

Nutrient co-limitation at the boundary of an oceanic gyre

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Nutrient limitation of oceanic primary production exerts a fundamental control on marine food webs and the flux of carbon into the deep ocean¹. The extensive boundaries of the oligotrophic sub-tropical gyres collectively define the most extreme transition in ocean productivity, but little is known about nutrient limitation in these zones¹⁻⁴. We conducted full factorial nutrient amendment experiments in the eastern boundary of the South Atlantic gyre and found extensive regions where supplying nitrogen or iron individually resulted in no significant phytoplankton growth over 48 hours, but adding both increased chlorophyll-a concentrations by up to ~40-fold, led to diatom proliferation, and reduced community diversity. Once nitrogen-iron co-limitation had been alleviated, addition of cobalt or cobalt-containing vitamin B₁₂ could further enhance chlorophyll-a yields up to 3-fold. Our results imply nitrogen-iron co-limitation is pervasive in the ocean, with other micronutrients also approaching co-deficiency. Such multi-nutrient limitations potentially increase phytoplankton community diversity.

From the results of nutrient-enrichment experiments performed to date, oceanic phytoplankton would appear to be proximally limited by the availability of either nitrogen (N) or iron (Fe)¹. Despite widespread observations of both nutrients being at low concentrations simultaneously², relatively little direct evidence exists for co-limitation of phytoplankton growth by these elements^{3,4}. Furthermore, field evidence for (co-)limitation by micronutrients other than Fe is sparse^{5,6}. Characterization and even definition of nutrient ‘co-limitation’ can be complex⁷⁻⁹ (Supplementary Discussion). However, the simplest case corresponds to two strictly essential nutrients (e.g., N and Fe) being concurrently drawn down to levels where only the supply of both in combination results in a significant biomass growth response. Such ‘simultaneous co-limitation’ occupies a midpoint in resource ratio space relative to single limitation and serial (or secondary) limitation^{10,11}, the latter representing the circumstance where a second nutrient only becomes limiting following addition of the first.

Considerations of such transitions in resource space remain largely theoretical^{8,10,12,13}, limiting our understanding of (co-)limitation in nature. An additional factor complicating widespread predictions of oceanic (co-)limitation relates to

reconciliation of operationally defined dissolved seawater nutrient concentrations with flexible phytoplankton demands¹. When evaluated within an appropriate framework^{1,8}, the clearest means of demonstrating oceanic nutrient (co-)limitation patterns and the associated short-term ecophysiological responses to nutrient re-supply are via direct testing in trace-metal-clean nutrient amendment bioassay experiments conducted with a factorial design. However, the logistical challenges associated with this approach have limited applications to few studies employing more than two nutrients¹.

To resolve potential (co-)limitation of phytoplankton communities by the three nutrients identified as most deficient in the South Atlantic gyre¹, we conducted 48 hour duration full-factorial N, Fe, and cobalt (Co) addition bioassay experiments throughout the SE Atlantic. This region receives relatively little dust input and is host to a marked productivity transition between the eastern boundary Benguela upwelling regime, a globally important fishery, and the South Atlantic oligotrophic gyre (Fig. 1)¹⁴. To elucidate the potential biochemical function of added Co, an additional N+Fe+Co-containing vitamin B₁₂ amendment was also conducted. Experiments were carried out on the German GEOTRACES cruise GA08, in December 2015 (Fig. 1a), with surface seawater collected using a towed trace-metal-clean sampling system, and shipboard incubations performed in triplicate and interpreted relative to untreated controls and the initial biogeochemical characterization of ambient seawater. Phytoplankton responses to nutrient amendment were assessed via changes in chlorophyll-a concentrations, flow cytometry cell counts of key phytoplankton groups, concentrations of diagnostic phytoplankton pigments, and nutrient-stress-specific active chlorophyll fluorescence measurements^{15,16}.

