

**I.O.S.**

**WAVES RECORDED AT DOWSING LIGHT VESSEL  
BETWEEN 1970 AND 1979**

by  
**B C H FORTNUM**

Data for May 1970 to April 1971  
November 1975 to October 1976  
and August 1977 to July 1979  
at position 53°34.0'N, 000°50' E

**Summary Analysis and Interpretation Report**

**Report No 126**

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B.C.H.FORTNUM

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at position  $53^{\circ}34.0'N$ ,  $000^{\circ}50.2'E$

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## 1. INTRODUCTION

### 1.1 Site description

The site at which the wave measurements were taken is shown on the map in figure 1.1. It is approximately 48 kilometres east of Spurn Point on the east coast of England, at position  $53^{\circ}34.0'N$ ,  $000^{\circ}50.2'E$ . It is moored in water of depth about 26 metres; however the sea-bed in its vicinity is very irregular, and several shoals are charted with water depths above them of between 6 and 20 metres. The tidal currents in the area are quite strong, the maximum current being about 2 knots (1 metre/sec), with directions of approximately  $150^{\circ}$  and  $330^{\circ}$ . Due to the shoals around the vessel the currents may be even stronger, and may occasionally cause either a real or an apparent increase in the steepness of the waves to an unusually high level. This effect would be most pronounced for short, low-period waves.

### 1.2 Description of measuring and recording systems

The vessel is fitted with a shipborne wave recorder as described in TUCKER (1956), which provides information about the water surface elevation. This information is recorded for a 12-minute period (usually) once every 3 hours by a pen on paper chart rolls. The expression for the correction factor which is applied to the wave height data is given in Appendix IV, together with a table of its values for various values of  $T_z$ , using appropriate values for the depth of the pressure transducer on the vessel (1.46 metres for the 1970/71 data, and 0.88 metres for the 1975/79 data). The correction factor is greater than 2 for  $T_z$  of 4.1 seconds and less for the 1970/71 data, and for  $T_z$  of 3.2 seconds and less for the 1975/79 data. Since the expression for the correction factor is rather inaccurate for these low values of  $T_z$ , data with  $T_z$  values of 3.5 seconds and less are not considered to be reliable (together with some of the data with  $T_z$  values between 3.5 and 4.0 seconds). However all the data have been included in the presentations in this report, with a warning shown on those figures which, if used without regard for this limitation on the data, might yield misleading information.

### 1.3 Details of calibration and maintenance

During the period covered by this report an NIO 4753 (valve) shipborne wave recorder was deployed on the lightvessel.

(i) 20 May 1970 - Instrument was overhauled and re-calibrated whilst the vessel was in dry dock at South Shields.

(ii) October 1971 - Repair and re-calibration of equipment were carried out following report of intermittent fault. The re-calibration showed that there had been a change in sensitivity of -5.6% (i.e. the accelerometers were under-reading the waves by this amount). It was not possible to find out whether the replacement component had contributed to this change, and therefore it was assumed that all data recorded up to the time of the intermittent fault were acceptable.

(iii) September/October 1975 - A replacement wave recorder was installed during the vessel's re-fit.

(iv) May/June 1977 - A re-calibration was carried out during the vessel's re-fit, and the change in sensitivity was found to be -3.2%. Vessel returned to Dowsing station on 23 September 1977.

(v) 12 October 1981 - A calibration was carried out when the vessel came off station, and the change in sensitivity was found to be +9.4%. It has been assumed that the change was linear with time.

#### 1.4 Wave data coverage and return (figures 1.4(A) - 1.4(L))

The periods covered by the data are 1 May 1970 to 30 April 1971, 1 November 1975 to 31 April 1971, and 1 August 1977 to 31 July 1979. For this period 158 of the 11688 possible chart records were either missing or classified as invalid, resulting in a data return of 98.7%; and 2.06% of the valid records were calms (see Appendix II). No attempt has been made to correct any bias which may have resulted from missing/invalid records, because of the uncertain reliability of available techniques. (Simple gap-filling by linear interpolation, up to a maximum of 7 consecutive records, has been carried out for the purpose of persistence calculations only: see section 3.6.) The approximate times when missing/invalid records occurred may be derived from the plots in figure 1.4 which show Hs as a time series. On these plots each vertical line represents a valid record, and the height of the line is proportional to the value of Hs for that record: therefore these plots also indicate the variation of Hs with time.

## 2. WIND DATA - COMPARISON WITH THE LONG-TERM AVERAGE

The meteorological station nearest to the wave measurement site is Spurn Head (53°35'N, 000°07'E) from which wind data have been analysed for the period January 1957 to July 1979. Winds approaching from directions which have very limited fetches associated with them have not been considered, so that only winds in the sector from 320° to 160° have been considered in this report (including a proportion of calms and variables). The data used are three-hourly synoptic wind speeds.

### 2.1 Monthly variation of wind speeds (figure 2.1)

For each month, the mean of the monthly means of wind speed is plotted. Only two months, January and December, have monthly means which are not distributed fairly evenly about their 'long-term' monthly means; these can be seen to be the two stormiest months in the period considered, and in both cases three out of the four monthly means are higher than the 'long-term' means.

### 2.2 Yearly variation of wind speeds (figures 2.2(A) - 2.2(C))

The year-to-year variability of wind conditions is illustrated in this figure. It shows, for each year, the maximum value of wind speed, and also the means of the next N highest wind speeds, where N = 5, 10, 20, 50, 100 (thus the highest 186 wind speeds are represented). Figure 2.2(A) shows that the wind speeds in 1970/71 lie in the top half of the range of wind speeds for the 14 year period; from figure 2.2(B) it can be seen that the very highest wind speeds in 1975/76 are towards the top of the range, and the mean is in the middle of the range of wind speeds for the 15 year period; and figure 2.2(C) shows that apart from the very highest wind speeds, the winds for 1977/79 are generally in the lower half of the wind speeds for the 15 year period. From these rather simple analyses, it may be deduced that the 4 years of wave data presented in this report were recorded during years which were not ex-

ceptionally and consistently stormy, although the very highest storms occurring during the 4-year period may have been amongst the highest in the 'long-term' period.

### 3. WAVE DATA - DESCRIPTION AND DISCUSSION OF THE PRESENTATIONS

Where figures show seasonal data, the seasons are defined as follows:

spring - March, April, May

summer - June, July, August

autumn - September, October, November

winter - December, January, February

The maximum value of Hs in these four years of data is 6.5 metres; the associated values of Tz and of Hmax(3hr) are also the highest in the data set, being 9.7 seconds and 12.2 metres.

#### 3.1 Statistics of variations of wave heights

##### 3.1.1 Monthly variation of Hs (figure 3.1.1)

For each month, the mean of the monthly means of significant wave height is plotted. February and December have the widest yearly variations in the values of mean Hs. This figure also shows that the highest wave conditions probably occurred during 1978/9 (as experienced in the months of January, February, March and December); and that the least severe conditions were during 1975/6 (only January, September and October are exceptions). This is in agreement with the conclusions drawn from figure 3.1.2 (see section 3.1.2).

##### 3.1.2 Yearly variation of Hs (figure 3.1.2)

The year-to-year variability of wave conditions is illustrated in this figure. It shows, for each year, the maximum value of Hs, and also the means of the next N highest values of Hs, where N = 5, 10, 20, 50, 100 (thus the highest 186 values of Hs are represented). The figure shows that, of the 4 years considered, the highest values of Hs were greatest during 1978/9, and least (marginally) during 1975/6. However, the yearly variation is not great.

#### 3.2 Statistics of wave heights

##### 3.2.1 Occurrence of Hs (figures 3.2.1.1-3.2.1.5)

The percentage occurrence of Hs is shown on histograms. The most frequently occurring values of Hs may be seen, from figure 3.2.1.5, to lie between 0.5 and 1.5 metres, accounting for 57% of the total.

##### 3.2.2 Exceedance of Hs and Hmax(3hr) (figures 3.2.2.1-3.2.2.5)

These graphs may be used to estimate the fraction of the time during which Hs was greater than, or less than, a given height. For instance, from figure 3.2.2.4 it may be seen that during winter the significant wave height exceeded 2.5 metres for approximately 19 per cent of the time.

### 3.3 Design wave heights

The methods used to calculate the design wave height (the most probable height of the highest wave with a return period of 50 years) are described in Appendix III. The difficulty in fitting one of the extreme-value distributions to this data set may be shown by reference to figure 3.3.2. The upper values of  $H_s$  (as shown by the crosses) appear from this figure to be limited in magnitude, possibly due to the shallow water depths in the area. Therefore a Fisher-Tippett III distribution, which assumes an upper limit to  $H_s$ , might be expected to fit the data best. A best fit of this distribution to all the data is found for  $A = 14$  metres, resulting in  $H_s(50) = 7.88$  metres.

However, since any of the long-term distributions will be greatly influenced by the highest data point, it is useful to examine it more closely. This data point includes the highest value of  $H_s$ , which is 6.46 metres; an increase of less than 1% in this value (well within the standard error in  $H_s$  of 6% which is inherent in the method of estimation, as described in Appendix II) would move this estimate from the 6.0-6.5 metres class to the 6.5-7.0 metres class. Therefore on figure 3.3.2 an additional point will appear in this top class, and also the probability associated with the next lowest class will be changed (these changes are shown by asterisks on figure 3.3.2). Within the limitations on the accuracy of  $H_s$ , this new data set is an equally valid representation of the wave climate at the site as is the original data set. Fitting a Fisher-Tippett I distribution to the whole of this new data set gives  $H_s(50) = 8.94$  metres. Since a small error in the estimate of the maximum  $H_s$  leads to a large difference in  $H_s(50)$ , the distributions described below have been chosen so that, visually, the top point has been given less weight than the others. (Note that the Fisher-Tippett III distribution leads to a much lower value of  $H_s(50)$  than do the other two distributions since it is more sensitive to the upper data points.)

#### 3.3.1 Weibull distribution of $H_s$ (figure 3.3.1)

The parameters of the Weibull distribution which most closely fits the data are  $A = 0.32$  metre,  $B = 1.12$  metres and  $C = 1.25$  (figure 3.3.1). Using this distribution, the value of  $H_s$  with a return period of 50 years is 8.5 metres. The value of  $T_z$  associated with this  $H_s$  is approximately 10 seconds, resulting in a value for the design wave height of 16.0 metres.

#### 3.3.2 Fisher-Tippett I distribution of $H_s$ (figure 3.3.2)

The parameters of the Fisher-Tippett I distribution which most closely fits the data are  $a = 1.45$  metre<sup>-1</sup> and  $b = 1.28$  (figure 3.3.2). Using this distribution, the value of  $H_s$  with a return period of 50 years is 9.1 metres. The value of  $T_z$  associated with this  $H_s$  is approximately 10 seconds, resulting in a value for the design wave height of 17.0 metres.

#### 3.3.3 Fisher-Tippett III distribution of $H_s$ (figure 3.3.3)

The parameters of the Fisher-Tippett III distribution which most closely fits the data are  $A = 14.0$  metres,  $B = 13.4$  metres and  $C = 14.1$  (figure 3.3.3). Using this distribution, the value of  $H_s$  with a return period of 50 years is 8.2 metres. The value of  $T_z$  associated with this  $H_s$  is approximately 9.5 seconds, resulting in a value for the design wave height of 15.5 metres.

### 3.3.4 Individual wave model (figure 3.3.4)

The value of steepness used in the wave-by-wave method of determining design wave heights (as described in Appendix III) is 1:18, and the inverse mean period is 0.194 Hz. Using these values and the Fisher-Tippett III parameters given in section 3.3.3, the data which appear in figure 3.3.4 are obtained; by interpolation the design wave height is found to be 16.1 metres. A higher value of design wave height is expected from this method than from the methods described above, for the reasons stated in Appendix III.

### 3.4 Statistics of wave periods

The percentage occurrence of  $T_z$  is shown on a histogram.

#### 3.4.1 Occurrence of $T_z$ (figures 3.4.1.1-3.4.1.5)

The most frequently occurring values of  $T_z$  in the data set lie between 4.0 and 5.0 seconds (31% of the total), and all values of  $T_z$  lie between 2.5 and 10.0 seconds (figure 3.4.1.5).

### 3.5 Statistics of wave height and period combined

These figures (sometimes called "scatter" plots) show the numbers of wave records having particular combinations of values of  $H_s$  and  $T_z$ . The numbers of wave records are presented as parts per thousand (the total number of valid observations being shown on each figure), except for those which would be less than one part per thousand; these are shown instead as single occurrences and are distinguished by being underlined.

#### 3.5.1 Occurrences of $H_s$ and $T_z$ combined (figures 3.5.1.1-3.5.1.5)

On these figures points of equal occurrences are joined by contour lines to give an indication of the bivariate probability distribution of  $H_s$  and  $T_z$ , and to illustrate the correlation between them. A wave "steepness" (as defined in Appendix III) can be calculated for each ( $H_s, T_z$ ) pair. A line is drawn on figure 3.5.1.5 showing a "steepness" of 1:12, which is the limiting "steepness" for the main body of the data. (Wave "steepnesses" as shown in this figure are less than the maximum of 1:7 for an individual wave, since  $H_s$  and  $T_z$  are parameters averaged over a number of waves most of which have steepnesses less than this maximum.)

### 3.6 Statistics of persistence of wave conditions

These figures show the means and standard deviations of the durations of storms and calms against each threshold value of  $H_s$ , and also the percentage of the total duration occupied by each event. Gaps in the data series of 7 or less records are filled (for the purpose of persistence calculations only) by linear interpolation; larger gaps are not filled, effectively reducing the series to a number of smaller sub-series, each with a correspondingly smaller total duration. 'Split' seasons (those in which the months are not consecutive) are not used in the persistence calculations. (For storms, the curves showing percentage of time occupied by the events are, for all practical purposes, the same as those showing percentage exceedance of  $H_s$  as described in section 3.2.2.)

### 3.6.1 Persistence of calms of Hs (figures 3.6.1.1-3.6.1.5)

Information about, for example, calms of Hs less than 0.9 metres at the Dowsing station during winter can be derived from figure 3.6.1.4. The mean duration of such calms was about 11.5 hours (with a standard deviation of 18 hours); they occupied about 18% of the total duration of 8652 hours, i.e. about 1560 hours; and therefore there were between 130 and 140 such calm events during this period. Since this represents 4 seasons, there were on average about 35 such events each winter.

### 3.6.2 Persistence of storms of Hs (figures 3.6.2.1-3.6.2.5)

Similar information can be derived for storms. For Hs of 3.9 metres during spring, figure 3.6.2.1 shows that the mean duration of such storms was approximately 56 hours (the fact that no value of standard deviation is shown for this value of Hs indicates that there was only one event). The total time occupied was approximately 1% of 6573 hours, which confirms that there was only one such storm event in the three springs analysed.

## 4. ACKNOWLEDGEMENTS

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## 5. REFERENCES

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## APPENDIX I

### Method of system calibration

I.1 Since there are two types of transducer in the shipborne wave recorder system, it is necessary to divide the calibration procedure into two sections. First the accelerometers are removed from the ship mountings and each is inserted into a rig which allows the transducer to be driven through a vertical circle of diameter 1 metre. The transducer is mounted in gimbals and maintains a vertical attitude during rotation. Two rotation rates are applied: 12 and 18 second periods which are derived from a crystal oscillator. The transducer is connected to the electronics unit in the usual way, and the calibration signal is displayed on the chart recorder. However, because a 1 metre 'heave' is small compared with the wave-heights usually experienced at sea, a precision amplifier (contained in the electronics) is switched into the circuit, converting the 1 metre into an apparent 10 metre signal. The output signal can then be read from the chart record and any corrections to instrument sensitivity made.

The pressure units cannot be easily subjected to a dynamic test since this requires the application of a sinusoidally-varying pressure. Therefore for routine re-calibration a static test is applied. Each pressure unit is fixed to the test rig and a series of discrete pressure levels is applied from a reservoir via a regulator valve. Each pressure level is set manually with the valve by reference to a precise pressure transducer contained within the calibrator unit. The output voltage of the transducer is monitored in the SBWR electronics unit and compared to the original laboratory calibration. Any changes in sensitivity are then compensated for by adjustment of the input amplifier gain.

Full re-calibrations are usually only possible when the ship comes into dock for its 3-yearly re-fit.

#### I.2 Monthly checks

All lightvessel crews are asked to drain water through the valve assemblies to ensure that no blockage prevents the water pressure being transmitted to the pressure sensors, and then to take a test record, on a monthly basis. The test record consists of a short length of pen-trace with all transducers turned off (electrically), followed by a few minutes recording with each transducer on its own. The record thus produced shows two heave records (one from each accelerometer) which should look broadly similar; and also the pressure traces, which may not agree so well, but when compared with other monthly test records should exhibit no systematic error. These tests are not direct checks on calibration accuracy but are often good indicators of a fault condition developing.





## APPENDIX II

### Definitions of wave parameters and method of analysis of wave chart data

II.1 The technique used to analyse the wave data was that proposed by TUCKER(1961) and DRAPER(1963), and reviewed by TANN(1976). A twelve-minute record is taken every three hours, and from this the following parameters may be derived.

II.1.1  $T_z$  - the mean zero up-crossing period. This is defined as the duration of the record divided by the number of zero up-crossings  $N_z$ . (A zero up-crossing is considered to occur when the trace crosses the mean line in an upward direction.)

II.1.2  $H_s$  - the significant wave height. This is defined as  $4\sigma$  where  $\sigma$  is the standard deviation of the record. (An estimate of  $\sigma$  is obtained from  $N_z$  and from the excursions of the two highest maxima, and of the two lowest minima, from the record mean.) For a narrow band random process this parameter approximates closely to the mean height of the highest one-third zero up-cross waves (see LONGUET-HIGGINS(1952)). Comparison between the two definitions is made by GODA(1970,1974). (A zero up-cross wave is defined as the portion of the wave record between two zero up-crossings, and its height is the vertical distance between the highest and lowest points on the wave.)

II.1.3  $H_{max}(3hr)$  - the most probable height of the highest zero up-cross wave in the recording interval of three hours. (This is derived from  $N_z$ , and the duration of the recording interval.) The parameter  $H_{max}(3hr)$ , which is the mode of the distribution, should not be confused with the expected height of the highest wave in three hours, which is the mean of the distribution. The mean of the distribution is typically 3 per cent higher than the mode (see TANN(1976)).

(A 'calm' record is one for which the sum of the height of the highest crest and of the lowest trough is less than 1 foot or 0.3 metre.)

### II.2 References

- DRAPER L 1963. Derivation of a 'design wave' from instrumental records of sea waves. Proceedings of the Institution of Civil Engineers 26, 291-304.
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## APPENDIX III

### Details of methods used for calculating design wave heights

III.1 By finding the long-term distribution of  $H_s$

III.1.1  $H_s$  is used as a measure of the "sea-state" (i.e. the intensity of wave activity), and it is sampled every 3 hours. It is assumed that a set of  $H_s$  data for one year, or an integral number of years, is representative of the wave climate.

For each value of  $H_s$ , the probability that this value will not be exceeded is calculated; this probability is then plotted against  $H_s$ . The axes are scaled according to a long-term distribution, so that data with a perfect fit would appear as a straight line on the diagram. This procedure is carried out using long-term distributions defined in the following ways

Weibull

$$\text{Prob}(H_s \leq h) = \begin{cases} 1 - \exp\left[-\left(\frac{h-A}{B}\right)^C\right], & \text{for } h > A \\ 0 & , \text{ for } h \leq A \end{cases}$$

where  $B$  and  $C$  are positive, and  $A$  represents a lower bound on  $h$ .

Fisher-Tippett I (first asymptote)

$$\text{Prob}(H_s \leq h) = \exp[-\exp(-ah+b)].$$

Fisher-Tippett III (third asymptote)

$$\text{Prob}(H_s \leq h) = \begin{cases} \exp\left[-\left(\frac{A-h}{B}\right)^C\right], & \text{for } h \leq A \\ 1 & , \text{ for } h > A \end{cases}$$

where  $B$  and  $C$  are positive, and  $A$  represents an upper bound on  $h$ . (See FISHER AND TIPPETT(1928) and GUMBEL (1958) for the derivations of these distributions.)

For each long-term distribution the best-fit straight line is drawn; this line is then extrapolated to the desired probability (see section III.1.2) and the corresponding value of  $H_s$  is read off as the "design sea-state".

III.1.2 To calculate the "sea-state" which will be exceeded only once in  $N$  years, a storm duration of  $D$  hours needs to be assumed. The probability that a randomly chosen time will be within this storm is then

$$\frac{D}{24 \times 365.25 \times N}$$

IOS uses  $D = 3$  hours (this choice is discussed in section III.1.5) which gives

$$\begin{aligned} \text{Probability} &= \frac{3.422 \times 10^{-4}}{N} \\ &= 6.845 \times 10^{-6} \quad \text{for } N = 50 \text{ years.} \end{aligned}$$

III.1.3 The value of  $T_z$  for the "design sea-state" is required before the highest wave in the storm can be calculated. This is derived from the bivariate distribution of  $H_s$  and  $T_z$  (figure 3.5.1.5). A line is drawn across this at the "design sea-state" value of  $H_s$  and the most likely value of  $T_z$  (the modal value) is then estimated using extrapolations of the probability contours.

III.1.4 The most probable value of the highest zero-up-cross wave in the storm is then derived by assuming that the heights of such waves follow a Rayleigh distribution whose probability density function is

$$\text{prob}(h) = \frac{2h}{(H_{rms})^2} \exp\left[-\left(\frac{h}{H_{rms}}\right)^2\right]$$

where  $H_{rms} \approx \frac{H_s}{\sqrt{2}}$ .

Exact theory is not available for zero-up-cross wave heights, but this distribution has been found to be an adequate fit to measured data. If there are  $n$  waves in the recording interval (3hr), then the probability that the highest wave,  $H$ , in three hours is less than  $h$  is

$$\text{Prob}(H \leq h) = \left\{ 1 - \exp\left[-\left(\frac{h}{H_{rms}}\right)^2\right] \right\}^n$$

with a corresponding probability density function

$$\frac{2n}{(H_{rms})^2} h \exp\left[-\left(\frac{h}{H_{rms}}\right)^2\right] \left\{ 1 - \exp\left[-\left(\frac{h}{H_{rms}}\right)^2\right] \right\}^{n-1}$$

The most probable value (the mode) of this probability density function is usually used and is given by

$$H_{max}(3hr) = H_{rms} \sqrt{\Psi}$$

where  $\Psi$  is a function of  $T_z$  which may be found using either figure 7 or equation 6.1-2 in TANN(1976).

III.1.5 In choosing the value of storm duration  $D$ , it should be noted that the effect of increasing  $D$  is to decrease the value of  $H_s$  for a given return period  $N$ . However, it also increases the ratio of  $H_{max}(3hr)$  to  $H_s$ . It is found that in practice these effects roughly cancel and typically the value of  $H_{max}(3hr)$  changes by only 3 per cent for a change of  $D$  from 3 to 15 hours. The choice of  $D$  is therefore not critical.

Many details of the above procedures may be found in TANN(1976).

III.2 By a wave-by-wave method

III.2.1 BATTJES(1970) shows that the probability that a randomly chosen wave will have a height H greater than h is

$$\text{Prob}(H>h) = \frac{\int_0^\infty \int_0^\infty R(h, H_s) T_z^{-1} p(T_z, H_s) dH_s dT_z}{\int_0^\infty \int_0^\infty T_z^{-1} p(T_z, H_s) dH_s dT_z} \dots\dots\dots(1)$$

where  $R(h, H_s)$  is the Rayleigh cumulative probability function and  $p(T_z, H_s)$  is the joint probability density function of  $H_s$  and  $T_z$ .

III.2.2 TANN makes the following suggestion in an unpublished manuscript. In order that values of  $H_s$  higher than those actually measured may be represented in the calculation of this probability, the values of  $H_s$  are assumed to have a long-term cumulative probability function  $F(H_s)$ , and a probability density function  $f(H_s) = F'(H_s)$ .

For each value of  $H_s$  throughout the long-term distribution, an average value of  $T_z^{-1}$  is used (denoted by  $\overline{T_z^{-1}(H_s)}$ ). It is defined as

$$\overline{T_z^{-1}(H_s)} = \int_0^\infty T_z^{-1} \frac{p(T_z, H_s)}{P(H_s)} dT_z$$

where  $P(H_s) = f(H_s)$ .

Therefore

$$\int_0^\infty T_z^{-1} p(T_z, H_s) dT_z = \overline{T_z^{-1}(H_s)}$$

which, when substituted into equation (1), allows the probability of exceedance to be written

$$\text{Prob}(H>h) = \frac{\int_0^\infty R(h, H_s) \overline{T_z^{-1}(H_s)} f(H_s) dH_s}{\int_0^\infty \overline{T_z^{-1}(H_s)} f(H_s) dH_s}$$

The value of  $\overline{T_z^{-1}(H_s)}$  used with each value of  $H_s$  is chosen to satisfy the condition of constant wave "steepness", where the relationship between "steepness"(1:s), water depth(d),  $H_s$  and  $T_z$  is

$$T_z = \sqrt{\frac{2\pi s H_s}{g} \coth\left(\frac{2\pi d}{s H_s}\right)}$$

The value for the steepness used in this report is given in section 3.3.4.

The long-term distribution used in the computation for this report is the Fisher-Tippett III extreme-value distribution, whose probability density function is

$$f(H_s) = \frac{C}{A-H_s} \left(\frac{A-H_s}{B}\right)^C \exp\left[-\left(\frac{A-H_s}{B}\right)^C\right]$$

The constants A,B,C are determined graphically as described in section III.1.1, and their values as used in this report are given in section 3.3.3.

III.2.3 Thus the probability of a wave exceeding each particular wave height may be found, and this probability may be converted into a return period of N years using the formula

$$N = \frac{1}{365.25 \times 24 \times 3600 \times T_{ave}^{-1} \times Prob}$$

where  $T_{ave}^{-1} = \frac{1}{\text{average period}}$

The value of the average wave period is contained in section 3.3.4. Since  $T_{ave}^{-1}$  is a non-analytic function of Prob, the simplest way of solving the problem is to calculate Prob for various values of h, calculate N for each of these values of Prob, and then interpolate to find the height h corresponding to the required value of N (in this case 50 years).

Whereas the method described in section III.1 assumes that the highest wave in a 50-year period will come from the most stormy 3-hour period in 50 years, the individual wave method takes into account the probability that storms other than the highest may provide the wave with a 50-year return period. Consequently the height of a 50-year wave as estimated by this method is likely to be greater than that estimated from the method of using a long-term distribution of Hs.

### III.3 References

- BATTJES J A 1970. Long-term wave height distribution at seven stations around the British Isles. National Institute of Oceanography, Internal Report No A44.
- FISHER R A AND TIPPETT L H C 1928. Limiting forms of frequency distribution of the largest or smallest member of a sample. Proceedings of the Cambridge Philosophical Society 24, 180-190.
- GUMBEL E J 1958. Statistics of Extremes. New York: Columbia University Press. 371 pp.
- TANN H M 1976. The estimation of wave parameters for the design of offshore structures. Institute of Oceanographic Sciences, Report No 23.

#### APPENDIX IV

##### The correction factor applied to wave heights

Two corrections need to be applied to the wave height data:

- (i) to compensate for the frequency response of the electronics; and
- (ii) to compensate for the hydrodynamic attenuation of the pressure fluctuations.

These are combined into a single correction factor C, which is dependent upon Tz, and upon the depth of the pressure transducer below the mean water level. The correction factor is

$$C = 0.83 \left\{ 1 + \frac{1}{77.44\omega^2} \right\}^{\frac{3}{2}} \exp \left[ \frac{2.5\omega^2 d}{g} \right]$$

where d is the depth of the pressure transducer,

$$\text{and } \omega = \frac{2\pi}{Tz}.$$

The value of Hs calculated by the method outlined in Appendix II is multiplied by C to obtain the corrected value of Hs.

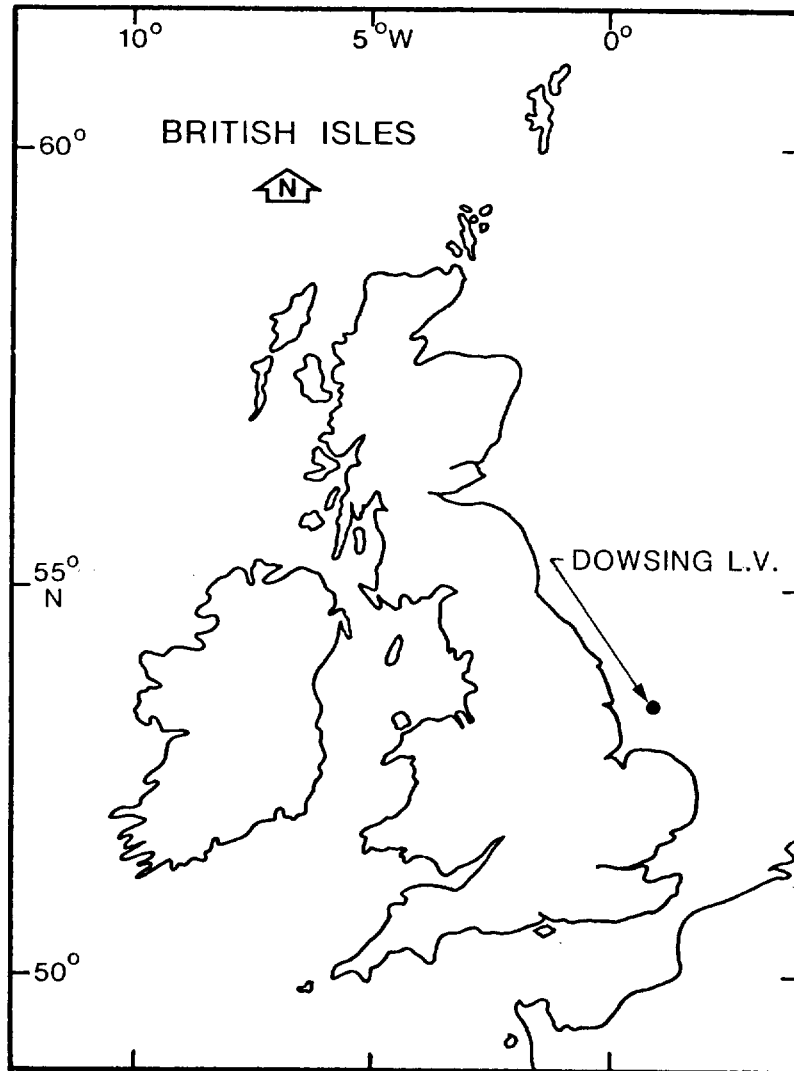
A table is given below showing the values of C for various values of Tz (and for each of the two values of d for the lightvessels on the Dowsing station during the period covered by this report:-

d = 1.46 metres for the 1970/71 data,  
and d = 0.88 metres for the 1975/79 data).

