of storms of Hs (figures 3.6.2.1-3.6.2.5)
Similar information can be derived for storms. For Hs of 3.5 metres during summer, figure 3.6.2.2 shows that the mean duration of such storms was approximately 4 hours (with a standard deviation of 3 hours); the total time occupied was a little over 1% of 4392 hours, i.e. about 50 hours; and therefore the number of such storm events in this period was 12 or 13. Since this represents two seasons, there were on average about 6 such events each summer.

4. ACKNOWLEDGEMENTS
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-5-  St. Gowan 75/78
5. REFERENCE

APPENDIX I

Method of system calibration

1.1 Since there are two types of transducer in the shipborne wave recorder system, it is necessary to divide the calibration procedure into two sections. First the accelerometers are removed from the ship mountings and each is inserted into a rig which allows the transducer to be driven through a vertical circle of diameter 1 metre. The transducer is mounted in gimbals and maintains a vertical attitude during rotation. Two rotation rates are applied: 12 and 18 second periods which are derived from a crystal oscillator. The transducer is connected to the electronics unit in the usual way, and the calibration signal is displayed on the chart recorder. However, because a 1 metre 'heave' is small compared with the wave-heights usually experienced at sea, a precision amplifier (contained in the electronics) is switched into the circuit, converting the 1 metre into an apparent 10 metre signal. The output signal can then be read from the chart record and any corrections to instrument sensitivity made.

The pressure units cannot be easily subjected to a dynamic test since this requires the application of a sinusoidally-varying pressure. Therefore for routine re-calibration a static test is applied. Each pressure unit is fixed to the test rig and a series of discrete pressure levels is applied from a reservoir via a regulator valve. Each pressure level is set manually with the valve by reference to a precise pressure transducer contained within the calibrator unit. The output voltage of the transducer is monitored in the SBWR electronics unit and compared to the original laboratory calibration. Any changes in sensitivity are then compensated for by adjustment of the input amplifier gain.

Full re-calibrations are usually only possible when the ship comes into dock for its 3-yearly re-fit.

1.2 Monthly checks
All lightvessel crews are asked to drain water through the valve assemblies to ensure that no blockage prevents the water pressure being transmitted to the pressure sensors, and then to take a test record, on a monthly basis. The test record consists of a short length of pen-trace with all transducers turned off (electrically), followed by a few minutes recording with each transducer on its own. The record thus produced shows two heave records (one from each accelerometer) which should look broadly similar; and also the pressure traces, which may not agree so well, but when compared with other monthly test records should exhibit no systematic error. These tests are not direct checks on calibration accuracy but are often good indicators of a fault condition developing.
APPENDIX II

Definitions of wave parameters and method of analysis of wave chart data

II.1.1 Tz - the mean zero up-crossing period. This is defined as the duration of the record divided by the number of zero up-crossings Nz. (A zero up-crossing is considered to occur when the trace crosses the mean line in an upward direction.)

II.1.2 Hs - the significant wave height. This is defined as $4\sigma$ where $\sigma$ is the standard deviation of the record. (An estimate of $\sigma$ is obtained from Nz and from the excursions of the two highest maxima, and of the two lowest minima, from the record mean.) For a narrow band random process this parameter approximates closely to the mean height of the highest one-third zero up-cross waves (see LONGUET-HIGGINS (1952)). Comparison between the two definitions is made by GODA (1970, 1974). (A zero up-cross wave is defined as the portion of the wave record between two zero up-crossings, and its height is the vertical distance between the highest and lowest points on the wave.)

II.1.3 Hmax(3hr) - the most probable height of the highest zero up-cross wave in the recording interval of three hours. (This is derived from Nz, and the duration of the recording interval.) The parameter Hmax(3hr), which is the mode of the distribution, should not be confused with the expected height of the highest wave in three hours, which is the mean of the distribution. The mean of the distribution is typically 3 per cent higher than the mode (see TANN (1976)).

(A 'calm' record is one for which the sum of the height of the highest crest and of the lowest trough is less than 1 foot or 0.3 metre.)

