Distribution of local $^{137}$Cs anomalies on the seafloor near the Fukushima Dai-ichi Nuclear Power Plant

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1. Introduction

The melt down at Fukushima Dai-ichi Nuclear Power Plant (F1NPP) resulted in radioactive material being released into the environment, with an estimated $3.5 \pm 0.7 \times 10^{15}$ Bq of $^{137}$Cs thought to have been discharged into the ocean following the melt down at Fukushima Dai-ichi Nuclear Power Plant (F1NPP). While efforts have been made to monitor seafloor radiation levels, the sampling techniques used cannot capture the continuous distribution of radionuclides. In this work, we apply in situ measurement techniques using a towed gamma ray spectrometer to map the continuous distribution of $^{137}$Cs on the seafloor within 20 km of the F1NPP. The results reveal the existence of local $^{137}$Cs anomalies, with levels of $^{137}$Cs an order of magnitude higher than the surrounding seafloors. The sizes of the anomalies mapped in this work range from a few meters to a few hundreds of meters in length, and it is demonstrated that the distribution of these anomalies is strongly influenced by meter scale features of the terrain.

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- Marine sediment
- Terrain features
- Towed gamma ray spectrometer
related to meter scale features of the seafloor terrain, and the concentrations of $^{137}$Cs are often more than an order of magnitude higher than in the surrounding regions. The existence of these anomalies should be taken into account when planning future survey efforts, and when considering the potential effects of $^{137}$Cs on marine ecology.

2. Instruments and methods

The instrument used in this work consists of a gamma ray spectrometer contained within a flexible rubber hose that is towed along the seafloor by a ship, as illustrated in Fig. 1 (Jones, 2001). The instrument, called the RESQ hose (RESQ: Radiometric Environment Survey and Quantification), is 8 m long with an external diameter of 0.145 m and weighs 135 kg in air and 115 kg in water. The compact size of the system allows it to be deployed from small fishing vessels and other ships of opportunity, which is particularly important for measurements in shallow waters near the shore that cannot be accessed using large research vessels. During operation, the system is attached to a wire that is used to lower it to the seafloor. Fig. 2 shows the device being lowered into the sea during a survey off Fukushima. The system has an internal battery that allows for up to 24 h continuous operation, and a data logging device that records the measurements of a depth sensor and a NaI(Tl) gamma ray scintillation spectrometer. The spectrometer has been calibrated to measure the gamma ray spectrum between 0.1 and 1.8 MeV over 1024 channels, and has a resolution of 6.9% at 0.662 MeV. The devices are covered using a rubber hose designed to reduce the risk of snagging, and provide protection from abrasion and impact damage during towing and handling on board the ship’s deck, while maintaining enough flexibility for the system to follow the undulations of the seafloor. The system is towed at velocities of between 2 and 3 knots and can be operated at depths of up to 500 m.

The device was deployed during 4 cruises between November 2012 and February 2013 to measure over 140 km of continuous radionuclide distribution along 10 transects within a 20 km radius of F1NPP, shortly after the lifting of government restrictions on access to the area on August 10 2012 (MEXT, 2012). Over 113,000 seafloor gamma spectra were measured at a sampling rate of 1 Hz. The data has been quantified, geo-referenced and smoothed using the methods described by the authors in Thornton et al. (2013). The levels of $^{137}$Cs have been determined through simulation using a Monte Carlo radiation transport model that computes the average concentration of the top 3 cm of the surface sediments, in accordance with sampling surveys (Kusakabe et al., 2013), based on the range of sediment types given in Table 1.