Tz(sec)	C	
	d=0.88m	d=1.46m
2.5	3.455	8.774
3.0	2.240	4.278
3.5	1.726	2.777
4.0	1.458	2.099
4.5	1.301	1.734
5.0	1.199	1.514
6.0	1.081	1.271
7.0	1.019	1.148
8.0	0.984	1.078
10.0	0.952	1.009
12.0	0.946	0.985
15.0	0.961	0.986
20.0	1.021	1.036



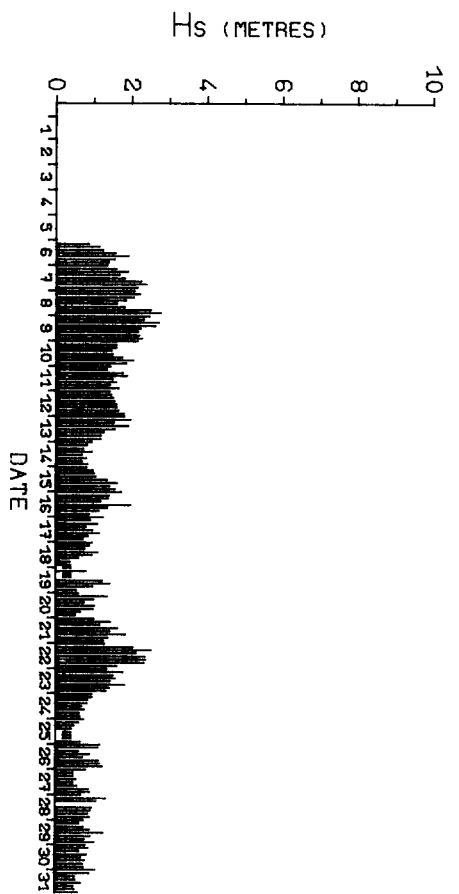




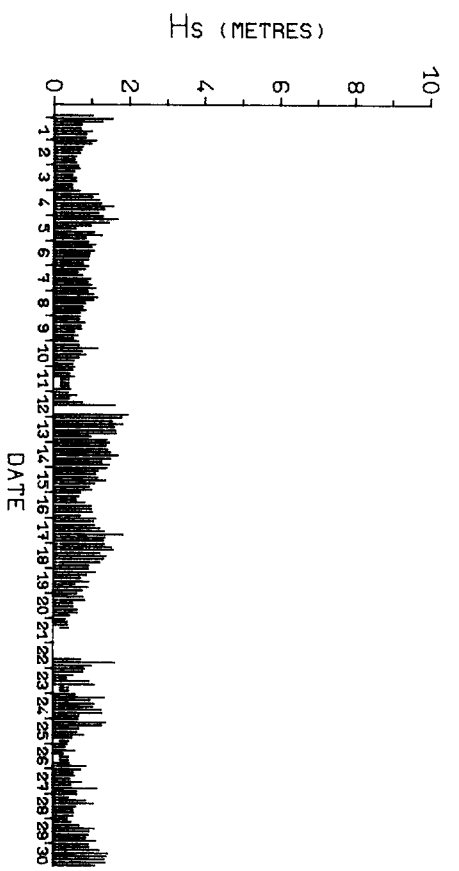
LOCATION MAP OF DOWSING WAVE RECORDER

FIG 1.1

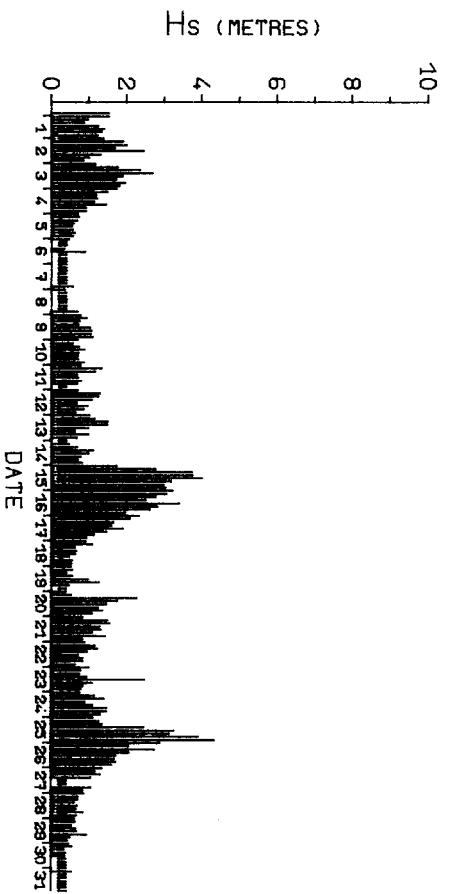
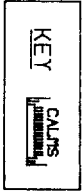




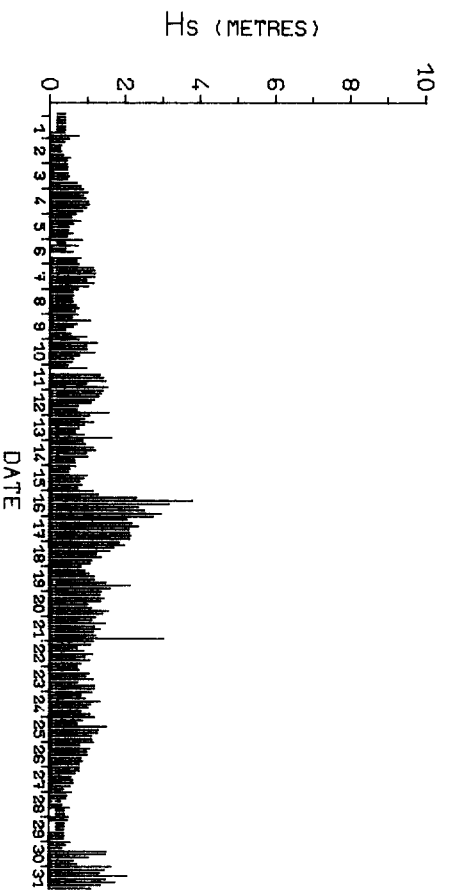
MAY 1970



JUN 1970



JUL 1970

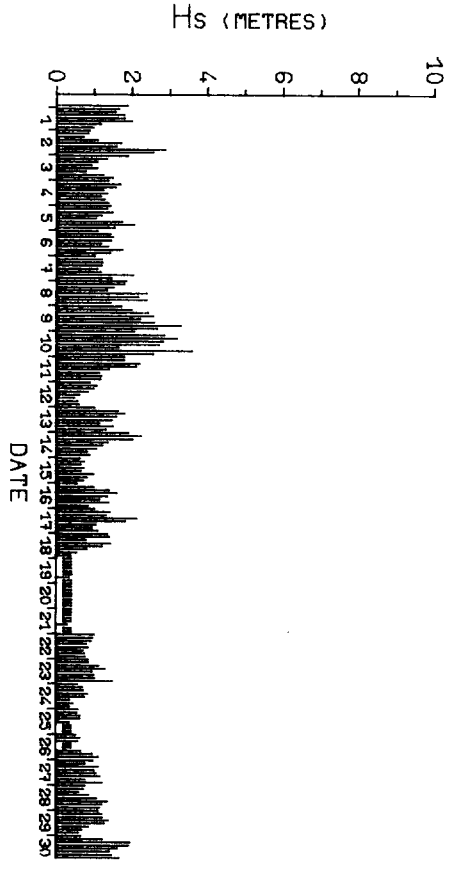


AUG 1970

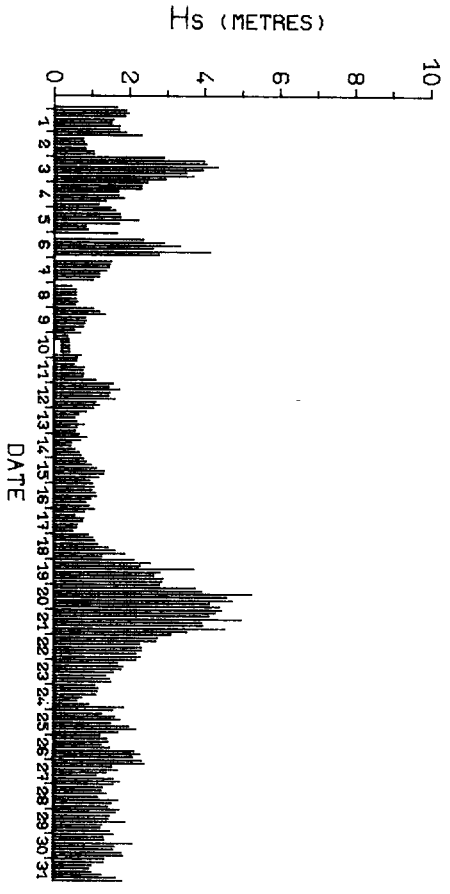
TIME SERIES OF HS

DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

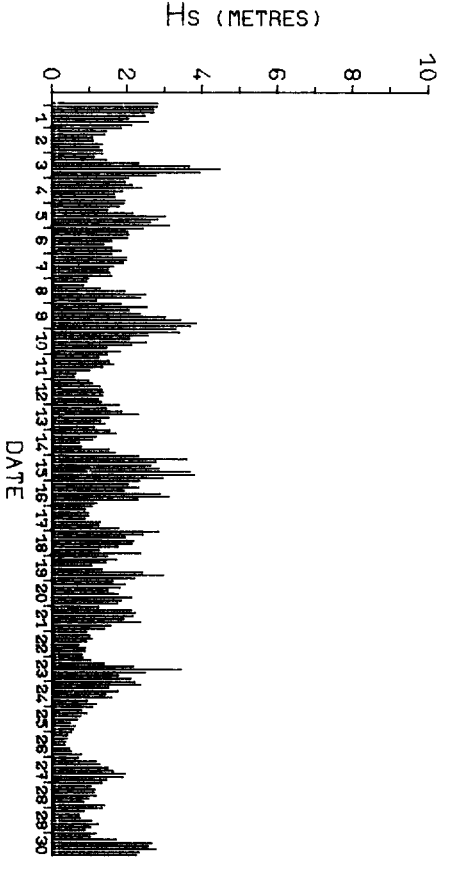
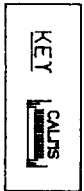
FIG 1.4(A)



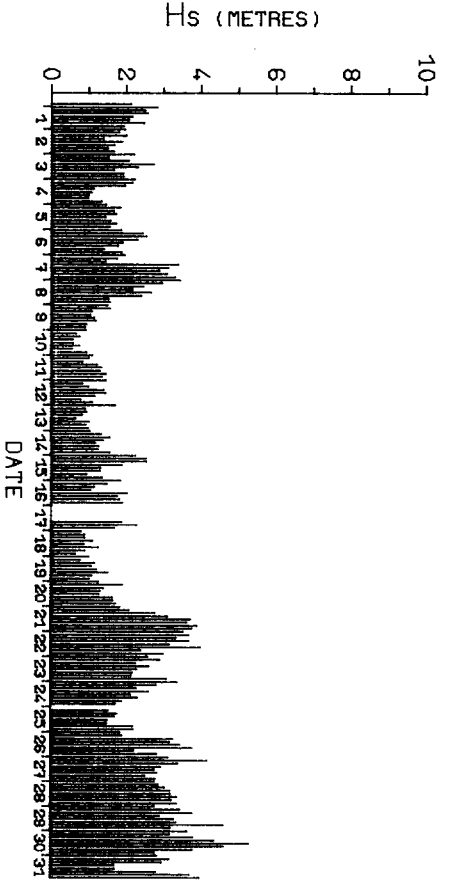
SEP 1970



OCT 1970



NOV 1970

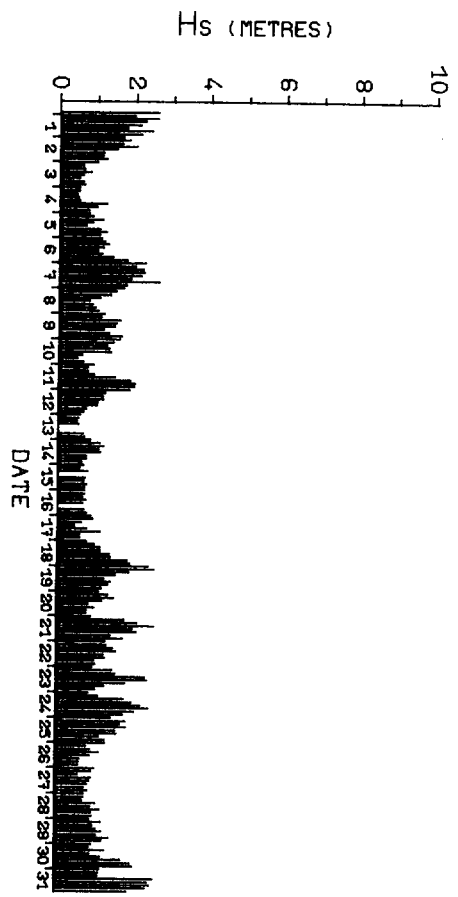


DEC 1970

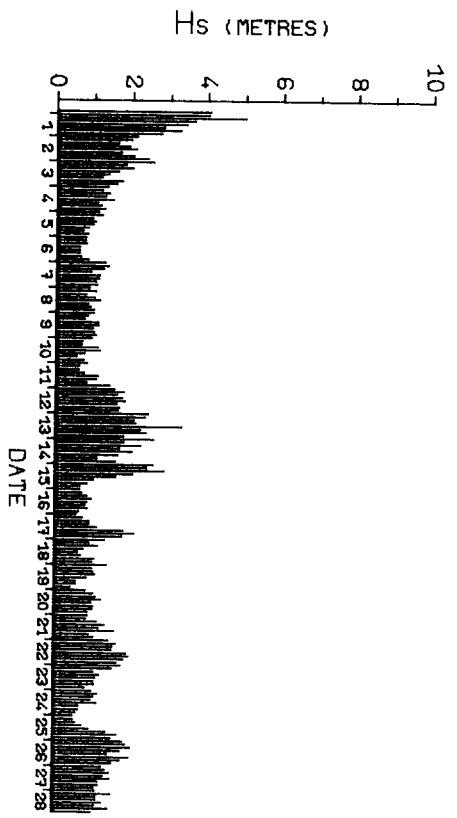
TIME SERIES OF HS

DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

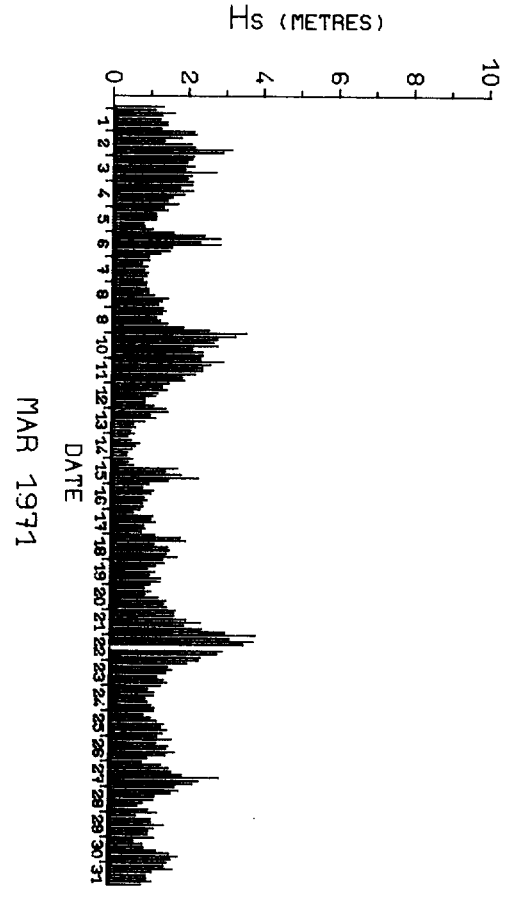
FIG 1.4(B)



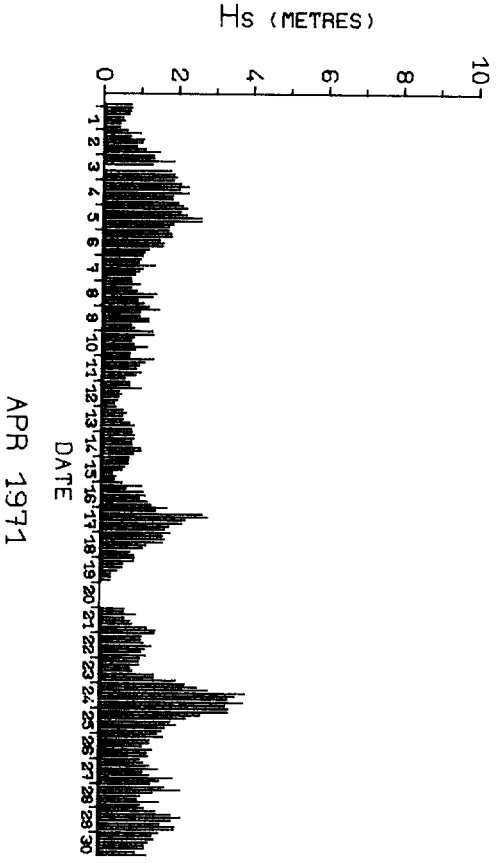
JAN 1971



FEB 1971



MAR 1971

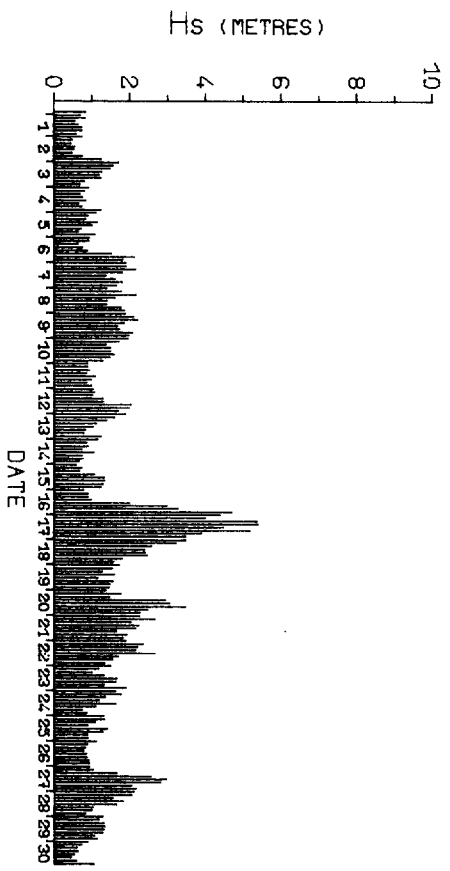


APR 1971

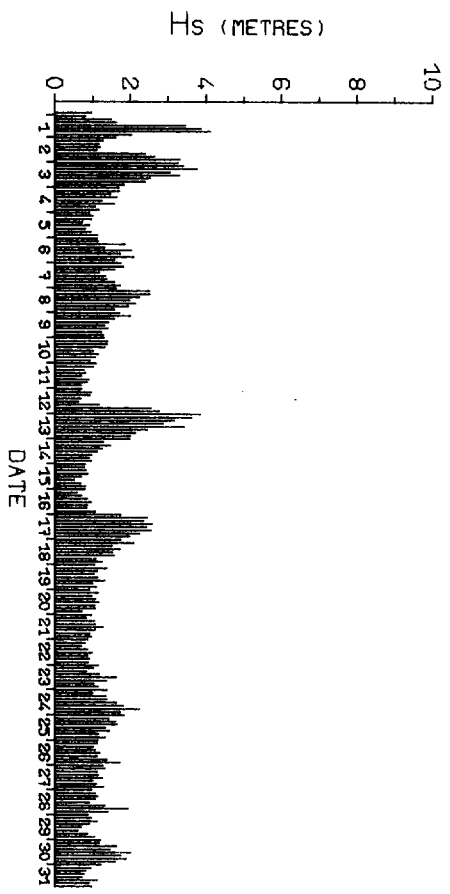
TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 1.4(C)

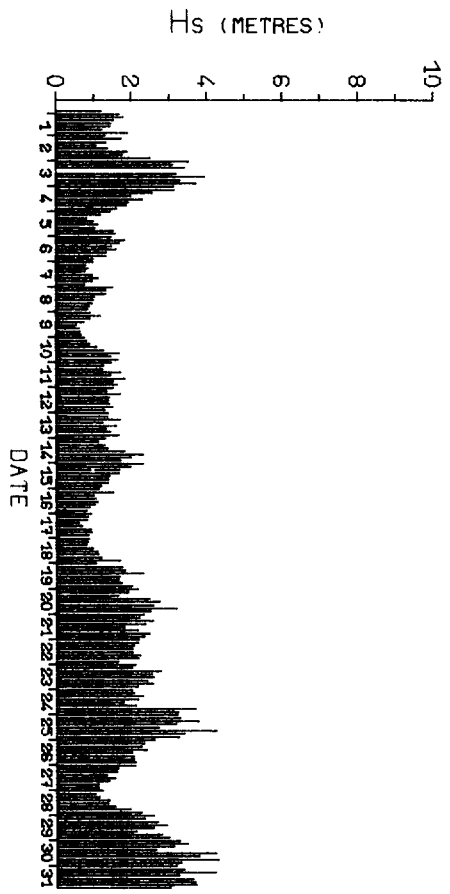


NOV 1975

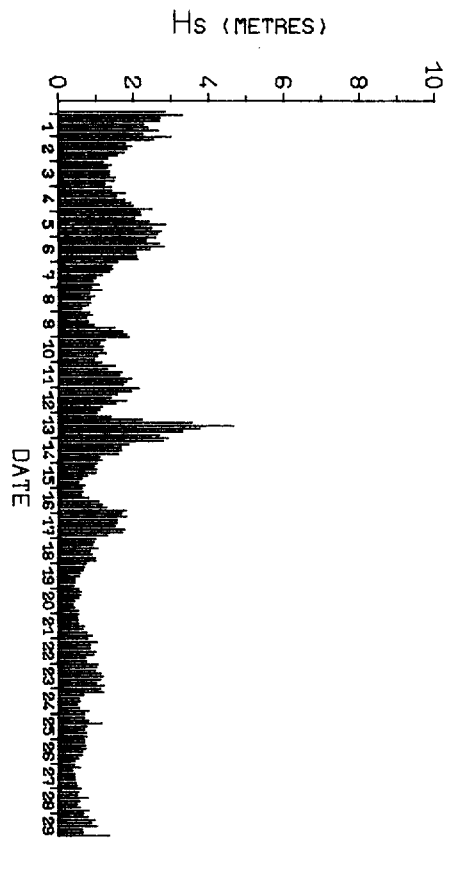


DEC 1975

KEY  
CALTS



JAN 1976

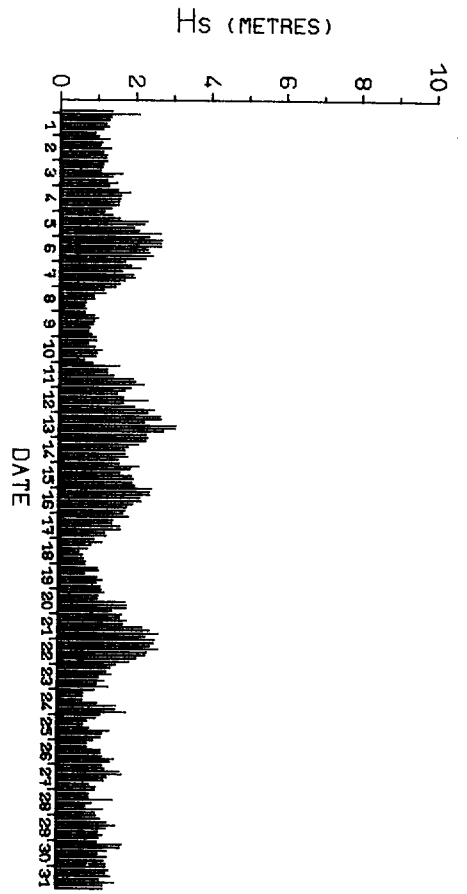


FEB 1976

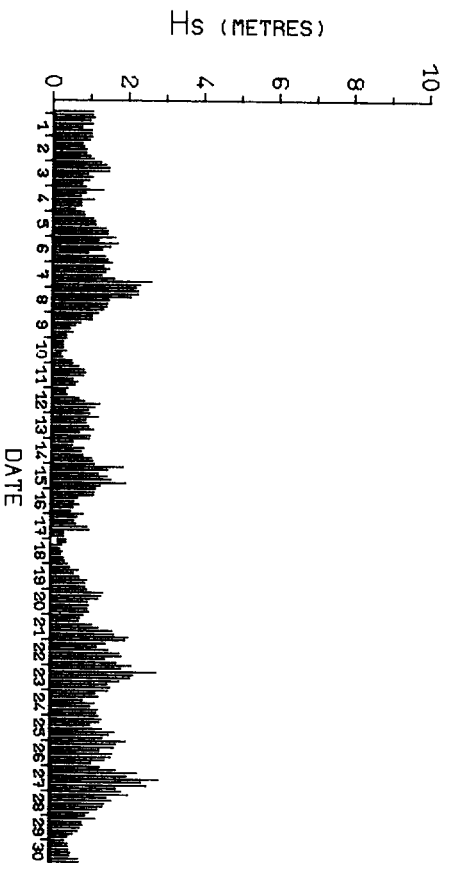
TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

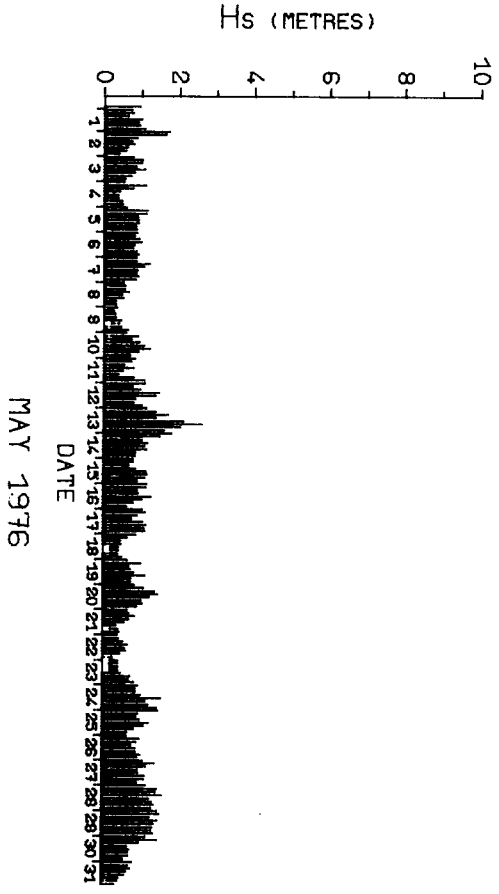
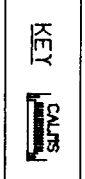
FIG 1.4(D)



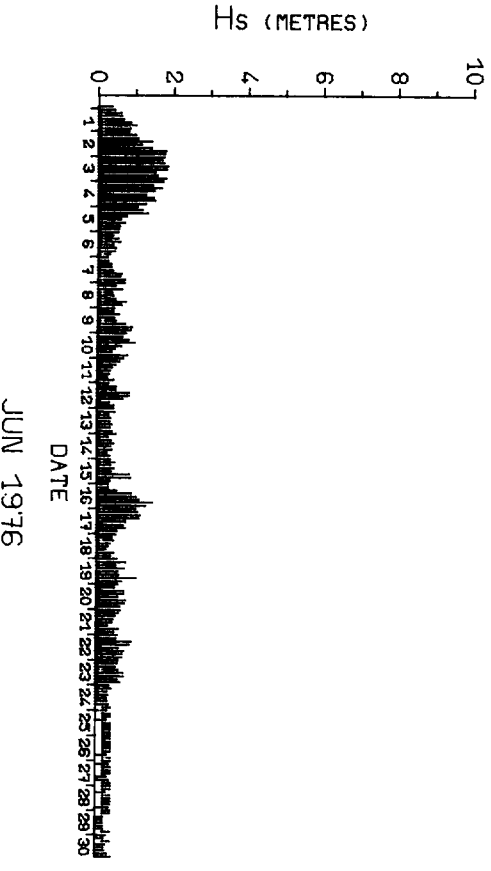
MAR 1976



APR 1976



MAY 1976

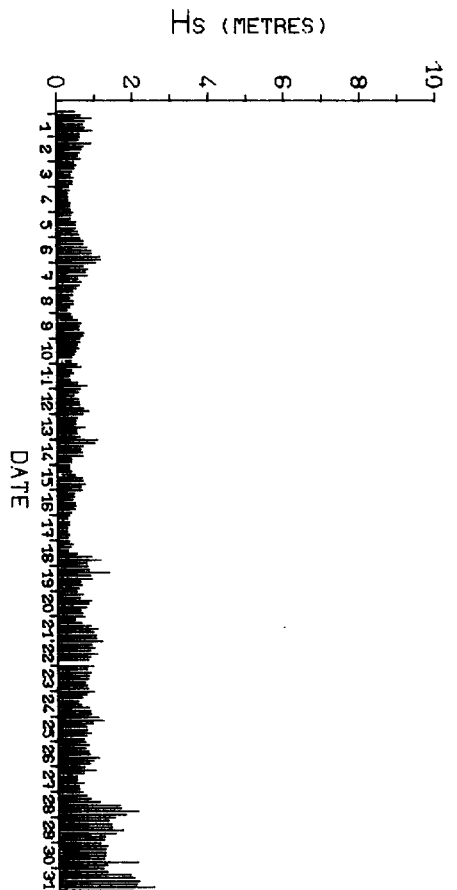


JUN 1976

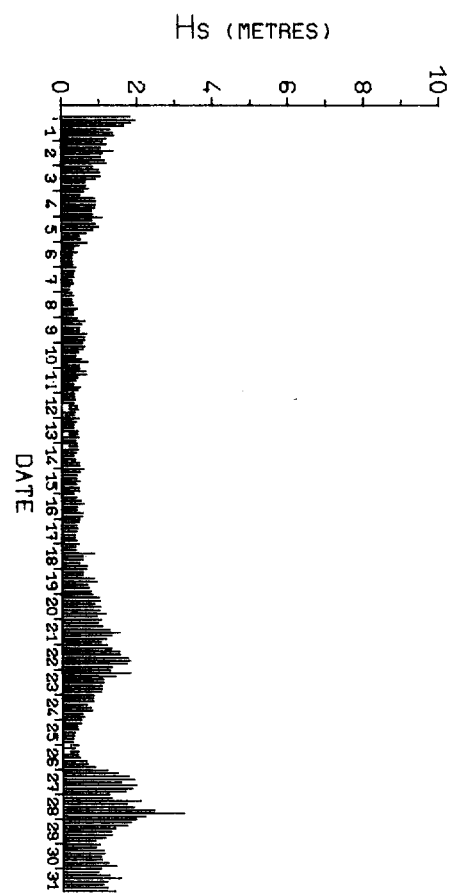
TIME SERIES OF HS

DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 1.4(E)