Away from fully nutrient replete coastal upwelling waters, bulk phytoplankton community responses demonstrated transitions between N and Fe (single/serial/co-)limitations (Fig. 1 and Extended Data Fig. 1). Aside from the coastal sites (Experiments 1 and 11), chlorophyll-a increased at two sites following amendment with N alone (Experiments 3 and 4), three sites exhibited responses that were consistent with serial limitation by N and Fe (Experiments 2, 5 and 6), and four experiments (Experiments 7–10) showed increases that were only significant following amendment with N+Fe. Together, these results imply widespread conditions at or approaching N-Fe co-limitation. At the clearly co-limited sites, accumulation of larger cells (approximately >2 μm) only occurred following amendment with at least N+Fe (Fig. 2a and Extended Data Fig. 2). In contrast, average cell counts of the cyanobacteria *Synechococcus* and *Prochlorococcus* typically exhibited no changes or reductions following supply of N+Fe, suggesting they were grazer-regulated and/or out competed by the larger cells^{17,18}. However, magnitudes of cellular fluorescence, indicative of pigmentation per cell, generally increased with N or N+Fe amendment for the prokaryotes, suggesting physiological recovery from initial nutrient limitation despite limited biomass accumulation (Fig. 2b and Extended Data Figs 3–5)¹⁷.

Responses to N+Fe amendment were most pronounced at the sites with strongest N-Fe co-limitation (Experiments 8–10). Responses to addition of N or Fe alone were not statistically significant at these locations, whereas chlorophyll-a biomass increased 21–38 times that of control samples in response to addition of N+Fe (Fig. 1). Although dramatic, the responses to N+Fe amendment in three of the experiments remained modest in comparison to N+Fe+Co or N+Fe+vitamin B₁₂ (2 nmol L⁻¹ added

Co; 100 pmol L⁻¹ B₁₂), where up to an additional 2-to-3-fold increase in chlorophyll-a biomass was observed (Experiments 7–9). Importantly, enhanced chlorophyll-a resulted from higher phytoplankton abundances, rather than increased cellular pigmentation alone (Fig. 2a and Extended Data Figs 2–5). Thus, over large oceanographic extents there was a clear upper limit on potential biomass accumulation, for at least the larger-celled community (Fig. 2), upon supply of N+Fe relative to that achievable with additional supply of Co or vitamin B₁₂.

The delicate balance of N-Fe co-limitation was clear in responses of the community-level physiological indicator, F_v/F_m (Fig. 1 and Extended Data Fig. 6)¹⁶. Experiments 6 and 8–10 were at or approaching co-limitation and had elevated initial F_v/F_m , characteristic of either nutrient replete, proximally N-limited, or N-Fe co-limited systems¹⁶. N amendment at these sites (alone or in combination with Co, but in the absence of Fe) resulted in significant F_v/F_m reductions. Such reductions result from greater Fe stress¹⁶ and match responses observed at N-Fe co-limited sites in the Equatorial Pacific³. Diurnal F_v/F_m cycles, including marked nocturnal decreases (Extended Data Fig. 7), also generally matched those previously observed in the co-limited Pacific³. Conversely, at Experiment 7 (also co-limited) low initial F_v/F_m and increases following Fe amendment presumably represented recovery from proximal physiological Fe stress, despite N-Fe co-limitation of biomass accumulation.

Qualitative community level biomass and F_v/F_m responses to N and/or Fe amendment were both predictable on the basis of observed seawater nutrient (dissolved N and Fe) concentrations, as illustrated on resource ratio plots (Fig. 3a–c)^{10,11}. Despite acknowledged complexity in phytoplankton responses to nutrient amendment (Supplementary Discussion), biomass accumulation could also be quantitatively reproduced using a relatively simple (semi-)empirical model that assumed a closed system, typical phytoplankton quotas, and minimizing or multiplicative forms of the Michaelis-Menten growth equation (Fig. 3d–f)⁸. Both qualitative and quantitative categorization of (co-)limitation based on our experiments were also strongly related to the ratio of ambient N and Fe concentrations (Fig. 4a, b). Transitions between differing single/serial/co-limitation for Fe, N and Co were thus reconcilable with the large-scale biogeochemical gradients observed across 1000s of km of the surface ocean.

