- Development of a vector-based 3D grain entrainment model
- with application to X-ray computed tomography (XCT)

scanned riverbed sediment

Hal Voepel^{1,2}, Julian Leyland¹, Rebecca Hodge², Sharif Ahmed^{3,4}, and David Sear¹

¹School of Geography & Environmental Science, University of Southampton,

University Road, Southampton SO17 1BJ, UK

²Department of Geography, Durham University,

Lower Mountjoy South Road, Durham DH1 3LE, UK

³μ-VIS X-Ray Imaging Centre, University of Southampton,

University Road, Southampton SO17 1BJ, UK

⁴Diamond Light Source Ltd, Harwell Science & Innovation Campus,

Fermi Ave, Didcot, Oxfordshire OX11 0DE, UK

13 Abstract

3

10

11

12

14

15

16

17

18

19

20

21

22

Sediment transport equations typically produce transport rates that are biased by orders of magnitude. A causal component of this inaccuracy is the inability to represent complex grain-scale interactions controlling entrainment. Grain-scale incipient motion has long been modelled using geometric relationships based on simplified particle geometry and two-dimensional (2D) force or moment balances. However, this approach neglects many complexities of real grains, including grain shape, cohesion and the angle of entrainment relative to flow direction. To better represent this complexity, we develop the first vector-based, fully three-dimensional (3D) grain rotation entrainment model that can be used to resolve any entrainment formulation in 3D, and which also includes the effect of matrix cohesion. To apply this model we use X-ray computed tomography to quantify the 3D structure of water-worked

river grains. We compare our 3D model results with those derived from application of a 2D entrainment model. We find that the 2D approach produces estimates of dimensionless critical shear stress (τ_{cr}^*) that are an order of magnitude lower than our 3D model. We demonstrate that it is more appropriate to use the c-axis when calculating 2D projections, which increases values of τ_{cr}^* to more closely match our 3D estimates. The 3D model reveals that the main controls on critical shear stress in our samples are projection of grains, cohesive effects from a fine-grained matrix, and bearing angle for the plane of rotation (the lateral angle of departure from downstream flow that, in part, defines the grain's direction of pivot about an axis formed by two contact points in 3D). The structural precision of our 3D model demonstrates sources of geometric error inherent in 2D models. By improving flow properties to better replicate local hydraulics in our 3D model, entrainment modelling of scanned riverbed grains has the potential for benchmarking 2D model enhancements.

Keywords: sediment transport; grain entrainment; computed tomography; fluvial geomorphology; 3D modelling

Introduction and problem statement

The initiation of transport of grains in a stream is dependent upon the driving bed shear stress (τ_b) crossing some critical threshold for entrainment: the critical shear stress, τ_{cr} , the dimensionless form of which is known as the Shields parameter, τ_{cr}^* (Shields, 1936). Although Shields (1936) defines τ_{cr}^* as the ratio of fluid force on the grain to the submerged weight of the grain, the comprehensive review by Buffington and Montgomery (1997) highlights that for natural systems there are other factors affecting entrainment. These factors include intergranular geometry (a function of grain shape, sorting and packing), fine sediment infiltration, and grain protrusion (fully defined below, but is a function of projection and exposure). Projection is the vertical distance from the local mean bed elevation to the top of the grain. Exposure is the vertical distance from the top of

any upstream particles blocking direct flow to the top of the grain. These recognised characteristics of natural systems led to a number of definitions linking the threshold of 53 motion to simple geometric properties and grain arrangements, whereupon entrainment models were developed using a 2D static force balance (Armanini and Gregoretti, 2005; Kirchner et al., 1990; Lamb et al., 2008; Naden, 1987; Wiberg and Smith, 1987) or a 2D static moment balance (Bridge and Bennett, 1992; Ling, 1995; Papanicolaou et al., 2002; Vollmer and Kleinhans, 2007; Wu and Jiang, 2007; Wu and Chou, 2003; Wu and Yang, 2004). These 2D entrainment models resulted in considerable variation of τ_{cr} estimates in the literature, mainly associated to the inherent variability in their derivation which leads to observed methodological biases between models (Buffington and Montgomery, 1997). Despite the acknowledged geometrical complexity of natural river systems, the approaches cited above make a number of significant simplifications in that they assume a given grain: (i) is spherical or ellipsoidal (or indeed, circular or elliptical in 2D), (ii) pivots about a single contact point in the downstream direction, (iii) is unaffected by mortaring by fine sediment (Barzilai et al., 2013; Hodge et al., 2013), and (iv) is not partially buried beneath other grains. A few 3D entrainment models have been developed that are capable of more realistic geometric arrangements addressing the reliance on a single contact point noted above (Ali and Dey, 2016; Dey, 1999; Nabi et al., 2013); nevertheless, they still rely on spherical approximations of grains with relatively simple particle arrangements when 70 compared to the natural conditions of riverbeds. Furthermore, the simplistic representation of grains as spheres can lead to further assumptions about simple ratios between 72 grain diameter and projection being suitable to characterise the exposure of the grain. A number of studies have highlighted the fact that existing geometric grain entrainment models are developed for idealised laboratory or theoretical conditions and so do not reproduce the conditions found in natural riverbeds. Miller et al. (1977) showed that despite using a carefully selected set of entrainment data from non-cohesive, uniform and well-rounded or spherical grains, considerable scatter still existed in the threshold curves derived from their measurements. Similarly, Buffington et al. (1992) noted the relatively poor performance of the Wiberg and Smith (1987) model of grain entrainment, and advocated extending it to include the effects of packing, grain arrangement (including partial burial) and protrusion induced flow acceleration as seen in natural systems.

Several studies address entrainment effects from grain exposure and projection through theoretical development (Kirchner et al., 1990; Wiberg and Smith, 1987) or flume experiments using natural river gravels (Fenton and Abbott, 1977), and from grain sheltering through remote sensing of bed topography (Measures and Tait, 2008). Measures and Tait (2008) identified two mechanisms of grain sheltering: direct sheltering, where adjacent particles prevent flow forces from directly acting upon the grain, and remote sheltering, where distant upstream particles modify local flow conditions in the vicinity of the grain. Empirical methods have been employed to identify the effects of friction angles and sorting (Johnston et al., 1998; Komar and Carling, 1991) as well as grain size and shape 91 (Komar and Li, 1986) on sediment entrainment. And whilst this research identified im-92 portant relationships between entrainment shear stress and granular properties of coarse 93 river sediment and their distributions, these studies are still based on relatively simple geometries. Simplified models described above do not take into account more complex grain characteristics such as particle arrangement and orientation as well as additional spatial angles that, along with pivot angle, better describe the full range of motion during grain entrainment.

To improve the performance of geometric grain entrainment models there is a need to better account for the complexities of natural systems, including addressing the funda-100 mental point that riverbed structures are three dimensional in nature. Existing models, 101 such as those identified above, treat the entrainment and resistance forces used in 2D 102 terms and/or use simple grain geometry and arrangements, giving rise to an oversimpli-103 fication of the problem. Similarly, current entrainment models do not account for the 104 cohesion forces that fine sediment can exert on grains in natural systems (Barzilai et al., 105 2013; Hodge et al., 2013). Until recently, 2D models have been a necessary simplification 106 because of the difficulty in measuring the 3D geometry of a water-worked bed of grains. 107 However, techniques such as high-throughput X-ray computed tomography (XCT) pro-108

vide a new opportunity for measuring 3D bed structure, providing a motivation for the development of new 3D models.

Application of any 3D entrainment model requires high resolution data of a grain ar-111 rangement from which to calculate grain parameters including size and geometry, spatial 112 orientation, exposed area, elevation and the proximity of grain-to-grain contact points. 113 Non-intrusive imaging methods such as high-throughput XCT have been used in Geo-114 sciences and Geomechanics for a number of years (Wildenschild and Sheppard, 2013; 115 Ahmed et al., 2016b; Callow et al., 2018), for example to explore pore networks (Pierret 116 et al., 2002), the breakthrough of solutes in porous media (Clausnitzer and Hopmans, 117 2000) or hyporheic zone sediment structure (Chen et al., 2009, 2010; Packman et al., 118 2006). Despite various hydrological applications (Wildenschild et al., 2002), and some 119 preliminary explorations using Magnetic Resonance Imaging (Haynes et al., 2009; Klein-120 hans et al., 2008), there are no published examples of using XCT to quantify river bed 121 sediment geometry in a way pertinent to grain-scale entrainment models. 122

This paper contributes two advances. The first is a novel 3D entrainment model that 123 employs vector mechanics to account for the 3D geometric structure of grains and their 124 surroundings as well as incorporating the effects of cohesive resisting forces. We chose 125 to use the 2D Kirchner et al. (1990) geometric grain entrainment model as a basis of 126 comparison with our 3D entrainment model. The second advance is developing a method 127 for collecting and processing high-resolution XCT data from sediment beds. We use the 128 new 3D model and XCT data to establish the τ_{cr}^* for individual grains within two river 129 bed samples taken from a prototype scale flume experiment: one of coarse grains and one 130 with additional cohesive sediments. We compare our results to τ_{cr}^* values calculated using 131 the Kirchner et al. (1990) model, and also evaluate the impact of the new cohesion term.

Grain rotation threshold of entrainment models

Here vector algebra is applied to a 3D moment balance to derive a physically-based thresh-134 old of motion entrainment model, which greatly simplifies the 3D modelling process when 135 compared with the oft-used trigonometric method. Our focus is exclusively on grain ro-136 tation as it is the primary mechanism of entrainment (Agudo et al., 2014) although it 137 should be recognised that all grains can undergo a mixture of entrainment modes throughout the sediment transport process (Ancey et al., 2006). Since our 3D model implements 139 XCT scanned images of bed sediment, it utilises spatial proximity of surrounding grains 140 and their contact points. Therefore, we choose an existing 2D grain protrusion-based 141 entrainment model to facilitate comparison with our 3D vector-based entrainment model. 142

Revisiting a 2D grain protrusion-based entrainment model

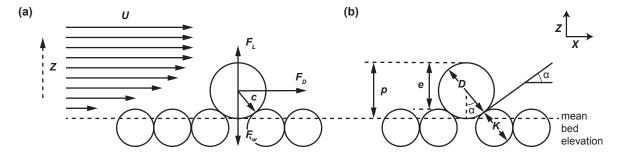


Figure 1: (a) Diagram of forces acting on a grain: F_L , F_D and F_W are the lift, drag and submerged weight forces, respectively, acting on the grain. U is the mean flow velocity at height Z above the mean bed elevation. (b) Idealised geometry for calculating grain projection (p) and exposure (e) as a function of the diameters of the test grain (D) and the bed particles (K) and the mean friction angle represented by the pivot angle (α) . Redrawn after Kirchner et al. (1990).