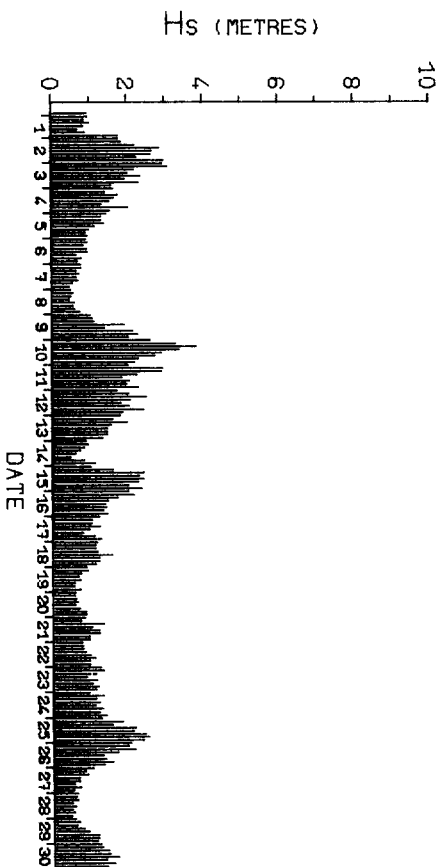


JUL 1976

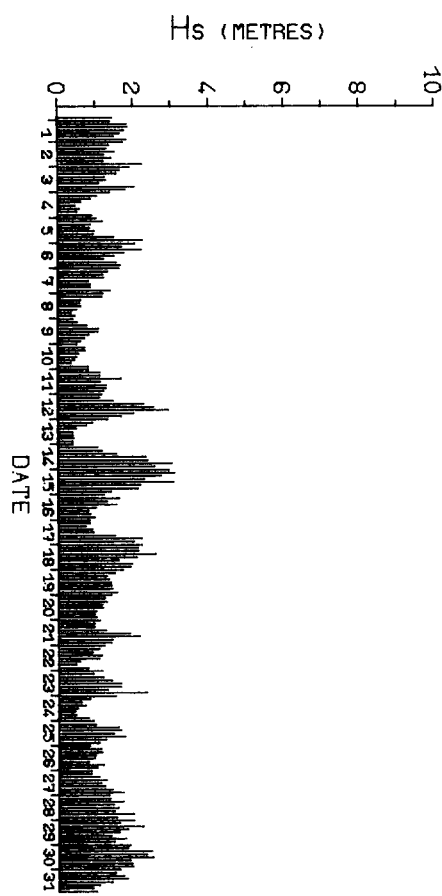


AUG 1976

KEY CALPS  

SEP 1976



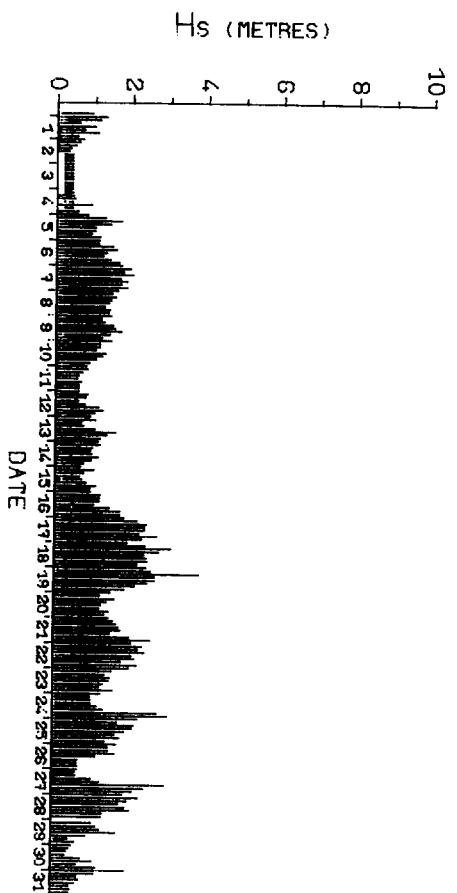
OCT 1976

TIME SERIES OF HS

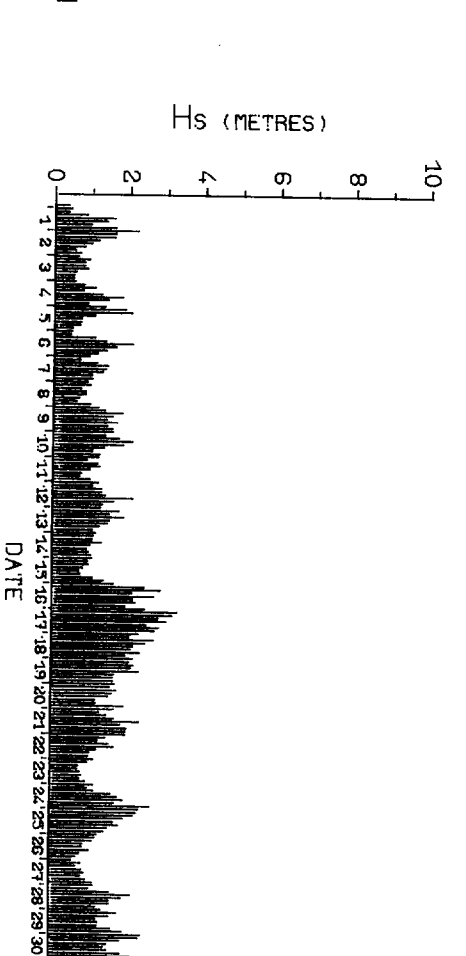
DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 1.4(F)



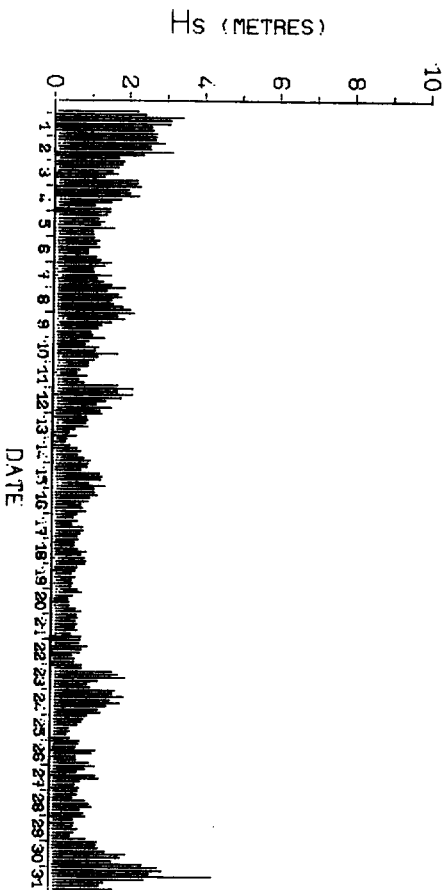


AUG 1977

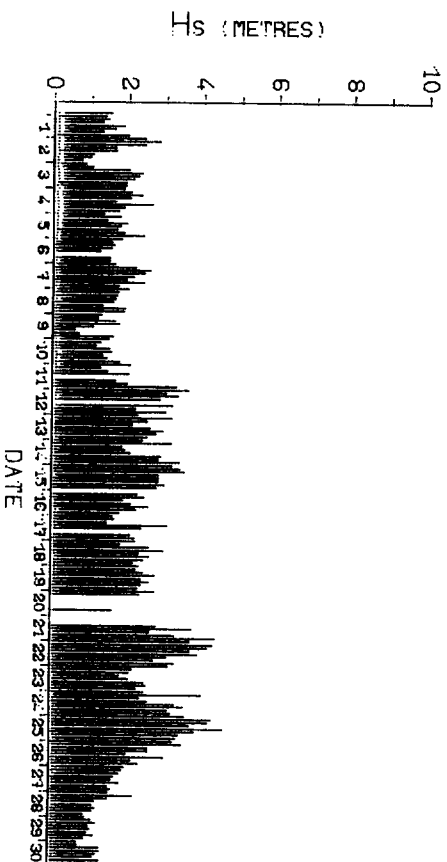


SEP 1977

KEY CALMS



OCT 1977

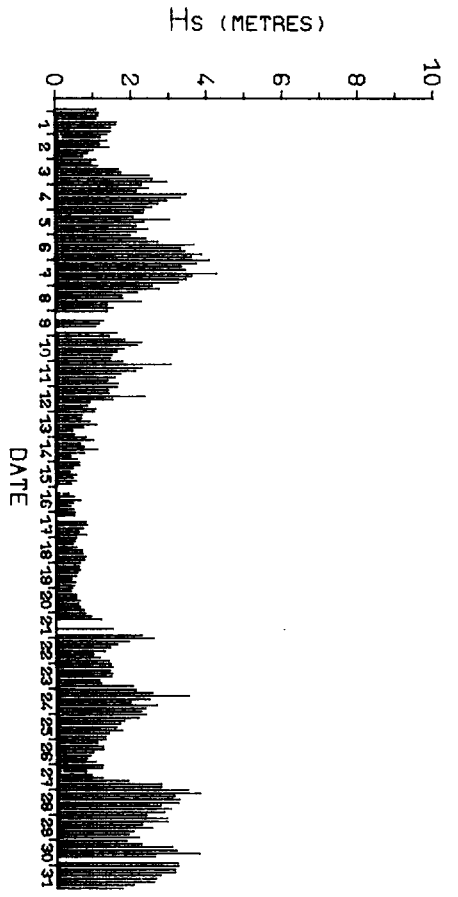


NOV 1977

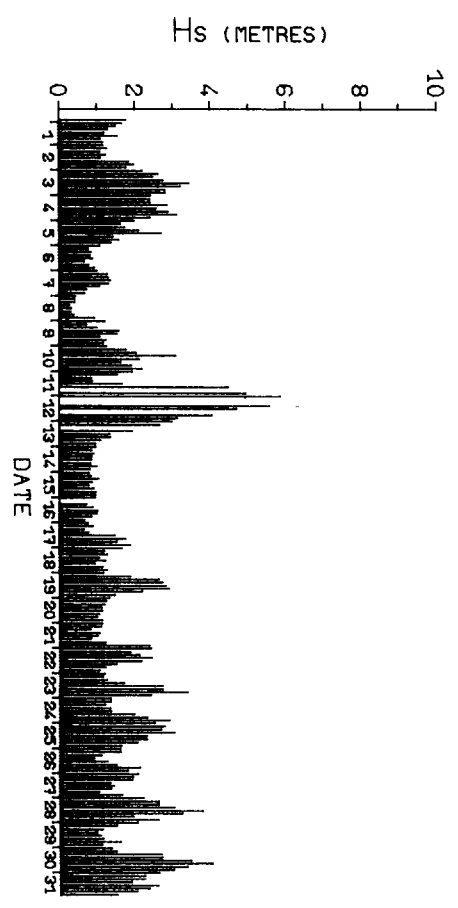
TIME SERIES OF HS

DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 1.4(G)

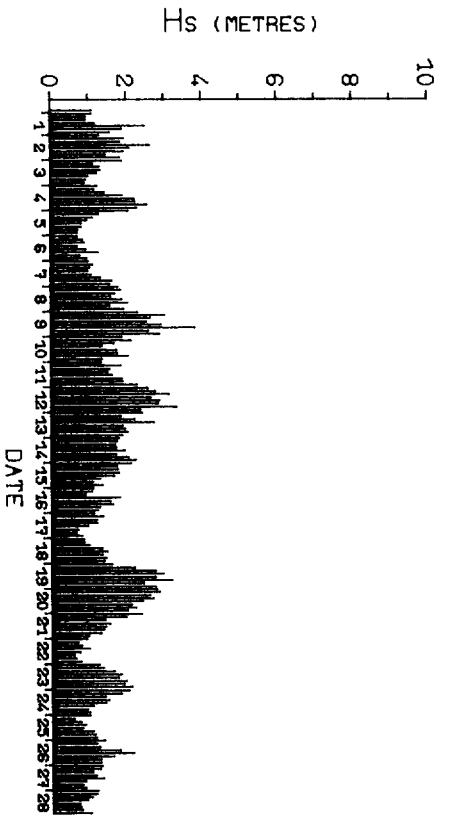


DEC 1977

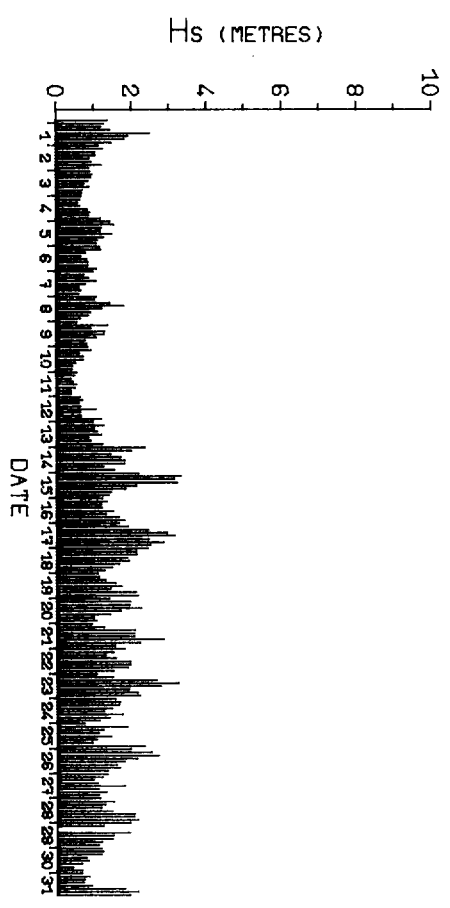


JAN 1978

KEY CALMS



FEB 1978

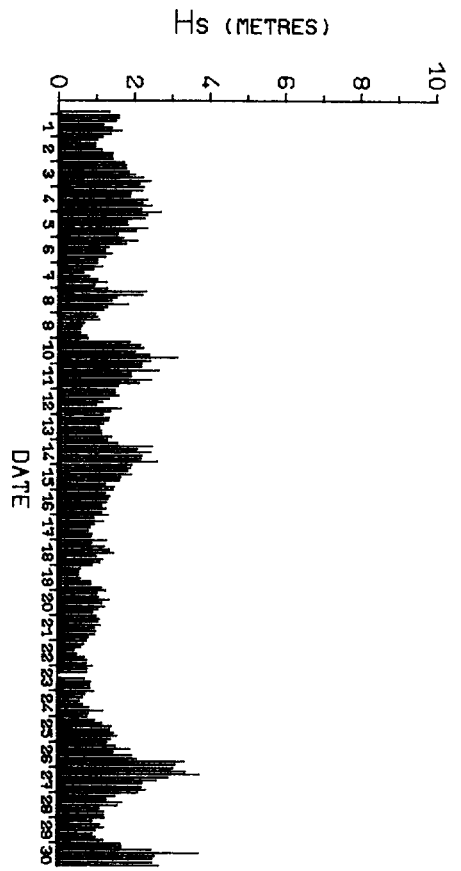


MAR 1978

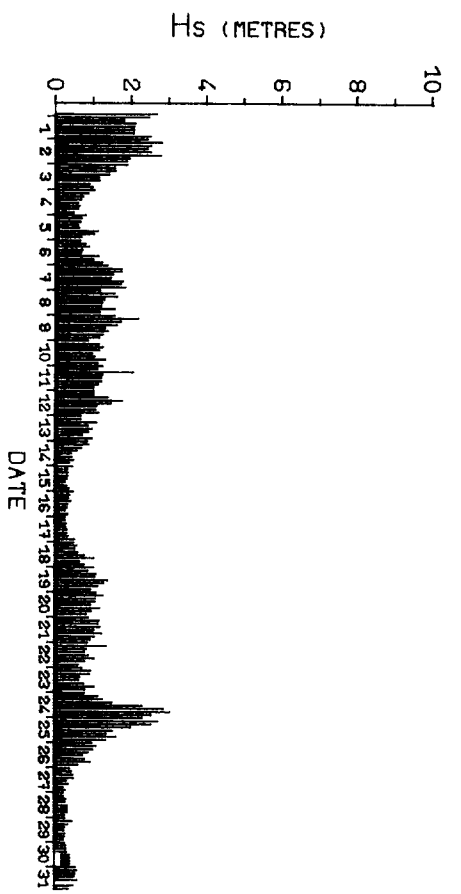
TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/765-10/76, 8/77-7/79

FIG 1.4(H)

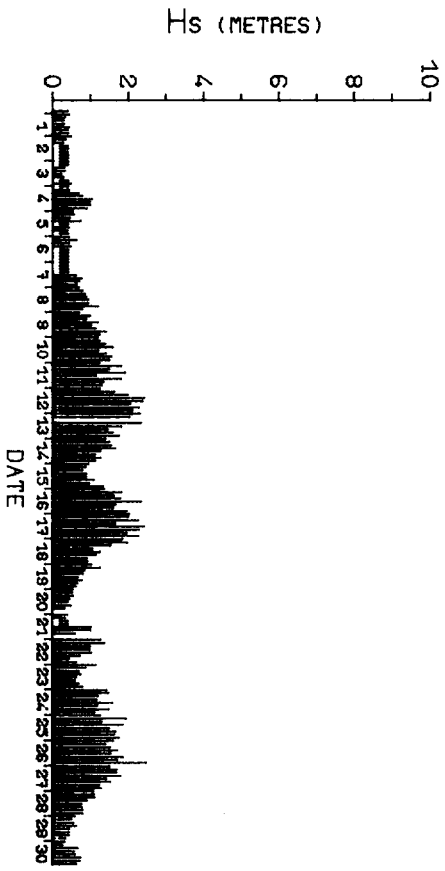


APR 1978

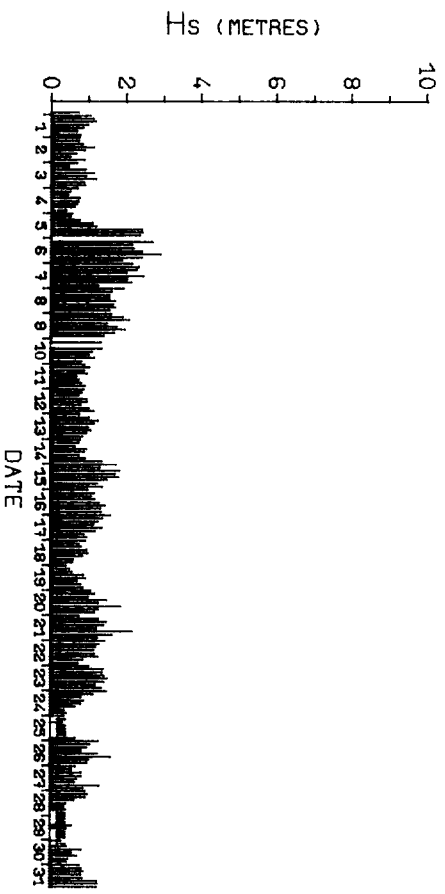


MAY 1978

KEY  
CALMS



JUN 1978

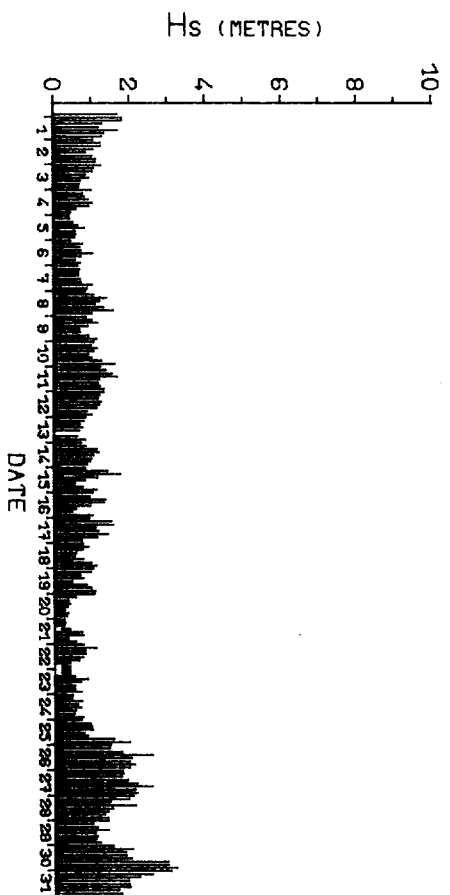


JUL 1978

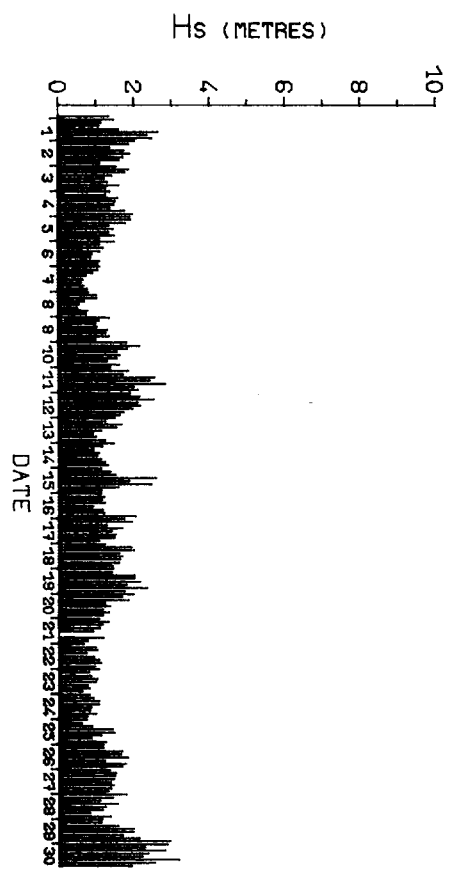
TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 1.4(I)

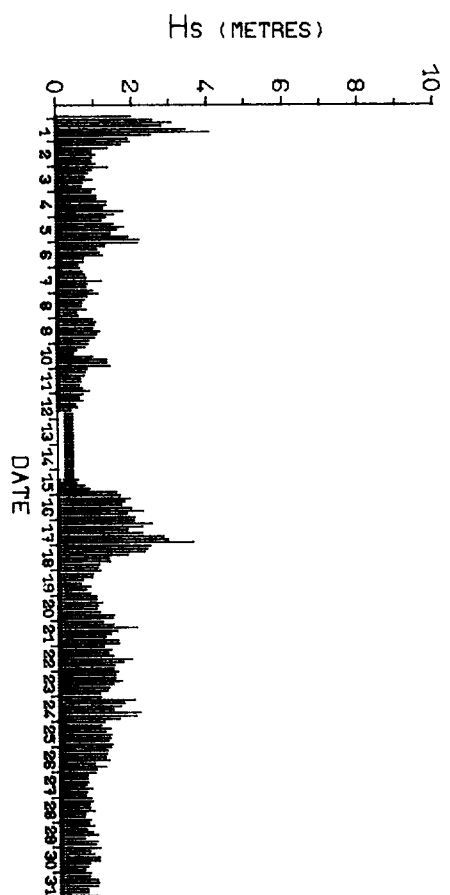


AUG 1978



SEP 1978

KEY CALS



OCT 1978

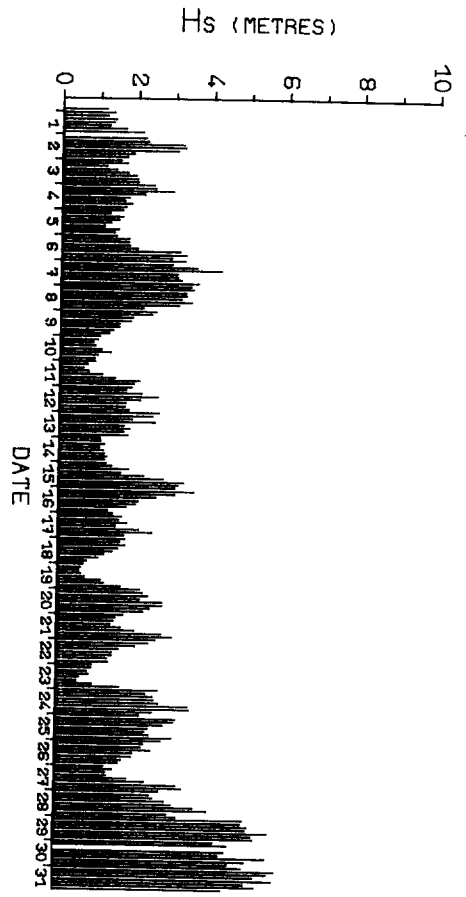


NOV 1978

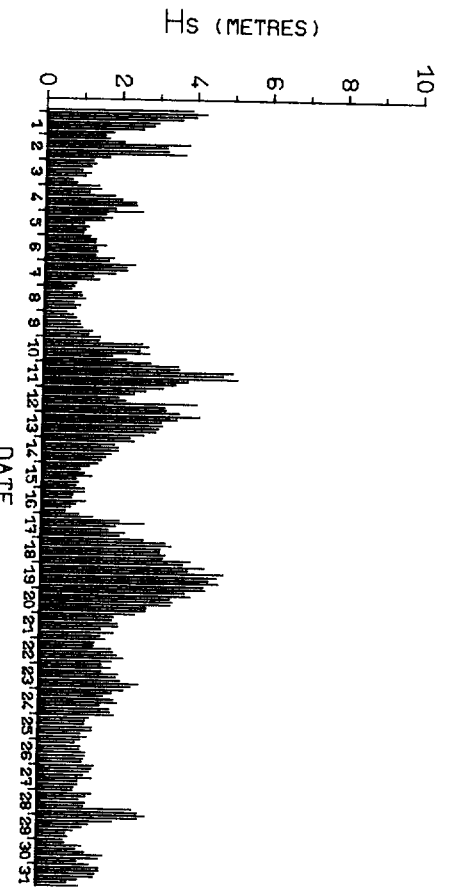
TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/765-10/76, 8/77-7/79

FIG 1.4(J)

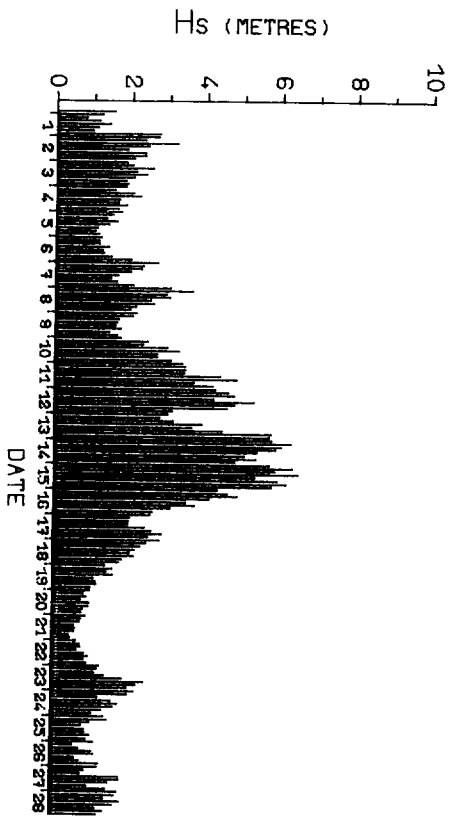


DEC 1978

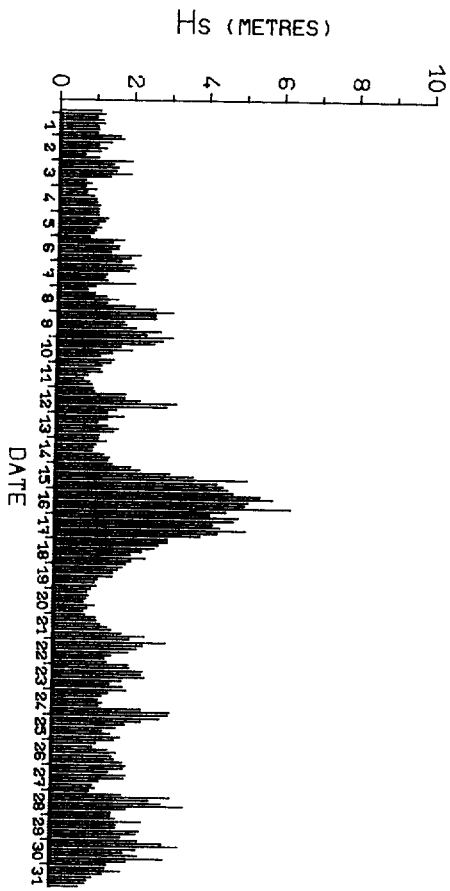


JAN 1979

KEY CALS



FEB 1979

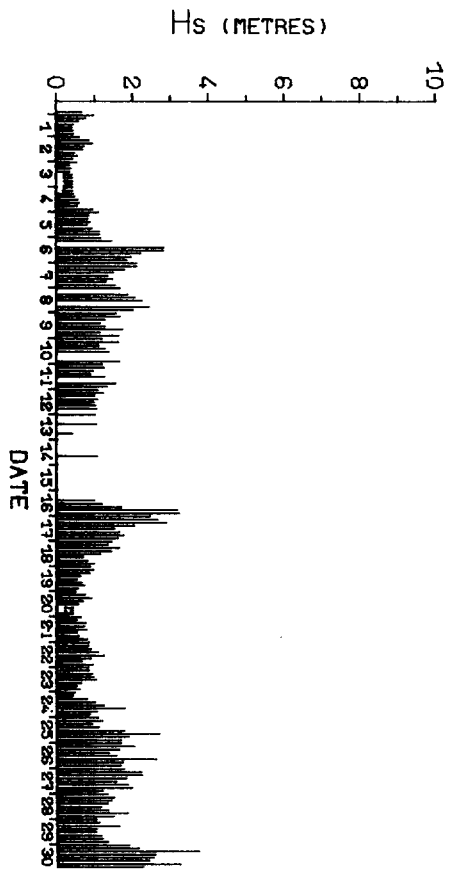


MAR 1979

TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/765-10/76, 8/77-7/79

FIG 1.4(K)