II.2 References


APPENDIX III

Details of methods used for calculating design wave heights

III.1 By finding the long-term distribution of Hs

III.1.1 Hs is used as a measure of the "sea-state" (i.e. the intensity of wave activity), and it is sampled every 3 hours. It is assumed that a set of Hs data for one year, or an integral number of years, is representative of the wave climate. For each value of Hs, the probability that this value will not be exceeded is calculated; this probability is then plotted against Hs. The axes are scaled according to a long-term distribution, so that data with a perfect fit would appear as a straight line on the diagram. This procedure is carried out using long-term distributions defined in the following ways:

* Weibull
  \[
  \text{Prob}(H_s \leq h) = \begin{cases} 
  1 - \exp\left[-\left(\frac{h-A}{B}\right)^C\right], & \text{for } h > A \\
  0 & \text{for } h \leq A 
  \end{cases}
  \]
  where B and C are positive, and A represents a lower bound on h.

* Fisher-Tippett I (first asymptote)
  \[
  \text{Prob}(H_s \leq h) = \exp[- \exp(-ah+b)].
  \]

* Fisher-Tippett III (third asymptote)
  \[
  \text{Prob}(H_s \leq h) = \begin{cases} 
  \exp\left[-\left(\frac{A-h}{B}\right)^C\right], & \text{for } h \leq A \\
  1 & \text{for } h > A 
  \end{cases}
  \]
  where B and C are positive, and A represents an upper bound on h. (See FISHER AND TIPPETT(1928) and GUMBEL (1958) for the derivations of these distributions.)

For each long-term distribution the best-fit straight line is drawn; this line is then extrapolated to the desired probability (see section III.1.2) and the corresponding value of Hs is read off as the "design sea-state".

III.1.2 To calculate the "sea-state" which will be exceeded only once in N years, a storm duration of D hours needs to be assumed. The probability that a randomly chosen time will be within this storm is then

\[
\frac{D}{24 \times 365.25 \times N}
\]
IOS uses $D = 3$ hours (this choice is discussed in section III.1.5) which gives

$$\text{Probability} = \frac{3.422 \times 10^{-4}}{N} = 6.845 \times 10^{-6} \text{ for } N = 50 \text{ years.}$$

III.1.3 The value of $T_z$ for the "design sea-state" is required before the highest wave in the storm can be calculated. This is derived from the bivariate distribution of $H_s$ and $T_z$ (figure 3.5.1.5). A line is drawn across this at the "design sea-state" value of $H_s$ and the most likely value of $T_z$ (the modal value) is then estimated using extrapolations of the probability contours.

III.1.4 The most probable value of the highest zero-up-cross wave in the storm is then derived by assuming that the heights of such waves follow a Rayleigh distribution whose probability density function is

$$\text{prob}(h) = \frac{2h}{(\text{Hrms})^2} \exp\left[-\left(\frac{h}{\text{Hrms}}\right)^2\right]$$

where $\text{Hrms} \approx \frac{H_s}{\sqrt{2}}$.

Exact theory is not available for zero-up-cross wave heights, but this distribution has been found to be an adequate fit to measured data. If there are $n$ waves in the recording interval (3hr), then the probability that the highest wave, $H$, in three hours is less than $h$ is

$$\text{Prob}(H<h) = \left\{1 - \exp\left[-\left(\frac{h}{\text{Hrms}}\right)^2\right]\right\}^n$$

with a corresponding probability density function

$$\frac{2n}{(\text{Hrms})^2} \text{h} \exp\left[-\left(\frac{h}{\text{Hrms}}\right)^2\right] \left\{1 - \exp\left[-\left(\frac{h}{\text{Hrms}}\right)^2\right]\right\}^{n-1}$$

The most probable value (the mode) of this probability density function is usually used and is given by

$$\text{Hmax (3hr)} = \text{Hrms} \sqrt{\Psi}$$

where $\Psi$ is a function of $T_z$ which may be found using either figure 7 or equation 6.1-2 in TANN(1976).