Prior to developing a 3D entrainment model, it is useful to establish the basis of a typical 2D grain protrusion entrainment model. In this instance the Kirchner et al. (1990) model is presented, although it is noted that any other model based on grain protrusion could be used to compare with our vector-based 3D entrainment model developed below. The key components of the 2D entrainment model are defined below, but for full details see Kirchner et al. (1990). Kirchner et al. (1990) used the 2D framework derived by Wiberg

and Smith (1987), where they account for bed slope in a static force balance used to 150 derive their entrainment model. Kirchner et al. (1990) presumed negligible bed slope for 151 their model by setting bed slope to zero. Here we use a trigonometric method to derive the same 2D entrainment model that Kirchner et al. (1990) used, except based on a static 153 moment balance of forces about a point of contact between two grains, which forms a rigid body centre of rotation in the entrainment model (head of vector C in Figure 1a). The 155 driving forces of entrainment, namely drag, F_D , and lift, F_L , act upon a circular particle and are functions of a logarithmic velocity flow profile, u(z), (Equation 1a). These forces 157 oppose a resistant force, namely submerged grain weight, F_W . A moment of force is the perpendicular distance from the centre of rotation for a rigid body to the line of action 159 of the force times the magnitude of that force. Summing all of the moments (positive 160 for clockwise rotation; negative for anticlockwise rotation) and setting their sum equal to 161 zero yields the static moment balance whereupon entrainment occurs (Equation 1b). The 162 entrainment forces of drag and lift depend on grain protrusion, which is defined by two 163 components: projection (p) of the top of the grain above mean bed elevation; and exposure 164 (e) of the top of the grain above local upstream maximum bed elevation (Kirchner et al., 165 1990, Appendix). It is theoretically assumed for the latter that the grain of interest, or 166 test grain of diameter, D, sits on a bed of uniform grains of median grain size, K_{50} , and 167 is a function of the mean friction angle, α , such that their geometric relationship is linked 168 (Figure 1b). Thus the three key equations that define the 2D Kirchner model are 169

$$u(z) = \sqrt{\frac{\tau_b}{\rho}} \kappa^{-1} f(z), \quad \text{for } f(z) = \begin{cases} \ln\left(\frac{z+z_0}{z_0}\right) & \text{if } z \ge 0, \\ 0 & \text{otherwise.} \end{cases}$$
(1a)

$$F_D \frac{1}{2} D \cos \alpha + F_L \frac{1}{2} D \sin \alpha - F_W \frac{1}{2} D \sin \alpha = 0, \tag{1b}$$

$$e = \frac{1}{2} [D - K_{50} + (D + K_{50}) \cos \alpha], \quad p = e + \frac{\pi}{12} K_{50},$$
 (1c)

where u(z) is flow velocity relative to height, z, above mean bed height at z=0, ρ is the density of water, τ_b is the boundary shear stress, $\kappa=0.407$ is van Karman's constant, f(z)

is a logarithmic profile, and $z_0 = K_{84}/10$ is a length scale. The drag force, F_D , is expressed as an integral of the product of dynamic pressure, $q = \frac{1}{2}\rho u(z)^2$, and the cross-sectional area of the grain, dA = w(z)dz, integrated over the protruding portion of the grain

$$F_{D} = \frac{C_{D}}{2} \rho \int_{p-e}^{p} w(z) u(z)^{2} dz$$

$$= \frac{C_{D}}{2} \frac{\tau_{b}}{\kappa^{2}} \int_{p-e}^{p} \sqrt{D^{2} - \left[2z - \left(2p - D\right)\right]^{2}} f(z)^{2} dz,$$
(2)

where C_D is the drag coefficient, and w(z) is the width of the grain cross section as a function of elevation z. The lift force, F_L , is the product of the dynamic pressure differential across the top and bottom of the grain and the planar cross-sectional area of the grain

$$F_{L} = \frac{C_{L}}{2} \rho A \left[u(p)^{2} - u(p-D)^{2} \right]$$

$$= \frac{C_{L}}{8} \frac{\tau_{b}}{\kappa^{2}} \pi D^{2} \left[f(p)^{2} - f(p-D)^{2} \right],$$
(3)

where C_L is a lift coefficient and A is the plan view cross sectional area of the grain. In Equations 2 and 3, Kirchner et al. (1990) used coefficient values of 0.40 and 0.20 for C_D and C_L , respectively. The submerged weight force, F_W , is defined as

$$F_W = \frac{1}{6} \left(\rho_s - \rho \right) g \pi D^3, \tag{4}$$

where ρ_s and ρ are the density of sediment and water, respectively, and g is gravitational acceleration of $9.81\,\mathrm{m\,s^{-2}}$. Substituting Equations (1a, 1c, 2–4) into the 2D moment balance, Equation 1b, Kirchner et al. (1990) expressed the τ_{cr} as

$$\tau_{cr} = \left(\rho_s - \rho\right) g\left(\frac{\pi D^3}{6}\right) \left\{ \frac{1}{\tan \alpha} \frac{C_D}{2\kappa^2} \int_{p-e}^p \sqrt{D^2 - \left[2z - \left(2p - D\right)\right]^2} f(z)^2 dz + \frac{\pi}{8} \frac{C_L}{\kappa^2} D^2 \left[f(p)^2 - f(p - D)^2\right] \right\}^{-1}.$$
 (5)

Developing a vector-based 3D grain rotation model

To account for the geometrical complexity of grains found in natural river systems, we use 186 a vector-based 3D moment balance approach for resolving τ_{cr} . The approach essentially adapts the components of fluid mechanics (Pritchard and Mitchell, 2015) used in the 188 2D critical shear model (Kirchner et al., 1990) for use within a 3D rigid body mechanics 189 framework (Beer et al., 2013) and also accounts for an additional resisting force associated 190 with cohesion. It is important to point out that the model being developed is fairly simple, 191 assuming a logarithmic velocity flow profile where all forces are configured in a simple 192 orientation relative to the system reference frame. Consequently, any additional driving 193 and resisting forces or fluvial mechanics modifications could subsequently be included as 194 components within the 3D vector-based framework (e.g. forces related to flow turbulence). 195 A 3D vector is written as $\mathbf{F} = F_x \hat{\mathbf{i}} + F_y \hat{\mathbf{j}} + F_z \hat{\mathbf{k}}$, where F_x , F_y , and F_z are the scalar 196 components, and $\hat{\mathbf{i}}$, $\hat{\mathbf{j}}$, and $\hat{\mathbf{k}}$ are the orthonormal unit vectors corresponding to the (x, y, z)197 axes of a Cartesian reference frame. To simplify the 3D entrainment model, the reference 198 frame is chosen such that force vectors are oriented along single vector components. Here 199 the stream flow velocity is defined in the direction of the x-axis, $u\hat{i}$ (m s⁻¹), and gravity 200 in the direction of the z-axis, $g\hat{\mathbf{k}}$ (9.81 m s⁻²). Unless otherwise stated, all vectors in this 201 section are given as 3×1 column vectors to ensure correctly applied vector calculations. 202 Typically, in the 2D case, grains are assumed to rotate in the direction of flow, and a 203 pivot angle is formed between the gravity vector and the vector extending from the grain's centre of mass to a single point of contact with an adjacent downstream grain. However, 205 in reality, grains do not pivot about a single contact point but rather they do so about 206 an axis of rotation (AOR), A, formed by the vector difference of a left contact vector, 207 \mathbf{CV}_L , and a right contact vector, \mathbf{CV}_R , between two contact points of adjacent grains (Figure 2a). Each contact vector extends from the grain's centre of mass to a contact 209 point with the adjacent grain. For illustration, think of a grain rotating within the saddle 210 formed by two downstream grains. All of the action of a pivoting grain occurs within 211 the plane of rotation (POR), which is orthogonal to the AOR (x'z'-plane in Figure 2a,b).

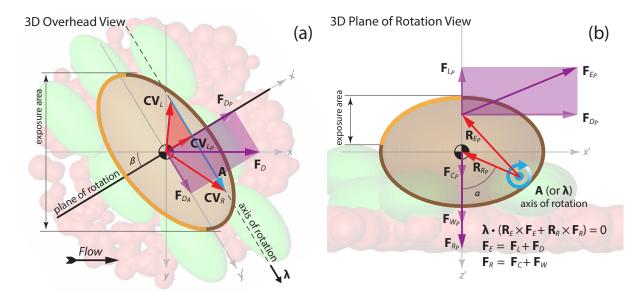


Figure 2: (a) An overhead image of a 3D ellipsoidal grain on a sediment bed is overlaid by a vector diagram shown in the xy-plane with z-axis going into the page. The AOR vector (\mathbf{A} , blue arrow) formed by left and right contact vectors (\mathbf{CV}_L and \mathbf{CV}_R , red arrows) that extend from the centre of mass of the grain to where the grain makes contact with two neighbouring surface grains (green stones). Vector \mathbf{A} forms a pivot for grain rotation when acted upon by a drag force (\mathbf{F}_D , violet arrow). The drag force resolves as two component vectors: one that is parallel to the AOR (\mathbf{F}_{D_A}) and one that resides within the POR (\mathbf{F}_{D_P}). Vectors that reside within the POR, denoted by the P subscript, contribute to pivoting the grain (e.g. \mathbf{F}_{D_P} and \mathbf{CV}_{L_P}). (b) The same grain viewed within the 3D POR reference frame (x', z') indicates the y'-axis is coming out of the page. The 3D POR view shows component vectors of resultant forces (\mathbf{F}_{E_P} and \mathbf{F}_{R_P} , violet arrows) and their positions (\mathbf{R}_{E_P} and \mathbf{R}_{R_P} , red arrows), which are the two moment arms for applied forces acting about the rigid body centre of rotation (\mathbf{A} , circular blue arrow). Angles for pivot, α , and bearing, β , are indicated. Tilt angle, γ , which is the departure angle of the rotation axis from the horizontal plane, is zero in the diagram and not shown.

The 3D orientation of the POR is induced by the rotation axis, where all of the 3D vector mechanics affecting rotation about an axis essentially collapses onto the 2D subspace we refer to as the POR.

Within any 3D axis rotational framework, an arbitrary vector, \mathbf{V} , in 3D vector space, may be resolved into two orthogonal component vectors: one vector parallel to the AOR, \mathbf{V}_A , and another vector that spans the POR, \mathbf{V}_P (note vector subscripts in Figure 2a,b). For example, the drag force, \mathbf{F}_D , shown in Figure 2a, is parallel to the downstream flow direction along the x-axis. Here drag force is resolved into two component vectors: one projected onto the AOR, \mathbf{F}_{D_A} , the other, projected onto the POR, \mathbf{F}_{D_P} . Of these two orthogonal component vectors, only the POR force projection affects entrainment of the grain since the AOR force projection is parallel to the rotation axis. As an analogy, an open door does not swing shut due to gravity, which is a vector parallel to the door's AOR along its hinges; you must apply an external, nonparallel force to push the door shut.

Figure 2 shows that all forces involved in grain entrainment act along multiple angles, 226 only one of which is the pivot angle. The overhead view of the grain (Figure 2a) shows 227 that the arrangement of the two contact points and the grain's centre of mass determines 228 a bearing angle (β , Figure 2a) and a tilt angle (γ , not shown) for the POR as well as a 229 pivot angle $(\alpha, \text{ Figure 2b})$ that lies within the POR. We define the bearing angle as the 230 angle formed between the POR and the direction of stream flow, and the tilt angle as the 231 angle formed between the AOR vector and the horizontal xy-plane. These two departure 232 angles, one from the downstream flow direction (bearing) and the other from gravity (tilt), 233 define the 3D spatial orientation of the POR as it begins its forward rotation (pivot) at 234 the threshold of entrainment. 235

To obtain the pivot angle, either contact vector plus the gravity vector must be projected 236 onto the POR before calculating the angle between these projected vectors. The 3×3 237 matrix that projects any vector onto the AOR is $\lambda(\lambda^T\lambda)^{-1}\lambda^T$, where λ is the unit vector 238 for the AOR (Figure 2a) and superscript T is the transpose of a vector. The projection 239 that is orthogonal to the AOR projection, which projects any vector onto the POR is 240 $\mathbf{I}_3 - \boldsymbol{\lambda}(\boldsymbol{\lambda}^T\boldsymbol{\lambda})^{-1}\boldsymbol{\lambda}^T$, where \mathbf{I}_3 is the 3×3 identity matrix. Hence, without loss of generality, 241 the projection of $\mathbf{C}\mathbf{V}_L$ onto the POR is given by $\mathbf{C}\mathbf{V}_{L_P} = (\mathbf{I}_3 - \boldsymbol{\lambda}(\boldsymbol{\lambda}^T\boldsymbol{\lambda})^{-1}\boldsymbol{\lambda}^T)\mathbf{C}\mathbf{V}_L$ 242 (Figure 2a). Similarly, the bearing angle is found by projecting the x-axis unit vector 243 $\hat{\mathbf{i}}$ onto an "un-tilted" vertical POR. Let $\boldsymbol{\lambda}_{XY}$ be $\boldsymbol{\lambda}$ where the $\hat{\mathbf{k}}$ scalar component is set 244 to zero and normalised. Then $\mathbf{I}_3 - \boldsymbol{\lambda}_{XY}^T (\boldsymbol{\lambda}_{XY}^T \boldsymbol{\lambda}_{XY})^{-1} \boldsymbol{\lambda}_{XY}$ is the matrix that projects any vector onto the vertical POR, and $\hat{\mathbf{i}}_{XY} = (\mathbf{I}_3 - \boldsymbol{\lambda}_{XY}^T (\boldsymbol{\lambda}_{XY}^T \boldsymbol{\lambda}_{XY})^{-1} \boldsymbol{\lambda}_{XY}) \hat{\mathbf{i}}$ is the projection of 246 the unit vector $\hat{\mathbf{i}}$ onto the vertical POR. Finally, we use the definition of the dot product to find the angles for pivot, α , bearing, β , and tilt, γ , which are given by

$$\alpha = \arccos\left(\frac{\mathbf{C}\mathbf{V}_{L_P} \cdot \mathbf{G}_P}{|\mathbf{C}\mathbf{V}_{L_P}||\mathbf{G}_P|}\right),\tag{6a}$$

$$\beta = \arccos\left(\frac{\hat{\mathbf{i}}_{XY} \cdot \hat{\mathbf{i}}}{|\hat{\mathbf{i}}_{XY}|}\right),\tag{6b}$$

$$\gamma = \arccos\left(\boldsymbol{\lambda} \cdot \boldsymbol{\lambda}_{XY}\right),\tag{6c}$$

where \mathbf{G}_P is the projection of gravity $g\hat{\mathbf{k}}$ onto the POR determined by the same projection used for contact vector projections, and |A| is the norm of arbitrary vector A. It is important to realise that the angles defined by Equations 6 are unnecessary for a vector-based 3D entrainment model. Nevertheless, since they are important to describe entrainment behaviour and to evaluate the 3D model, these angles are calculated once threshold of entrainment is determined.