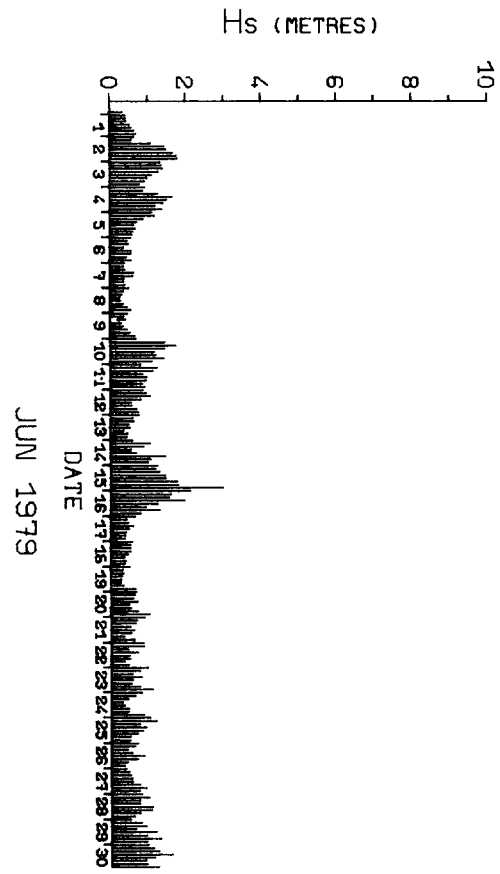


APR 1979

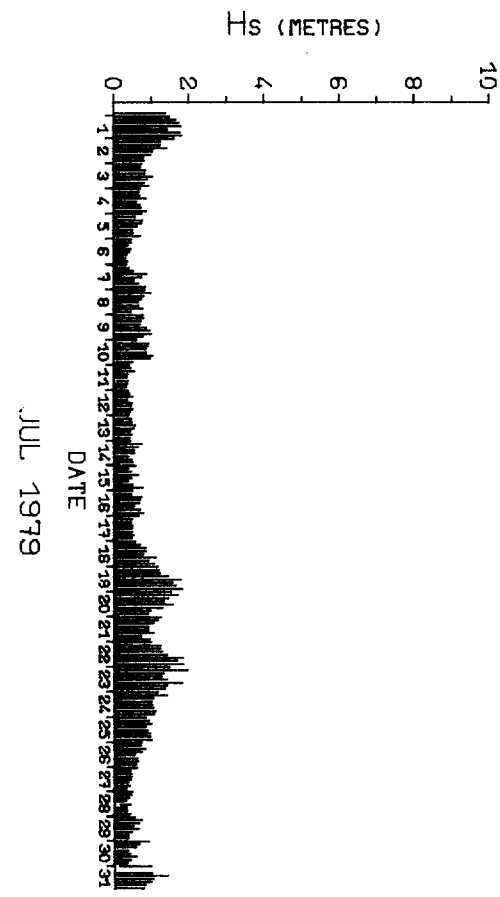


MAY 1979

KEY CALTS



JUN 1979



JUL 1979

TIME SERIES OF HS

DOWSING LV 5/70-4/71, 11/765-10/76, 8/77-7/79

FIG 1.4(L)

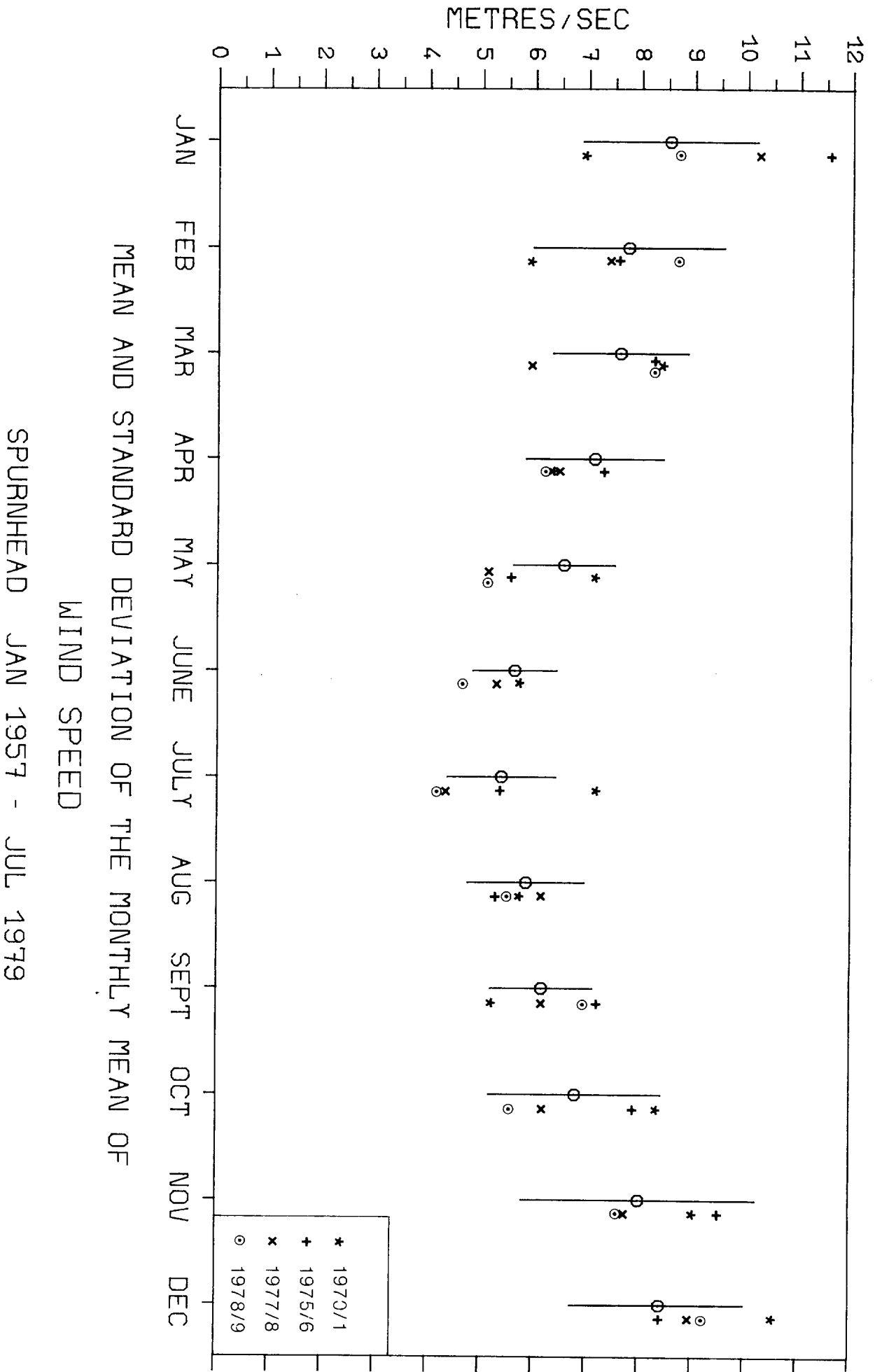
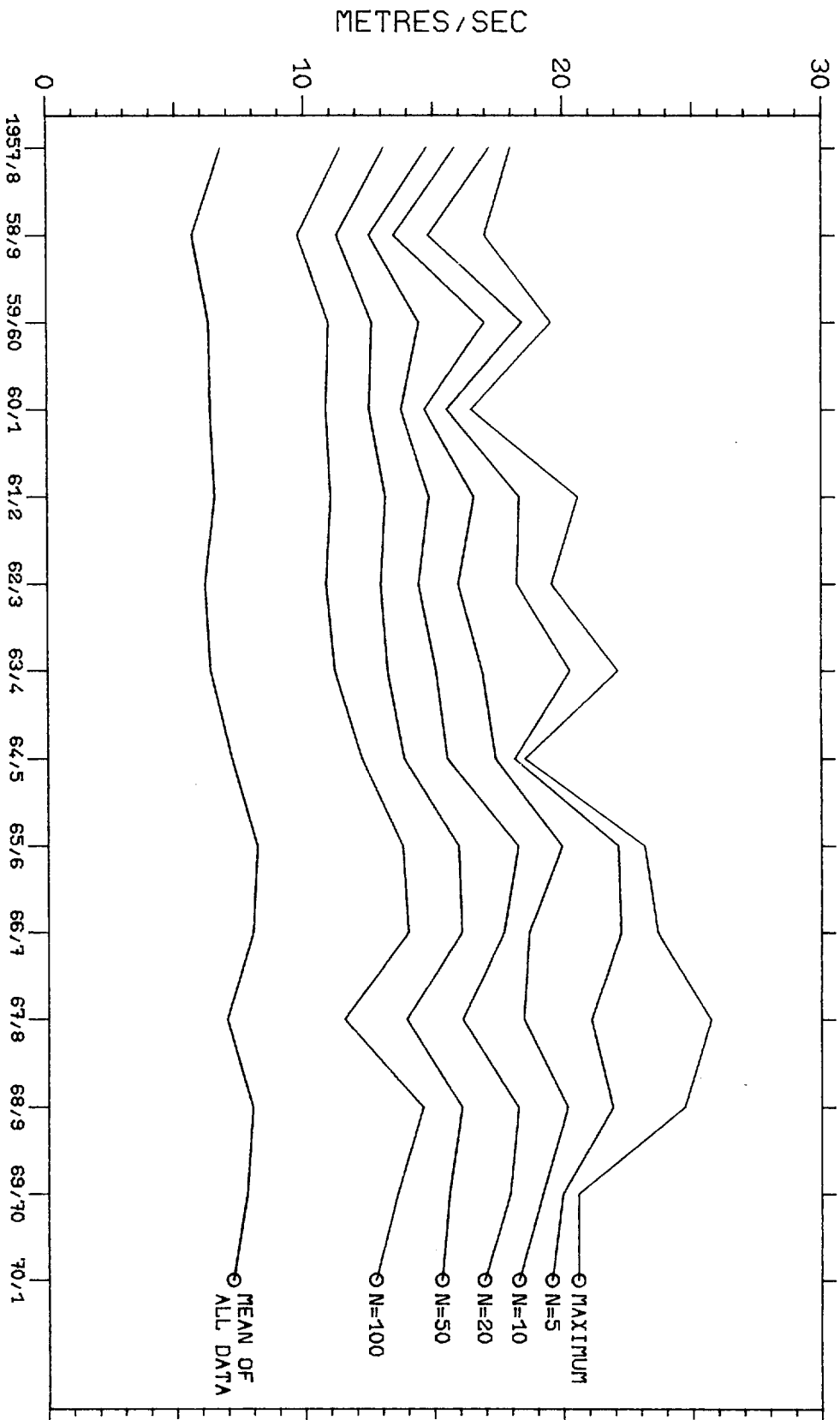


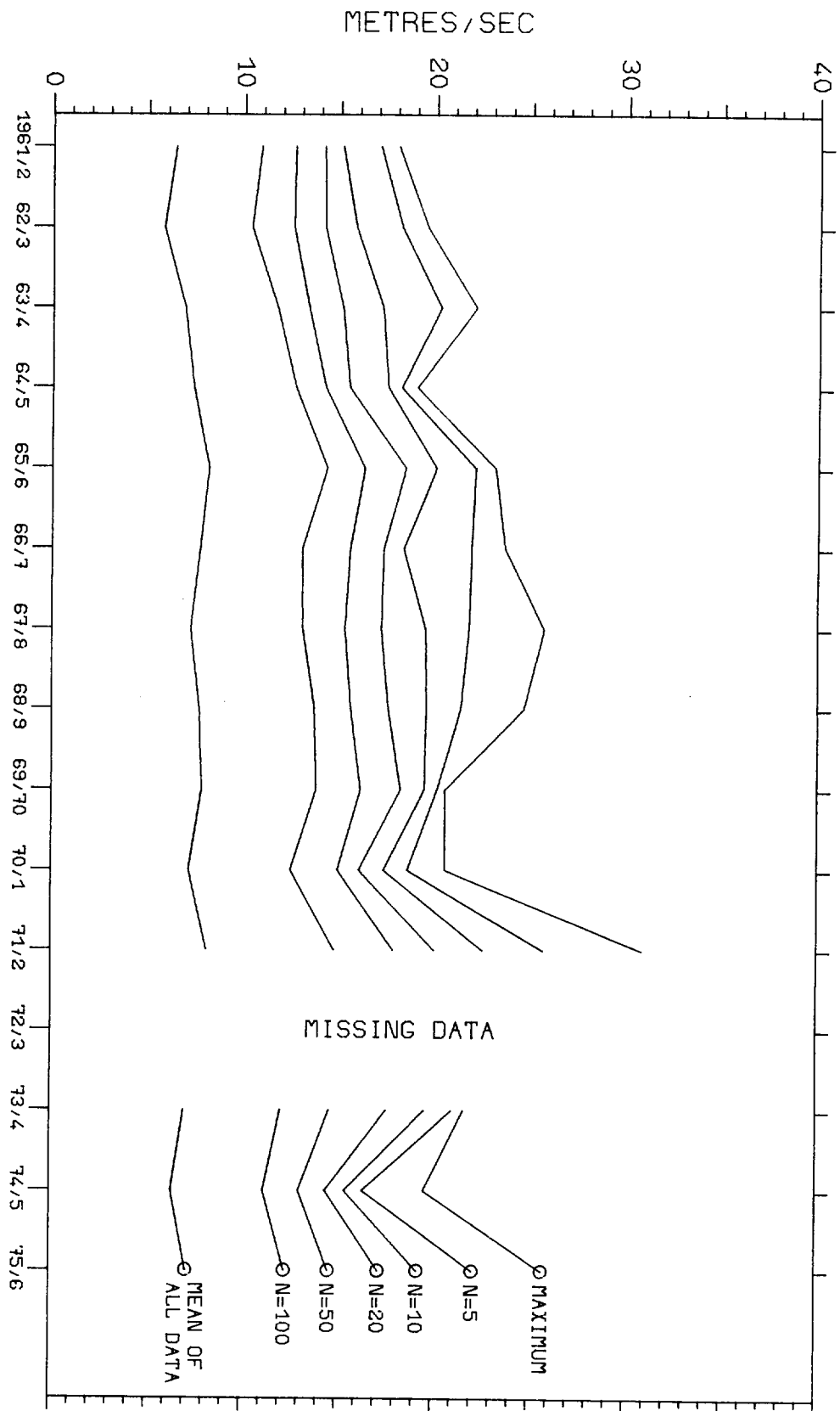
FIG 2.1



MEAN OF N LARGEST VALUES OF WIND SPEED  
SPURNHEAD MAY 1957 - APR 1971

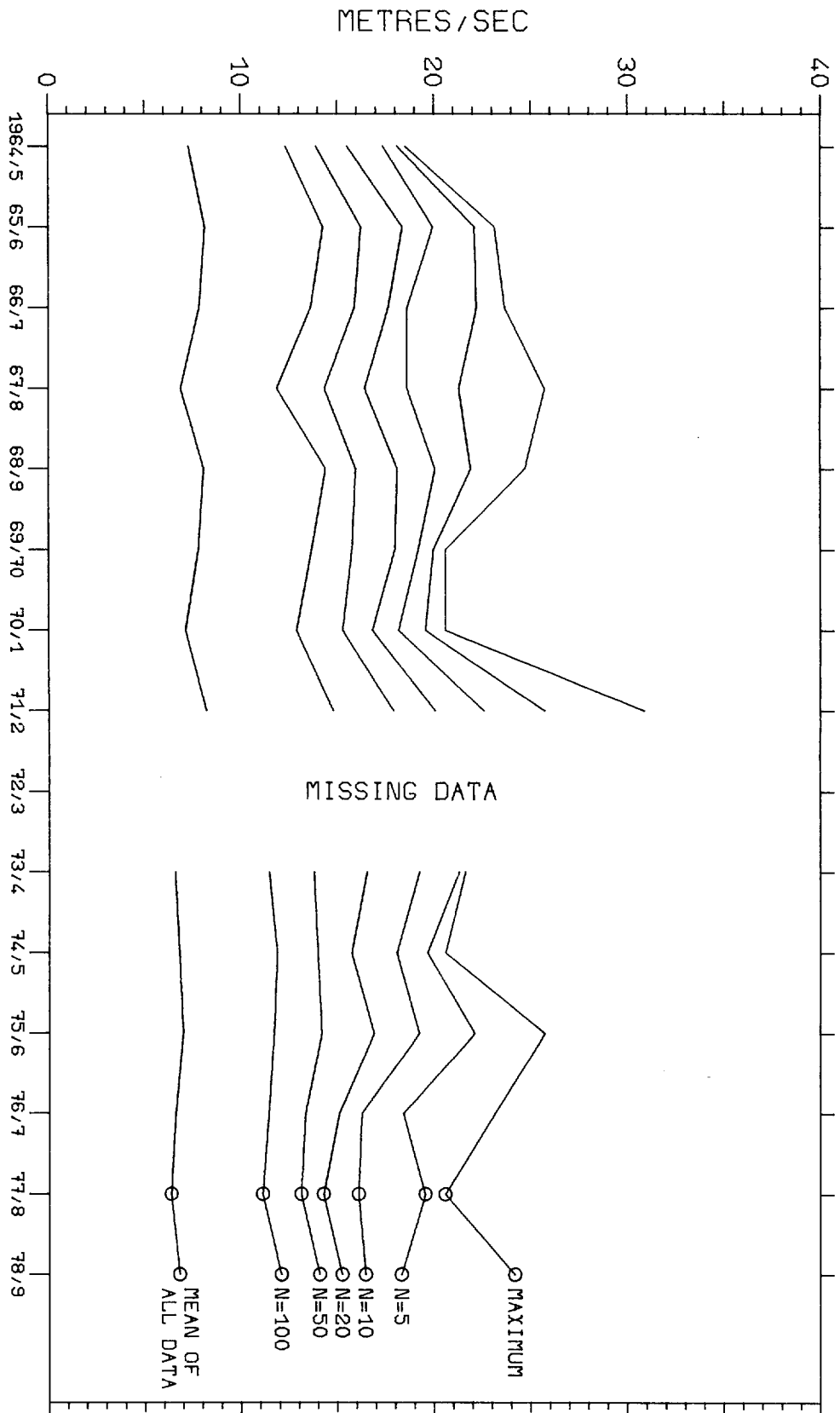
FIG 2.2(A)





MEAN OF N LARGEST VALUES OF WIND SPEED  
 SPURNHEAD NOV 1961 - OCT 1976

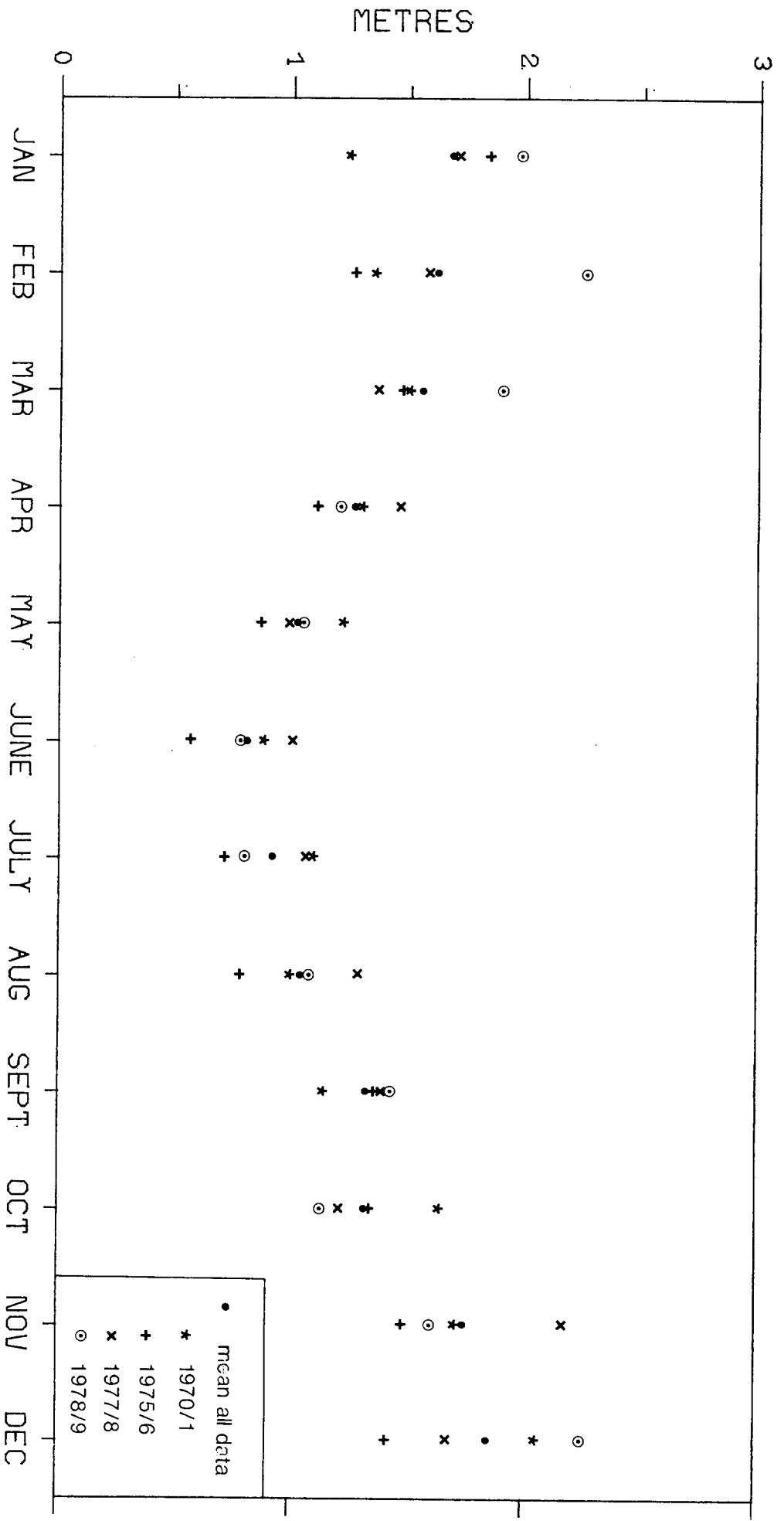
FIG 2.2(B)



MEAN OF N LARGEST VALUES OF WIND SPEED

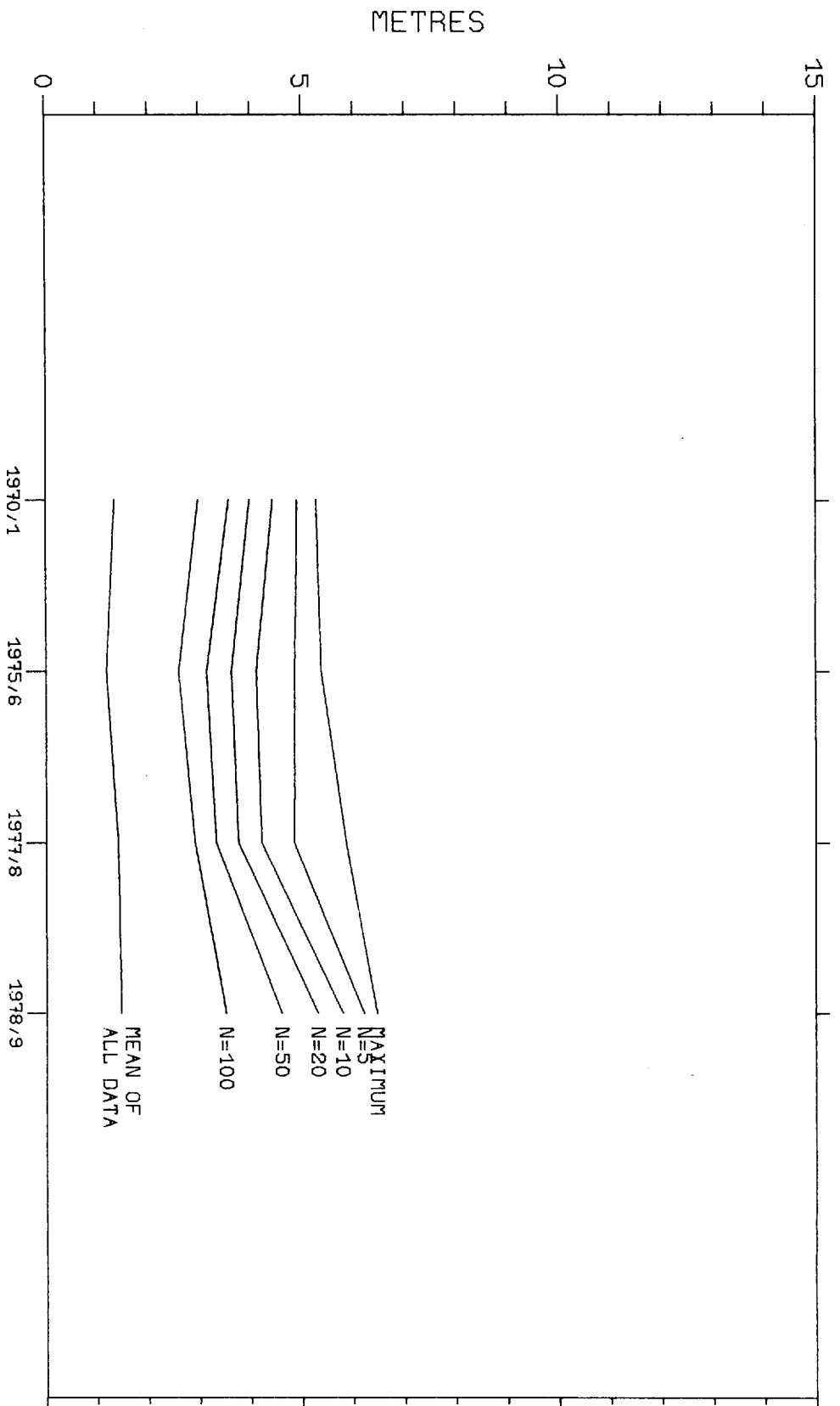
SPURNHEAD AUG 1964 - JUL 1979

FIG 2.2(C)



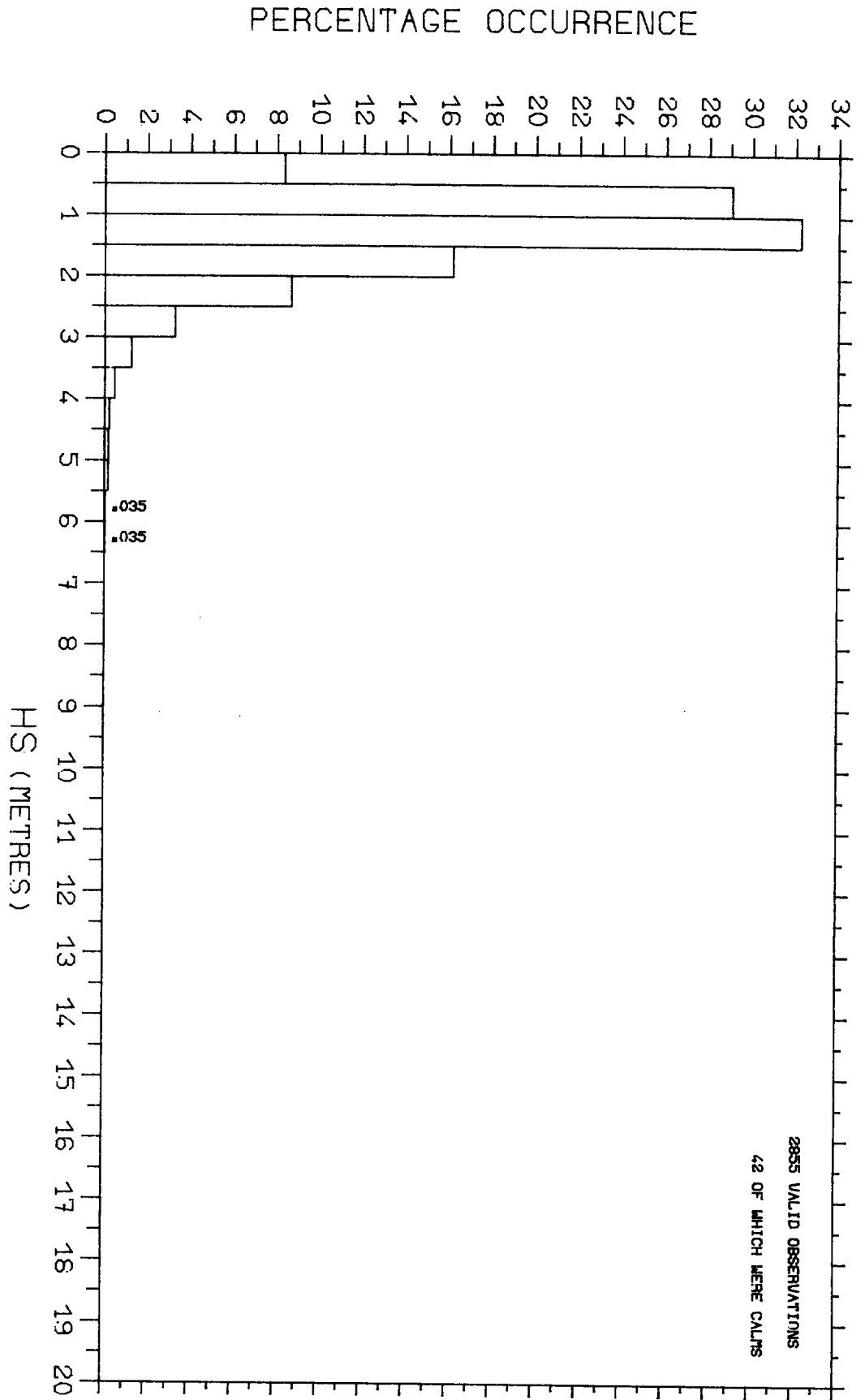
MONTHLY MEANS OF Hs FOR EACH YEAR  
 DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.1.1



MEAN OF N LARGEST VALUES OF  $H_s$   
 DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.1.2



PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71,11/75-10/76,8/77-7/79 SPRINGS