III.1.5 In choosing the value of storm duration $D$, it should be noted that the effect of increasing $D$ is to decrease the value of $H_s$ for a given return period $N$. However, it also increases the ratio of $\text{Hmax (3hr)}$ to $H_s$. It is found that in practice these effects roughly cancel and typically the value of $\text{Hmax (3hr)}$ changes by only 3 per cent for a change of $D$ from 3 to 15 hours. The choice of $D$ is therefore not critical.

Many details of the above procedures may be found in TANN(1976).
III.2 By a wave-by-wave method

III.2.1 Battjes (1970) shows that the probability that a randomly chosen wave will have a height H greater than h is

\[
\text{Prob}(H > h) = \frac{\int_0^\infty \int_0^\infty R(h, Hs) Tz^{-1} p(Tz, Hs) dHs \, dTz}{\int_0^\infty \int_0^\infty Tz^{-1} p(Tz, Hs) dHs \, dTz}
\]

where \( R(h, Hs) \) is the Rayleigh cumulative probability function and \( p(Tz, Hs) \) is the joint probability density function of Hs and Tz.

III.2.2 Tanaka makes the following suggestion in an unpublished manuscript. In order that values of Hs higher than those actually measured may be represented in the calculation of this probability, the values of Hs are assumed to have a long-term cumulative probability function \( F(Hs) \), and a probability density function \( f(Hs) = F'(Hs) \).

For each value of Hs throughout the long-term distribution, an average value of \( Tz^{-1} \) is used (denoted by \( Tz^{-1}(Hs) \)). It is defined as

\[
Tz^{-1}(Hs) = \int_0^\infty Tz^{-1} \frac{p(Tz, Hs)}{P(Hs)} \, dTz
\]

where \( P(Hs) = f(Hs) \).

Therefore

\[
\int_0^\infty Tz^{-1} p(Tz, Hs) \, dTz = Tz^{-1}(Hs)
\]

which, when substituted into equation (1), allows the probability of exceedance to be written

\[
\text{Prob}(H > h) = \frac{\int_0^\infty R(h, Hs) Tz^{-1}(Hs) f(Hs) \, dHs}{\int_0^\infty Tz^{-1}(Hs) f(Hs) \, dHs}
\]

The value of \( Tz^{-1}(Hs) \) used with each value of Hs is chosen to satisfy the condition of constant wave "steepness", where the relationship between "steepness" (1/s), water depth (d), Hs and Tz is

\[
Tz = \sqrt{\frac{2 \pi g Hs}{3 \coth \left( \frac{2 \pi d}{3 Hs} \right)}}.
\]

The value for the steepness used in this report is given in section 3.3.4.

The long-term distribution used in the computation for this report is the Fisher-Tippett III extreme-value distribution, whose probability density function is

\[
f(Hs) = \frac{C}{A - Hs} \left( \frac{A - Hs}{B} \right)^C \exp \left[ - \left( \frac{A - Hs}{B} \right)^C \right].
\]

The constants A, B, C are determined graphically as described in section III.1.1, and their values as used in this report are given in section 3.3.3.
III.2.3 Thus the probability of a wave exceeding each particular wave height may be found, and this probability may be converted into a return period of $N$ years using the formula

$$N = \frac{1}{365.25 \times 24 \times 3600 \times T_{ave}^{-1} \times \text{Prob}}$$

where $T_{ave}^{-1} = \frac{1}{\text{average period}}$.

The value of the average wave period is contained in section 3.3.4. Since $T_{ave}^{-1}$ is a non-analytic function of Prob, the simplest way of solving the problem is to calculate $\text{Prob}$ for various values of $h$, calculate $N$ for each of these values of $\text{Prob}$, and then interpolate to find the height $h$ corresponding to the required value of $N$ (in this case 50 years).

Whereas the method described in section III.1 assumes that the highest wave in a 50-year period will come from the most stormy 3-hour period in 50 years, the individual wave method takes into account the probability that storms other than the highest may provide the wave with a 50-year return period. Consequently, the height of a 50-year wave as estimated by this method is likely to be greater than that estimated from the method of using a long-term distribution of $H_s$.