Our observations of widespread nutrient co-limitation suggest an interaction between the biogeochemical setting and the extant phytoplankton community. Simultaneous biological depletion of multiple nutrients provides a setting for co-limitation, and potentially drives a subsequent reinforcing biological response. All co-limited sites we identified were host to a diverse phytoplankton assemblage of both prokaryotes and eukaryotes (Fig. 2 and Extended Data Fig. 8), and although more oligotrophic, generally had higher diversity than sites under single N limitation (Fig. 1 and Extended Data Fig. 9). Enhanced diversity under resource co-limitation is predicted on the basis of increased niche dimensionality¹¹ and has been observed in both terrestrial and lake systems^{19,20}. Diverse communities are expected to tend towards community-level co-limitation because of differences across, and plasticity within, taxa for their stoichiometric requirement for the shared limiting nutrients^{11–13}. Reciprocally, environments where multiple nutrients are simultaneously low favour diversity by encouraging a spectrum of mechanisms for accessing each fraction of total nutrient pools¹¹. For example, following near-complete exhaustion of inorganic N and the most accessible dissolved Fe species, specialist acquisition strategies allow

149 progressive use of different chemical forms of these nutrients, including organically
150 bound pools or other physicochemical species. Thus alongside physical forcing²¹ and
151 top-down ecological control, gradients in heterogeneous nutrient pools at
152 biogeochemical transitions⁴ implicitly favour diversity and community nutrient co-
153 limitation^{12,13}. Consistent with this hypothesis, experimental amendment with N+Fe,
154 and more so N+Fe+Co (or vitamin B₁₂) significantly reduced diversity at the N-Fe co-
155 limited sites (Fig. 2a and Extended Data Fig. 9) presumably through reducing niche
156 dimensions to those favouring diatoms (Fig. 2c and Extended Data Fig. 8)²².

158 Addition of multiple nutrients within N-limited gyre systems typically produces less
159 dramatic chlorophyll-a increases than we observed¹⁵. Strong niche exclusion of
160 bloom-forming diatoms, tighter grazer control, and/or lower maximal growth rates of
161 extant populations may mute overall biomass increases to nutrient amendment within
162 these central gyre systems. In contrast, responses to N+Fe amendment closely
163 resembled diatom responses to Fe-only amendment within N replete ($>10 \mu\text{mol L}^{-1}$)²³
164 proximally Fe-limited ocean systems^{18,23}, where addition of Fe fulfils nutritional
165 requirements (i.e., Fig. 3b). Secondary chlorophyll-a biomass responses to Co
166 amendment have also been observed in some Fe-limited regions^{1,5}. Although likely
167 system-dependent, our observations show that overall biomass responses can also be
168 serially restricted by Co following addition of N+Fe alone at the boundaries between
169 N and Fe limited regions.

171 Like Fe, Co was relatively enriched in the ancient ocean in which algae evolved²⁴,
172 potentially contributing to its obligate requirement in many phytoplankton². However,
173 whilst the largest cellular sinks of N and Fe in phytoplankton are relatively well
174 established², greater ambiguity exists for Co. Two principle functions for Co in
175 phytoplankton have been elucidated: as a cofactor in carbonic anhydrase (CA), and
176 vitamin B₁₂, a cofactor in several enzymes². Less well characterized are possible roles
177 in phosphatases, acyltransferases and hydratases²⁵. Assigning a Co requirement to
178 specific biochemical roles is complicated as Zn or Cd can potentially substitute for
179 Co-CA², whilst many phytoplankton can grow without vitamin B₁₂ (Ref. 26), albeit at
180 some resource cost²⁷. Our results support Co responses linked to both B₁₂ and B₁₂-
181 independent roles (compare Experiments 5 and 10 in Fig. 1). However, statistically
182 indistinguishable chlorophyll-a responses and similar diagnostic pigment assemblages
183 between N+Fe+Co and N+Fe+B₁₂ amendments in Experiments 7–9 support a more
184 widespread vitamin B₁₂ role for the added Co within the most rapidly responding
185 taxa. This suggests tighter coupling between Co availability and vitamin B₁₂
186 production in the South Atlantic relative to previous observations in the coastal
187 Southern Ocean where B₁₂ additions, but not Co, were stimulatory⁶. Disassembly of
188 the supplied vitamin B₁₂, resulting in purposeful/inadvertent Co liberation and
189 subsequent incorporation into CA, cannot be ruled out. However, contrasting
190 responses between N+Fe+Co and N+Fe+vitamin B₁₂ in Experiment 5 at least
191 suggested this was not always the case, since Co additions stimulated additional
192 growth when added in combination with N+Fe, whereas B₁₂ did not. Thus, whilst
193 Co/vitamin B₁₂ availability clearly had a widespread impact on achievable biomass
194 yield, resolving the biochemical function of added Co and extrapolating observations
195 of such serial limitation to the in situ condition^{8–10}, remains difficult at this stage
196 (Supplementary Discussion).