FIG 3.2.1.1

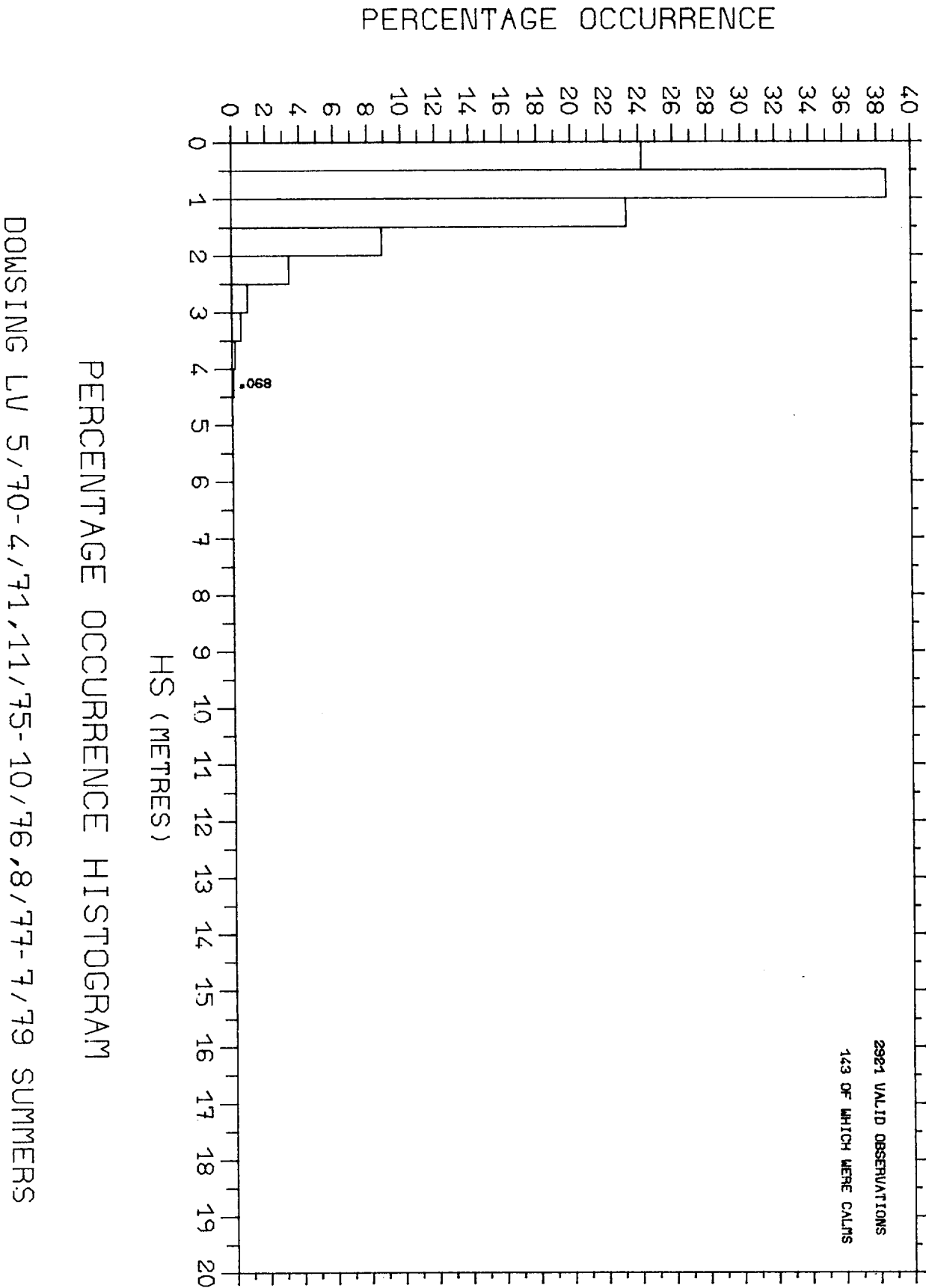
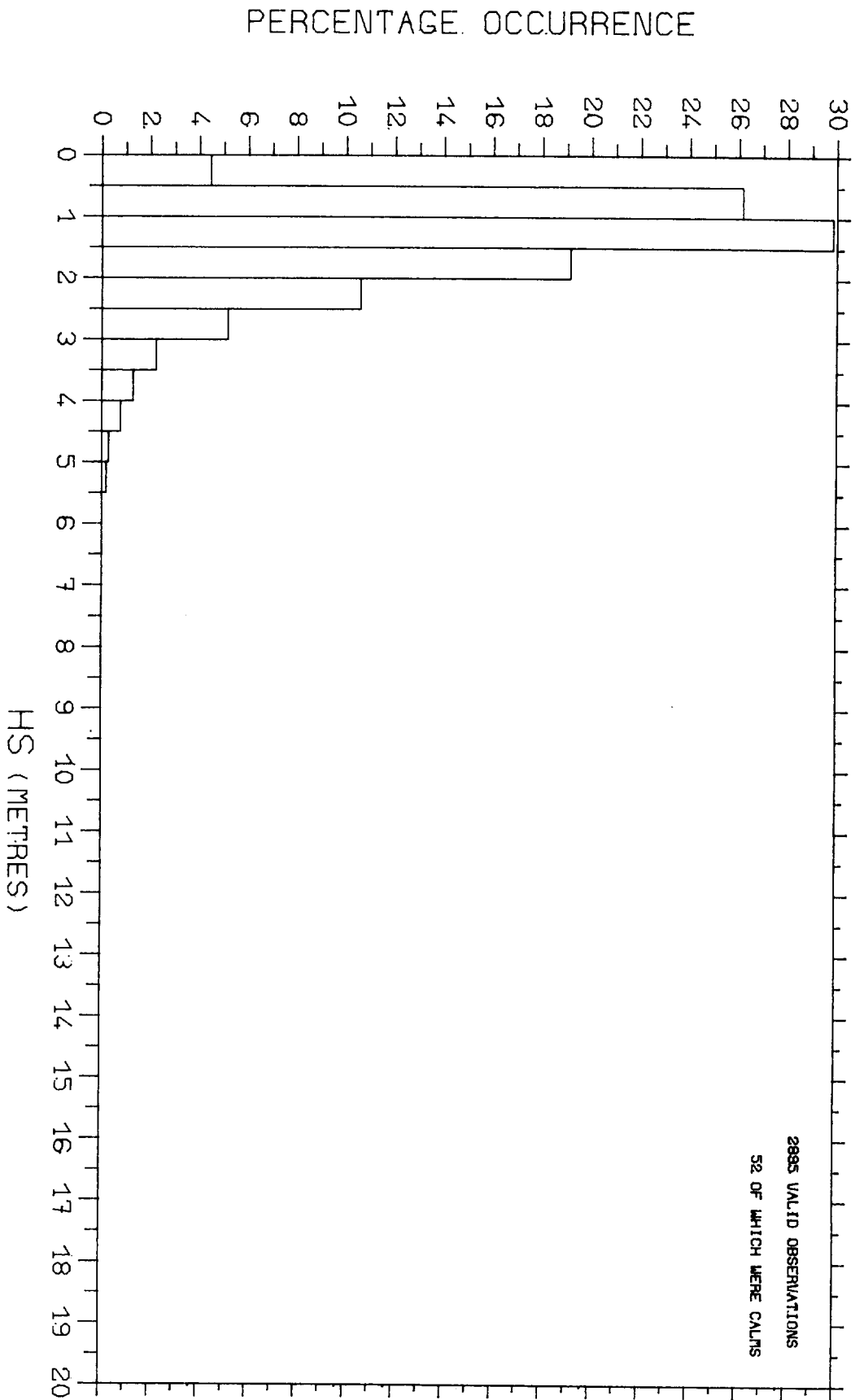


FIG 3.2.1.2

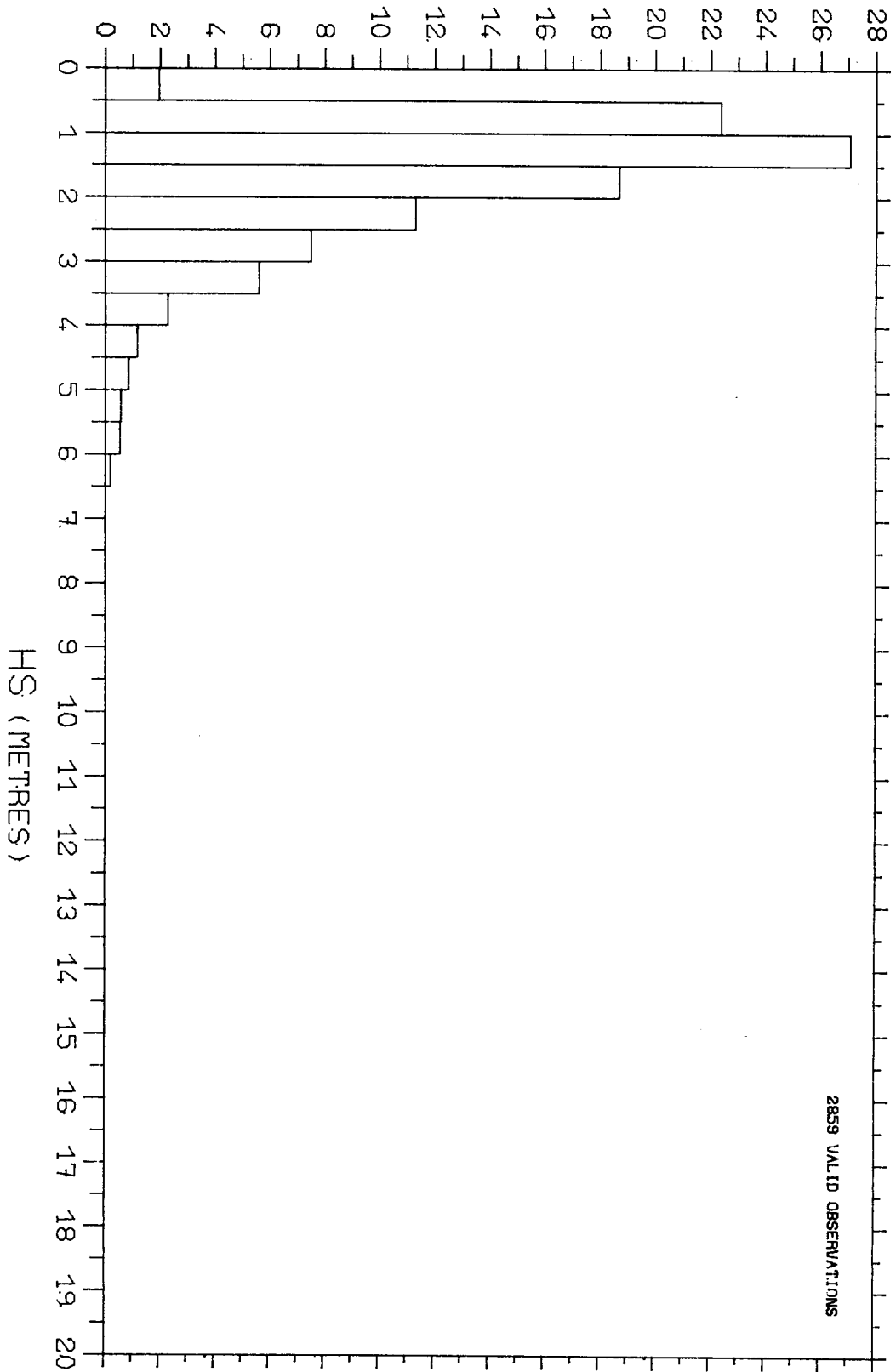


PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71,11/75-10/76,8/77-7/79 AUTUMNS

FIG 3.2.1.3

PERCENTAGE OCCURRENCE

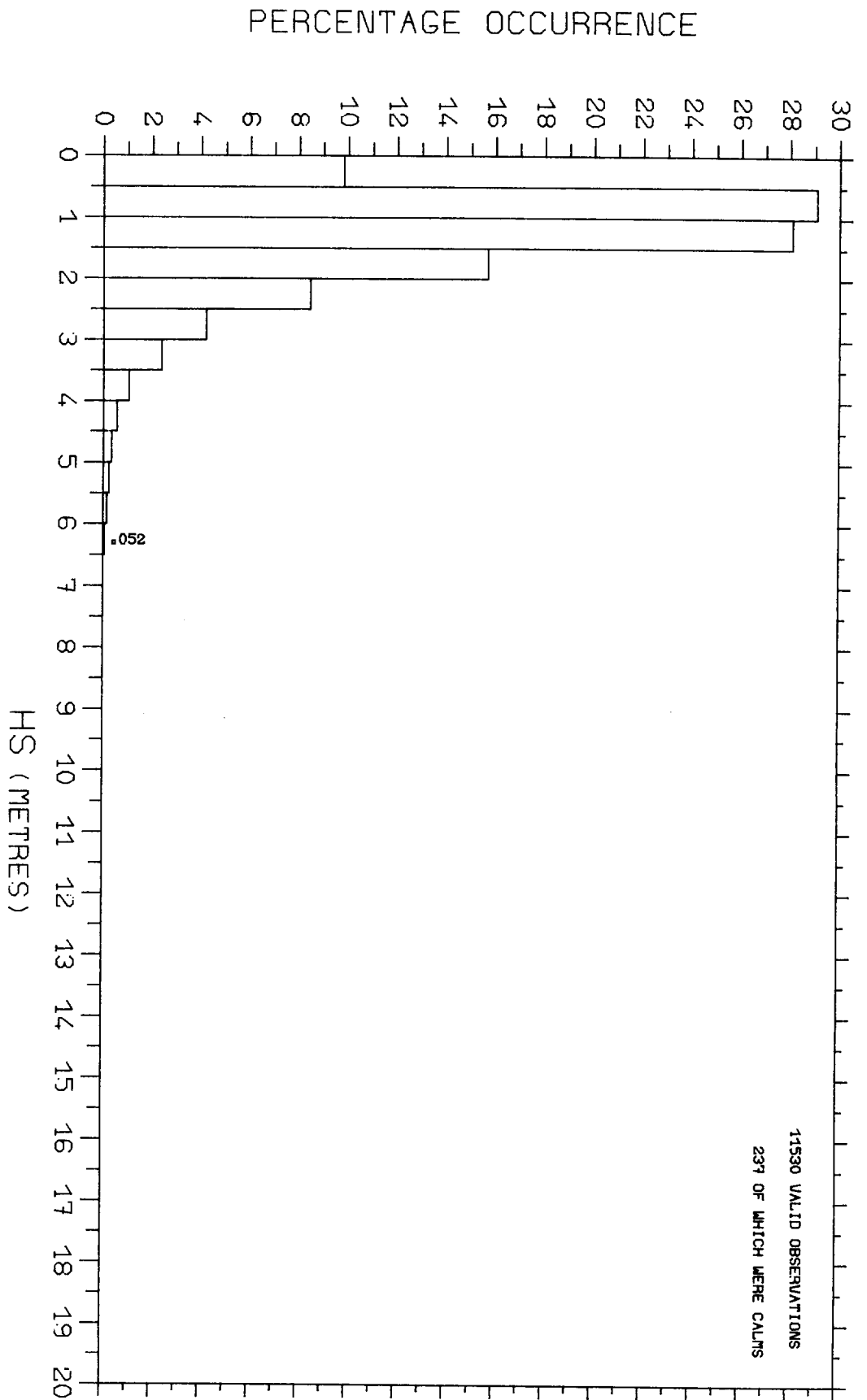


PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71,11/75-10/76,8/77-7/79 WINTERS

FIG 3.2.1.4

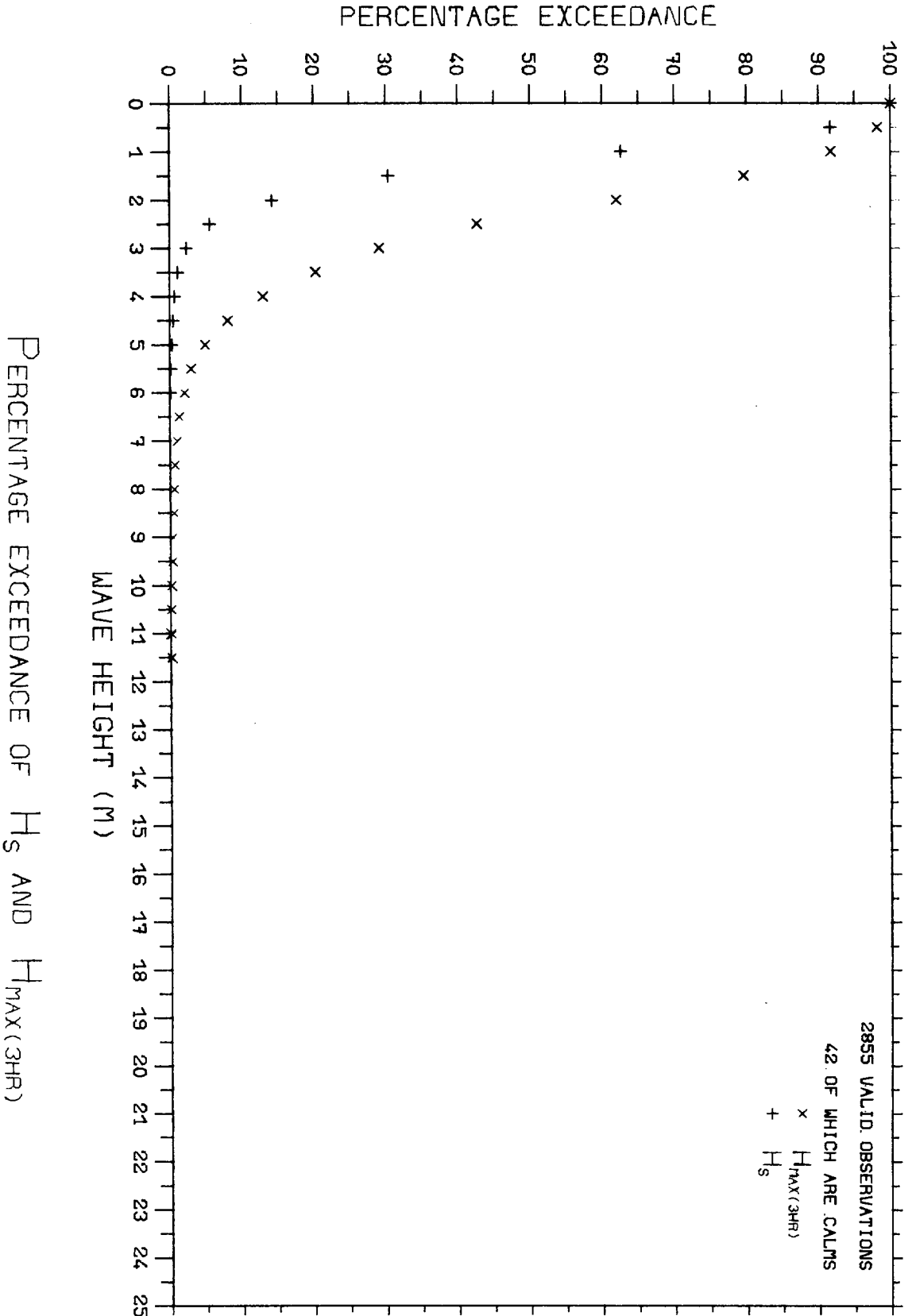




PERCENTAGE OCCURRENCE HISTOGRAM

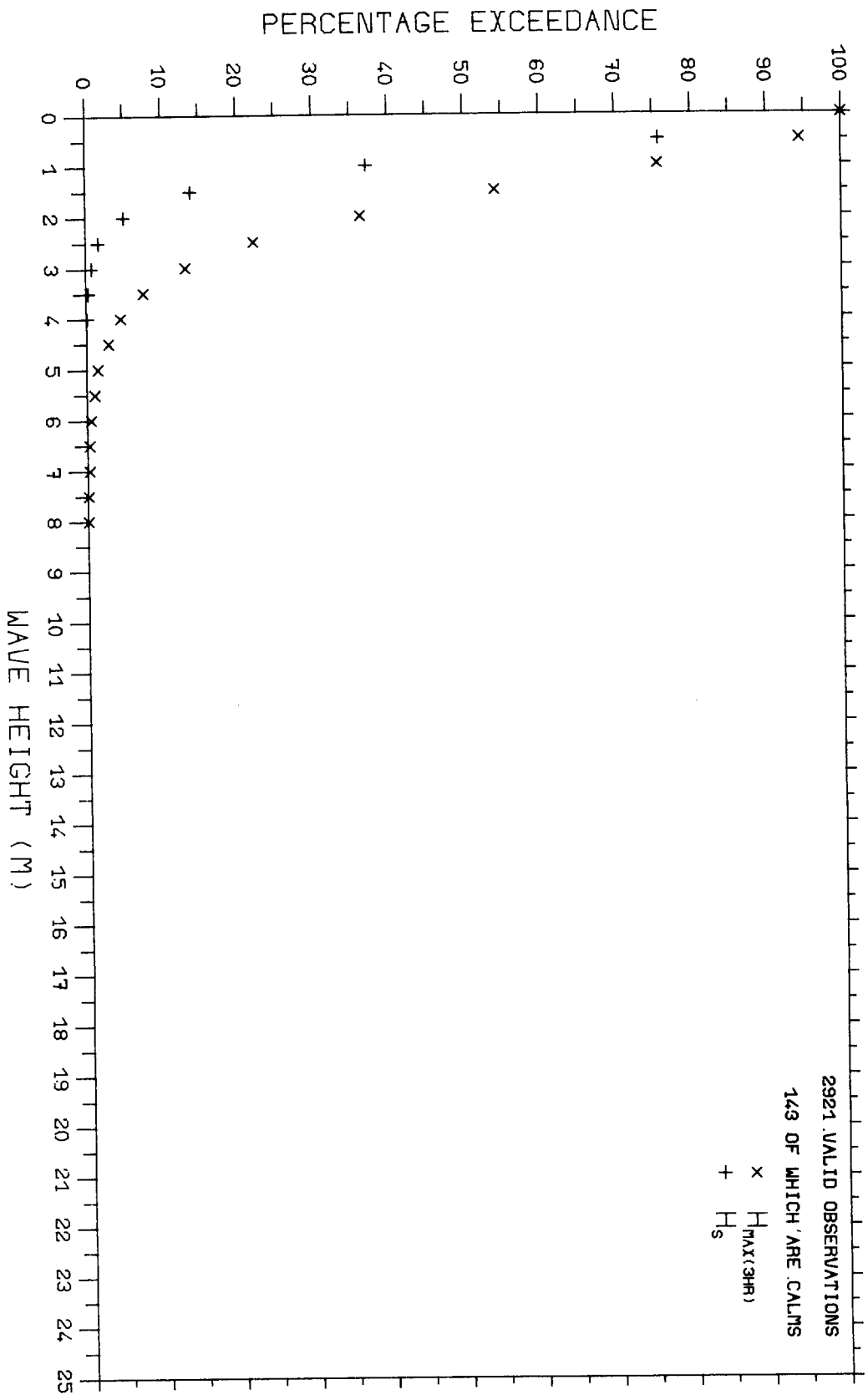
DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.2.1.5



DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - SPRINGS

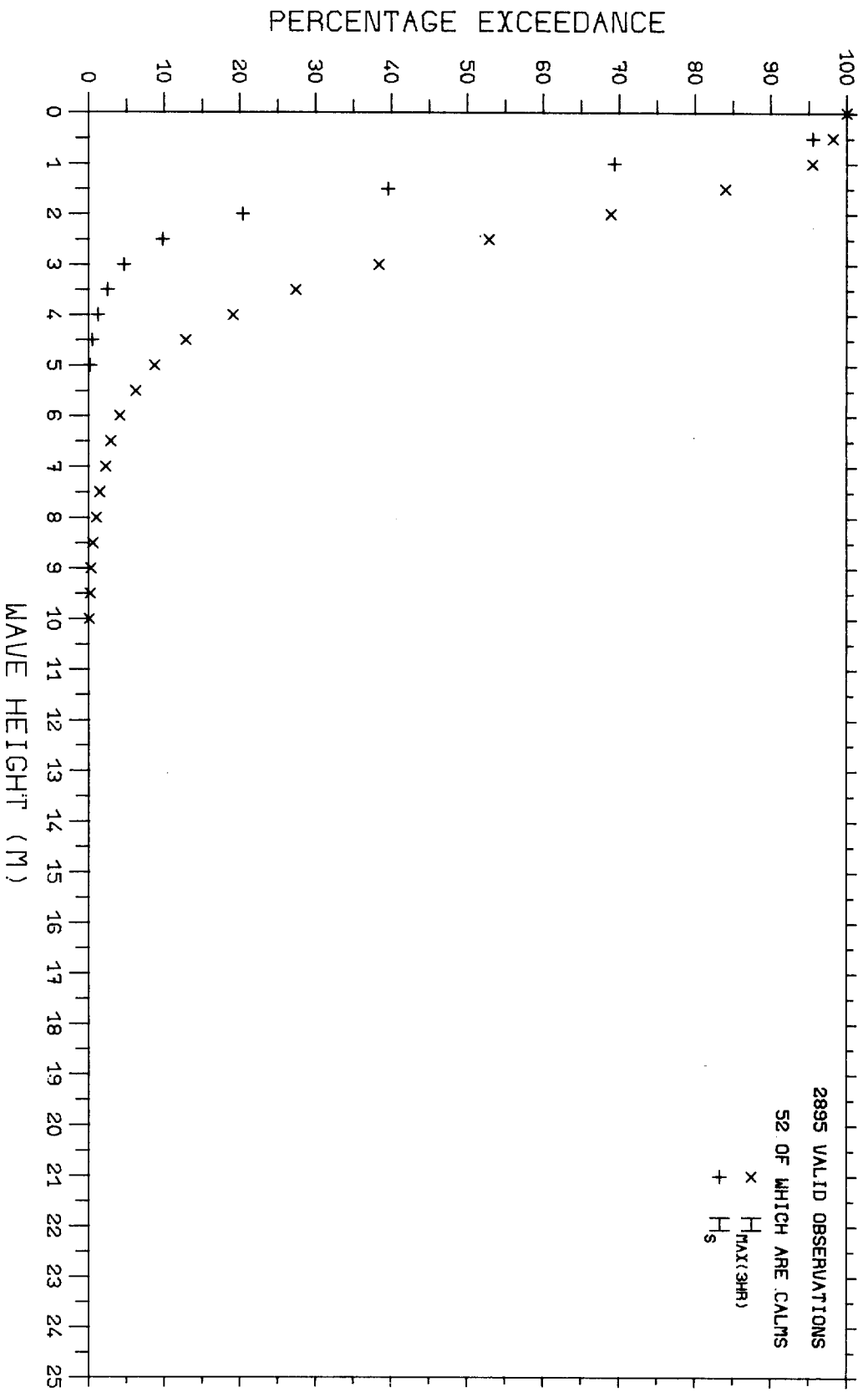
FIG 3.2.2.1



PERCENTAGE EXCEEDANCE OF  $H_s$  AND  $H_{MAX(3HR)}$

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - SUMMERS

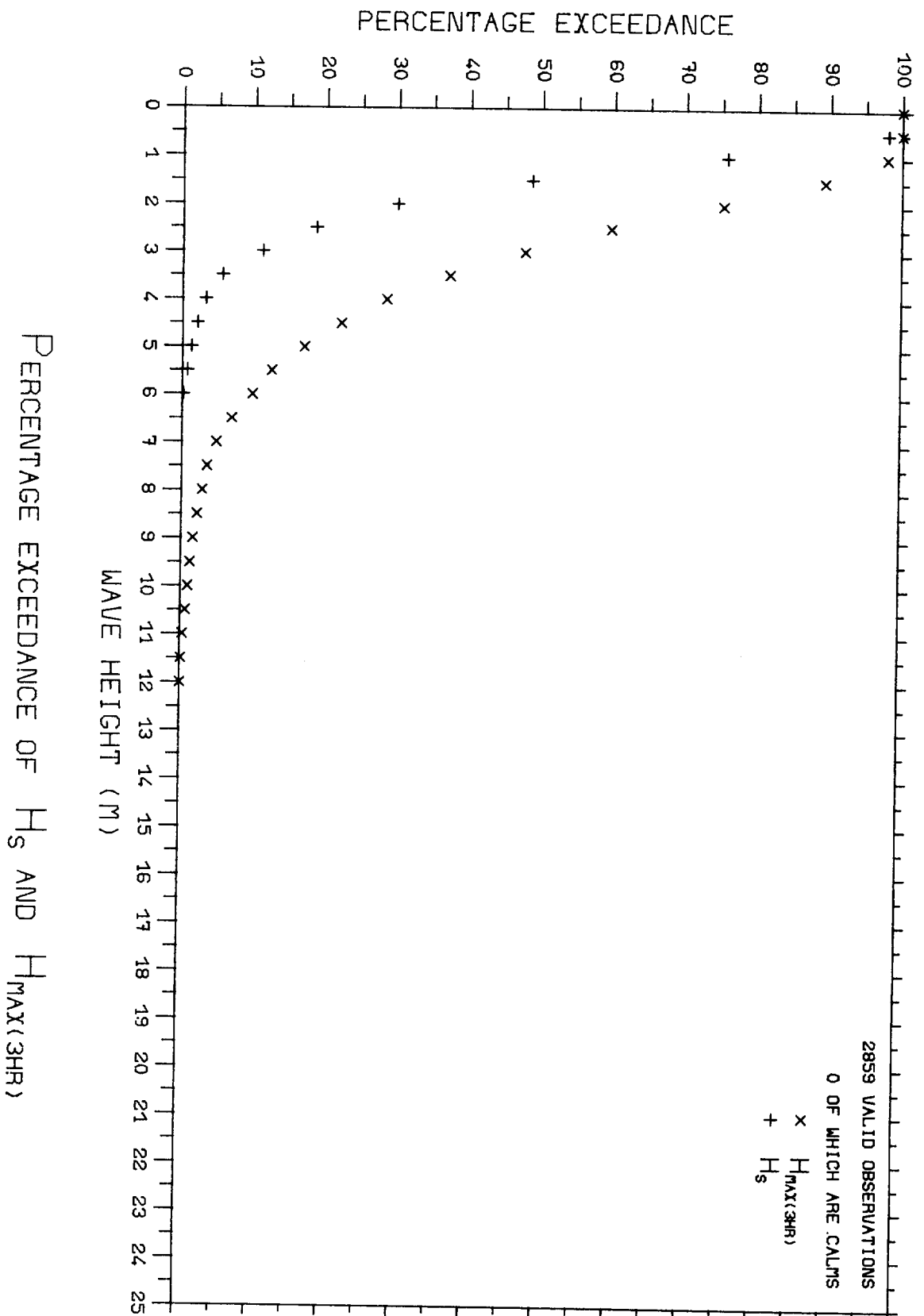
FIG 3.2.2.2



PERCENTAGE EXCEEDANCE OF  $H_s$  AND  $H_{MAX(3HR)}$

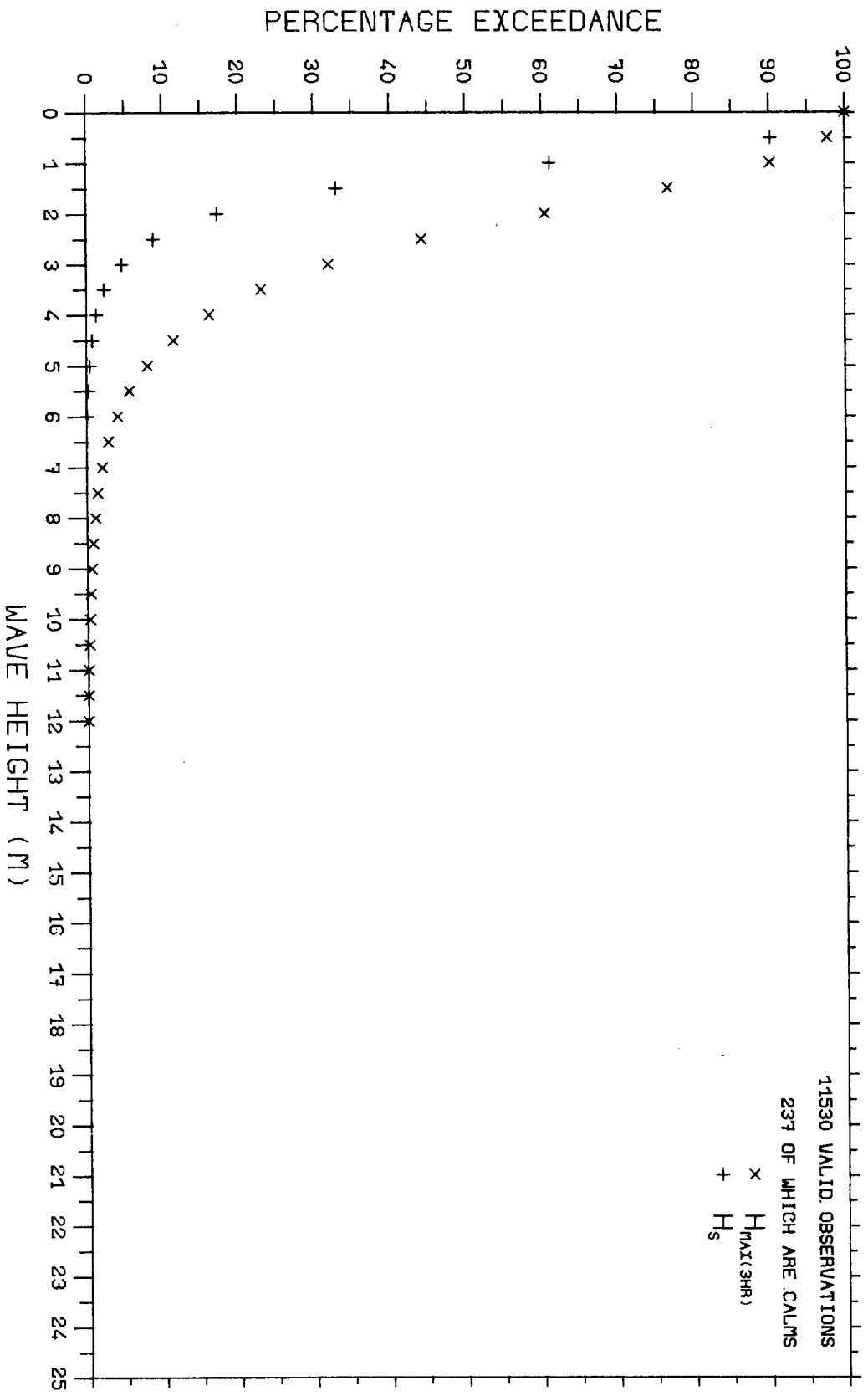
DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - AUTUMNS