3. Results

Fig. 3 shows the continuous distribution of $^{137}$Cs measured in Bq/kg (wet weight), where the colors indicate the mean values for the range of sediments modeled. The spatial resolution of the map has been optimized to satisfy a 1σ statistical measurement uncertainty of 5% of the measured value at each point. This is achieved using an inverse distance weighted window function with a 100 m limit imposed on the minimum resolution of the map, beyond which measurement uncertainty is allowed to increase. In areas with high levels of $^{137}$Cs, the resolution of the map increases accordingly, where average $^{137}$Cs levels of 250, 500, and 1000 Bq/kg would lead to resolutions of about 76, 38, and 19 m, respectively, with some variation depending on the local distribution of $^{137}$Cs. The measurements show that the levels of $^{137}$Cs are relatively high within 4 km of the coastline, averaging 292 Bq/kg ($\sigma_v = 351$ Bq/kg), where $\sigma_v$ is the standard deviation of the measurements made in the area. Over 65% of the measurements in this region yield >200 Bq/kg, where the levels south of

![Fig. 1. Illustration of a ship towed gamma ray spectrometer (adapted from Jones (2001)).](image)

![Fig. 2. Deployment of the towed gamma ray spectrometer off Fukushima.](image)

![Table 1](image)

<table>
<thead>
<tr>
<th>Sediment type</th>
<th>$\rho_{\text{bulk}}$ (g/cm$^3$)</th>
<th>Water content (% by mass)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Coarse sand</td>
<td>2.03</td>
<td>19.5</td>
</tr>
<tr>
<td>Fine sand</td>
<td>1.98</td>
<td>22.7</td>
</tr>
<tr>
<td>Very fine sand</td>
<td>1.91</td>
<td>25.4</td>
</tr>
<tr>
<td>Silty sand</td>
<td>1.83</td>
<td>29.6</td>
</tr>
<tr>
<td>Sandy silt</td>
<td>1.56</td>
<td>44.8</td>
</tr>
<tr>
<td>Sandy-silt-clay</td>
<td>1.58</td>
<td>43.7</td>
</tr>
<tr>
<td>Clayey silt</td>
<td>1.43</td>
<td>53.8</td>
</tr>
<tr>
<td>Silty clay</td>
<td>1.42</td>
<td>54.9</td>
</tr>
<tr>
<td>Inorganic (% by mass)</td>
<td>SiO$_2$: 59.86, Al$_2$O$_3$: 12.99, Fe$_2$O$_3$: 4.42, CaO: 2.14, MnO: 1.03, P$_2$O$_5$: 0.13</td>
<td></td>
</tr>
<tr>
<td>Organic (% by mass)</td>
<td>C: 5.28, N: 0.57, H: 0.75, O: 3.01</td>
<td></td>
</tr>
</tbody>
</table>
**Fig. 3.** $^{137}$Cs levels on the seafloor within 20 km of the F1NPP. Measurements totaling over 140 km in length were made using a towed gamma ray spectrometer that was deployed during 4 cruises (6 days of measurement) between November 2012 and February 2013. The levels are determined for the top 3 cm of the surface sediments, where the average of the values calculated for the sediment types shown in Table 1 are taken, and are indicated by the color bar in the top right of the figure. Several of the $^{137}$Cs anomalies (shown in Fig. 4) exceed the upper range of the color bar. The regions labeled A–D correspond to the profile plots in Fig. 5.