198 Large-scale ocean circulation and biogeochemical interactions set the conditions for
199 spatial patterns of nutrient (co-)limitation in the ocean^{1,28}. Sub-surface ratios of two

200 nutrients, such as N and Fe, can thus provide a useful index for position in N-Fe
201 resource space²⁸. Deep waters feeding major upwelling zones have a high N:Fe ratio
202 and phytoplankton growth depletes Fe before N. In contrast, surface waters in the
203 cores of stratified gyre systems have a low N:Fe ratio, resulting from heavily
204 restricted N resupply and, presumably, input of Fe from aerosols. Transitions between
205 these regimes define a shift in resource ratio space (Figs 1b and 3a–c), and therefore
206 potential for N-Fe co-limitation. A previous study found overlap of N and Fe stress
207 biomarkers within *Prochlorococcus* ecotypes across a transition in N:Fe ratios in the
208 Pacific⁴ and our results suggest that, at the whole-community level, diversity in
209 phytoplankton requirements and a spectrum of acquisition strategies further broadens
210 the co-limitation zone^{8,12,13}.

211
212 In contrast to sub-surface ratios that partly dictate relative supply²⁸, assuming steady
213 state, ratios of measured residual N and Fe concentrations in the surface ocean reflect
214 the end point of biological uptake and hence the competition for these potentially
215 limiting resources (Supplementary Discussion). The range of N:Fe ratios measured at
216 experimentally-determined co-limited sites thus provide an empirical means for
217 predicting N-Fe co-limitation at large spatial scales (Fig. 4a). Surface nutrient fields
218 from a complex global biogeochemical model²⁹ predict only ~2% of the surface ocean
219 to fall within the stoichiometric range for N:Fe where we found direct evidence for N-
220 Fe co-limitation, with a further 12% predicted as serially limited and hence
221 approaching co-limitation, mostly distributed in between upwelling and gyre regions
222 (Fig. 4c). In contrast, analysis of the available surface ocean N and Fe data suggests
223 that co-limitation may actually be ~4-fold more prevalent than these model
224 predictions (Fig. 4c and Extended Data Fig. 10). Regions of co-limitation may thus
225 represent a key feature of low latitude ocean that is under-represented in the global
226 models we rely on for projecting the impact of climate change. The abrupt transitions
227 between N and Fe limitation that occur within current models likely reflect the
228 omission of sufficient diversity and physiological plasticity (e.g., related to variable
229 nutrient demands and acquisition traits) within simulated phytoplankton
230 communities³⁰.

231
232 Nutrient inputs to the ocean are projected to change¹. Modified aerosol inputs, altered
233 stratification and wind stress, and the redox status of the upwelling regimes
234 characterizing eastern boundary currents could all directly impact nutrient fluxes and
235 stoichiometry at gyre margins¹. We find that processes of co-limitation^{8,10}, by N and
236 Fe as well as additional nutrients such as Co^{6,8,26}, may be crucial in determining the
237 responses of phytoplankton community structure and productivity to such forcing,
238 particularly at regional scales. Accordingly, recognition of multi-nutrient serial/co-
239 limitation^{4,8,9} and better representation of the underlying processes within ocean
240 models will thus lead to more realistic projections of feedbacks regulating climate and
241 marine food webs.

Figure legends

Figure 1. Nutrient limitation in the SE Atlantic. **a**, Cruise track and locations of bioassay experiment sites. For scale, the distance between the north and south zonal transects is ~3,000 km. **b**, Section of interpolated N:Fe ratios measured on the CoFeMUG cruise (orange line in **a**)¹⁴. Dark blue=Fe deficient relative to N, white=near equal deficiency, red=N deficient. Large black and grey symbols indicate data within the range we found N-Fe co-limitation and secondary limitation respectively. **c–k**, Phytoplankton responses to nutrient amendment in Experiments 2–10. Dots indicate replicate treatment bottles; bar heights and lines indicate the mean and range, respectively (n=3). Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, n=3). Horizontal lines indicate initial values. Amendment label colour indicates F_v/F_m was significantly increased (red) or reduced (blue) relative to the control (ANOVA and Tukey HSD $p \leq 0.05$, n=3). Co saturation of cation transporter sites in Experiment 8 could have induced Fe stress and the F_v/F_m reduction².