Our 3D entrainment model uses four forces acting upon a grain: the drag force, \mathbf{F}_D , the lift force, \mathbf{F}_L , the submerged weight of the grain, \mathbf{F}_W , and a cohesive force, \mathbf{F}_C , due to fine-grained matrix making contact with the particle (Figure 2b). The resultant forces for entrainment $\mathbf{F}_E = \mathbf{F}_D + \mathbf{F}_L$ and for resistance $\mathbf{F}_R = \mathbf{F}_W + \mathbf{F}_C$, are each applied at positions \mathbf{R}_E and \mathbf{R}_R , respectively. Each position vector extends from anywhere along the AOR (either contact point is most convenient) to the line of action of the applied force (Figure 2b). The static moment balance about an AOR is given as the scalar triple product

$$\lambda \cdot (\mathbf{R}_E \times \mathbf{F}_E + \mathbf{R}_R \times \mathbf{F}_R) = \lambda \cdot (\mathbf{R}_E \times (\mathbf{F}_D + \mathbf{F}_L) + \mathbf{R}_R \times (\mathbf{F}_W + \mathbf{F}_C)) = 0.$$
 (7)

The dot product in Equation 7 is the scalar resolute of the resultant moment vector, $\sum \mathbf{R} \times \mathbf{F}$, in the direction of the AOR unit vector, λ . Recall that vector $\mathbf{R} \times \mathbf{F}$ is orthogonal to the 2D plane spanned by vectors \mathbf{R} and \mathbf{F} . Thus, the scalar resolute in
Equation 7 is equivalent to collapsing all of the \mathbf{R} and \mathbf{F} vectors onto the POR as vector components (Figure 2b), calculating the resultant of all their moments, and then measuring its magnitude, yielding a signed scalar value (see Appendix A for a geometric proof).
Subsequent subsections give details for the entrainment and resistance force vectors and
their corresponding position vectors given in Equation 7.

271 Entrainment force components for the 3D model

We assume non-slip conditions at the bed and a logarithmic velocity profile with increasing elevation, z, up the water column. Flow velocity is zero at the mean bed elevation. Both drag and lift forces are functions of the velocity profile, which is given by

$$u(z) = \begin{cases} \kappa^{-1} \sqrt{\frac{\tau_b}{\rho}} \ln\left(\frac{z+z_0}{z_0}\right) & z \in Z \ge 0, \\ 0 & \text{otherwise,} \end{cases}$$
 (8)

where $\kappa = 0.407$ is van Karman's constant, τ_b is the boundary shear stress in N m⁻², ρ is water density assumed to be $1000 \,\mathrm{kg} \,\mathrm{m}^{-3}$, $z_0 = D_{84}/10$ is a bed roughness length scale, reference elevation, z = 0, starts at the local mean bed elevation, which is taken over a distance of D_{84} upstream and downstream of the grain, and $z \in Z$ are discrete nonnegative elevations above the mean bed. Drag force in vector format is given by

$$\mathbf{F}_D = F_D \hat{\mathbf{i}} = E F_{adj} \frac{C_D}{2} \rho \sum_{z \in Z} u(z)^2 A_{\perp}(z) \hat{\mathbf{i}}, \tag{9a}$$

$$EF_{adj} = 1 - \max\left\{0, \frac{F_u - EF}{F_u}\right\},\tag{9b}$$

where C_D is the drag coefficient (assumed to be 0.91 for natural sediment after Schmeeckle et al. (2007) and will be used in both 2D and 3D models for model comparison), u(z) is the logarithmic velocity profile defined in Equation 8, $A_{\perp}(z)$ is the cross-sectional area of the grain perpendicular to flow as a function of discrete elevation height increments $z \in Z$ above the mean bed elevation, $z \in Z$ above the mean bed elevation, $z \in Z$ above the sheltering effects of upstream grains, and EF_{adj} is the necessary adjustment for the portion of EF already accounted for by F_u .

To estimate exposure factor, EF, for any given grain we update the concept of calculating sphere protrusion areas (Yager et al., 2007) to include a weighted average of multiple area ratios over several viewing angles of the bed. To better account for the turbulent nature of streamflow, we assume that flow velocity (an instantaneous vector we term \mathbf{V}) influencing the drag force could interact with the grain at any angle between horizontal and vertical relative to the bed, but that only the horizontal component of said flow is relevant to the resulting drag force ($\mathbf{V}\sin\theta$, where θ is the angle from the overhead view). The exposure factor, EF, is the weighted average of exposure area ratios given by

$$EF \approx \frac{\sum_{\theta \in \Theta} ER_{\theta} \sin^2 \theta}{\sum_{\theta \in \Theta} \sin^2 \theta},$$
 (10)

where ER_{θ} is the exposure area ratio at viewing angle θ , which is defined as the exposed area of the grain viewed from angle θ divided by the area of the grain viewed from the 297 same angle but with no other grains present, and Θ is the collection of equidistant angles from 0° at the overhead view of the bed to 90° at the downstream horizontal view of 299 the bed (see Appendix B for physical derivation of EF and EF_{adj}). There are four noteworthy observations about Equation 10: (i) normalising the exposed-view area by 301 total-view area yields a dimensionless metric, (ii) the metric is physically consistent with 302 the product of dynamic pressure and the cross-sectional grain area, $\frac{1}{2}\rho u^2 A$, used in drag 303 force calculations, (iii) we do not need to know velocities or grain areas, as they cancel across numerator and denominator, and (iv) the weights sum to unity as required for any 305 weighted mean. Lift force in vector format is given by

$$\mathbf{F}_L = -F_L \hat{\mathbf{k}} = -\frac{C_L}{2} \rho A_{\parallel} \left(u(z_T)^2 - u(z_B)^2 \right) \hat{\mathbf{k}}, \tag{11}$$

where C_L is the lift coefficient assumed to be 0.20 for natural grains (Schmeeckle et al., 2007), A_{\parallel} is the overhead cross-sectional area of the grain parallel to flow, and $u(z_T)$

and $u(z_B)$ are the logarithmic velocity profile values at the top and bottom of the grain, respectively.

The position vector for entrainment, $\mathbf{R}_E = R_{E_x} \hat{\mathbf{i}} + R_{E_y} \hat{\mathbf{j}} + R_{E_z} \hat{\mathbf{k}}$, extends from the AOR to the intersection of perpendicular lines of action for lift and drag forces (see Figure 2b). 312 The line of action for the lift force runs vertically through the grain centre of mass. Due 313 to the location and distribution of the velocity profile, however, the line of action for the 314 drag force is horizontally situated somewhere above the grain's centre of mass. To find 315 the elevation along the velocity profile for the horizontal line of action for drag force, we 316 need to find the midpoint of distributed drag force across the face of the grain. By using 317 the drag force distribution as a weighting mechanism, the elevation for the horizontal line 318 of action for drag force is found by calculating 319

$$z_E = \frac{\sum_{z \in Z} z \cdot u(z)^2 A_{\perp}(z)}{\sum_{z \in Z} u(z)^2 A_{\perp}(z)},$$
(12)

 $_{\rm 320}$ where z_E is the elevation above the mean bed located along the velocity profile.

Resistant force components for the 3D entrainment model

There are two resisting forces: a submerged grain weight and a cohesive force due to contact with any fine-grain matrix. Submerged weight force in vector format is given by

$$\mathbf{F}_W = F_W \hat{\mathbf{k}} = (\rho_s - \rho) g V_s \hat{\mathbf{k}}, \tag{13}$$

where $\rho_s - \rho$ is the submerged density of sediment grain assumed to be $1650 \,\mathrm{kg} \,\mathrm{m}^{-3}$, g is the gravitational constant of $9.81 \,\mathrm{m} \,\mathrm{s}^{-2}$, and V_S is the grain volume (m³). Cohesive force in vector format is given by

$$\mathbf{F}_C = F_C \hat{\mathbf{k}} = \eta A_S f_S \hat{\mathbf{k}},\tag{14a}$$

where η is the cohesive force per unit area (N mm⁻²), A_S is the surface area of the grain (mm²), and f_S is the fraction of the grain surface area in contact with matrix fines. An empirical power law model was derived from a cohesive tensile force experiment where a 25 mm diameter marble was pulled with a handheld force gauge from several binary mixtures of sand and clay (p < .001, $R^2 = 0.54$ (log), see Appendix C for details). This model is given by

$$\eta_{\phi} = a D_{\phi}^{b} F^{c}, \quad \text{for } \phi \in \Phi,$$
(14b)

where η_{ϕ} is the force per area power law (N mm⁻²) for grains with median axes within matrix sand size range ϕ , D_{ϕ} is the mean for matrix sand grain size range ϕ , F is the fraction of clay in the sand-clay matrix, and the log-linearised model regression coefficients are a=0.002579, b=-0.5332, and c=0.2328. Sieving analysis rarely yields a uniform distribution of mass across all sand size classifications. As such, we use Equation 14b to estimate a mass-weighted mean of cohesive force across grain sizes. Metric grain size ranges for sand classifications are $\Phi = \{(0.0625, 0.125], (0.125, 0.25], (0.25, 0.5], (0.5, 1], (1, 2]\}$ mm, and the set of means for each grain size range is $D_{\Phi} = \{0.0938, 0.188, 0.375, 0.75, 1.5\}$ mm. The mass-weighted mean force per area η used in Equation 14a is given by

$$\eta = \frac{\sum_{\phi \in \Phi} \eta_{\phi} m_{\phi}}{\sum_{\phi \in \Phi} m_{\phi}},\tag{14c}$$

where η_{ϕ} is the force per area from Equation 14b and m_{ϕ} is the sediment mass for size range $\phi \in \Phi$ determined from sieving analysis of the sediment used in XCT scanning. Equation 14c is necessary since the cohesive tensile force experiments were performed on binary mixtures of a narrow range of grain sizes for sand mixed with clay whereas sand in the cohesive matrix from the flume experiments consists of a broad range of grain sizes for sand mixed with clay. Each force contribution from the narrow grain size range, η_{ϕ} , is weighted by the mass of grains, m_{ϕ} , that were sieved within that size range.

The position vector for resistance, $\mathbf{R}_R = R_{R_x} \hat{\mathbf{i}} + R_{R_y} \hat{\mathbf{j}} + R_{R_z} \hat{\mathbf{k}}$, extends from either contact point on AOR to the centre of mass of the grain (see Figure 2b). Observe that \mathbf{R}_R is just

 $-\mathbf{CV}_{L_P}$ (or the equivalent vector $-\mathbf{CV}_{R_P}$).