**Fig. 4.** Distribution of local $^{137}$Cs anomalies on the seafloor. The locations where the levels of $^{137}$Cs are a factor of 5 and a factor of 10 higher than the average of measurements made within a 2 km radius of each point are shown in orange and red, respectively. The regions highlighted with a black border correspond to anomalies where the maximum levels of $^{137}$Cs measured are >5000 Bq/kg. The anomalies along the measured transects vary from a few meters to several hundreds of meters in length, and their distribution is strongly influenced by features of the underwater terrain.
the plant average 326 Bq/kg ($\sigma_v = 441$ Bq/kg), which is slightly higher and more variable than the levels north of the plant, which average 249 Bq/kg ($\sigma_v = 168$ Bq/kg). The highest levels of $^{137}$Cs were recorded 1–2 km south of the plant with an average of 438 Bq/kg ($\sigma_v = 867$ Bq/kg). The large values of the standard deviations illustrate the strong variations in the levels of $^{137}$Cs observed. The $^{137}$Cs levels decrease further out from shore averaging 69 Bq/kg ($\sigma_v = 73$ Bq/kg) between 4 and 12 km from the coastline, with less than 3% of the measurements yielding >200 Bq/kg. The highest levels of contamination this distance from the shore average 128 Bq/kg ($\sigma_v = 73$ Bq/kg) between 8 and 10 km south of the plant. Beyond 12 km, the levels of $^{137}$Cs increase to average 144 Bq/kg ($\sigma_v = 163$ Bq/kg), with over 20% of the measurements yielding >200 Bq/kg. The highest $^{137}$Cs levels at this distance are between 0 and 4 km north of F1NPP, averaging 218 Bq/kg ($\sigma_v = 270$ Bq/kg). The observation that the concentrations of $^{137}$Cs near the shore are higher south of the plant is consistent with sampling surveys and may be related to the high concentration of $^{137}$Cs in seawater that flowed south from the plant following the accident (Kawamura et al., 2011; Masumoto et al., 2012; Miyazawa et al., 2012). The distribution further out to sea is also consistent with the results of sampling surveys, and is thought to be a function of the types of marine sediment found on the seafloor. The area up to 12 km from the shore is dominated by rocky outcrops (Fukushima Prefecture, 1996; Aoyagi and Igarashi, 1999), and the areas further out consist mainly of fine silty clays, which cesium has a high affinity for (Lieser et al., 1986; Lieser and Steinkopff, 1988; Creemers et al., 1988; Cornell, 1993; Boretzen and Salbu, 2002; IAEA, 2004).

While the measurements are consistent with the findings of sampling surveys, they also reveal the existence of a number of local anomalies in the levels of $^{137}$Cs, which to date have not been captured by sampling. Fig. 4 shows the locations where the levels of $^{137}$Cs are a factor of 5, and a factor of 10 higher than the average values of measurements made within a 2 km radius of each point. Although these anomalies account for only 0.9% of the measurements made, 30% of these measurements have $^{137}$Cs levels >1000 Bq/kg, and all measurements >1000 Bq/kg in this work were made in these anomalies. The size of the anomalies varies from a few meters to several 100 m in length, and their distribution is strongly influenced by local features of the terrain. Anomalies have been consistently found at the bases of vertical features of the terrain, as seen in the examples in Fig. 5, which show the levels of $^{137}$Cs measured together with the depth of the seafloor (the vertical axis of the depth profiles has been exaggerated for clarity of presentation). The margins represent the range of possible levels of $^{137}$Cs taking into account the effects of different types of sediment.

Fig. 5. Profiles of $^{137}$Cs measurements made in the regions labeled A–D in Figs. 3 and 4. The shaded regions in the top graphs represent the range of uncertainty in the $^{137}$Cs levels, taking into account the combined effects of different types of seafloor sediments on detector response and the statistical uncertainty in the measurements of the detector itself. The dark lines show the mean values, which are plotted in the color maps that show the depth profiles where the measurements were made. The vertical axis of the depth profiles has been exaggerated for clarity of presentation. A strong influence of terrain features on the distribution of $^{137}$Cs is observed.
on the measurements of the detector, determined through Monte Carlo simulations, and the standard deviation in the measurements of the detector itself, as described in our previous work (Thornton et al., 2013). The upper bounds represent the levels determined for silty clays, and the lower bounds represent the levels determined for coarse sands, where all other sediment types modeled lie between these two boundaries. The dark lines represent the mean values for the range of sediments modeled, and correspond to the values shown in Fig. 3. The areas of seafloor shown in Fig. 5A and B both have high levels of $^{137}$Cs recorded at the bases of vertical terrain features. The region in Fig. 5A, labeled A in Figs. 3 and 4, is an 8 m high southward facing feature of the terrain located 5.9 km from shore and 3.7 km north of F1NPP. While the levels of $^{137}$Cs on top of the feature average $65 \pm 9$ Bq/kg (where the range of values represents measurement uncertainty), the average level of $^{137}$Cs measured at its base within 320 m of the feature is $524 \pm 63$ Bq/kg, with a maximum value of $985 \pm 118$ Bq/kg in this patch. Another anomaly was mapped a few 100 m further on from the feature. The patch is 70 m in length, and averages $651 \pm 77$ Bq/kg with a maximum $^{137}$Cs level of $1,432 \pm 173$ Bq/kg. The results show that the terrain strongly influences the level of $^{137}$Cs, with more than an order of magnitude difference in the levels measured on top and at the base of the feature. Similar observations were made for the seafloor shown in Fig. 5B, which is a 3.5 m high feature located 3.2 km east of F1NPP. This region serves as a boundary for the radiation levels, with the seafloor on the plant side of the feature, 1.7–3.2 km from the plant along this transect averaging $446 \pm 62$ Bq/kg, and the levels on the other side of the feature, 3.4–4.9 km from the plant yielding an average of $133 \pm 17$ Bq/kg. The level of $^{137}$Cs at the base of the feature has a maximum value of $2276 \pm 266$ Bq/kg, with an average of $1534 \pm 175$ Bq/kg over the 70 m long patch. The seafloor in Fig. 5C shows an anomalous depression located 10.3 km east and 0.7 km north of the F1NPP. The highest level of $^{137}$Cs in this patch is $1190 \pm 136$ Bq/kg, with an average of $508 \pm 58$ Bq/kg over the 105 m length of the anomaly measured along the transect. Here the size of the depression bounds the dimensions of the anomaly, and it is clear that features of the terrain influence not only the distribution, but also the size of the anomalies identified in this work.