Figure 2. Example ecophysiological responses to nutrient amendment at a N-Fe co-limited site (Experiment 8). **a**, Flow cytometry cell counts (relative units—mean counts, n=3, have been normalized to control for each cell type). Nano=nanophytoplankton (approximately $>2 \mu\text{m}$); Pico=picophytoplankton (approximately $<2 \mu\text{m}$); *Syn*=*Synechococcus*; *Pro*=*Prochlorococcus*. The superimposed scatter plot is the Exponential Shannon Wiener diversity Index (ESWI)²⁵, calculated using cell counts (dots and line) or pigment-derived community (crosses; see **c**). **b**, Mean fluorescence per cell (relative, as in **a**). **c**, Pigment-derived taxonomic contributions to total chlorophyll-a in ambient waters ('Initial') and after selected nutrient amendments. Percentage contributions of diatoms are labelled.

Figure 3. Factorial nutrient limitation scenarios. **a–c**, Initial position of seawater nutrient concentrations and movement in resource space following experimental amendments (symbols defined in Fig. 1a). Background colour represents growth rate predicted using a minimizing Michaelis-Menten equation^{8,10} (darker green=higher growth rate). Concentrations at nutrient limited sites generally follow a theoretical interspecific N-Fe tradeoff curve^{11,12}. Dashed line indicates the theoretical transition from N to Fe limitation using assumed average phytoplankton quotas¹. Solid lines define the envelope of N:Fe ratios where we found simultaneous N-Fe co-limitation. Experiment 7 is here classified as serially Fe-N limited, given the physiological response to Fe supply (Fig. 1h). The all-red dot represents nutrient concentrations measured at a high N, Fe-limited experimental site in the Equatorial Pacific, completing the N-Fe limitation sequence observed in the SE Atlantic (Supplementary Table S1). **d–f**, Simulated nutrient utilization with phytoplankton stimulation using assumed-average phytoplankton nutrient quotas, maximum growth rates of 2.5 day^{-1} , half-saturation concentrations (for growth) set to $0.25 \mu\text{mol L}^{-1}$ for N and scaled for Fe and Co using average quotas, and a factor of 150 to convert phytoplankton carbon to chlorophyll-a. **d–e**, example simulated drawdown and chlorophyll increases compared to measured chlorophyll-a concentrations at 48h (symbols representing individual bottle replicates, n=3, and line indicating the range). **f**, Predicted vs. measured growth for all experiment simulations using same parameterizations. Dotted line=1:1 and solid line indicates the least squares linear regression (P value for two-tailed F test).

Figure 4: Predicting oceanic co-limitation using N:Fe ratios. **a**, discrete categorization of limitation and **b**, continuous scale based on a derived '(co-)limitation index' (Supplementary Discussion), as a function of observed N:Fe concentrations. Point symbols are as defined in Figures 1 and 3. Lines indicate the least squares linear regression (P value for two-tailed F test). **c**, Global co-limitation prediction using N:Fe generated by a biogeochemical model with available observational data over plotted. Grey and yellow grid cells/observations indicate data within the serially/co-limited N:Fe range represented in **a** and **b**. Thresholds where N and Fe concentrations are both characterized as replete, regardless of the N:Fe ratio, have been applied (maximum concentrations measured under co-/serial limitation). For all panels the units of calculated N:Fe are $\mu\text{mol}:\text{nmol}$.

Data availability

All data from the current study are available from the corresponding author on reasonable request.