III.3 References


APPENDIX IV

The correction factor applied to wave heights

Two corrections need to be applied to the wave height data:
(i) to compensate for the frequency response of the electronics; and
(ii) to compensate for the hydrodynamic attenuation of the pressure fluctuations.
These are combined into a single correction factor $C$, which is dependent upon $T_z$, and upon the depth of the pressure transducer below the mean water level. The correction factor is

$$C = 0.83 \left(1 + \frac{1}{77.44 \omega^2} \right)^{\frac{3}{2}} \exp \left[ \frac{2.5 \omega^2 d}{g} \right]$$

where $d$ is the depth of the pressure transducer,
and $\omega = \frac{2\pi}{T_z}$.

The value of $H_s$ calculated by the method outlined in Appendix II is multiplied by $C$ to obtain the corrected value of $H_s$.

A table showing the values of $C$ for various values of $T_z$ (and with $d = 1.62$ metres for the lightvessel on the St. Gowan station) is given below.

<table>
<thead>
<tr>
<th>$T_z$(sec)</th>
<th>$C$</th>
</tr>
</thead>
<tbody>
<tr>
<td>2.5</td>
<td>11.224</td>
</tr>
<tr>
<td>3.0</td>
<td>5.076</td>
</tr>
<tr>
<td>4.0</td>
<td>2.311</td>
</tr>
<tr>
<td>4.5</td>
<td>1.871</td>
</tr>
<tr>
<td>5.0</td>
<td>1.610</td>
</tr>
<tr>
<td>6.0</td>
<td>1.327</td>
</tr>
<tr>
<td>7.0</td>
<td>1.185</td>
</tr>
<tr>
<td>8.0</td>
<td>1.104</td>
</tr>
<tr>
<td>10.0</td>
<td>1.025</td>
</tr>
<tr>
<td>12.0</td>
<td>0.996</td>
</tr>
<tr>
<td>15.0</td>
<td>0.974</td>
</tr>
<tr>
<td>20.0</td>
<td>1.040</td>
</tr>
</tbody>
</table>
LOCATION MAP OF ST. GOWAN WAVE RECORDER

FIG 1.1
TIME SERIES OF HS


FIG 1.4(A)

DATE

NOV 1975

DATE

OCT 1975

DATE

SEP 1975

DATE

AUG 1975

HS (METERS)

0 1 2

3 4 5 6 7 8 9 10

HS (METERS)

0 1 2

3 4 5 6 7 8 9 10

HS (METERS)

0 1 2

3 4 5 6 7 8 9 10

HS (METERS)

0 1 2

3 4 5 6 7 8 9 10

HS (METERS)

0 1 2

3 4 5 6 7 8 9 10
TIME SERIES OF HS

MAR 1976

DATE

FEB 1976

DATE

JAN 1976

DATE

DEC 1975

DATE
FIG 10.4(C)
TIME SERIES OF HS

JUL 1976
DATE

JUN 1976
DATE

MAY 1976
DATE

APR 1976
DATE

Hs (METERS)
TIME SERIES OF HS

AUG 1977

DATE

JUL 1977

DATE

JUN 1977

DATE

MAY 1977

DATE
TIME SERIES OF HS

DEC 1977

NOV 1977

OCT 1977

SEP 1977

DATE

DATE

DATE

DATE

KEY


FIG 1.4 (F)
FIG 1.e (H)
TIME SERIES OF HS

AUG 1978

JUL 1978

JUN 1978

MAR 1978

DATE

DATE

DATE

DATE

HS (METERS)

HS (METERS)

HS (METERS)

HS (METERS)

KEY CURVE
FIG 1.c(1)


TIME SERIES OF HS

DEC 1978

OCT 1978

NOV 1978

SEP 1978

Date

Date

Date

Date

Hs (Metres)

Hs (Metres)

Hs (Metres)

Hs (Metres)
Fig 2.4

Milford Haven Jan 1965 - Dec 1978

Wind Speed

Mean and standard deviation of the monthly mean of

1978
1977
1976/6

METRES/SEC

JAN FEB MAR APR MAY JUNE JULY AUG SEPT OCT NOV DEC
FIG 2.2(A)