In addition to terrain related anomalies, anomalies have also been identified in areas with no noticeable features of the seafloor. The seafloor shown in Fig. 5D, located 1.6 km east of the F1NPP between 0.2 km and 2.2 km south of the plant, has particularly high levels of $^{137}$Cs averaging $528 \pm 67$ Bq/kg. The highest level of $^{137}$Cs recorded in this region is $40,152 \pm 3998$ Bq/kg, with two other locations nearby where the levels of $^{137}$Cs measured are $>5000$ Bq/kg. These anomalies are extremely local with only a few meters of the seafloor affected along the measured transects. While measurements $>5000$ Bq/kg account for only 0.0012% of the measurements made in this work, the possibility of high concentration particulate matter being present in these areas must be considered. It is clear that extensive sampling of the identified regions is necessary to determine the cause of these anomalies.

4. Discussion

While it has been demonstrated that the well documented affinity of $^{137}$Cs to fine-grained sediments determines the overall distribution of $^{137}$Cs on the seafloor (Otosaka and Kobayashi, 2013), it has also been pointed out that sediment mineralogy alone cannot completely account for the spatial distributions observed along the east coast of Japan (Kusakabe et al., 2013). With regard to this point, the influence of the original distribution of $^{137}$Cs in the water column has been identified as a potential cause by Oikawa et al., (2013), who describe a scenario for rapid downward migration of $^{137}$Cs in the water column. While the majority of $^{137}$Cs in the water column is known to be in the form of dissolved ions (Stanners and Aston, 1981; Nies et al., 1991; Knapinska-Skiba

![Fig. 6. A conceptual model of the mechanisms thought to be responsible for the terrain related anomalies identified in this work. (A and C) The illustrations represent a snapshot of the situation immediately after the input of $^{137}$Cs, showing locally homogenous (on the scale of a few 100 m) distribution of $^{137}$Cs bound to particles in the water column, which sink to the seafloor due to gravity processes. (B and D) The illustrations show the same area some years later, i.e. when measurements were made. Sediments with high concentrations of $^{137}$Cs have been dispersed in areas exposed to underwater currents, whereas in areas shielded from underwater currents by features of the terrain the high concentration sediments remain locally confined. The depth profile in the illustration has been exaggerated for clarity of presentation.](image-url)
et al., 1994; Lujaniene et al., 2004, 2010), it has been shown that once $^{137}$Cs is incorporated into particulate matter, it is rapidly removed from seawater to the bottom sediment with settling velocities ranging from 29 to 190 m day$^{-1}$ (Fowler et al., 1987; Kusakabe et al., 1988). The unusual sedimentary environment resulting from suspended load carried back from land by the tsunami may also have contributed to rapid removal of nuclides from seawater (Kusakabe et al., 2013).