Calculation of au_{cr} for the 3D entrainment model

For any given grain, the centre of mass and coordinates of all nearby contact points are 340 used to construct all vectors and metrics. A set of critical shear stress values, $\{\tau_{cr}\}$, is 341 calculated for each stone using every viable pair of contact vectors, $(\mathbf{CV}_L, \mathbf{CV}_R)$. Each 342 viable pair of contact vectors induces a unique set of two position vectors, \mathbf{R}_E and \mathbf{R}_R , 343 used in the entrainment model (Equation 7), where their positions extend from either 344 contact point to the grain's centre of mass for the resistance vector and to the horizontal 345 line of action (Equation 12) for the entrainment vector. For each grain, entrainment 346 computation starts by calculating the resistance force vector, $\mathbf{F}_R = \mathbf{F}_W + \mathbf{F}_C$, for the 347 entrainment model. Then a list of all viable pairs of contact vectors, $(\mathbf{CV}_L, \mathbf{CV}_R)$, and 348 their corresponding position vectors is constructed. The viability of each pair of contact vectors is established using four tests: (i) the POR passing through the grain's centre of 350 mass must intersect the AOR between the pair of contacts, (ii) the POR must not tilt so severely that both contact vectors are on the same side of a POR-aligned vertical plane 352 containing the gravity vector, (iii) the POR bearing angle is bounded by $|\beta| < 90^{\circ}$, and (iv) any contact pair must be eliminated from the list if a third contact point is positioned 354 such that it prevents the grain from forward rotation. From all viable contact pairs, the smallest critical shear value in the set, $\tau_{cr} = \min\{\tau_{cr}\}$, and its corresponding contact pair, 356 $(\mathbf{CV}_L, \mathbf{CV}_R)$, are denoted as the threshold of motion solution for that grain.

The most efficient way to calculate the scalar triple product, $\mathbf{\lambda} \cdot \mathbf{R} \times \mathbf{F}$, is to express it
as a 3 × 3 matrix with the AOR unit vector, $\mathbf{\lambda} = \lambda_x \hat{\mathbf{i}} + \lambda_y \hat{\mathbf{j}} + \lambda_z \hat{\mathbf{k}}$ in the first row, the
position vector, $\mathbf{R} = R_x \hat{\mathbf{i}} + R_y \hat{\mathbf{j}} + R_z \hat{\mathbf{k}}$, in the second row, and the resultant force vector,

 $\mathbf{F} = F_x \hat{\mathbf{i}} + F_y \hat{\mathbf{j}} + F_z \hat{\mathbf{k}}$ in the third row, and then evaluate its determinant

$$egin{align*} oldsymbol{\lambda} \cdot \mathbf{R} imes \mathbf{F} = egin{align*} \lambda_x & \lambda_y & \lambda_z \ R_x & R_y & R_z \ F_x & F_y & F_z \ \end{bmatrix} = \lambda_x \Big(R_y F_z - R_z F_y \Big) - \lambda_y \Big(R_x F_z - R_z F_x \Big) + \lambda_z \Big(R_x F_y - R_y F_x \Big). \end{split}$$

Recognising that we need one determinant matrix for each position vector in the equation for a static moment balance about an AOR, $\lambda \cdot \sum \mathbf{R} \times \mathbf{F} = 0$, our 3D moment balance (Equation 7) yields the following scalar equation

$$\begin{vmatrix} \lambda_{x} & \lambda_{y} & \lambda_{z} \\ R_{Ex} & R_{Ey} & R_{Ez} \\ F_{D} & 0 & -F_{L} \end{vmatrix} + \begin{vmatrix} \lambda_{x} & \lambda_{y} & \lambda_{z} \\ R_{Rx} & R_{Ry} & R_{Rz} \\ 0 & 0 & F_{W} + F_{C} \end{vmatrix} = 0$$
 (15a)

$$\lambda_x \left[-R_{E_y} F_L + R_{R_y} \left(F_W + F_C \right) \right] +$$

$$\lambda_y \left[R_{E_z} F_D + R_{E_x} F_L - R_{R_x} \left(F_W + F_C \right) \right] - \lambda_z R_{E_y} F_D = 0$$
(15b)

Algorithmic solutions, using iterations of boundary shear stress (see Appendix D), should be calculated using the scalar triple product form, Equations 7 or 15a, until the scalar moment reaches zero whereas analytical solutions should use the scalar equation form given by Equation 15b. Factoring by force components with all resistant force terms placed on the right hand side of Equation 15b and all entrainment force terms on the left hand side, we solve for the boundary shear stress yielding the critical shear equation

$$\tau_{cr} = \frac{\Lambda_R (\rho_s - \rho) g V_S}{\Lambda_D \mathscr{F}_D(A_\perp) + \Lambda_L \mathscr{F}_L(A_\parallel)} + \frac{\Lambda_R \eta A_S f_S}{\Lambda_D \mathscr{F}_D(A_\perp) + \Lambda_L \mathscr{F}_L(A_\parallel)},\tag{15c}$$

366 where

$$\Lambda_R = \lambda_y R_{R_x} - \lambda_x R_{R_y}, \quad \Lambda_D = \lambda_y R_{E_z} - \lambda_z R_{E_y}, \quad \Lambda_L = \lambda_y R_{E_x} - \lambda_x R_{E_y},$$

$$\mathscr{F}_D(A_\perp) = E F_{adj} \frac{C_D}{2\kappa^2} \sum_{z \in Z} \left[\ln \left(\frac{z + z_0}{z_0} \right) \right]^2 A_\perp(z), \text{ and}$$

$$\mathscr{F}_L(A_\parallel) = \frac{C_L}{2\kappa^2} A_\parallel \left(\left[\ln \left(\frac{z_T + z_0}{z_0} \right) \right]^2 - \left[\ln \left(\frac{z_B + z_0}{z_0} \right) \right]^2 \right).$$

For clarity, we partition the critical shear formula (Equation 15c) into two components: 369 the first term is the 3D equivalent to the 2D entrainment model of Kirchner et al. (1990) 370 and the second term is the cohesive force component due to particle contact with the fine-371 grain matrix. The Λ 's are the linear combinations of scalar components, in units of length 372 (m), consisting of products of components from the AOR unit vector and position vectors; 373 they represent the 3D geometry of grain-to-grain rotation for the rigid body mechanics 374 of the entrainment model. Scalar force component functions, \mathscr{F}_D and \mathscr{F}_L , are attributed 375 to the continuum mechanics for drag and lift forces, respectively, and are shown here 376 as functions of cross-sectional grain areas (m²) perpendicular and parallel to streamflow, 377 respectively. Terms in the denominator are in units of cubic length (m³) whilst those in 378 the numerator are in units of force-length (N m) making the τ_{cr} in units of force per area 379 $(N m^{-2})$. Dimensionless τ_{cr} is calculated as $\tau_{cr}^* = \tau_{cr}/[(\rho_s - \rho)gD]$ for grain diameter D.

381 Application of the model to XCT scanned river grains

We used XCT to image an extracted section of an analogue riverbed, which was waterworked in a prototype scale flume. The 3D image was then processed to derive grain metrics, which were used to apply the new 3D entrainment model.

Development and extraction of analogue riverbed sediments

A prototype scale riffle-pool sequence was constructed in a large ($60 \text{m} \log \times 2.1 \text{m} \text{wide} \times 0.7 \text{m}$ deep) flume at the University of Southampton's Chilworth Hydraulics facility (Fig-

ure 3a), using river gravels with a grain size distribution that was carefully matched 388 through sieving to those found in a natural riffle-pool system at Bury Green Brook, UK 389 (mean D_{50} range is 25 mm to 34 mm; (Hodge et al., 2013)). Prior to water-working, a 390 series of five 250 mm diameter metal mesh baskets were buried with the rim $1.5 \times D_{50}$ 391 below the surface of the bed, allowing the surface grains of the bed to move without being 392 impeded. Water-working of the gravels was achieved by slowly raising the flow rate at 393 the start of each flume run to a critical flow level (when several painted D_{50} stones placed on the riffle mobilised) where it was sustained for fifteen minutes prior to reducing the 395 flow to a half critical level for six hours. Powell et al. (2016) show that the majority of bed adjustment occurs over this timescale. The mean \pm standard deviation of half 397 critical flow pooled over all flume runs was $6.26 \times 10^{-2} \pm 5.97 \times 10^{-3} \,\mathrm{m}^3 \,\mathrm{s}^{-1}$ with half 398 critical flow for individual flume runs ranging from $5.44 \times 10^{-2} \pm 1.87 \times 10^{-2} \,\mathrm{m}^3\,\mathrm{s}^{-1}$ to 399 $7.03 \times 10^{-2} \pm 2.02 \times 10^{-2} \,\mathrm{m}^3 \,\mathrm{s}^{-1}$. For this work, we examined two end-member extremes in grain size distribution from the pool tail: (R1) coarse gravels (minimum grain size 401 = 4 mm) with no fines and (R2) the same coarse sediment with the additional input of 402 sands and clay $(1-4 \mu m)$. The clay was continuously added to the inlet flow to achieve 403 a steady concentration of $500\,\mathrm{mg}\,\mathrm{L}^{-1}$ throughout the run, measured using a conductivity 404 meter with a predetermined calibration. In both runs the final sediment was framework 405 supported. At the end of each run, the baskets were carefully extracted and briefly sub-406 merged in liquid wax to fix the in situ spatial arrangement of bed grains prior to transport 407 to the scanner (Figure 3b). 408

409 XCT imaging, reconstruction and registration

The two baskets of sediment were imaged using a micro-focus Nikon Metrology μ CT scanner at the μ -VIS X-Ray Imaging Centre, University of Southampton, UK. The image scanner has a 450 kVp X-ray source and a 2048 pixel Curved Linear Detector Array (CLDA). Data were acquired using an electron accelerating potential of 440 kVp and a tube current of 922 μ A. 951 equiangular projections were acquired through 360° with

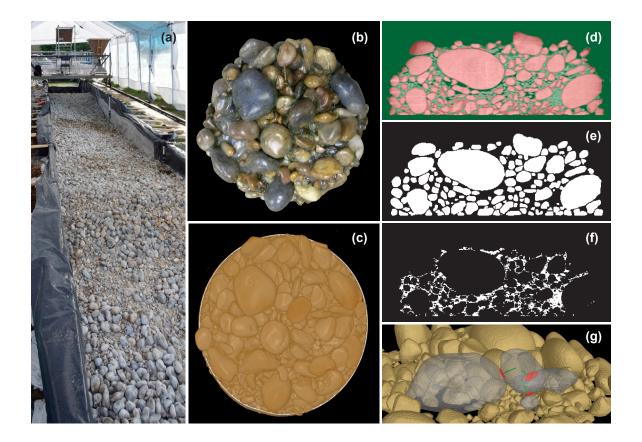


Figure 3: (a) The Chilworth Hydraulic Facility outdoor flume showing the constructed prototype scale riffle-pool sequence. Flume dimensions are $60\,\mathrm{m} \times 2.1\,\mathrm{m} \times 0.7\,\mathrm{m}$. (b) An extracted basket that has been fixed by submergence in hot wax prior to XCT scanning. (c) An output 3D rendered image. (d) Binary segmentation process showing coarse grains (in red) and fine-grain matrix, air and wax (in green). (e) A slice through the binary image stack of coarse grains. (f) A slice through the binary image stack of fine-grain matrix. (g) Transparent mock-up of four grains (in grey) showing centre of mass-to-contact point vectors of a single object grain (partial green arrows) and their grain-to-grain contact points (shown as red patches for clarity).