1. Moore C. M. *et al.* Processes and patterns of oceanic nutrient limitation. *Nat. Geosci.* **6**, 701–710 (2013).
2. Morel, F. M. M., Milligan A. J., & Saito, M. A. “Marine Bioinorganic Chemistry” in *The oceans and marine geochemistry*, (Treatise on Geochemistry **6**, 2003).
3. Behrenfeld, M. J. *et al.*, Controls on tropical Pacific Ocean productivity revealed through nutrient stress diagnostics. *Nature* **442**, 1025–1028 (2006).
4. Saito, M. A. *et al.*, Multiple nutrient stresses at intersecting Pacific Ocean biomes detected by protein biomarkers. *Science* **345**, 1173–7 (2014).
5. Saito, M. A., Rocap, G. & Moffett J. W., Production of cobalt binding ligands in a *Synechococcus* feature at the Costa Rica upwelling dome. *Limnol. Oceanogr.* **50**, 279–290 (2005).
6. Bertrand E. M. *et al.*, Vitamin B12 and iron colimitation of phytoplankton growth in the Ross Sea. *Limnol. Oceanogr.* **52**, 1079–1093 (2007).
7. Arrigo, K. R., Marine microorganisms and global nutrient cycles. *Nature* **437**, 343–348 (2005).
8. Saito, M. A., Goepfert, T. J. & Ritt, J. T. Some thoughts on the concept of colimitation: Three definitions and the importance of bioavailability. *Limnol. Oceanogr.* **53**, 276–290 (2008).
9. Harpole W. S. *et al.*, Nutrient co-limitation of primary producer communities. *Ecol. Lett.* **14**, 852–862 (2011).
10. Sperfeld, E., Raubenheimer, D. & Wacker A., Bridging factorial and gradient concepts of resource co-limitation: Towards a general framework applied to consumers. *Ecol. Lett.* **19**, 201–215 (2016).
11. Tilman, D. *Resource competition and community structure* (Princeton Univ. Press, 1982).
12. Schade, J. D. *et al.*, A conceptual framework for ecosystem stoichiometry: balancing resource supply and demand. *Oikos* **109**, 40–51 (2005).
13. Danger, M., Daufresne, T., Lucas, F., Pissard, S. & Lacroix, G., Does Liebig’s law of the minimum scale up from species to communities? *Oikos* **117**, 1741–1751 (2008).
14. Noble, A. E. *et al.*, Basin-scale inputs of cobalt, iron, and manganese from the Benguela-Angola front to the South Atlantic Ocean. *Limnol. Oceanogr.* **57**, 989–1010 (2012).
15. Moore, C. M. *et al.*, Relative influence of nitrogen and phosphorus availability on phytoplankton physiology and productivity in the oligotrophic sub-tropical North Atlantic Ocean. *Limnol. Oceanogr.* **53**, 824–834 (2008).
16. Behrenfeld, M. J. & Milligan, A. J., Photophysiological Expressions of Iron Stress in Phytoplankton. *Annu. Rev. Mar. Sci.* **5**, 217–46 (2013).
17. Mann, E. L. & Chisholm, S. W. Iron limits the cell division rate of *Prochlorococcus* in the eastern equatorial Pacific. **45**, 1067–1076 (2000).
18. Landry, M. R. *et al.*, Biological response to iron fertilization in the eastern equatorial Pacific (IronEx II). I. Microplankton community abundances and biomass. *Mar. Ecol. Prog. Ser.* **201**, 27–42 (2000).
19. Harpole, W. S. *et al.*, Addition of multiple limiting resources reduces grassland diversity. *Nature* **537**, 93–96 (2016).
20. Interlandi, S. J. & Kilham, S. S., Limiting resources and the regulation of diversity in phytoplankton communities. *Ecology* **82**, 1270–1282 (2001).