Fig. 6 shows a conceptual model for the mechanisms thought to be responsible for the observed relation between features of the terrain and the high level $^{137}$Cs anomalies recorded in this work. Fig. 6A and C show two snapshots of the situation shortly after contamination, where the underwater currents flow in opposite directions normal to a vertical feature of the terrain. Field observations of currents at a depth of 20 m, 30 km off F1NPP by Miyazawa et al. (2012) indicate that the mean currents in the region are relatively weak with velocities typically less than 0.4 m/s. Diurnal cycling of the currents along the north–south direction occurs due to wind effects, and simulations performed in their study demonstrate that tidal currents and river discharge flows also have a moderate impact on the transport of dissolved $^{137}$Cs. Since rapidly sinking particles are thought to be responsible for transport of $^{137}$Cs from the water column to the sediments (Fowler et al., 1987; Kusakabe et al., 1988; Oikawa et al., 2013; Kusakabe et al., 2013), it is reasonable to assume that the horizontal distribution of sinking $^{137}$Cs particles in the water column would be relatively homogeneous over the scale of a few 100 m at any moment in time. Since the downward flux of sinking particulate matter is largely dependent on the dimensions of the particles and gravity processes (Gardner and Walsh, 1990; Klaas and Archer, 2002; Buesseler et al., 2007), it follows that the distribution of contaminated particles reaching the seafloor at any moment would also be relatively homogenous. Fig. 6B and D show the corresponding situation 2 years later, roughly when the observations were made. In the illustrations, the contaminated particulate matter in the water column has subsided, and contaminated sediments in areas exposed to underwater currents have been remobilized and dispersed (Otosaika and Kobayashi, 2013). In areas where the seafloor is shielded from currents by the terrain, even though particle re-suspension would allow for some vertical and horizontal mixing (Gardner et al., 1985), the range of horizontal motion would be limited, with a tendency for pockets of contaminated fine-grained sediments to remain confined due to the energy lowering effects of the terrain and its influence on the local patterns of flow (Kennish, 2001). While it is necessary to verify the model through analysis of sediments sampled in the affected areas, the implications of the model are that the levels of $^{137}$Cs in these anomalies are likely to remain relatively unchanged over the timescales of a few years due to the effects of the local terrain on sediment transport. The influence of such features of the terrain should be considered together with other factors that can influence the distribution of $^{137}$Cs in the marine environment, such as secondary contamination from ground water and river inlets (Yoshida and Kanda, 2012; Nagao et al., 2013).

5. Conclusions

The measurements made in this work have revealed the existence of several $^{137}$Cs anomalies on the seafloor within 20 km of F1NPP. A strong correlation between the size and distribution of anomalies and features of the terrain has been demonstrated, with anomalies consistently found at the bases of vertical features of the terrain where the pockets of sediments are sheltered from underwater currents. It is clear from the results of this study that fine, meter scale features of the seafloor terrain play a significant role in determining the distribution of $^{137}$Cs on the seafloor within 20 km of the F1NPP.

Based on the size and distribution of the anomalies mapped in this work, it can be said that the density of sampling points required to survey this region effectively using a standard grid based approach would be impractical and the costs associated with such an effort would be prohibitive. It is clear that a more targeted approach to sampling based on prior screening using in situ measurement techniques is necessary. The approach described in this work should be combined with wide area acoustic surveys to determine the distribution of fine-grained sediments off F1NPP. Direct observations of the nature of the seafloor in the areas identified are necessary together with targeted sampling efforts in order to achieve a better understanding of the processes involved for the different types of anomaly identified in this work, i.e. (1) 100 m scale terrain related anomalies, and (2) more localized meter scale anomalies showing no correlation with features of the terrain. It is hoped that the results described can help focus future survey and recovery efforts, and so advance our understanding of the potential effects of the accident on the marine environment.

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