4 frames per projection taken to reduce noise. An exposure time of 60 ms was used. A detailed discussion of XCT artefacts and corrections is presented in Clausnitzer and 416 Hopmans (2000). 3D reconstruction of the projections was performed using NITRO, a 417 quasi-Newton and trust region differentiable optimisation algorithm found in Digi XCT 418 (Digisens, 2014; More and Sorensen, 1983; Nocedal and Wright, 1999). The resulting 419 isometric voxel (the 3D extension of a 2D pixel) size is 600 µm. Interactive volume regis-420 tration was performed using VGStudio Max 2.1 (Graphics, 2011). Three reference points 421 were used to rotate and align the 3D datasets orthogonal to the original directions of 422 flow and gravity, which is an important stage for later calculations of grain geometry and 423

exposure area in relation to downstream flow direction (Figure 3c).

⁴²⁵ Image processing and model application

After reconstruction and registration, the XCT data consists of a 3D array of voxels whose greyscale intensity is representative of the material density at that point. Image 427 segmentation was undertaken to split the image into coarse grains (larger than 4 mm; 428 Figure 3d, in red), and finer grains of sands and clay (termed 'matrix' and applicable 429 only in R2) along with air/wax (Figure 3d, in green), and was performed in Fiji ImageJ 430 (Schindelin et al., 2012). Segmentation is performed using a semi-automated classification 431 process in which the user sequentially defines known regions of air and wax (background), 432 grains, and matrix, to train the classification algorithm, resulting in binary masks of each 433 type of material (Figure 3e,f). Despite the larger sand grains in the matrix being greater 434 than the image resolution, it was not possible to individually identify them as it takes 2-4 435 voxel widths at a 600 μm resolution to describe any grain with even rudimentary detail; 436 consequently they were included in the general matrix class. Following classification, an 437 initial separation algorithm was applied to the grain dataset to isolate individual grains. 438 An adverse effect of image segmentation and separation is the inherent surface erosion 439 over all of the separated particles in the image. The extent of erosion is determined by a 440 number of factors including but not limited to scan quality, material density differences, segmentation steps, and the separation algorithm. No contact actually occurs between 442 separated particles in the images.

Two sets of 2D image slices (each set composes a 3D image stack), a binary image stack of coarse grains (Figure 3e) and a greyscale image stack of fine-grain matrix (similar to Figure 3f), were imported into MATLAB (MathWorks, 2018) where four categories of grain metrics were extracted prior to entrainment calculations: (i) grain characteristics consisting of unique grain identification, centroid, volume, surface area, maximum grain elevation and its coordinates, (ii) principal axis lengths and their spatial orientation, (iii) grain-to-grain relationships consisting of coordinates of contact points with surrounding grains,

their minimum separation distances, and identification numbers of all contact grains, and
(iv) matrix image processing to remove air pores and to calculate the fractional areas of
matrix fines in contact with coarse sample grains. Imaging software packages typically
have functions to extract characteristics such as labelling particles and obtaining their
centroids, areas, and bounding boxes. Herein we used MATLAB with the Image Processing Toolbox and the Parallel Processing Toolbox to compute metrics. All code developed
is available online at https://github.com/NERCPATCheS/VectorEntrainment3D.

XCT scan data imported into MATLAB do not intrinsically have any information about 458 scaling required to make real spatial measurements. To obtain spatial dimensions from 459 array calculations for distance, area, or volume, the calculated results must be multiplied 460 by the voxel resolution, squared resolution, or cubed resolution, respectively. Dividing 461 any of these converted spatial dimension by the appropriate voxel resolution converts 462 the spatial dimension for each voxel back into an index value; this facilitates coding of 463 entrainment model formulas where indices are required (Equations 8, 9a, 11, and 12). 464 For our MATLAB coding, we left coordinates, such as centroids, as index values to avoid 465 rounding errors in moment calculations using Equation 7, where we used a binary search algorithm rather than the analytical solution (Equation 15c) to calculate our results (see 467 Appendix D). Note that the Λ 's are dimensionless when Equation 15c is used with indices to find analytical solutions of grain entrainment for XCT scanned images. 469

Individual grain characteristics must first be established before particle group relation-470 ships can be derived. The surface area of each grain was calculated by eroding the 3D 471 binary image of grains with a 3D ball structural element to create labelled perimeters for each grain of one voxel thickness (Russ, 2011). Tabulating the surface and internal 473 labelled voxels for each grain allowed the surface area and volume to be readily calcu-474 lated. The maximum elevation of each labelled grain was also recorded. Axis length and 475 orientation of image particles is usually calculated using ellipsoidal fitting (Ahmed et al., 476 2016a; Ge et al., 2005), but this method performs poorly when grains are more angular, 477 as they often are in natural systems. Instead, we performed a Principle Components 478

Analysis (PCA) on all voxel coordinates of individual grains to produce axis lengths and 479 grain orientation metrics. Contact points were found using an efficient search algorithm 480 that calculates the minimum sums of squared distances between all surface voxels of an 481 object grain and nearby surface voxels of adjacent grains. Then a threshold distance is 482 used during entrainment calculations to identify contact points in close enough proximity 483 for consideration as a potential contact stones. Entrainment results were sensitive to 484 small increments in threshold values as the number of stones having viable entrainment calculations increased from very few to nearly all surface stones. An example of contact 486 points is shown in a transparent mock-up image (Figure 3g), which illustrates a group of grains (grey) shown with partial contact vectors (green) that extend from the centre of 488 mass of an object particle (centre grain) to the contact points (red patches) of its neigh-489 bouring grains. Greyscale image thresholding was used on the matrix images to segment 490 pore space (i.e. darker greyscale values) from the fine-grain matrix, which converts the 491 remaining matrix into a binary image. The labelled image of surface grains were then 492 used with the binary matrix image to determine the amount surface area in contact with 493 matrix voxels, and a contact area fraction, f_S , was calculated for each grain. 494

Once grain metrics were computed and tabulated, we used the scalar triple product form of our 3D entrainment model (Equation 7) along with a binary search algorithm (see Appendix D) to calculate τ_{cr} for each grain. Then the corresponding pivot angle and two POR angles for bearing and tilt were calculated (Equation 6). Finally, τ_{cr}^* was calculated from the iterated solution of τ_{cr} for each grain that experienced sufficient shear stress for entrainment. The 3D model was run both with and without the cohesion term on the R2 sample to demonstrate how the cohesion term affects estimates of τ_{cr}^* when fines are present.

To provide a comparison of our vector-based 3D entrainment model with the 2D model of Kirchner et al. (1990), 2D metrics were also computed for all surface grains, denoted as any grain visible when looking directly down onto the R1 sample from overhead. Vertical 2D image slices were selected for each surface grain whereby each slice was parallel to

downstream flow and contained the (x, y, z) coordinates for the 3D grain's centre of mass (see Figure 3e). Each slice was used to obtain the (x, z) coordinates for the contact 508 point and the grain tops for each object grain and its upstream neighbouring grain. 509 The 2D moment arm vector $\mathbf{R} = \Delta x \hat{\mathbf{i}} + \Delta z \hat{\mathbf{k}}$ was calculated, where Δx and Δz are 510 the distances from the contact point to the grain's centre of mass. Grain diameters (b-axis and c-axis) were determined by the 3D PCA routine outlined above. D_{50} and 512 D_{84} were obtained from sieving analysis of the R1 sample, and are used to calculate the empirical protrusion parameters and the local bed area, respectively. Empirical protrusion 514 parameters were calculated by direct measurements from the vertical 2D image slices: projection, p, was defined as the vertical distance from the top of the object grain to the 516 mean bed elevation that was calculated for the 3D results; and exposure, e, was defined 517 as the vertical distance from the top of the object grain to the top of the upstream 518 neighbouring grain. Theoretical protrusion parameters for projection, p, and exposure, e, 519 were calculated for 2D entrainment based on both b- and c-axis lengths (Equation 1c). 520 The local bed area was defined as the D_{84} distance upstream and downstream of the 521 3D bounding box for the grain, the latter of which was found using image processing 522 functions. The mean bed elevation is thus the average elevation of all surface elevations 523 within this bounded area. 524

Results

Our new vector-based 3D grain entrainment model and the protrusion-based 2D grain entrainment model of Kirchner et al. (1990) were used to calculate pivot angle, α , projection, p, and dimensionless critical shear stress, τ_{cr}^* , for 28 viable surface grains from the R1 sample, allowing comparisons to be drawn between 2D and 3D models. To ensure that 2D versus 3D treatment of geometry and protrusion were compared rather than the input data alone, we ran the 2D model based on empirical protrusion in addition to those based on theoretical protrusion (Equation 1c). Empirical 2D parameters were measured p and e values taken directly from processed XCT images. We further assessed 2D-to-3D ratios of model parameters for α , p, and τ_{cr}^* where the 2D models were based on b-axis, c-axis and empirical measurements for p and e. Calculations of τ_{cr}^* , both with and without the cohesive force given by the second term in Equation 15c, were made for 40 grains in the R2 sample, allowing direct assessment of cohesive effects. Specific grains from the R1 sample were selected to highlight entrainment metric differences under various complicated grain arrangements and orientations. We used both samples to examine whether water-worked grains lay flat. Finally, we explored the relationship between τ_{cr}^* and p both without and with cohesive force using R1 and R2, respectively.

Comparison of the 2D and 3D grain entrainment models

R1 sample 2D entrainment metrics calculated using theoretical protrusion parameters based on b- and c-axis grain lengths, 2D entrainment metrics using empirical protrusion measurements, and our 3D entrainment metrics are shown in Table 1.

Distributions for α , p, and τ_{cr}^* were lognormal for all 2D and 3D entrainment metrics. 546 Statistics associated with lognormal distributions are the geometric mean, $\hat{\mu}_G$, and standard deviation, $\hat{\sigma}_G$, which are given in Figure 4 (Limpert et al., 2001; McAlister, 1879). 548 Quantile intervals (QIs) for lognormal values, $\hat{\mu}_G * \hat{\sigma}_G = [\hat{\mu}_G \times \hat{\sigma}_G, \hat{\mu}_G \div \hat{\sigma}_G]$, are the geo-549 metric counterpart to $\hat{\mu} \pm \hat{\sigma}$ used for normal values (Limpert et al., 2001) and are given for 550 comparison of dispersion about $\hat{\mu}_G$. A random variable X is said to have a lognormal dis-551 tribution if $\log(X)$ has a normal distribution, denoted $\log(X) \sim N(\mu, \sigma^2)$, with mean μ 552 and variance σ^2 (Casella and Berger, 2001). Since the 2D-to-3D ratios were also lognormal, 553 we can test equality of our 3D metrics with each set of 2D metrics through 2D/3D = 1, and take its log-transform, $\log(2D) - \log(3D) = 0$. Since $\log(2D) - \log(3D) \sim N(\mu_d, \sigma_d^2)$ 555 with mean $\mu_d = E(\log 2D - \log 3D)$ and variance $\sigma_d^2 = Var(\log 2D - \log 3D)$, we can use a paired sample t-test to check differences between various 2D models and our 3D model 557 by H_0 : $\mu_d = 0$ vs H_A : $\mu_d \neq 0$. A paired difference is a blocking method that removes correlated variation due to particle-to-particle variability (e.g. size, shape) allowing only 559 model differences to be tested (Ott and Longnecker, 2001).

Table 1: Entrainment metrics for 2D and 3D models calculated for the R1 sample ($D_{50} = 23.0 \text{ mm}$, $D_{84} = 37.4 \text{ mm}$).