FIG 3.2.2.3



PERCENTAGE EXCEEDANCE OF  $H_s$  AND  $H_{MAX(3HR)}$   
 DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - WINTERS

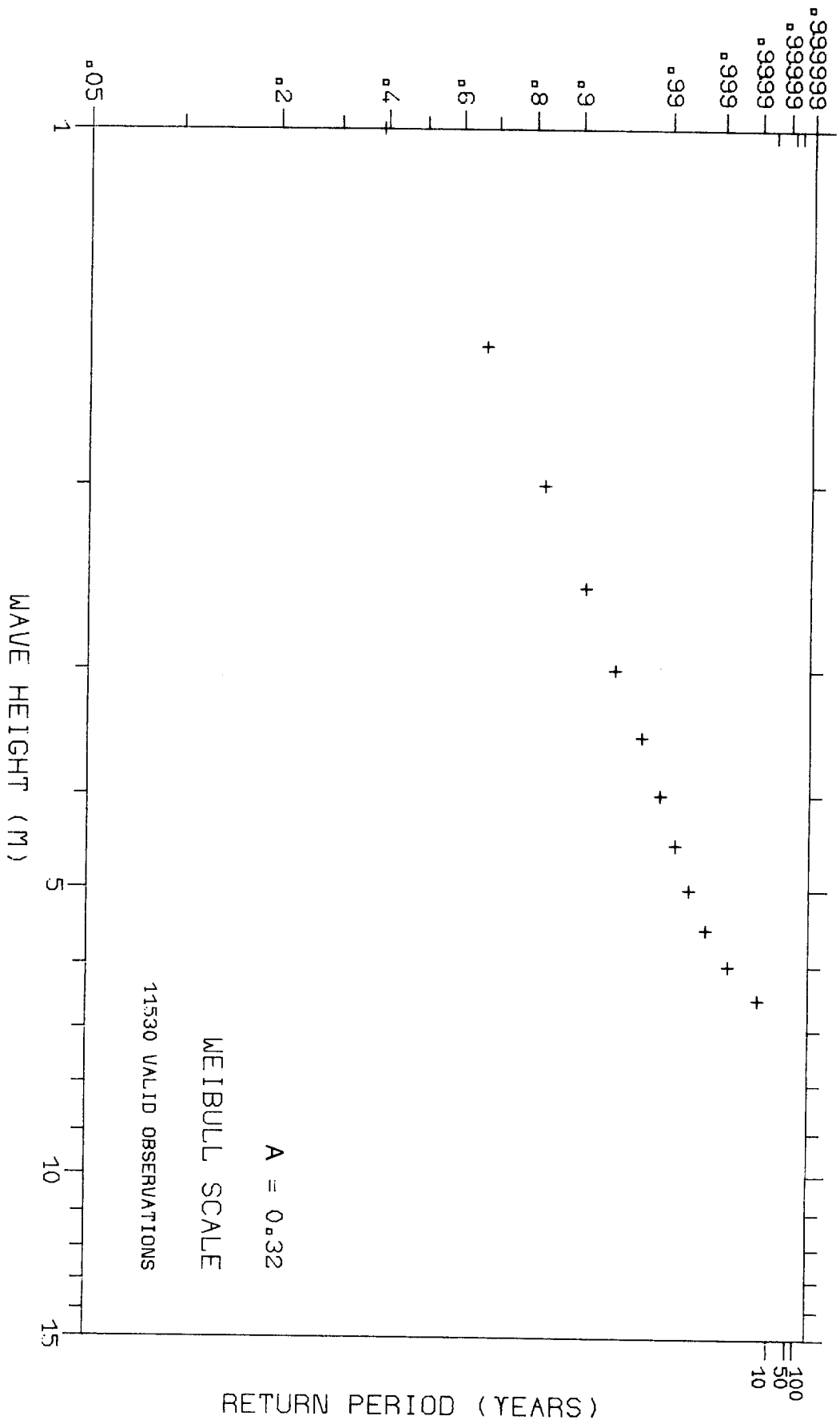
FIG 3.2.2.4



PERCENTAGE EXCEEDANCE OF  $H_s$  AND  $H_{\text{MAX}(3\text{HR})}$   
DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.2.2.5

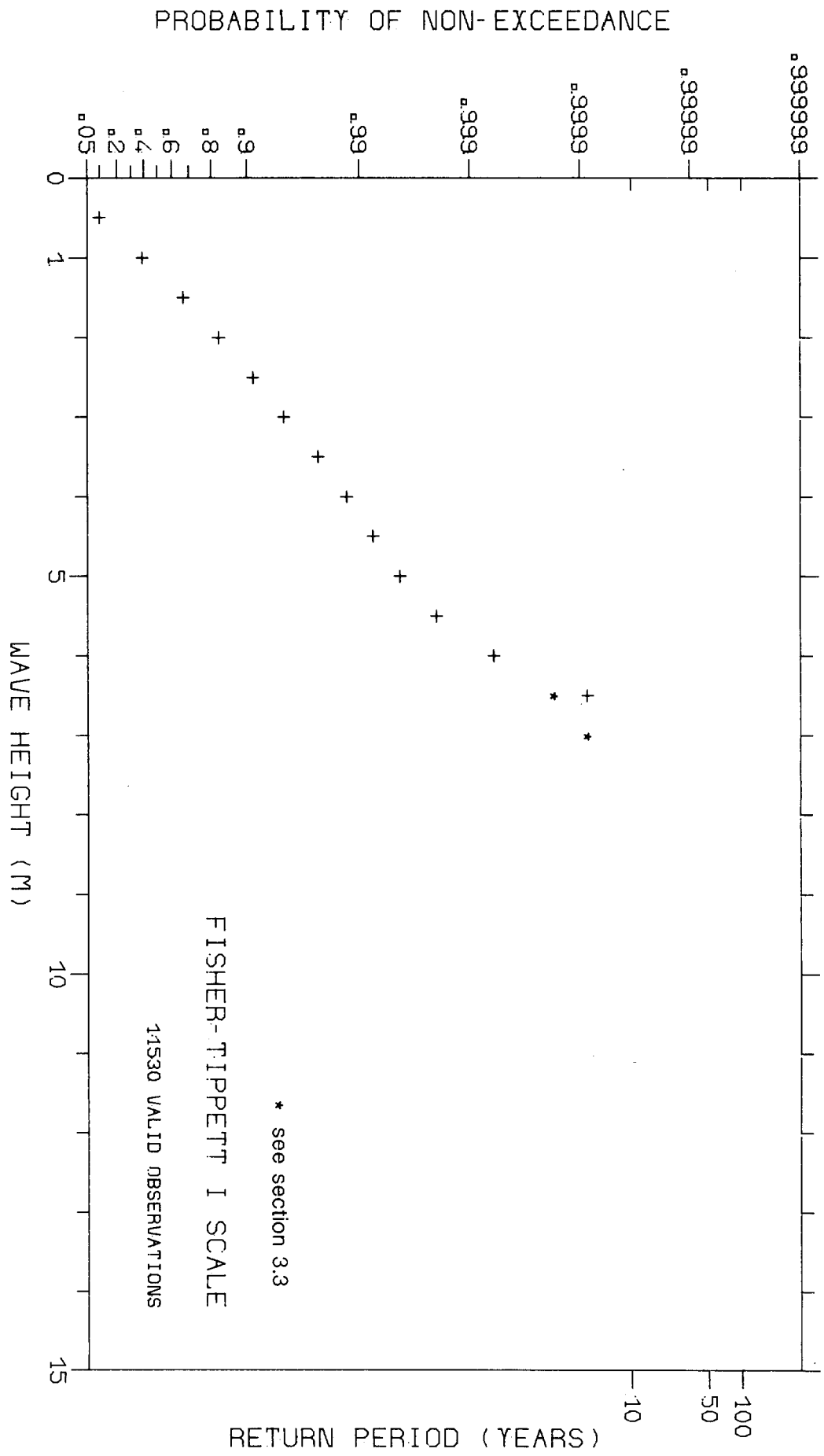
PROBABILITY OF NON-EXCEEDANCE



CUMULATIVE DISTRIBUTION OF WAVE HEIGHT, HS

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.3.1

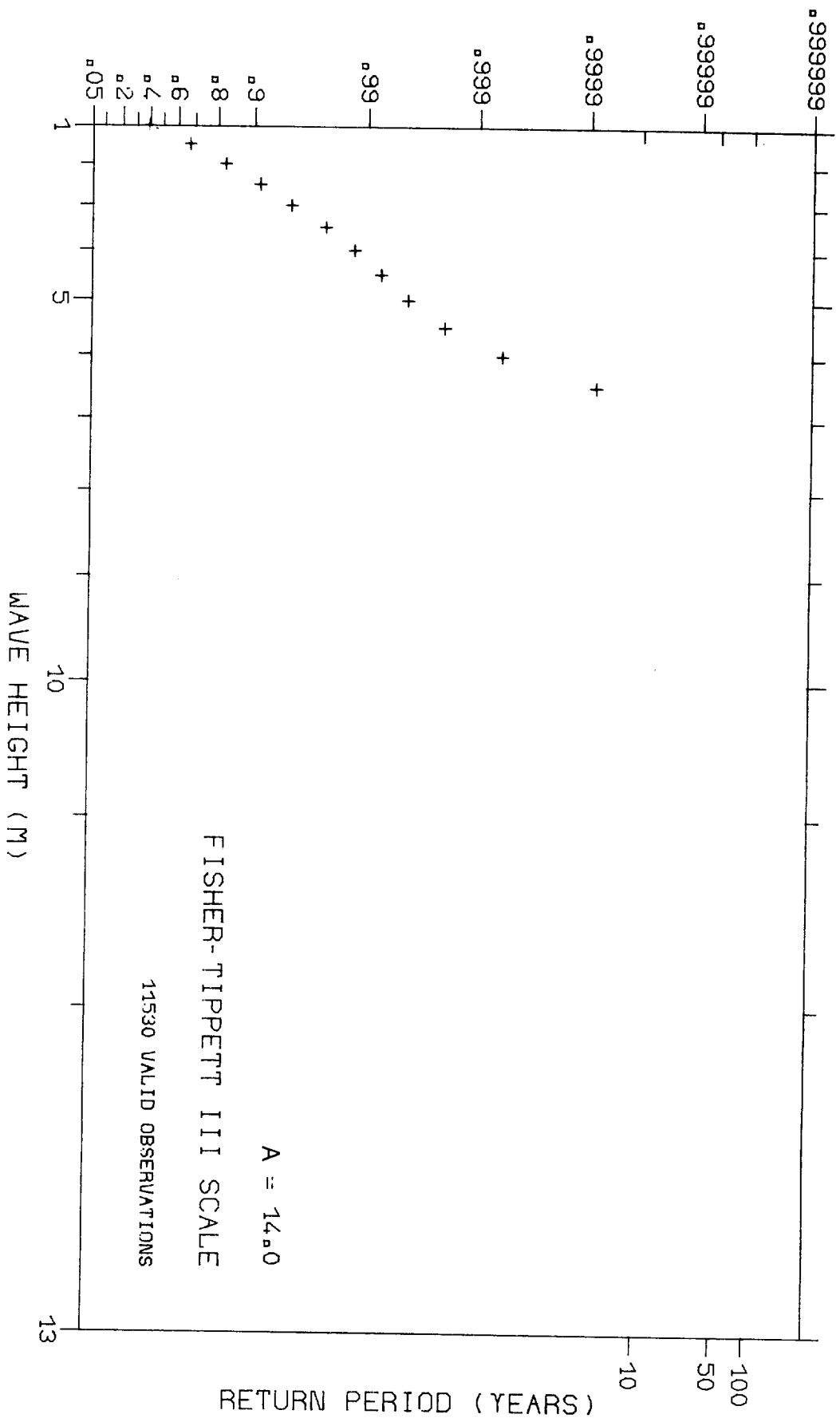


CUMULATIVE DISTRIBUTION OF WAVE HEIGHT, HS  
 DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.3.2

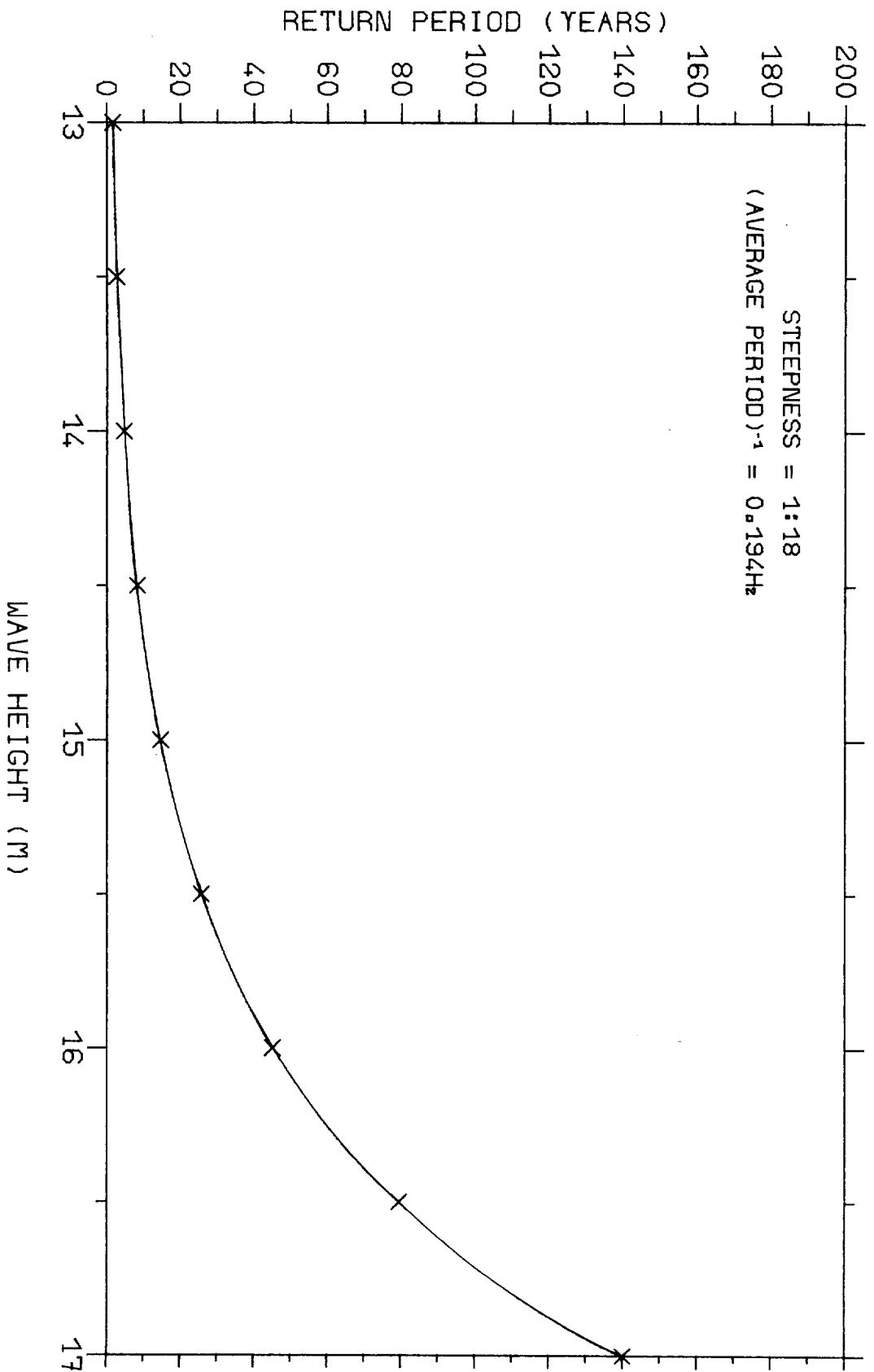


PROBABILITY OF NON-EXCEEDANCE



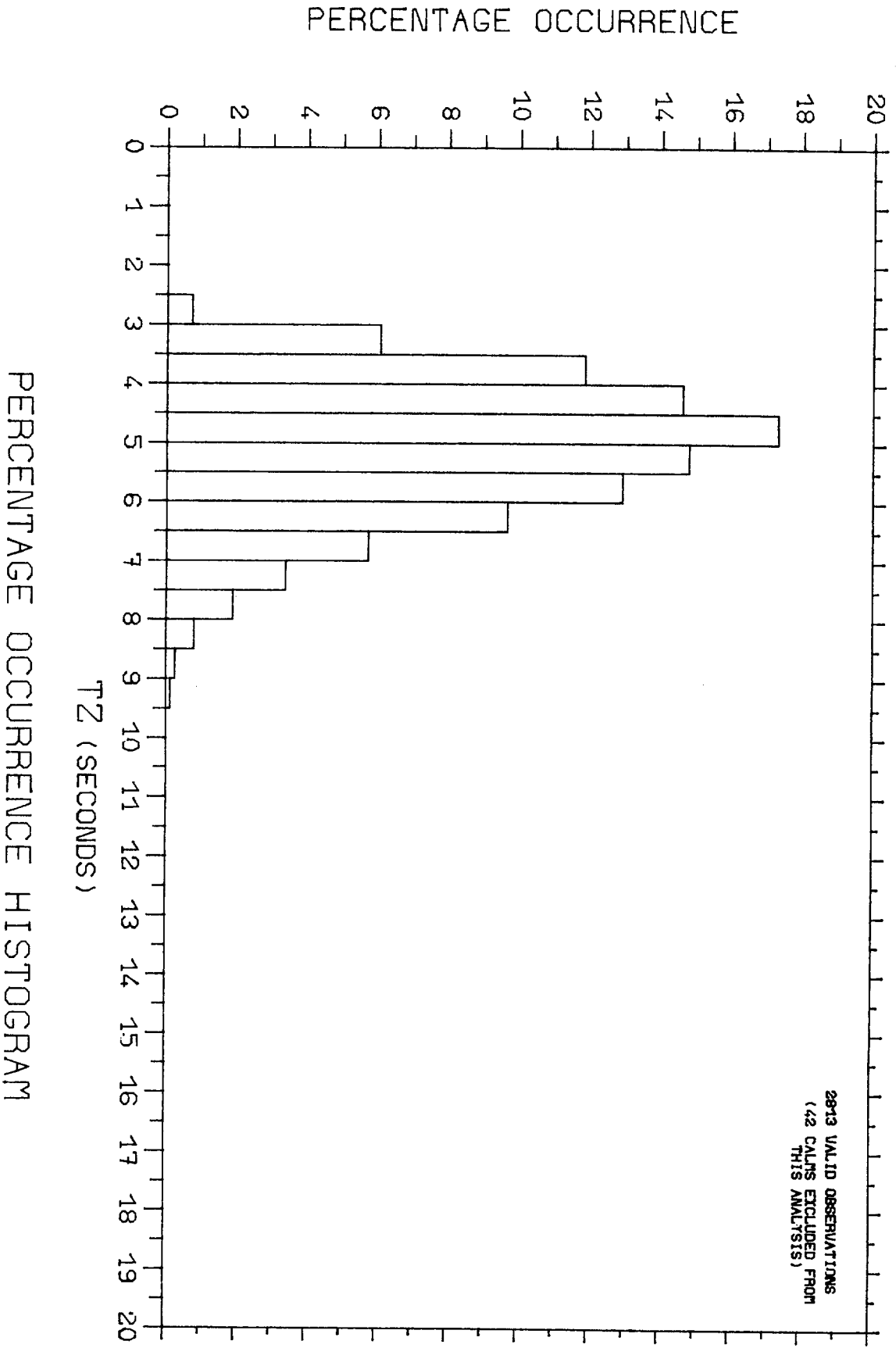
DOMSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