21. Li, W. K. W., Macroecological patterns of phytoplankton in the northwestern North Atlantic Ocean. *Nature* **419**, 154–157 (2002).
22. Edwards, K. F., Klausmeier, C. A. & Litchman, E. Evidence for a three-way trade-off between nitrogen and phosphorus competitive abilities and cell size in phytoplankton. *Ecology* **92**, 2085–2095 (2011).
23. Boyd, P. W. *et al.*, Mesoscale iron enrichment experiments 1993–2005: synthesis and future directions. *Science* **315**, 612–617 (2007).
24. Saito, M. A., Sigman, D. M. & Morel, F. M. M. The bioinorganic chemistry of the ancient ocean: The co-evolution of cyanobacterial metal requirements and biogeochemical cycles at the Archean-Proterozoic boundary? *Inorganica Chim. Acta.* **356**, 308–318 (2003).
25. Martinez, S., Yang, X., Bennett, B., Holz, R. C. A cobalt-containing eukaryotic nitrile hydratase. *Biochim. Biophys. Acta - Proteins Proteomics* **1865**, 107–112 (2017).
26. Croft, M. T., Lawrence, A. D., Raux-Deery, E., Warren, M. J. & Smith, A. G. Algae acquire vitamin B12 through a symbiotic relationship with bacteria. *Nature* **438**, 90–3 (2005).
27. Bertrand E. M. *et al.*, Methionine synthase interreplacement in diatom cultures and communities : Implications for the persistence of B12 use by eukaryotic phytoplankton. *Limnol. Oceanogr.* **58**, 1431–1450 (2013).
28. Ward, B. A., Dutkiewicz, S., Moore, C. M. & Follows, M. J. Iron, phosphorus, and nitrogen supply ratios define the biogeography of nitrogen fixation. *Limnol. Oceanogr.* **58**, 2059–2075 (2013).
29. Tagliabue, A., *et al.*, How well do global ocean biogeochemistry models simulate dissolved iron distributions? *Global Biogeochem. Cycles* **30**, 149–174 (2016).
30. Göthlich, L. & Oschlies, A. Phytoplankton niche generation by interspecific stoichiometric variation. *Global Biogeochem. Cycles* **26**, 1–8 (2012).

Supplementary Information

Methods, tables, and additional discussion can be found in the Supplementary Information file.

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Author Contributions

T.J.B. conceived, designed and carried out the study, analyzed the data, and wrote the first draft of the manuscript. C.M.M. and T.J.B. worked on subsequent drafts and improved the data analysis. E.P.A. co-led the research cruise and oversaw the nutrient analyses. A.E. oversaw the flow cytometry analyses. I.R. and T.J.B. analyzed the trace metal concentrations. E.M.B. contributed to interpretation of results. A.T. provided, and helped interpret, the PISCES2 model output. All authors commented on the manuscript.

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Extended Data Figure legends

Extended Data Figure 1: Phytoplankton responses to nutrient amendment at near coastal sites. **a–l**, Chlorophyll-a biomass, community, and F_v/F_m changes in Experiments 1 (**a–f**) and 11 (**g–l**). Dots represent treatment replicates, bars indicate the mean, and lines represent the range. Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$). N was excluded from factorial due to high ambient N concentrations (determined on-ship).

Extended Data Figure 2: Responses of the nanophytoplankton community in the bioassay experiments. Grey data points represent cell counts in replicate treatment bottles; bar heights and lines indicate the mean and range, respectively ($n=3$; units: $\times 1000$ cells mL^{-1}). Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, $n=3$; n.s.=‘not significant’). Horizontal lines indicate initial cell counts. Red data points represent chlorophyll-a fluorescence per cell and blue data points represent total nanophytoplankton chlorophyll-a fluorescence, i.e. cell counts \times cellular chlorophyll fluorescence (both have arbitrary units with different scales, lines indicate the range).

Extended Data Figure 3: Responses of the picophytoplankton community in the bioassay experiments. Grey data points represent cell counts in replicate treatment bottles; bar heights and lines indicate the mean and range, respectively ($n=3$; units: $\times 1000$ cells mL^{-1}). Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, $n=3$; n.s.=‘not significant’). Horizontal lines indicate initial cell counts. Red data points represent chlorophyll-a fluorescence per cell and blue data points represent total picophytoplankton chlorophyll-a fluorescence, i.e. cell counts \times cellular chlorophyll fluorescence (both have arbitrary units with different scales, lines indicate the range).

Extended Data Figure 4: Responses of *Synechococcus* in the bioassay experiments. Grey data points represent cell counts in replicate treatment bottles; bar heights and lines indicate the mean and range, respectively ($n=3$; units: $\times 1000$ cells mL^{-1}). Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, $n=3$). Horizontal lines indicate initial cell counts. Red data points represent chlorophyll-a fluorescence per cell and blue data points represent total *Synechococcus* chlorophyll-a fluorescence, i.e. cell counts \times cellular chlorophyll fluorescence (both have arbitrary units with different scales, lines indicate the range).