Gr	Grain	Axis	Axis Length (mm)	(mm)	Prc	Projection	$p \pmod{p}$	n)		Critical S	Shear, τ	*	Pivot,	α (°)	\overline{POR}	(0)
No		a-axis	b-axis	c-axis	2D-b	2D-c	2D-е	3D	2D-b	2D-c	2D-e	3D	2D	3D	β	>
_	-	59.4	46.3	34.0	35.1	25.8	21.0	22.80	0.073	0.175	0.199	0.184	59.9	90.4	4.4	15.0
2	2	80.8	54.6	46.8	49.8	43.1	19.8	22.20	0.031	0.069	0.166	0.848	43.9	25.6	84.2	7.7
ಣ	က	30.5	28.1	16	12	5.2	12.6	15	0.486	1.469	0.451	0.134	82.2	11.6	80.9	15.5
4	4	45.0	37.8	20.3	40.6	24.0	19.2	19.20	0.019	0.095	0.259	0.071	26.6	34.1	11.0	13.4
ರ	ಬ	43.1	25.2	18.8	11.1	7.4	3.6	7.20	0.525	0.926	2.421	0.730	80.5	123.8	28.3	12.4
9	9	33.3	27.2	22.3	25.2	21.1	19.2	19.2	0.063	0.149	0.137	0.044	47	18.9	34.4	14.6
7	-	29.9	26.5	21.1	9.1	6.3	12	15	0.709	1.132	0.501	0.262	86.8	97.3	47.4	10
∞	∞	78.1	40.0	29.7	21.2	14.9	19.2	21.00	0.229	0.398	0.284	0.168	77.8	52.2	38.2	10.9
6	6	27.6	20.7	13.4	11.3	9.9	13.2	13.80	0.445	1.067	0.471	0.171	72.9	59.3	25.0	13.6
10	10	43.6	39.5	24.1	36.4	23.2	20.4	21.00	0.041	0.154	0.297	0.065	45.0	30.8	4.1	4.1
11	11	25.1	19.0	15.6	24.9	21.5	10.8	10.80	0.005	0.015	0.035	0.061	5.2	12.9	18.4	22.9
12	12	36.8	32	20.4	36.6	25.3	4.8	9.6	0.014	0.055	1.574	0.544	18.4	72.9	58	4
13	13	23.7	17.7	13.1	3.4	1.1	1.2	2.40	2.672	17.421	14.194	3.420	0.06	81.6	50.9	5.5
14	14	45.5	30.1	22.7	14.8	10.4	14.4	14.4	0.353	0.605	0.357	0.282	78.7	8.77	43.3	4.8
15	16	78.9	73.5	44.5	60.2	37.0	8.4	10.20	0.037	0.121	0.758	0.175	53.1	162.4	16.1	56.6
I	17	33.0	21.3	15.1	-15.8	-16.0	1.2	1.80	8	8	14.215	5.187	161.6	164.6	0.6	43.1
16	18	36.4	30.5	20.1	30.8	21.5	13.8	15.00	0.037	0.132	0.306	0.232	38.3	17.5	9.92	10.5
17	19	44.5	41.1	27	22.1	13.5	15	15.6	0.215	0.446	0.401	0.226	77.3	54.6	54.6	8.2
18	23	43.5	30.0	17.3	17.2	0.6	9.9	7.80	0.253	0.711	1.066	0.615	73.1	78.8	33.7	4.0
I	24	26.6	18.6	7.9	1.2	-3.4	2.4	2.40	13.457	8	5.282	5.695	97.1	43.6	82.4	42.8
19	25	18.0	13.3	10.1	4.4	2.5	4.8	5.40	1.814	4.128	1.608	0.961	7.62	51.9	48.0	41.7
20	26	37.1	27.7	16.3	16.4	8.9	5.4	09.9	0.257	0.719	1.282	0.941	71.6	93.3	51.7	13.9
21	27	16.6	15.4	12.4	10.3	8.1	1.2	1.20	0.430	0.767	12.229	8.575	65.2	92.8	32.7	31.0
22	28	69.3	6.09	41.1	55.5	38.4	7.2	8.40	0.028	0.088	0.538	1.672	43.3	22.3	81.3	25.1
23	30	55.4	38.9	27.2	20.8	13.6	1.8	8.40	0.230	0.440	5.790	0.247	77.3	173.6	39.4	33.5
I	32	28.8	28.1	20.7	-15.5	-15.8	0.0	09.0	8	8	8	23.635	160.7	141.3	1.8	7.1
I	36	54.2	29.6	24.4	-1.6	-3.1	1.2	4.20	8	8	14.875	2.835	114.6	13.2	81.5	16.5
24	39	31.0	29.2	17.3	10.1	3.9	12.0	12.60	0.639	2.170	0.522	0.271	88.0	47.8	49.8	14.7
25	41	35.8	21.2	15.0	5.1	2.0	9.0	09.0	1.483	5.913	49.791	20.209	0.06	125.9	35.7	1.5
26	51	26.5	17.8	14.6	10.3	8.1	1.2	1.80	0.486	0.787	8.917	7.1111	70.3	112.6	11.3	32.1
27	22	22.8	16.5	13.0	14.6	11.8	2.4	2.40	0.156	0.369	4.121	2.290	53.1	17.7	59.9	2.6
28	69	29.0	18.9	12.2	21.6	15.4		3.60	0.042	0.162	7.129	2.278	32.7	23.5	31.6	44.5
Value	s in b	Values in bold font indicate grains illustrated in	indicate	grains illu	strated	in Figure		6. Blank Grain	n No are	not listed	ed in Fig	in Figure 4 due	to infinite 2L	nite 2D	τ_{cr}^* values	es.

Grain sizes were determined by the PCA method. Metrics labelled '2D-b', '2D-c' and '2D-e' are calculated from b-axis, c-axis and empirically derived lengths (from XCT images), respectively. Grain No is used in Figure 4; Grain ID is used in Figure 6.

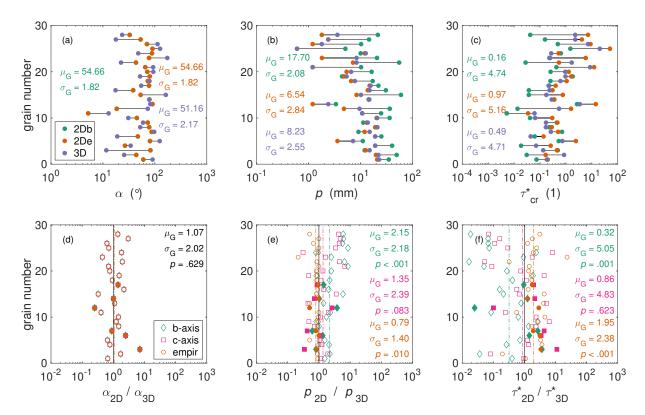


Figure 4: (a-c) Parameters for 2D Kirchner model using b-axis theoretical (2Db, green) and empirically derived (2De, orange) p and e values compared with our 3D entrainment model (3D, violet) across 28 individual grains in the R1 sample for: (a) pivot angle, α (some symbols hidden), (b) projection, p, and (c) dimensionless critical shear stress, τ_{cr}^* . (d-f) Parameter ratios 2D-to-3D of individual surface grains using the 2D Kirchner model for theoretical b-axis (green diamonds), c-axis (fuchsia squares), and empirical (orange circles) p and e values. Ratios are shown for: (d) pivot angle, $\alpha_{\rm 2D}$ -to- $\alpha_{\rm 3D}$, (e) projection, $p_{\rm 2D}$ -to- $p_{\rm 3D}$, and (f) dimensionless critical shear stress, $\tau_{\rm 2D}^*$ -to- $\tau_{\rm 3D}^*$. Geometric mean, $\hat{\mu}_G$, and standard deviation, $\hat{\sigma}_G$, are shown along vertical lines for geometric means (coloured dot-dash) relative to unity (solid black). p-values for paired sample t-test of lognormal ratios are indicated. Solid symbols highlight Grain IDs 3, 6, 7, 12, 14, and 19, which are analysed further below and in Figure 6. Four surface grains were excluded as obtuse 2D pivot angles generates infinite τ_{cr}^* values when using Equation 5.

Here we examine parameter values for α , p, and τ_{cr}^* (Figure 4a-c) using entrainment calculations from the 2D b-axis theoretical method (2Db, green), the 2D empirical method (2De, orange), and our 3D vector entrainment method (3D, violet). Comparison p-value results from paired sample t-tests of log-transformed 2D-to-3D parameter ratios for α , p, and τ_{cr}^* (Figure 4d-f) are from the 2D b-axis theoretical method (green diamonds), the 2D c-axis theoretical method (fuchsia squares), and the 2D empirical method (orange circles), each compared with our 3D vector method. Four R1 sample grains were excluded from Figure 4 due to infinite 2D τ_{cr}^* values, which occurs whenever pivot angles are high enough

to generate negative projections (Equation 1c).

The pivot angles, α , have aggregate similarity as inferred by a geometric mean of $\hat{\mu}_G = 1.07$ 570 for the α_{2D} -to- α_{3D} ratio (Figure 4d). There is moderate scatter around the geometric 571 mean, $\hat{\sigma}_G = 2.02$, because whereas the 2D model assumes that the grain pivots around a 572 single contact point in the downstream direction, the 3D model allows for pairs of contact 573 points in a wide range of orientations. The 2D (b, c, & e) and 3D QIs for α are [30.0, 99.5] 574 and [23.6, 111.0], respectively (Figure 4a). This implies that 3D entrained grains can pivot 575 about a wide range of pivot angles within the POR, as the POR itself diverges from flow 576 direction (the bearing angle) and gravity (the tilt angle). For any given grain, there is no 577 systematic pattern as to whether a α_{2D} or α_{3D} value is higher.

Projections calculated for 2Db are greater than those for 3D, with a geometric mean for 579 p_{2D} -to- p_{3D} ratios of $\hat{\mu}_G = 2.15$ and only five grains having $p_{2D}/p_{3D} < 1$. This difference is amplified in the calculation of τ_{cr}^* , with a relatively high 2Db projection leading to 581 relatively lower τ_{cr}^* values; hence τ_{2D}^* -to- τ_{3D}^* ratios have a geometric mean of $\hat{\mu}_G = 0.32$ 582 with considerable scatter relative to the mean, $\hat{\sigma}_G = 5.05$ (Figure 4f). The 2Db and 3D 583 QIs for p are $[8.5 \,\mathrm{mm}, 36.8 \,\mathrm{mm}]$ and $[3.2 \,\mathrm{mm}, 21.0 \,\mathrm{mm}]$, respectively (Figure 4b), differing 584 by +163% and +75% at the lower and upper bounds, respectively. This discrepancy at 585 the endpoints is reflected in the corresponding 2Db and 3D QIs for τ_{cr}^* at [0.033, 0.750] 586 and [0.105, 2.328], respectively, differing by an average of about -112% at the endpoints. 587

The difference in 2Db and 3D projection values reflects the fact that the 2Db projection is calculated under the assumption that the grain is sitting on a uniform bed of D_{50} grains (Equation 1, Kirchner et al. 1990), and does not include the increased sheltering effect of larger grains, something that is explicitly accounted for with our 3D exposure factor. In the 2Db model, the entrained grain is assumed to be circular, with a diameter equal to the b-axis of the original grain. We also calculated the 2D p and τ_{cr}^* values using the c-axis instead, 2Dc, which reflects the fact that water-worked grains are typically arranged with their c-axis vertical (Komar and Li, 1986), and thus c-axis may be more relevant for calculating grain protrusion used to obtain entrainment forces of drag and lift.