FIG 3.3.3



RETURN PERIOD v. WAVE HEIGHT - INDIVIDUAL WAVE MODEL  
 DOWNSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

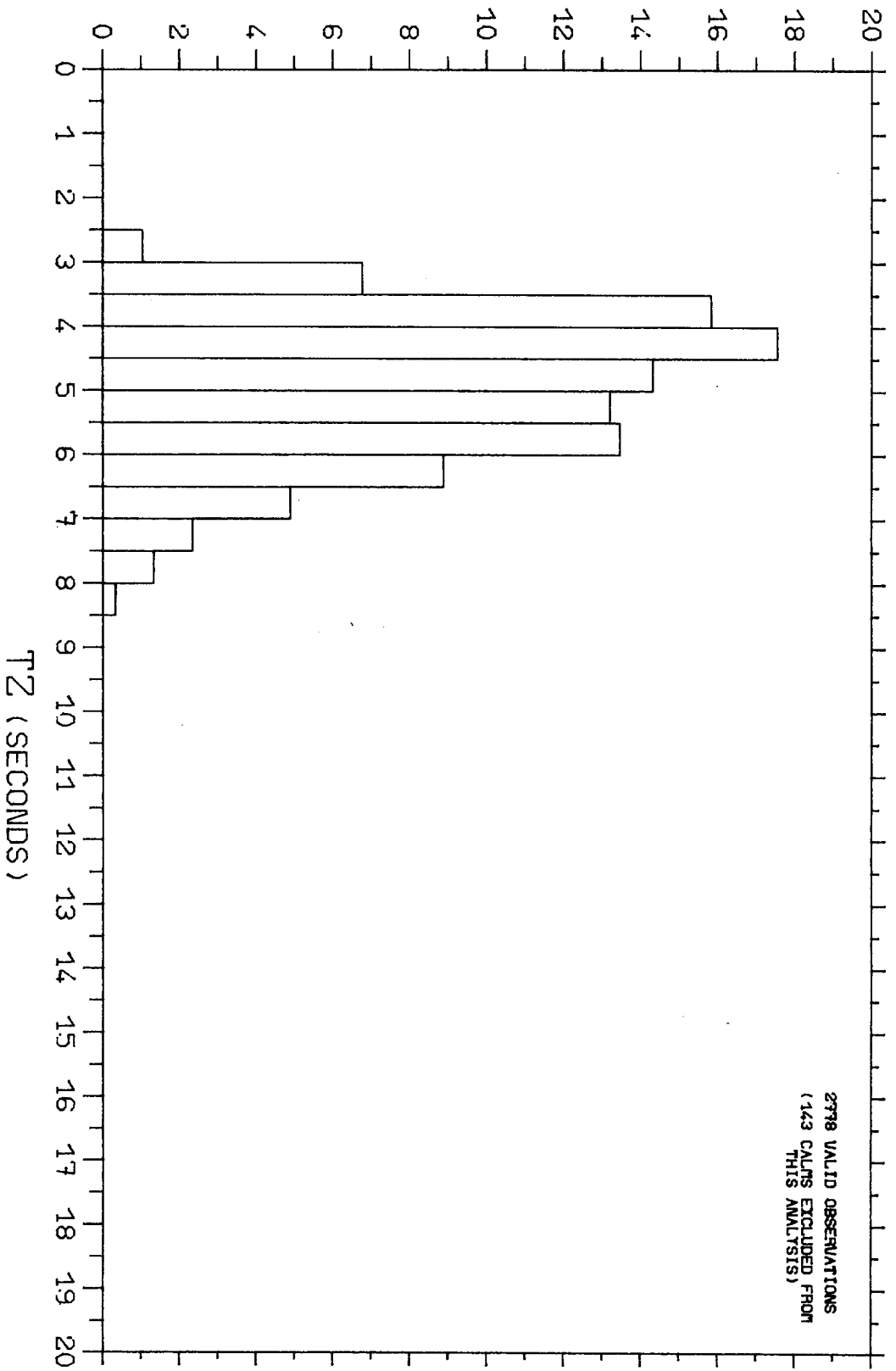
FIG 3.3.4



DOWSING LV 5/70-4/71,11/75-10/76,8/77-7/79 SPRINGS

FIG 3.4.1.1

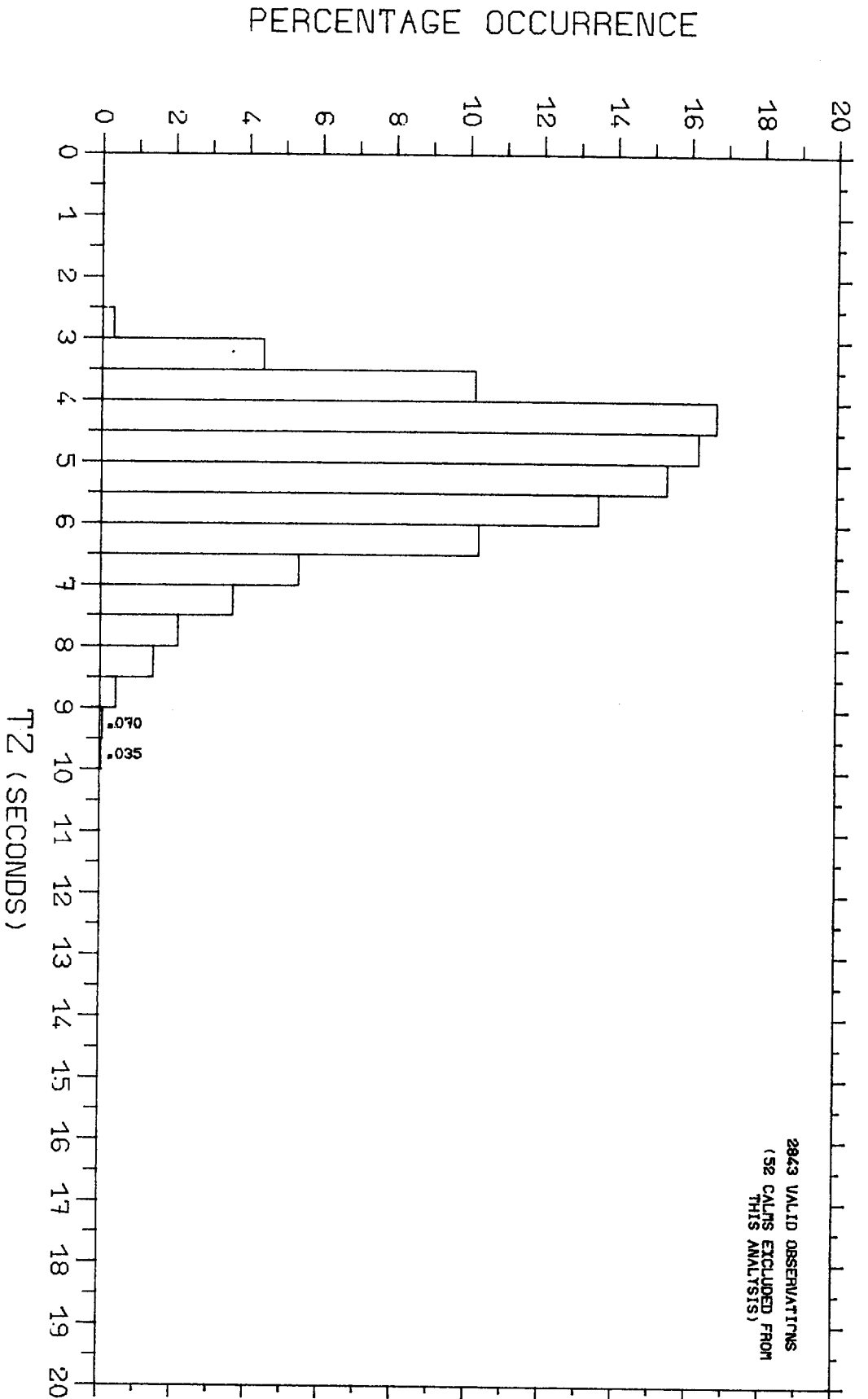
PERCENTAGE OCCURRENCE



PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 SUMMERS

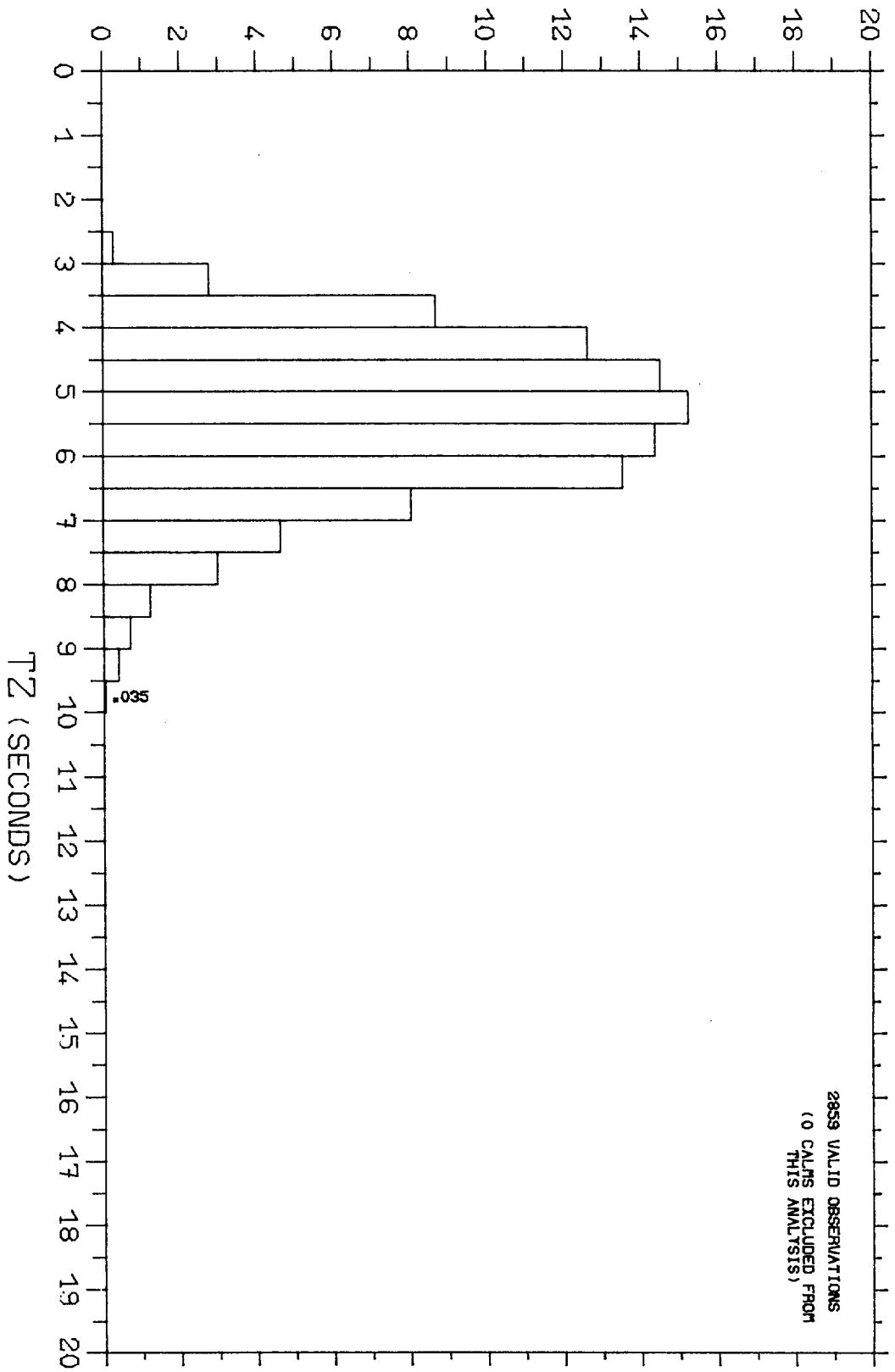
FIG 3.4.1.2



DOMSING LV 5/70-4/71,11/75-10/76,8/77-7/79 AUTUMNS

FIG 3.4.1.3

PERCENTAGE OCCURRENCE

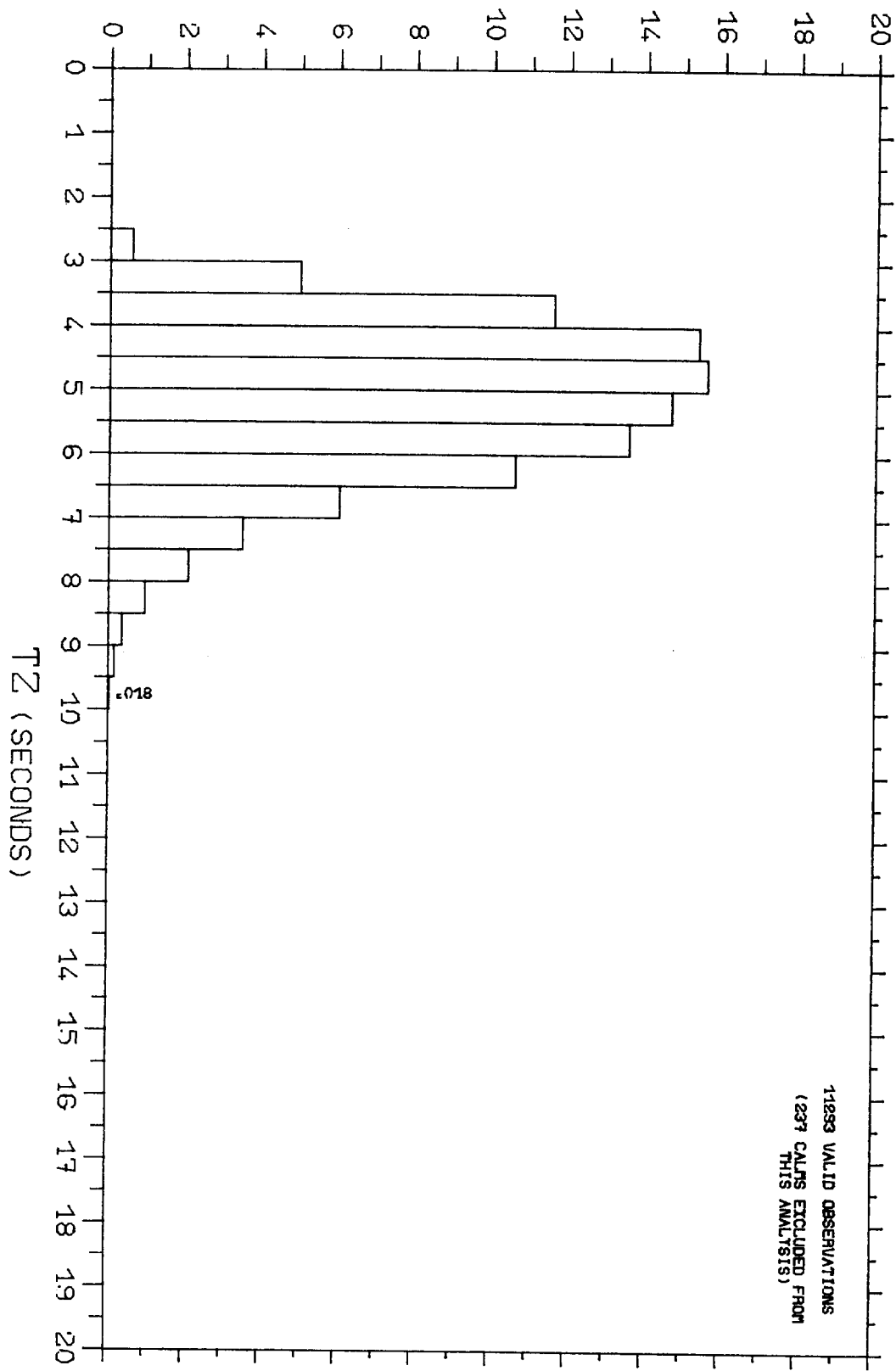


PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 WINTERS

FIG 3.4.1.4

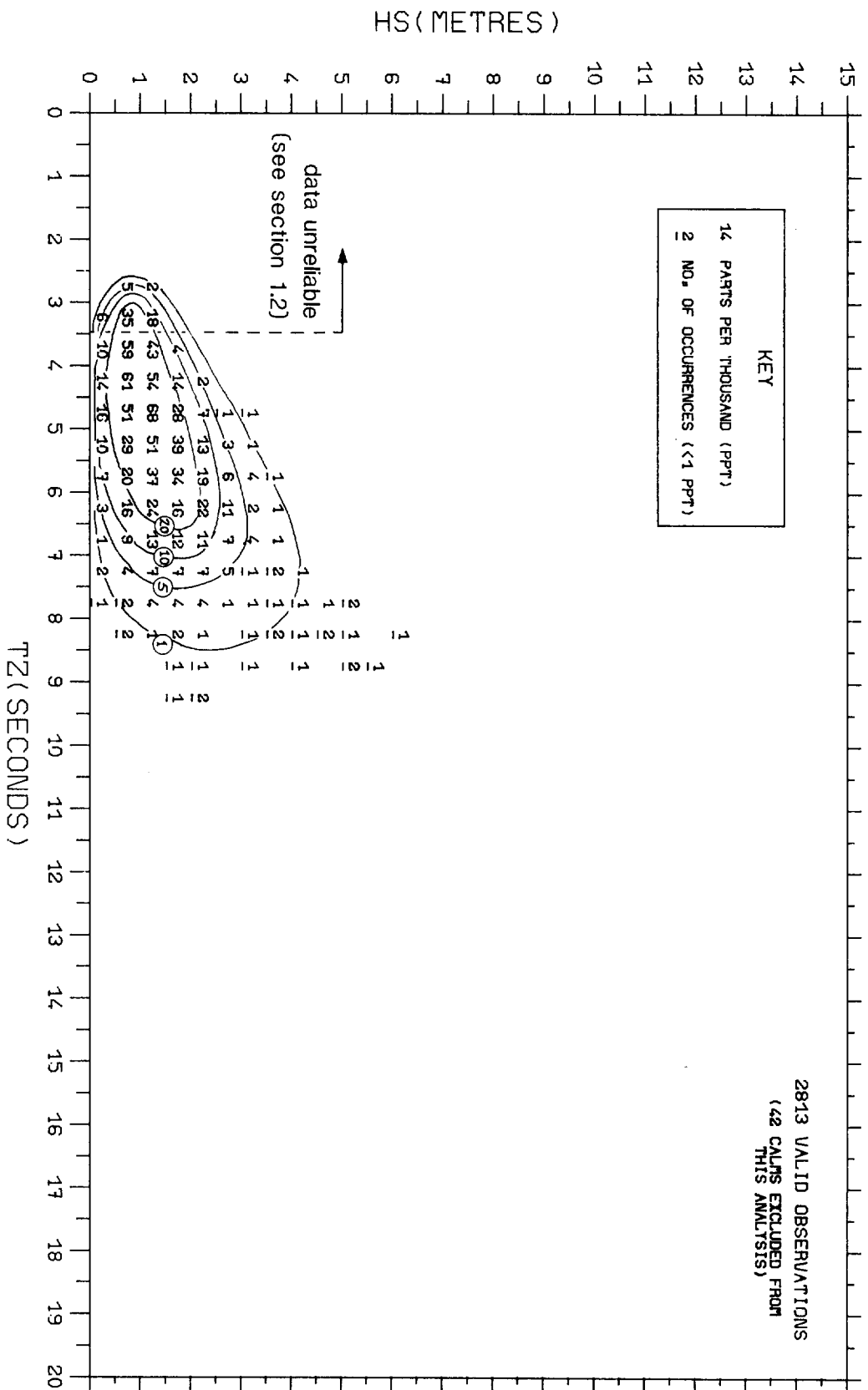
PERCENTAGE OCCURRENCE



PERCENTAGE OCCURRENCE HISTOGRAM

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

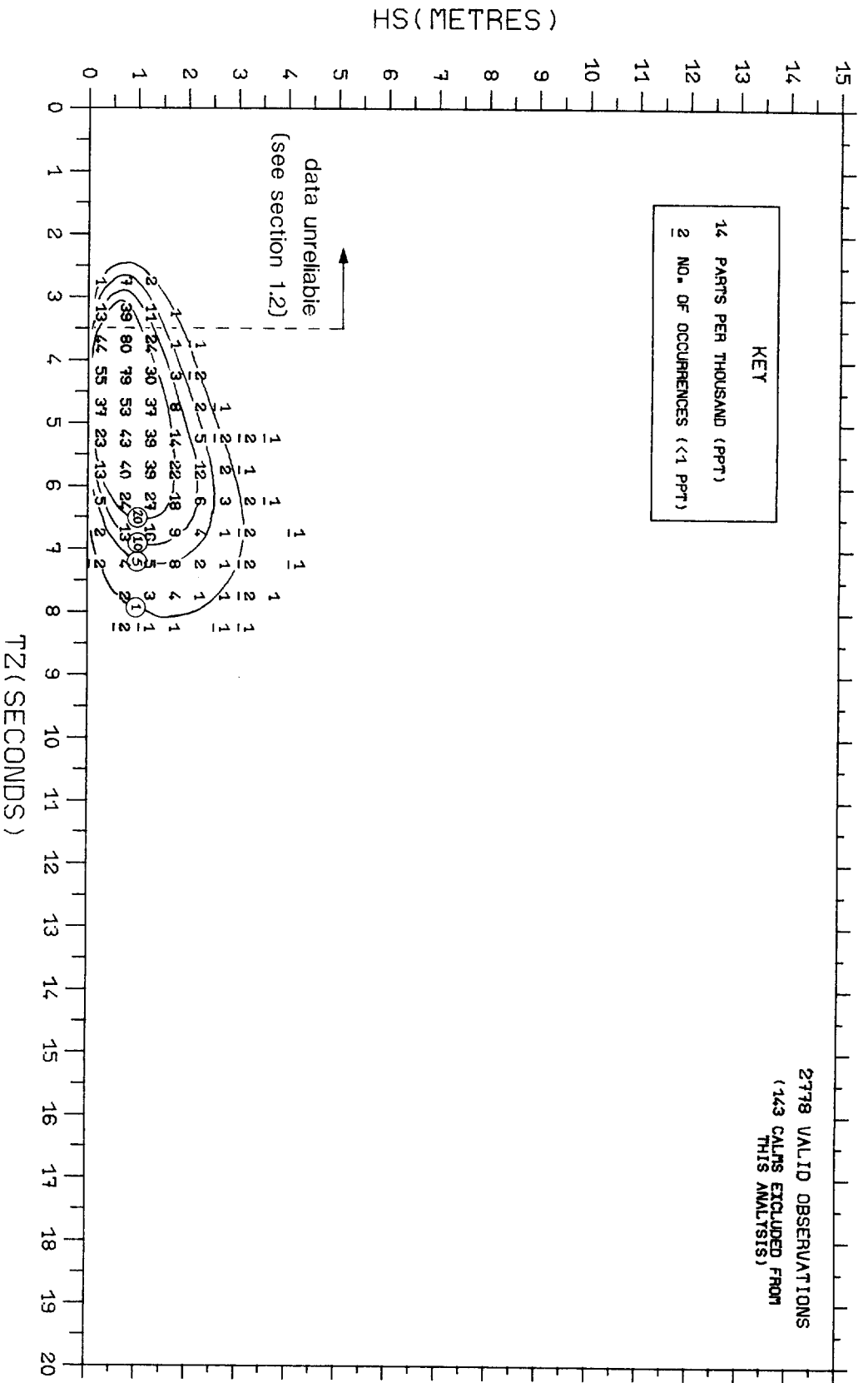
FIG 3.4.1.5



DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - SPRINGS

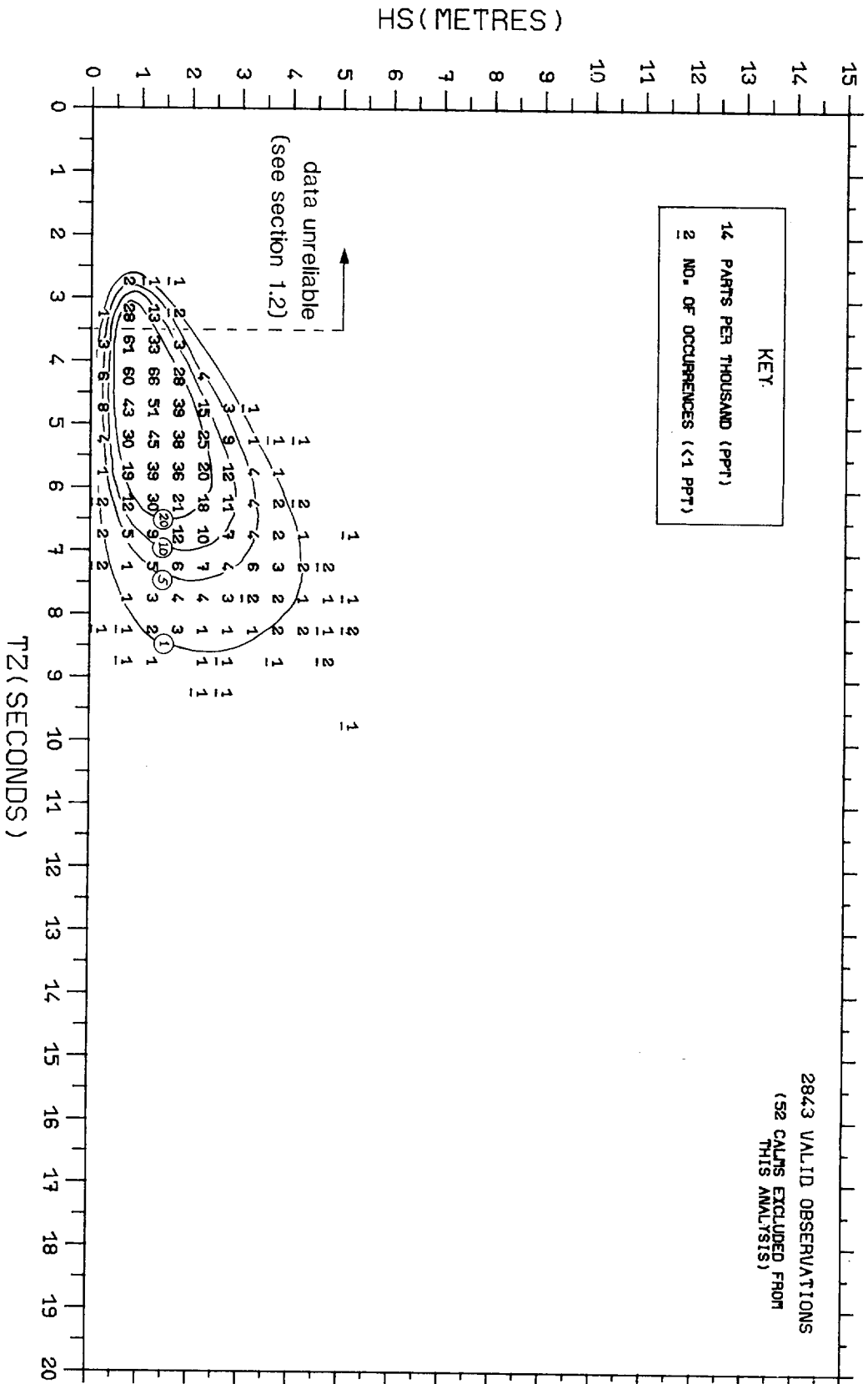
FIG 3.5.1.1





DOMSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - SUMMERS

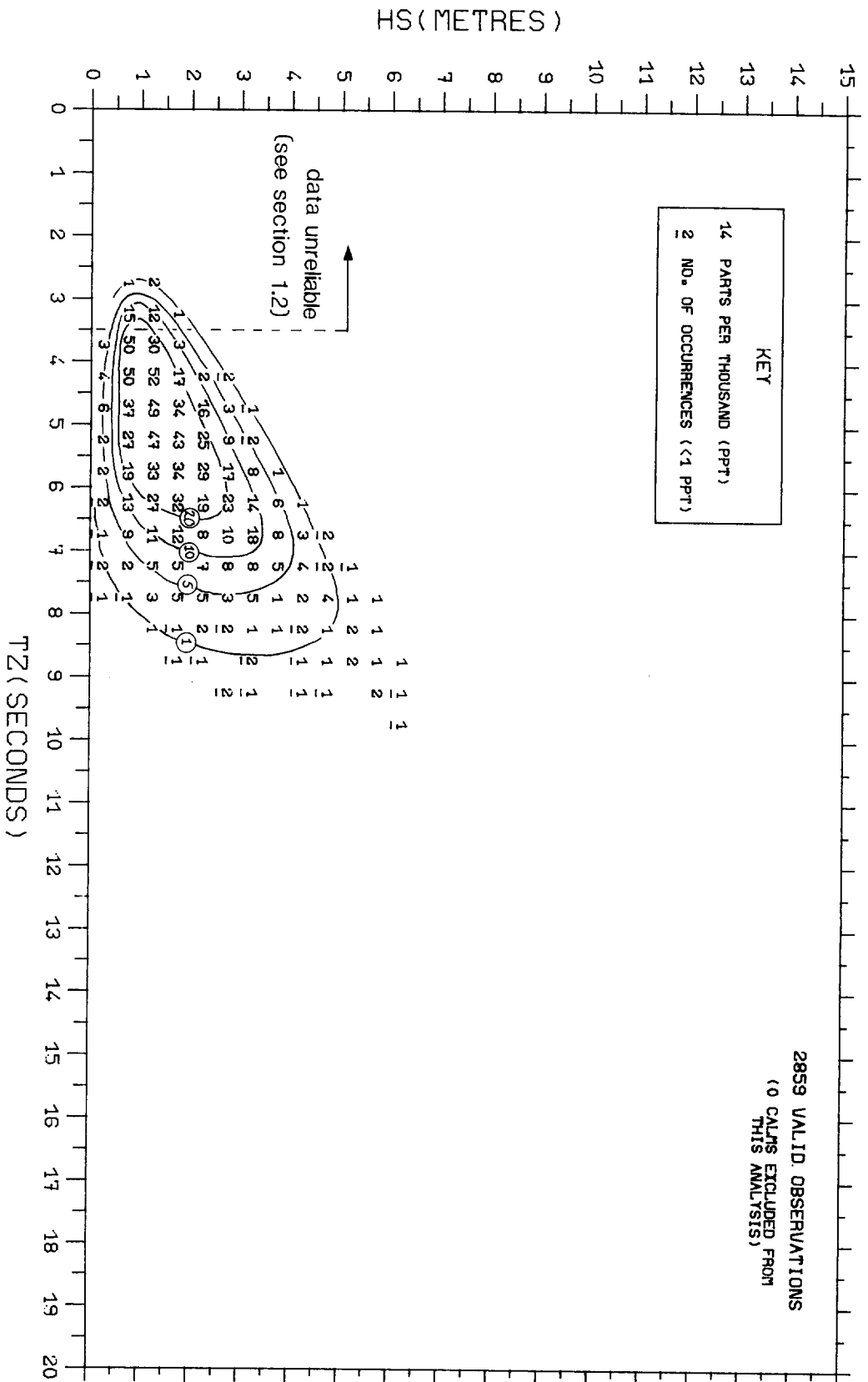
FIG 3.5.1.2



SCATTER PLOT OF HS AND TZ

DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - AUTUMNS

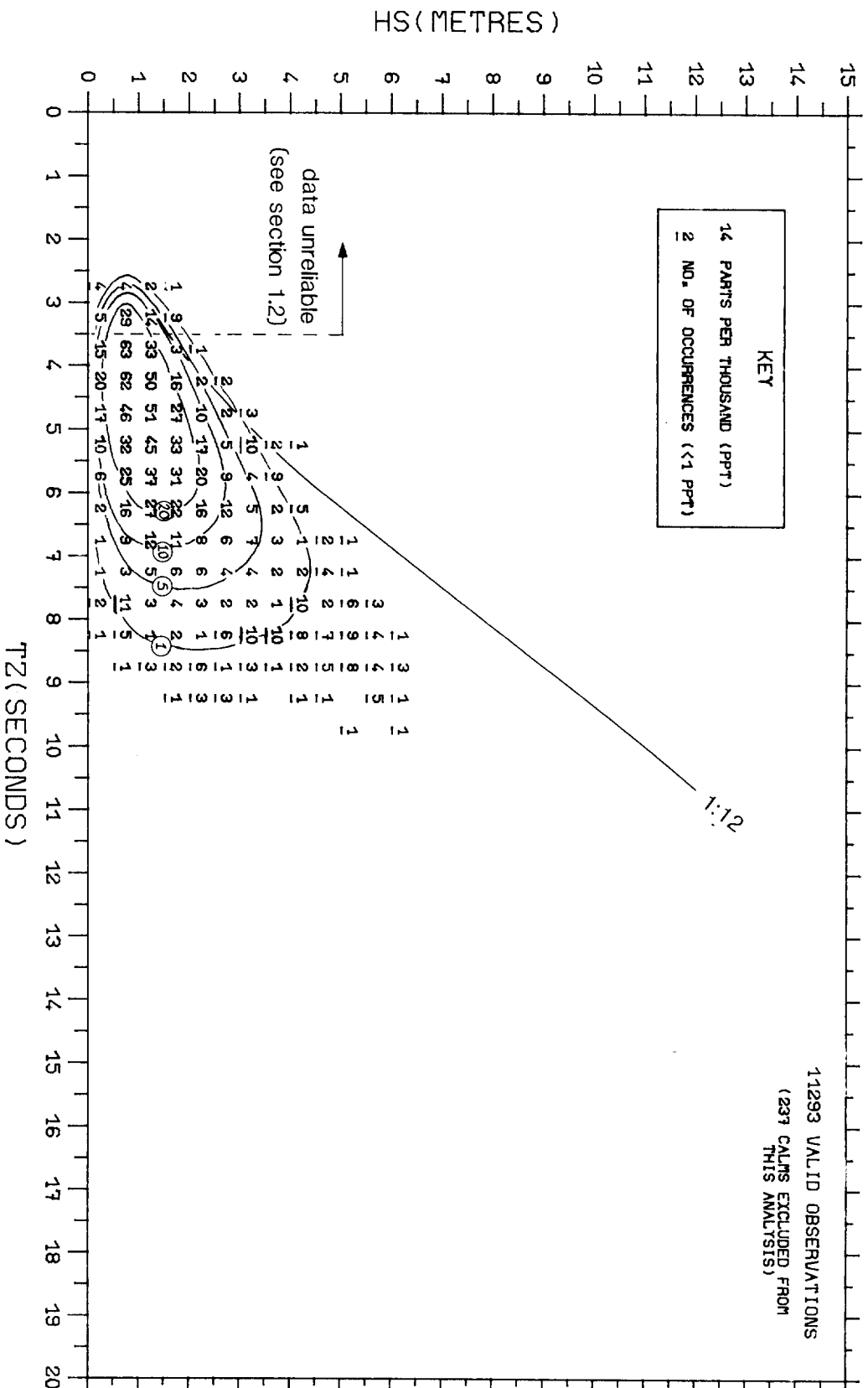
FIG 3.5.1.3



SCATTER PLOT OF HS AND TZ

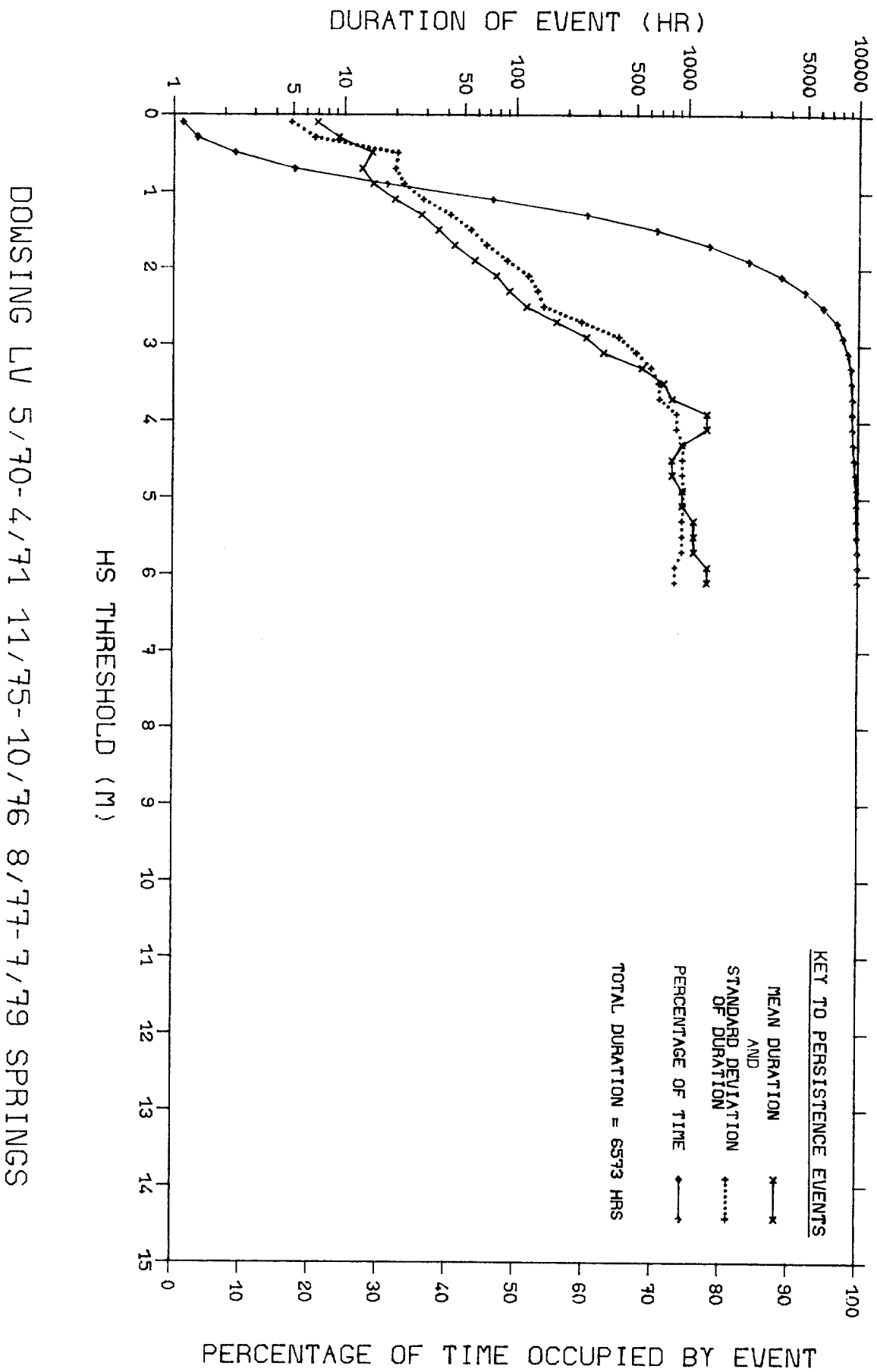
DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79 - WINTERS