MILFORD HAVEN JAN 1965 - DEC 1978
MEAN OF N LARGEST VALUES OF WIND SPEED
FIG. 3.4.1


MONTHLY MEAN OF Hs FOR EACH YEAR

<table>
<thead>
<tr>
<th>Year</th>
<th>Jan</th>
<th>Feb</th>
<th>Mar</th>
<th>Apr</th>
<th>May</th>
<th>Jun</th>
<th>Jul</th>
<th>Aug</th>
</tr>
</thead>
<tbody>
<tr>
<td>1978</td>
<td>*</td>
<td>*</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1977</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1975/6</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Mean all years - •

METRES
Mean of N Largest Values of H°
ST. GOMAN LU 1975/6, 1977, 1978 - AUTUMNS

PERCENTAGE OCCURRENCE HISTOGRAM

HS (METERS)

PERCENTAGE OCCURRENCE

PERCENTAGE OCCURRENCE HISTOGRAM

HS (METERS)

PERCENT OCCURRENCE

100 OF WHICH WERE CALMS
8000 VALID OBSERVATIONS
 FIG 3.2.201


PERCENTAGE EXCEEDANCE OF $H_S$ AND $H_{MAX(3HR)}$

WAVE HEIGHT (M)

PERCENTAGE EXCEEDANCE

24 OF WHICH ARE CALMS
2149 VALID OBSERVATIONS
FIG 3.2.2.2


PERCENTAGE EXCEEDANCE OF $H_s$ AND $H_{MAX(3HR)}$

WAVE HEIGHT (M)

PERCENTAGE EXCEEDANCE

59 OF WHICH ARE CALMS
1999 VALID OBSERVATIONS
FIG 3.2.2,3


PERCENTAGE EXCEEDANCE OF Hs AND HMAX(3HR)

WAVE HEIGHT (M)
FIG 3.2.0.4


PERCENTAGE EXCEEDANCE OF Hs AND Hmax (3HR)

WAVE HEIGHT (M)

PERCENTAGE EXCEEDANCE

2098 VALID OBSERVATIONS
FIG 3.2°2.5
ST. COLUMBIA AUG 1975 - JUL 1977, 1978

PERCENTAGE EXCEEDANCE OF Hs AND HMAX (3HR)

WAVE HEIGHT (m)

PERCENTAGE EXCEEDANCE

100
90
80
70
60
50
40
30
20
10
0

100 VALID OBSERVATIONS
Cumulative distribution of wave height HS

8500 valid observations

Weibull scale

A = 0.24

Fig. 3.3.1
FIG 3.3.2
CUMULATIVE DISTRIBUTION OF WAVE HEIGHT, HS

RETURN PERIOD (YEARS)

WAVE HEIGHT (M)

8050 VALID OBSERVATIONS
FISHER-TIPPER I SCALE

PROBABILITY OF NON-EVENTANCE

6666°
6666°
6666°
Figure 3.3

CUMULATIVE DISTRIBUTION OF WAVE HEIGHT, HS

WAVE HEIGHT (M)

RETURN PERIOD (YEARS)

8050 VALID OBSERVATIONS
FISHER-TIPPETT III SCALE
A = 34°0

PROBABILITY OF NON-EXCEEDANCE

66°0
666°0
6666°0
66666°0
666666°0
RETURN PERIOD vs WAVE HEIGHT - INDIVIDUAL WAVE MODEL

AVEO (M)

RETURN PERIOD (YEARS)

AVEO PERIOD) = 0.150
STEEPNESS = 1:18
FIG 3.4.10.1

ST. GOMAN LV 1975-6, 1977, 1978 - SPRINGS

PERCENTAGE OCCURRENCE HISTOGRAM

TZ (SECONDS)
ST. GOMAN LV 1975/6, 1977, 1978 - WINTER

PERCENTAGE OCCURRENCE HISTOGRAM

TZ (SECONDS)