Using the c-axis reduces the p_{2D} -to- p_{3D} ratio mean from $\hat{\mu}_G = 2.15 \, \text{mm}$ to $\hat{\mu}_G = 1.35 \, \text{mm}$ 597 (Figure 4e), and almost triples the τ_{2D}^* -to- τ_{3D}^* ratio mean from $\hat{\mu}_G = 0.32$ to $\hat{\mu}_G = 0.86$ 598 (Figure 4f). Model 2Dc had a geometric mean and standard deviation of $\hat{\mu}_G = 11.1 \,\mathrm{mm}$ 599 and $\hat{\sigma}_G = 2.44 \,\mathrm{mm}$ for p, and $\hat{\mu}_G = 0.425$ to $\hat{\sigma}_G = 4.54$ for τ_{cr}^* (not shown in Figure 4). The 600 2Dc model QIs for p and τ_{cr}^* are [4.6 mm, 27.1 mm] and [0.094, 1.932], respectively. The p 601 endpoint discrepancy between 2D and 3D QIs decreases from $+163\,\%$ and $+75\,\%$ for 2Db 602 values to +41% and +29% for 2Dc values at the lower and upper bounds, respectively. Similarly the τ_{cr}^* endpoint discrepancy between 2D and 3D QIs increases from -214%604 and $-210\,\%$ for 2Db values to $-12\,\%$ and $-21\,\%$ for 2Dc values. This reduction in p and τ_{cr}^* QI endpoint discrepancy from 2Db to 2Dc models, each relative to the vector-based 606 3D model, suggests 2D calculations using the c-axis may be more appropriate than using 607 the b-axis. 608

As an attempt to place 2D calculations on an equal footing with our 3D entrainment 609 calculations, we used direct measurements of projection, p, and exposure, e, to derive an 610 empirical entrainment model, 2De. p_{2D} -to- p_{3D} ratios for 2De are 27 % smaller than those 611 for 3D (Figure 4e) with geometric metrics of $\hat{\mu}_G = 0.79$ and $\hat{\sigma}_G = 1.40$, respectively. This 612 small difference nearly doubles τ_{cr}^* values for 2De relative to our 3D model (Figure 4f) 613 with geometric metrics for τ_{2D}^* -to- τ_{3D}^* ratios of $\hat{\mu}_G = 1.95$ and $\hat{\sigma}_G = 2.38$, respectively. 614 The 2De model QIs for p and τ_{cr}^* are [2.3 mm, 18.5 mm] and [0.187, 4.982], respectively 615 (Figure 4bc). The p endpoint discrepancy between 2D and 3D QIs increased from -40%616 and $-13\,\%$ for 2De values to $+41\,\%$ and $+29\,\%$ for 2Dc values at the lower and upper 617 bounds, respectively. Similarly the τ_{cr}^* discrepancy between 2D and 3D QIs decreases from $+78\,\%$ and $+114\,\%$ for 2De values to $-12\,\%$ and $-21\,\%$ for 2Dc values. The 2D-to-3D 619 discrepancy shift from 2De to 2Dc is a much larger swing than that from 2Db to 2Dc, the former having a respective total percent change of 81 % and 42 % for lower and upper 621 bound of p and 90 % and 135 % for that of τ_{cr}^* . 622

Paired sample t-test of log-transformed 2D-to-3D parameter ratios were performed to evaluate mean ratio differences from unity of various 2D methods and our 3D method (Figure 4d-f). There were insignificant differences (at the 5% level) in α ratios across all methods. p_{2D} -to- p_{3D} and τ_{2D}^* -to- τ_{3D}^* ratios for 2Dc are insignificantly different from unity whilst those for 2Db and 2De have significant differences. We infer from paired ratio t-test that 2D estimates of τ_{cr}^* using the c-axis (p = .623) is a more suitable proxy for 3D grain entrainment calculations compared to the commonly used b-axis (p = .001) or using empirical measurements (p < .001).

Effect of the cohesion term on critical shear stress

We use the data from the R2 sample, which contained additional fine sediment, to in-632 vestigate the effect of our cohesion term in the 3D entrainment model (Equation 15c), 633 by calculating τ_{cr}^* with and without the cohesion term included. On average, the cohe-634 sion term approximately doubles the aggregate τ_{cr}^* (Figure 5), but there is a considerable 635 spread in the data. Eight out of 40 surface grains are unaffected by cohesion, whereas for one grain (stone no. 35) τ_{cr}^* increases nearly ten fold when adding the cohesion term. This 637 variation is because different grains in the sample have different matrix contact areas, with some grains on the surface of our sample being in contact with no matrix, whereas those 639 that are buried deeper might have more stone surface area contact with matrix fines. The scatterplot of $\tau_{F_c=0}^*/\tau_{F_c\neq 0}^*$ ratios suggests this to be the case as the grain numbers shown 641 are ordered from highest to lowest surface elevations (Figure 5b).

Analysis of specific exemplar grains from the XCT image

To demonstrate the veracity of the new 3D model we look in detail at six exemplar grains from the R1 sample. Table 1 lists 2D and 3D entrainment calculations for these exemplar stones (shown in bold font). Figure 6 shows the 3D representation of these grains from the XCT image, each grouped with their left and right contact grains. Qualitative characteristics of each exemplar stone, relative to theoretical protrusion values based on b-axis calculations, are given in Table 2 and briefly discussed below.

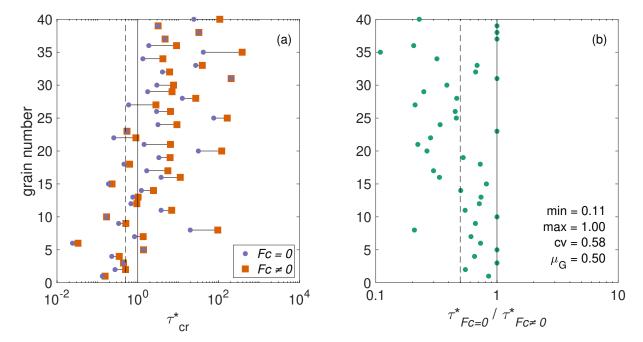


Figure 5: The effect of the cohesion term in our 3D entrainment model (Equation 15c) on resultant τ_{cr}^* for R2 (with fines): (a) comparison of 3D modelling across individual grains with cohesive effects (orange squares) and without cohesive effects (violet circles), and (b) ratio of results without cohesion to those with the cohesion term added. Descriptive statistics for minimum, maximum, coefficient of variation, and geometric mean are shown along with vertical lines for geometric mean (dashed) relative to unity (solid black).

Grains 3 and 7 have 3D τ_{cr}^* values that are considerably less than the 2D equivalents, but for different reasons. Grain 3 shares a downstream contact point between the 2D and 3D 651 models, but the additional contact point for the 3D model (the right contact) lies below 652 the grain, resulting in a low pivot angle of 11.6° in the 3D model relative to high pivot 653 angle of 82.2° in the 2D model. Alternatively, the lower 3D value of τ_{cr}^* for grain 7 is 654 attributed to the larger projection of 15.0 mm for 3D compared to 9.14 mm for 2D, as 655 the pivot angle used for the 2D and 3D models are similar. Grain 12 requires a higher τ_{cr}^* 656 using the 3D model compared to the 2D implementation (over 38 times higher than 2D). 657 The combined effect of much lower projection ($\sim 74\%$ relative difference) and much larger 658 pivot angle (about 3 times) for the 3D model relative to the 2D model contributes to the 659 large in τ_{cr}^* difference. The two downstream grains visibly making contact with Grain 660 12 are not accounted for in the 2D model since neither are included in the 2D vertical 661 plane passing through the Grain 12 centre of mass. Grains 6, 14 and 19 have similar τ_{cr}^* 662 values for the 2D and 3D entrainment models, however we note that this is not because

Table 2: Qualitative comparison of metrics for select R1 sample grains shown in Figure 6.

				Grai	n ID		
Metric	Comparison	3	6	7	12	14	19
Critical Shear, τ_{cr}^*	2D > 3D	X		X			
	2D < 3D				X		
	$2D \approx 3D$		X			X	X
Projection, p	2D > 3D				X		
	2D < 3D	X		X			
	$2D \approx 3D$		X			X	X
Pivot Angle, α	2D > 3D	X	X				X
	2D < 3D				X		
	$2D \approx 3D$			X		X	
Bearing Angle, β	$\beta < 38.2^{\circ}$		X				
	$\beta \geqslant 38.2^{\circ}$	X		X	X	X	X
$\overline{\text{Tilt Angle}}, \gamma$	$\gamma < 13.6^{\circ}$			X	X	X	X
	$\gamma \geqslant 13.6^{\circ}$	X	X				
Shared Contact	Yes	X	X	X			X
Between 2D & 3D	No				X	X	
Complicated	Yes	X				X	X
Arrangement	No		X	X	X		

Median angle values for R1 are used to partition high and low bearing and tilt angle values. Comparisons are based on b-axis calculations (2D-b in Table 1).

the 2D model accurately accounts for entrainment geometry, rather interacting metrics counterbalance one another to coincidentally arrive at similar estimates. For example 665 grain 14 has a wide angle between contact points, reducing the distance between the 666 centre of mass and the AOR in the 3D model relative to the 2D, but this reduction in 667 length of both position vectors is offset by a moderately high bearing angle. Grains 6 668 and 19 have lower pivot angles in the 3D model implementation, but these are offset by 669 relatively high bearing angles. These exemplar grains illustrate how 3D grain orientation 670 combined with contact grain arrangement affects adjustments in τ_{cr}^* values in ways that 671 are far too complex for 2D modelling.

Do water-worked grains lay flat?

There has been some research into whether grains are preferentially deposited with either the b- or c-axis aligned with gravity (Komar and Li, 1986). Our XCT derived grain

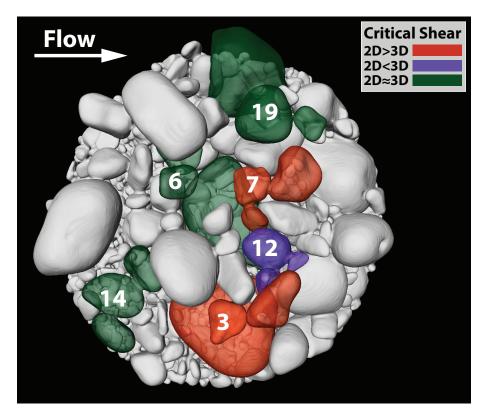


Figure 6: 3D visualisation of surface grains for the R1 sample showing examples where τ_{cr}^* is estimated as being: (i) higher by the 3D model compared to the 2D model (Grain ID 12), (ii) lower by the 3D model compared to the 2D model (Grain IDs 3 and 7), and (iii) approximately equal for both 2D and 3D models (Grain IDs 6, 14 and 19).

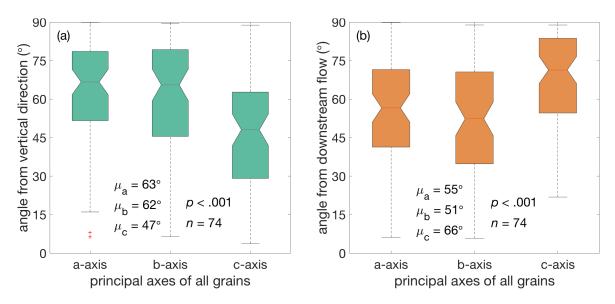


Figure 7: Distribution of angle separation between principal a-, b- and c-axis from: (a) the vertical direction (green boxes), and (b) the downstream flow direction (orange boxes), for all surface grains in the R1 and R2 samples, where 0° represents a perfect alignment of a given principal axis with the named direction on the y-axis of the box plot.

geometries from the R1 and R2 samples suggest that there is considerable variation in 676 our water-worked grains (Figure 7) but a general preference for them to be positioned 677 with their c-axis more closely aligned to vertical than either of the other two principal 678 axes. The boxplot in Figure 7a indicates a c-axis alignment with gravity is within about 679 45° for about half of all grains whilst the a-axis and b-axis fall within about 65° for 680 about half the grains. Similarly, results in Figure 7b show the c-axis alignment with the 681 downstream direction are up to as much as 75° for about half of the grains. This result further supports our earlier conjecture that c-axis lengths would be a better proxy upon 683 which to base grain entrainment calculations than b-axis lengths.

Grain projection relationship to au_{cr}^*

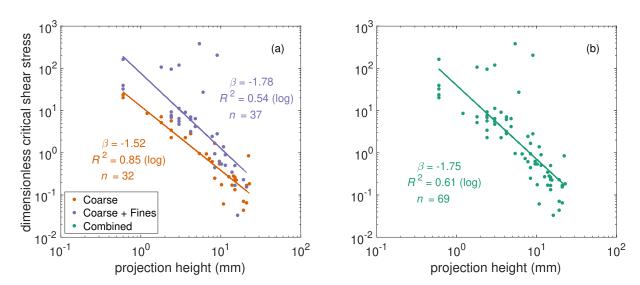


Figure 8: Power law models of τ_{cr}^* versus projection height of the entrained grain for: (a) coarse stones from the R1 sample (orange circles) and coarse plus added fines from the R2 sample (violet circles), and (b) stones from combined R1 and R2 samples (green circles). Model regressions for R1 and R2 samples were significantly different (p < .001). Critical shear stress for coarse plus fines scales with projection as -1.78 similar to that of the combined sets as -1.75.