FIG 3.5.1.4



DOWSING LV 5/70-4/71, 11/75-10/76, 8/77-7/79

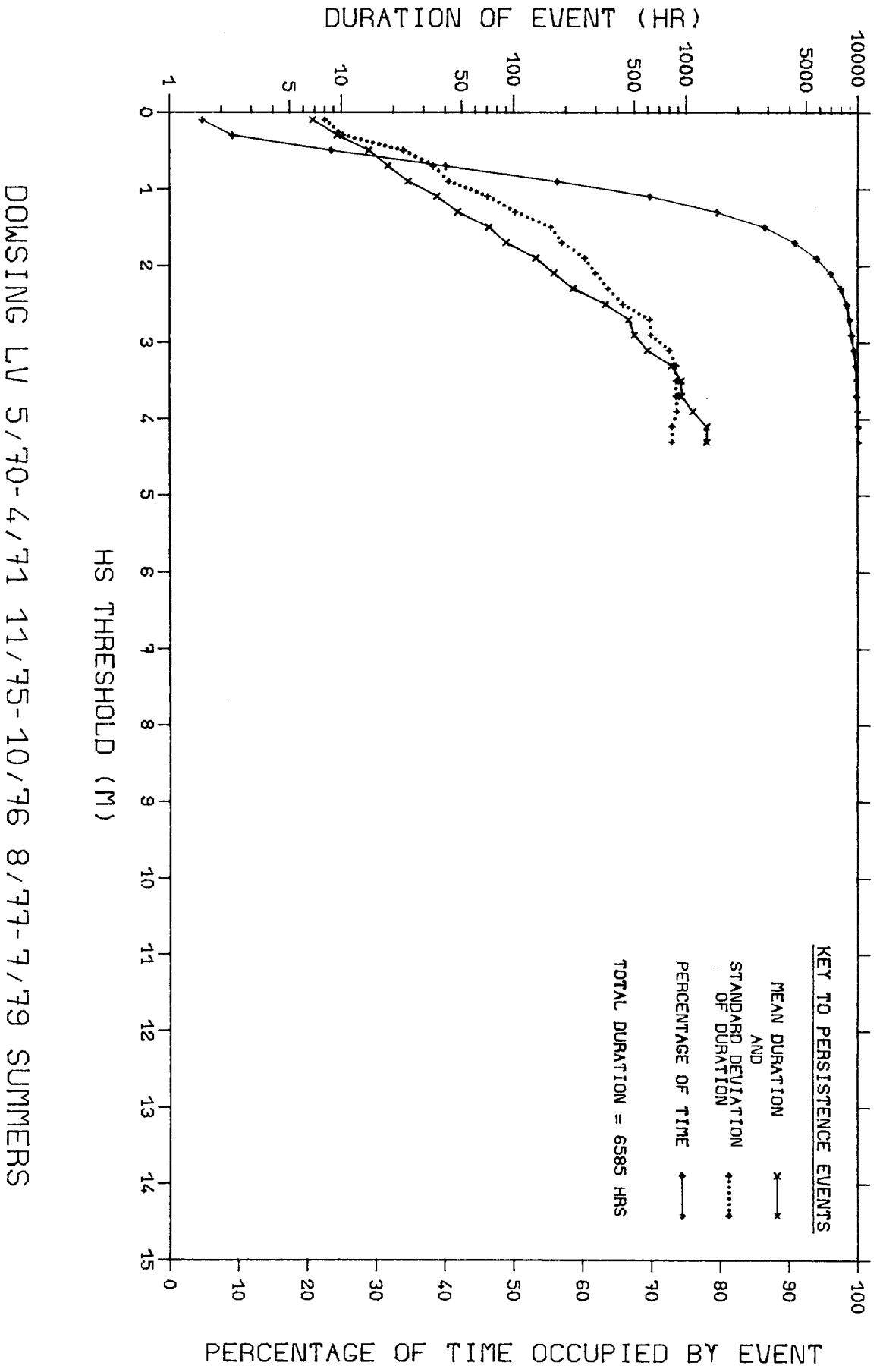
FIG 3.5.1.5



DOMSING LV 5/70-4/71 11/75-10/76 8/77-7/79 SPRINGS

CALMS

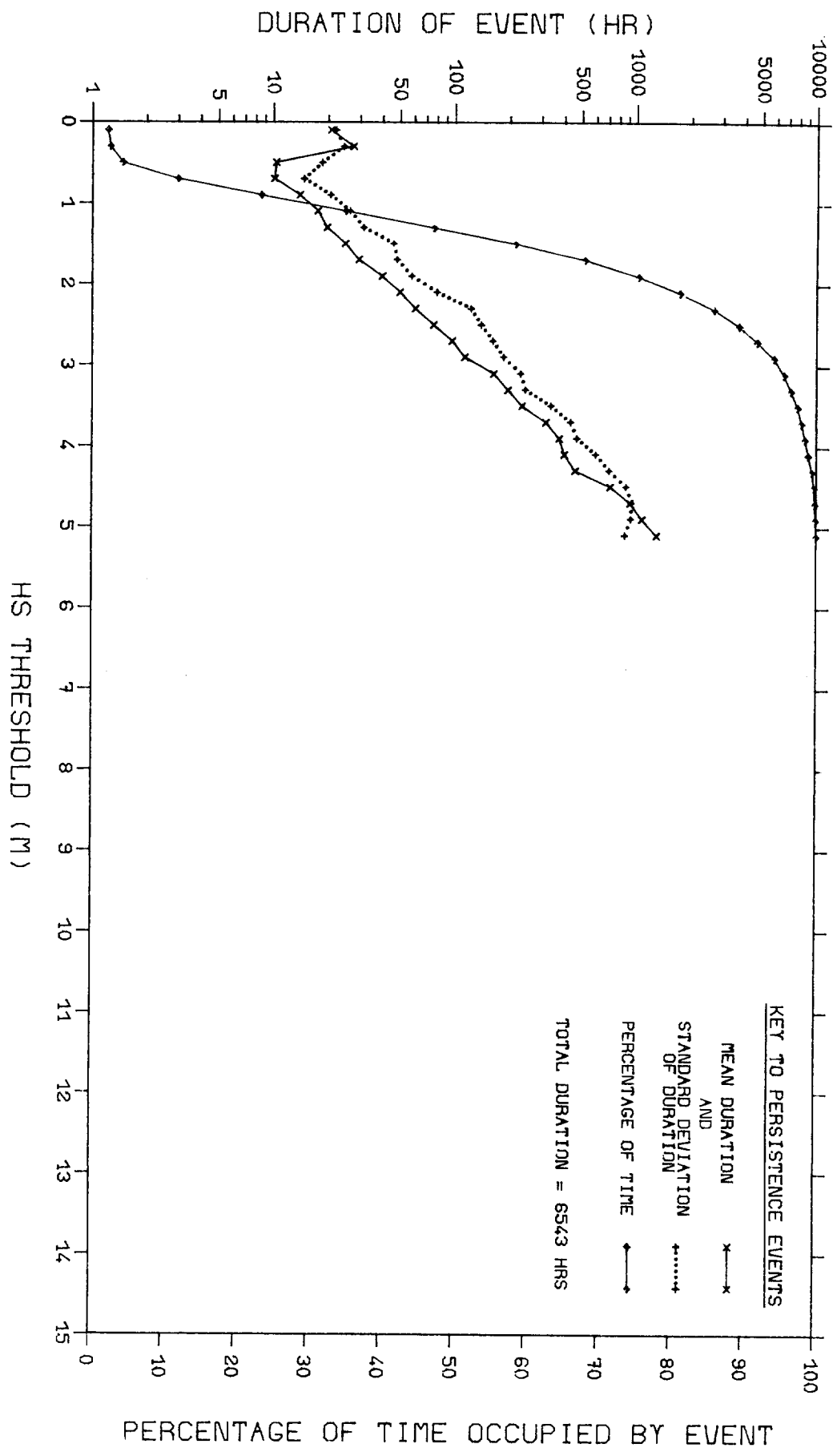
FIG 3.6.1.1



DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79 SUMMERS

CALMS

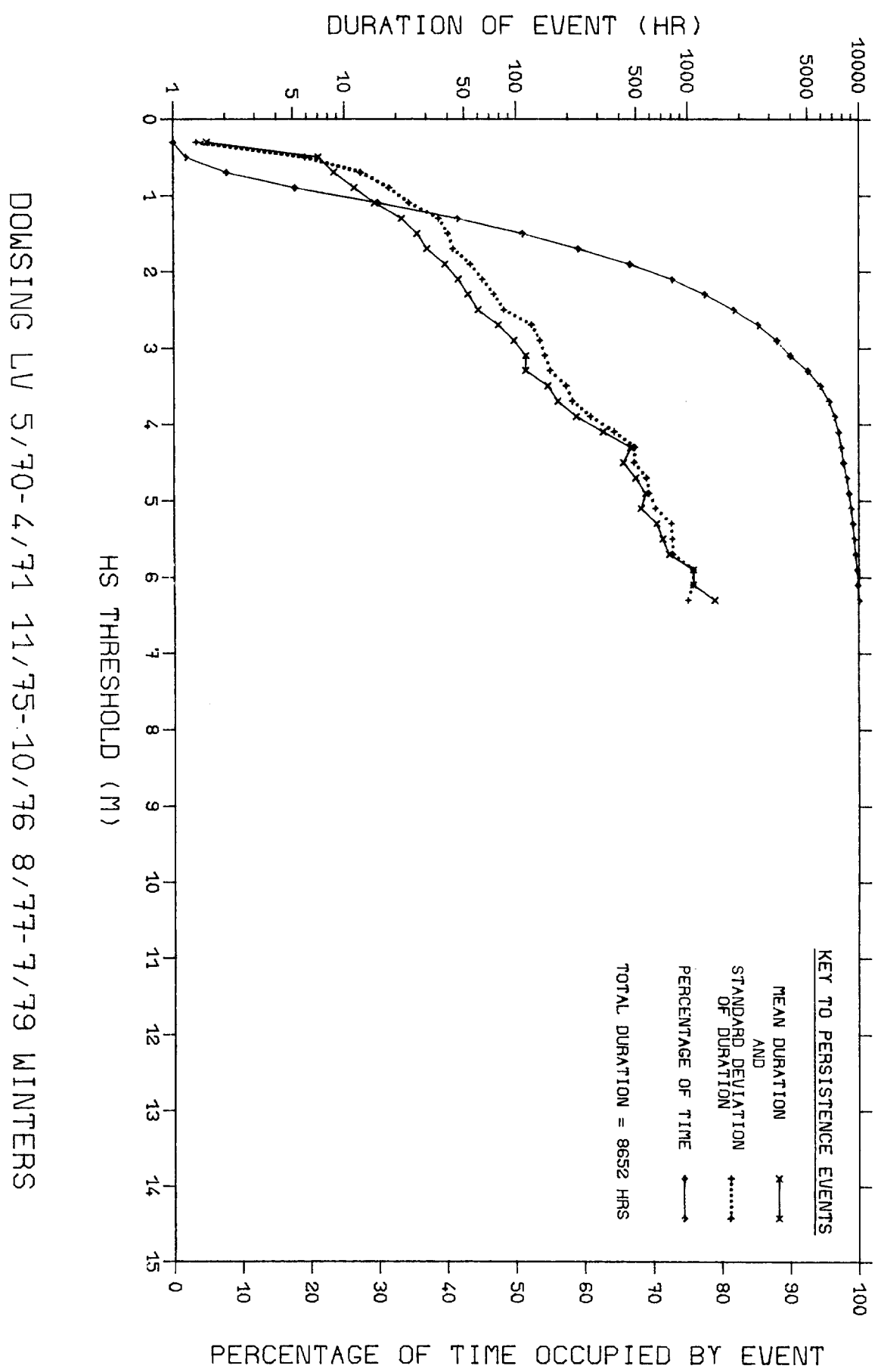
FIG 3.6.1.2



DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79 AUTUMNS

CALMS

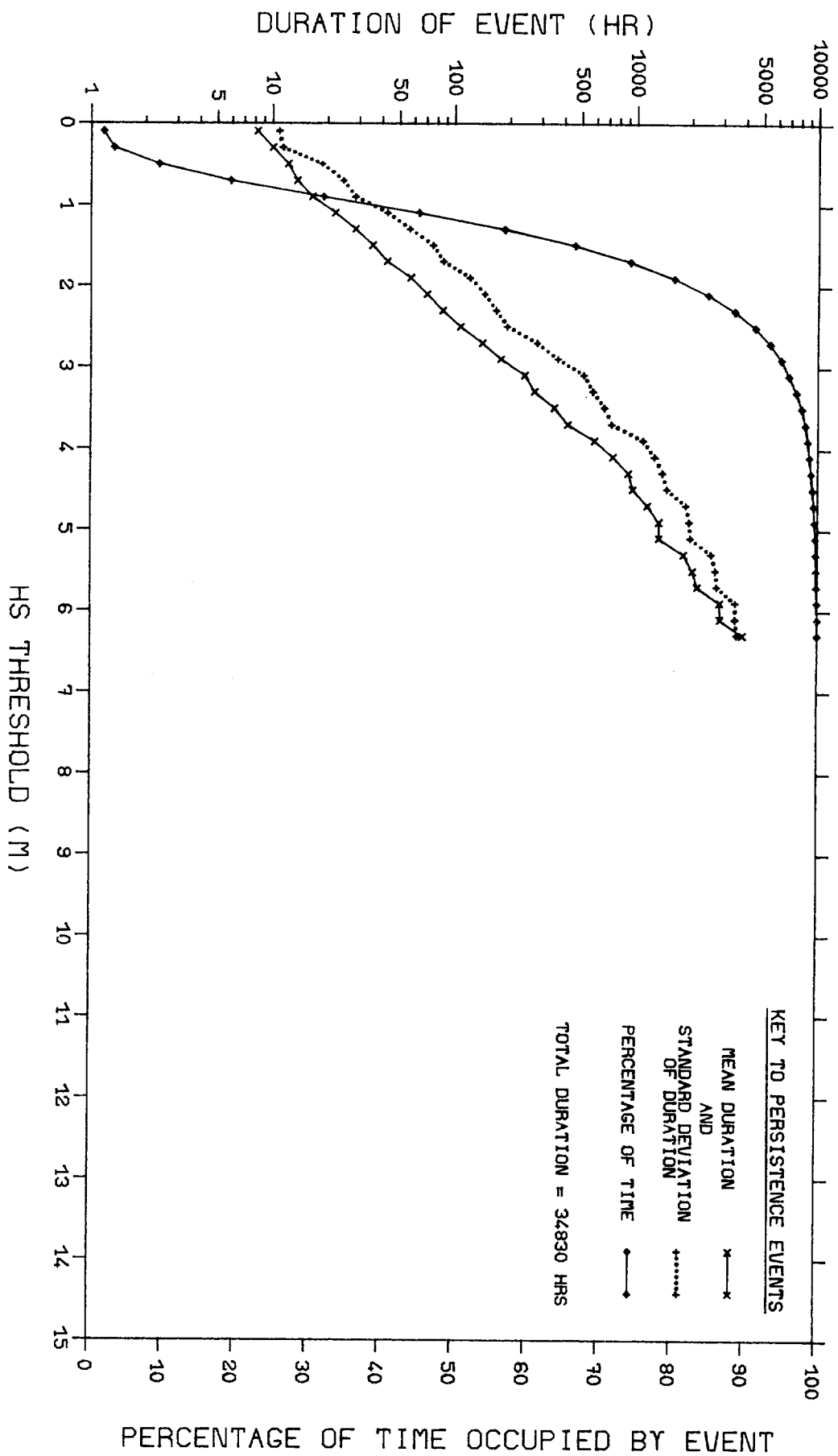
FIG 3.6.1.3



DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79 WINTERS  
 CALMS

FIG 3.6.1.4

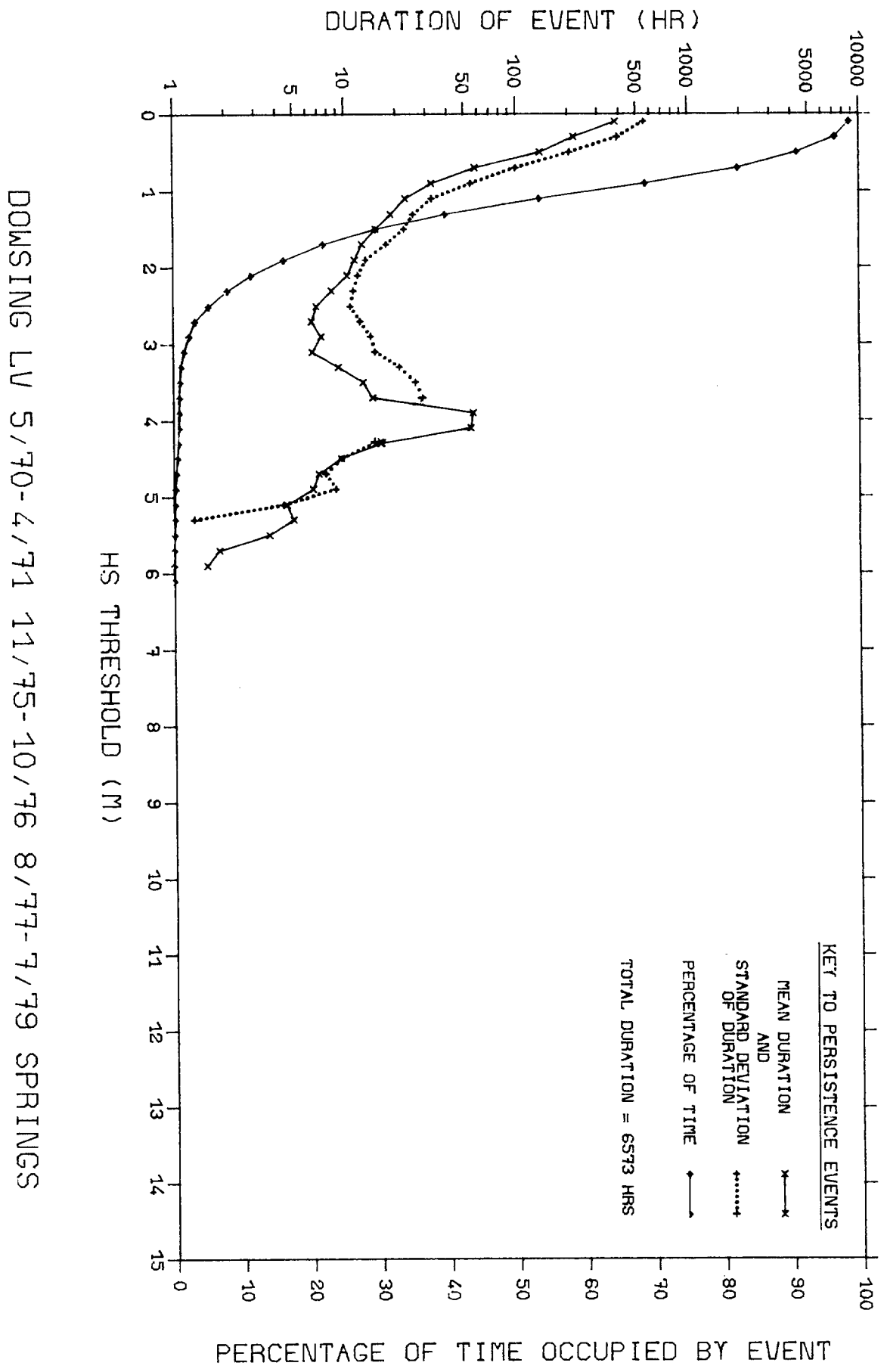




DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79

CALMS

FIG 3.6.1.5



DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79 SPRINGS STORMS

FIG 3.6.2.1

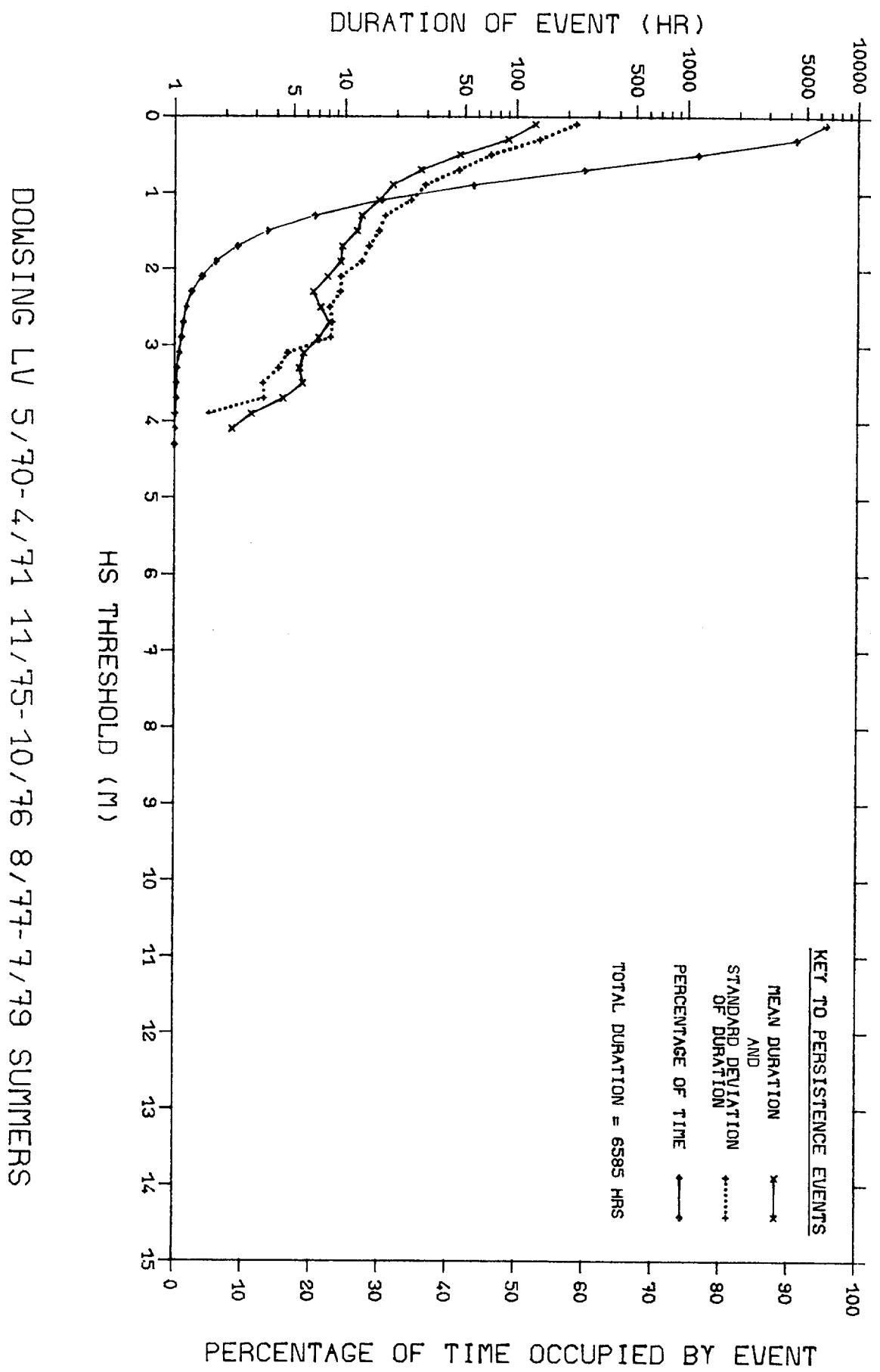


FIG 3.6.2.2

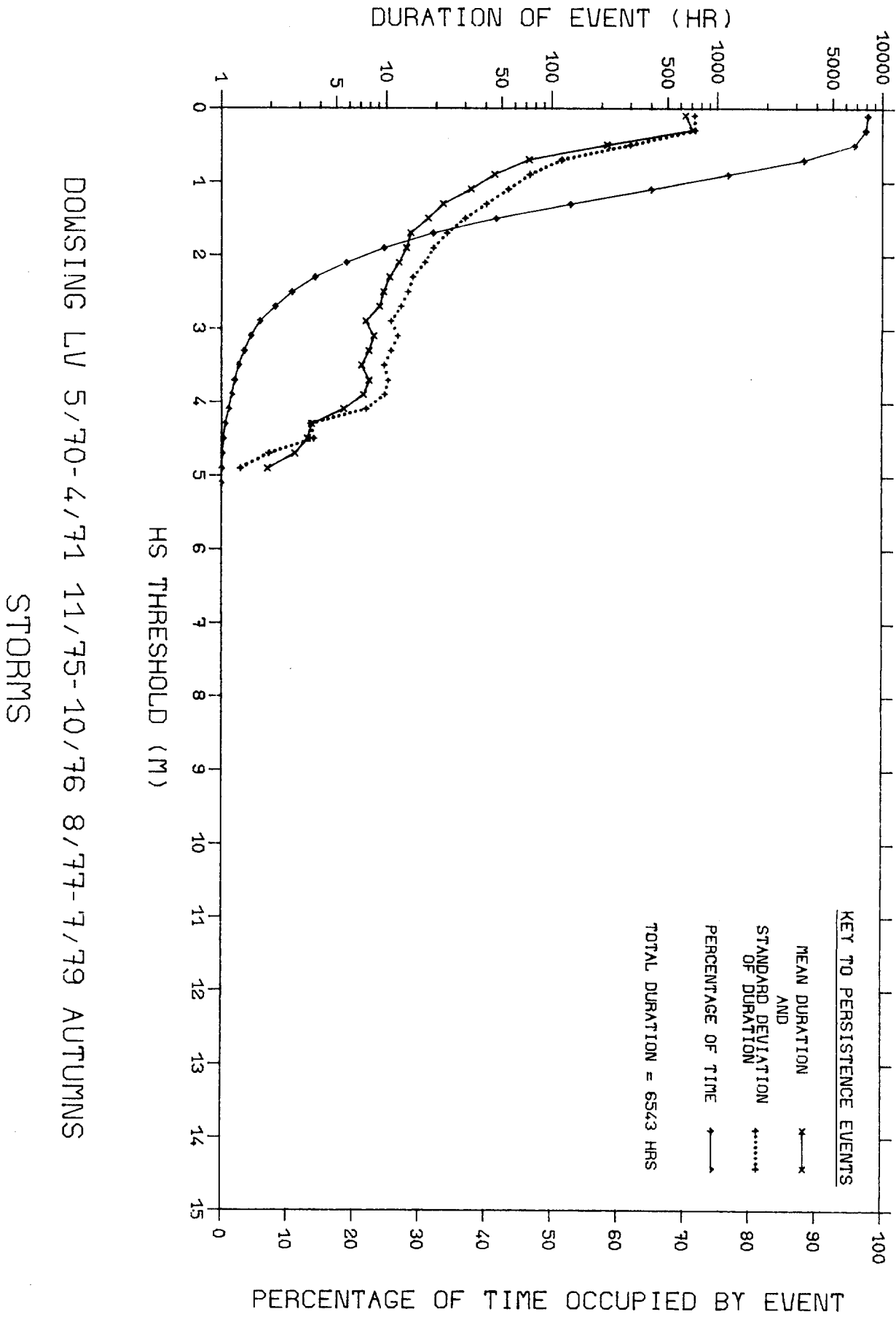
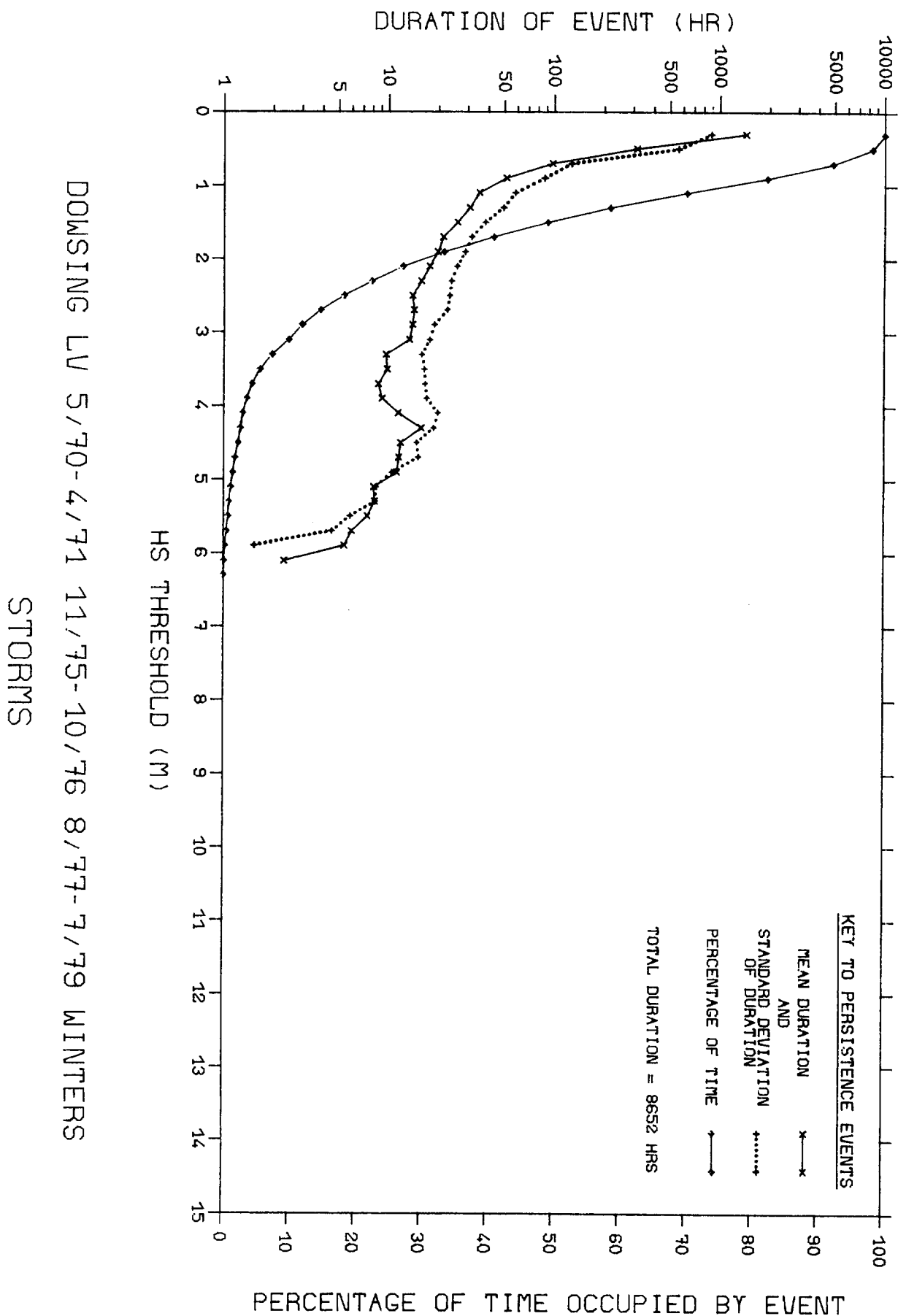


FIG 3.6.2.3



DOWNSING LV 5/70-4/71 11/75-10/76 8/77-7/79 WINTERS

STORMS

FIG 3.6.2.4