Extended Data Figure 5: Responses of *Prochlorochoccus* in the bioassay experiments. Grey data points represent cell counts in replicate treatment bottles; bar heights and lines indicate the mean and range, respectively ($n=3$; units: $\times 1000$ cells mL^{-1}). Statistically indistinguishable means are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, $n=3$). Horizontal lines indicate initial cell counts. Red data points represent chlorophyll-a fluorescence per cell and blue data points represent total *Prochlorochoccus* chlorophyll-a fluorescence, i.e. cell counts \times cellular chlorophyll fluorescence (both have arbitrary units with different scales, lines indicate the range).

Extended Data Figure 6: F_v/F_m responses to nutrient treatment. Data points represent measurements from replicate treatment bottles; bar heights and lines

indicate the mean and range, respectively. Statistically indistinguishable means are labelled with the same letter (ANOVA and Tukey HSD $p \leq 0.05$, $n=3$; n.s.=‘not significant’). Horizontal lines indicate initial conditions. Changes in F_v/F_m between initial (t=0h) and control (t=48h) time points likely reflect differential relaxation of PSII down regulation/PSII repair.

Extended Data Figure 7: Diel cycles in F_v/F_m measurements in offshore waters. **a–r**, Diel cycles; grey dots=individual F_v/F_m (F_v'/F_m' during daytime) measurements and blue line=100 point moving average. Data was blank-corrected using a mean blank value for all offshore surface waters. Light blue boundaries=range generated when the blank is increased or reduced by the standard deviation of the measured blank values. Red line=photosynthetically available radiation (PAR). **s**, Map showing the data collection locations in relation to bioassay experiments.

Extended Data Figure 8: CHEMTAX-derived community assemblages (fractional contribution to total chlorophyll-a). **a–h**, Initial waters from Experiments 1–7 and 11. **i–j**, Initial waters and selected treatments from Experiment 9 (**i**) and 10 (**j**).

Extended Data Figure 9: Exponential Shannon Wiener diversity Indices for the experiments. Indices calculated using flow cytometry cell counts (grey dots represent treatment replicates, bars represent the mean, and lines represent the range) or pigment-derived community (black dots; $n=1$ and where available). Statistically indistinguishable means for FCM-derived ESWI are labelled with the same letter (ANOVA and Fisher PLSD $p \leq 0.05$, $n=3$). Horizontal lines indicate initial conditions.

Extended Data Figure 10: Potential large-scale distribution of oceanic N-Fe co-limitation. **a**, Global surface ocean as predicted using simulated nutrient fields from an ocean biogeochemical model run (PISCES2) (Ref. 29); co-limited regions (yellow grid cells) are assigned to grid cells with an N:Fe ratio falling in the range of N-Fe co-limited experiments (see Figure 4a, b); N-Fe or Fe-N serially limited regions (i.e., those approaching N-Fe co-limitation, grey grid cells) are assigned to grid cells with a N:Fe ratio falling in the range of N-Fe or Fe-N serially limited experiments. Large black dots show the locations where additional evidence of secondary/co-limitation between N and Fe has been found (see Supplementary Table S2 for details). Crosses are locations where nutrient enrichment experiments have been performed and found evidence for N (blue crosses) or Fe (red crosses) limitation (from synthesis by Ref. 1). **b**, Observational N:Fe data gridded at the same resolution as the model. Observational Fe data (Ref. 29) have been combined with interpolated World Ocean Atlas (WOA) nitrate for location and month of the dissolved Fe measurement. **c**, Vertical domain of N:Fe ratios for a section of measured nutrient concentrations through the South Atlantic in austral summer (extended version of Figure 1b; CoFeMUG cruise¹⁴). In the central gyre, N supply from deeper waters is restricted by surface stratification whilst subsurface waters are Fe-deficient relative to N, resulting from N remineralization and Fe scavenging. Large black dots indicate data points where the measured N:Fe ratio was in the range we found N-Fe co-limitation; grey dots=within bounds of measured secondary N-Fe or Fe-N limitation. For **a** and **c**, thresholds where N and Fe concentrations are both characterized as replete, regardless of the N:Fe ratio, have been applied; these are the maximum N or Fe concentrations in Supplementary Table S1 where serial or co-limitation was found.