PERCENTAGE OCCURRENCE

THIS ANALYSIS
(2) VALUES EXCLUDED FROM
20% LVLD OBSERVATIONS

PERCENTAGE OCCURRENCE HISTOGRAM

TZ (SECONDS)

PERCENTAGE OCCURRENCE

THIS ANALYSIS
1400 CASES EXCLUDED FROM
SPACE VULG OBSERVATIONS
FIG 3.5.2.1

ST GONAN LU 1975/6, 1977, 1978 - SPRINGS

SCATTER PLOT OF HS AND T2

TZ (SECONDS)

(see section 1.2)

data unreliable

NO. OF OCCURRENCES (12 PPT)
14 PARTS PER MILLION (ppm)

KEY

2095 VALID OBSERVATIONS

THIS ANALYSIS
24 CLARA ETC. 1982 - FROM

METRES
FIG 3.5.1.2

ST. GOMAN LU 1975/6, 1977, 1978 - SUMMERS

SCATTER PLOT OF HS AND T2

(T2(SECONDS)

0 1 2 3 4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20

data unreliable
(see section 1.2)

HS(METRES)

1 2 3 4 5 6 7 8 9 10 11 12 13 14 15

KEY

1440 VALID OBSERVATIONS
11 PARTS PER THOUSAND (PPT)

NO. OF OCCURRENCES (1 PPT)

19 CRAIG (this algorithm)

CRAIG EXCLUDED FROM

1440 VALID OBSERVATIONS
FIG 3.5.1.3

ST. COLUMAN LV 1975-6, 1977, 1978 - AUTUMNS

SCATTER PLOT OF HS AND TZ

TZ (SECONDS)

HS (METRES)

KEY

NO. OF OCCURRENCES (OUT)
14 PARTS PER THOUSAND (PPT)

(See section 1.2)

Data unreliable

1831 VALID OBSERVATIONS

(22 CLAMS EXCLUDED FROM THIS ANALYSIS)
ST. GOWAN LU 1975/6, 1977, 1978 - SPRINGS

FIG. 3.6.1.1

DURATION OF EVENT (HR)

HS THRESHOLD (M)

PERCENTAGE OF TIME OCCUPIED BY EVENT

TOTAL DURATION = 6385 HRS

PERCENTAGE OF TIME

MEAN DURATION AND STANDARD DEVIATION

KEY TO PERSISTENCE EVENTS
FIG 3a.6.10

CALS

ST. COWAN LU 1975/6, 1977, 1978 - AUTUMNS

HS THRESHOLD (M)

PERCENTAGE OF TIME OCCUPIED BY EVENT

TOTAL DURATION = 5694 HRS

MEAN DURATION

KEY TO RESISTANCE EVENTS

PERCENTAGE OF TIME

STANDARD DEVIATION

DURATION OF EVENT (HR)
Fig 3.6.2.4

Storms

St. Gowan Lu 1975/6, 1977, 1978 - Springs

HS Threshold (m)

Percentage of Time Occupied by Event

Total Duration = 6588 Hrs

Percentage of Time

Standard Deviation

Mean Duration

Key to Persistence Events
FIG 3.6.2.3
STORMS
ST. GOMAN LU 1975/6, 1977, 1978 - AUTUMNS

HS THRESHOLD (M)

PERCENTAGE OF TIME OCCUPIED BY EVENT

TOTAL DURATION = 3594 HRS

--- DURATION OF EVENT
--- PERCENTAGE OF TIME
--- STANDARD DEVIATION
--- MEAN DURATION

KEY TO PERSISTENCE EVENTS
FIG 3.6.2.4

STORMS

ST. GOWAN LU 1975/6, 1977, 1978 - WINTERS

H5 - THRESHOLD (M)

PERCENTAGE OF TIME OCCUPIED BY EVENT

TOTAL DURATION = 4173 HRS

- PERCENTAGE OF DURATION
- STANDARD DEVIATION
- MEAN DURATION

KEY TO PERSISTENCE EVENTS

DURATION OF EVENT (HR)