Grain protrusion (and likely projection, p) is known to have a fairly strong correlation with τ_{cr}^* (Fenton and Abbott, 1977). Data from both the coarse R1 and coarse plus fines R2 samples show a strong to moderate power law relationship between τ_{cr}^* and p with 85% and 54% explained variation in $\log(\tau_{cr}^*)$, respectively (Figure 8a). Dummy-variable regression of linearised power law models combines R1 and R2 data for testing linear model differences (Draper and Smith, 1998). To perform this test, let $\log (\tau_{cr}^*) = \alpha + \beta \log (p) + \epsilon$ be the linearised power law for: (i) Coarse, and (ii) Coarse + Files models. Let dummy variable Z=1 for the coarse grains model and let Z=0 otherwise. Then the combined model that allows us to simultaneously test for equal slope, $\alpha \equiv \log (\alpha')$, and scaling, β , parameters is $\log (\tau_{cr}^*) = \alpha_1 + \alpha_2 Z + \beta_1 \log (p) + \beta_2 Z \log (p) + \epsilon$, or written in parsed form

$$\log (\tau_{cr}^*) = \begin{cases} (\alpha_1 + \alpha_2) + (\beta_1 + \beta_2) \log (p) + \epsilon & \text{Coarse } (Z = 1), \\ \alpha_1 + \beta_1 \log (p) + \epsilon & \text{Coarse } + \text{Fines } (Z = 0), \end{cases}$$

where the error term assumption is $\epsilon \sim N(0, \sigma^2)$. Using a partial F-test to test the 696 hypothesis H_0 : $\alpha_2 = 0$ and $\beta_2 = 0$ vs H_A : $\alpha_2 \neq 0$ or $\beta_2 \neq 0$, model parameters are 697 significantly different (p < .001). The power law models suggest τ_{cr}^* is higher for grains in 698 contact with matrix fines, and yet τ_{cr}^* differences between models decreases with increased projection. τ_{cr}^* for coarse plus fines scales faster ($\beta = -1.78$) than coarse grains alone 700 $(\beta = -1.52)$. This is likely the result of less surface contact with matrix fines as the 701 grains project further from the river bed. Cohesion comparison (Figure 5) suggests this 702 might be the case as the lowest grain numbers in the plot are stones positioned at the highest elevations. Plotting the combined R1 and R2 samples (Figure 8b) suggests matrix 704 fines may play a dominant role as the scaling parameter ($\beta = -1.75$) is similar to R2.

Discussion

Sources of geometric error inherent in the 2D model

The wide-ranging differences in entrainment parameter values between the 2D and 3D models are primarily attributed to a number of geometric factors: (i) the 2D model is a special case of the 3D model with a range of motion restricted to a 2D flow-parallel plane, and where the AOR coincides with a single contact point, (ii) in the 3D model,

the AOR (and therefore the POR) orientation is induced by variations in contact vector 712 pair arrangements affecting entrainment results, (iii) variations of position vector lengths 713 induced by contact pair arrangements affect the magnitude of the position vector and 714 its contribution to the moment balance, and (iv) changes in force vector magnitude and 715 direction induced by variations in POR orientation also affect moment balance. 716 factors affect how forces are geometrically transferred through the POR into the rotational 717 mechanics of entrainment, and each are discussed below by considering a 2D entrainment model with fixed force and position vectors in static equilibrium about a single contact 719 point (Figure 1a). We first show that the 2D model is a special case of the 3D model, case (i) above, and then we proceed to use it in a geometric argument to show the remaining 721 cases. When combined, these factors show the geometric biases inherent in a 2D model. 722 Given the careful consideration in our attempt to match the continuum mechanics for 723 2D and 3D entrainment models, we suggest these 2D geometric biases as the primary difference between 2D and 3D entrainment models. 725

For this discussion, consider a vector-based 2D entrainment model with fixed vectors in 726 static equilibrium about a single contact point (Figure 1a). Consider the xz-plane as the POR for the 2D model (shown in Figure 1a). Now by the right hand rule for reference 728 frames, the AOR has unit vector $\lambda = -\hat{\mathbf{j}}$. Reversing the direction of the single contact vector, **C**, in Figure 1a, the position vector for the 2D model is $\mathbf{R} = -\frac{1}{2}D\sin\alpha\hat{\mathbf{i}} + \frac{1}{2}D\cos\alpha\hat{\mathbf{k}}$. 730 The force vectors for drag, lift and submerged grain weight are $\mathbf{F}_D = F_D \hat{\mathbf{i}}$, $\mathbf{F}_L = F_L \hat{\mathbf{k}}$ 731 and $\mathbf{F}_W = -F_W \hat{\mathbf{k}}$, respectively. Using the matrix determinant method to calculate the 732 scalar triple product, we arrive at the same 2D moment balance as Equation 1b, thereby showing that the 2D entrainment model is just a special case of the 3D model restricted 734 to a 2D subspace in the form of a flow-parallel plane.

Now the POR of our fixed 2D model contains all of the force vectors involved in the 2D moment balance. By parsing 2D force vectors into those with horizontal lines of action (drag force) and those with vertical lines of action (lift and submerged weight), we can examine the effects of bearing and tilt angles separately. By increasing the bearing angle,

 β , from the 2D position, thus reflecting a changing pair of contact vectors in 3D, the 740 increased POR deflection from downstream flow (x'-axis in Figure 2a) has no effect on 741 the contribution from vertical forces since those vectors remain within the POR. However, the drag force is effectively reduced by a factor of $\cos \beta$ as the POR swings away from the 743 horizontal line of action (vector \mathbf{F}_{D_P} in Figure 2a). For the moment to remain in balance, the boundary shear stress must be increased to compensate for the loss in driving force 745 being projected onto the POR, which is caused by an increased bearing angle. Conversely, increasing tilt angle, γ , from the 2D position for the POR has no effect on contribution 747 of the horizontal force on entrainment, but the tilted POR does affect how vertical force contributes to overall entrainment, as vertical forces are effectively reduced by a factor of 749 $\cos \gamma$. Again, this requires an adjustment in the boundary shear stress for the moment 750 to remain in balance. However, since the resultant of vertical forces, $(F_L - F_W)\hat{\mathbf{k}}$ largely 751 cancelled out by their difference, any increase in bed shear stress contributes more towards 752 the magnitude of drag force in the total entrainment, $\mathbf{F}_E = \mathbf{F}_D + \mathbf{F}_L$, which is also partly 753 attributed to the fact that drag and lift forces scale proportionally to their coefficients, 754 $C_D = 0.91$, and $C_L = 0.20$, respectively. 755

By the definition of cross product of position and force vectors, $\mathbf{R} \times \mathbf{F} = \|\mathbf{R}\| \|\mathbf{F}\| \sin \theta \hat{\mathbf{n}}$, 756 the magnitude of the moment, $\|\mathbf{R} \times \mathbf{F}\|$, is reduced whenever the magnitude of either 757 position or force vector is reduced, or whenever the angle between their vectors diverges 758 from an orthogonal orientation. Consider the location of the AOR shown in Figure 2b 759 (blue spiral arrow), which shows the AOR passing through the body of the stone rather 760 than rotating at its edge. Acute angles between contact vectors places the AOR near the edge of the stone (e.g. Grain 7 in Figure 6) whilst very obtuse angles places the AOR 762 much closer to the stone's centre of mass (e.g. Grain 14). Increasing the angle separating 763 contact vectors on the same grain reduces the length of both position vectors (Figure 2b) 764 as the AOR shifts towards the centre of mass thus affecting the magnitude of the moment 765 balance for resisting and driving forces. 766

Consider again increasing the bearing angle of the POR from a 2D position, which reduces

the magnitude of the drag force vector projection on the POR. The reduced drag force 768 projection (Figure 2a,b) results in an entrainment vector, $\mathbf{F}_{E_P} = \mathbf{F}_{D_P} + \mathbf{F}_{L_P}$, that changes 769 in both magnitude and direction within the POR (i.e. \mathbf{F}_{D_P} decreases as \mathbf{F}_{L_P} remains 770 fixed). As bearing angle increases, the angle between the entrainment vector, \mathbf{F}_{E_P} , and 771 its position vector, \mathbf{R}_{E_P} , (shown in Figure 2b) increases from slightly acute to slightly obtuse as the drag force projection decreases in magnitude. This result corresponds to 773 increasing angle θ in the cross product $\|\mathbf{R}_{E_P}\| \|\mathbf{F}_{E_P}\| \sin \theta \hat{\mathbf{n}}$ from an acute angle to the maximum $\|\mathbf{R} \times \mathbf{F}\|$ value when vectors become orthogonal only to decreasing again with 775 increasing obtuse angle. Any of these changes in vector magnitude, direction, or angle between vectors requires adjustments in the boundary shear stress to bring the moment 777 back into static equilibrium. 778

When combined, these geometric factors affecting moment balance form the bulk of entrainment parameter differences between 2D and 3D entrainment models (see Figure 4).

Moreover, using a vector-based 3D framework to calculate static moments from scanned
images of the actual grain structure yields precise results with respect to the rigid body
portion of the 3D entrainment model. Thus, any 2D-to-3D differences in entrainment
values on the same grain are primarily attributed to a measurable, geometric bias, which
is inherent in any 2D model. The preceding arguments also demonstrate friction angles
obtained in the field are actually an amalgamation of angles for pivot, bearing, and tilt.

Sources of error in the 3D entrainment model

Inaccuracies in the 3D model are primarily due to: (i) simplifying assumptions for the fluid mechanics of the 3D model, (ii) errors related to XCT scanning and image processing, (iii) missing grains and contact points from the images, and (iv) the rule set used to define viable pairs of contact vectors. We attempted to minimize 2D-to-3D model differences with respect to fluid mechanics used in the 3D model so that model differences would primarily be attributed to the rigid body mechanics. Indeed, since fluid mechanics developed for our 3D entrainment model are a discrete analogue of the continuous func-

tions used in the 2D Kirchner et al. (1990) model, they are mechanistically comparable.

In examining both models using the coarse sample of R1, without cohesive influence, and
using the same drag and lift coefficients, we are effectively able to compare the influence
of a 3D framework relative to a 2D framework by controlling for continuum influences.

Errors attributed to XCT scanning and image processing includes quality and resolution 799 of the scan, segmentation algorithm and steps taken to segment grains from matrix, 800 and separation methods and the unavoidable grain erosion that ensues. Although their 801 cumulative error is relatively minor, when compared with geometric errors, they can lead 802 to missing grains and contact points, which can only be quantified through a detailed error 803 analysis that is beyond the scope of this work. Smaller grains are usually the culprits in 804 this regard since larger stones tend to retain their shape and proximity to neighbouring 805 grains post-separation. 806

We developed a rule set for viable contact pairs based on fairly relaxed assumptions with respect to the limits of POR orientation. More investigations into these limits using carefully controlled experiments should establish a physically rigorous set of rules that are likely to be more restrictive. Differences between the new and current rule sets would reveal any error that exists with the current rule set.

Grain orientation and representative axis

Our results show that, for our samples, the Kirchner 2D model typically predicts a lower τ_{cr}^* than our vector-based 3D model. The lower 2D prediction can be explained by two simplifications that are made by the Kirchner et al. (1990) model. First, grain geometry is simplified down to a case of a circular grain sitting on a bed of uniform circular grains. Second, grains are assumed to pivot around a single point in the downstream direction. The Kirchner model, and many similar 2D models, represents grains as a circle with a diameter equal to the grain's b-axis. However, natural grains are almost never spherical, instead having a variety of complex forms that are usually represented as an ellipsoid with

principal axes: a (longest), b (intermediate) and c (shortest). The variety in grain size and shape mean that the protrusion of any grain is determined by the grain's orientation and position relative to those of its neighbouring grains.

Numerous studies have evaluated the preferential orientation of water-worked grains in 824 both experimental and field scenarios. Aberle and Nikora (2006) found that for flume 825 derived armour surfaces, the majority of grains were oriented with their long axis aligned 826 with the flow, which was attributed to interaction dynamics between the bed and constant 827 discharge in the experiment. In contrast, the largest grains, which did not move during 828 water working and armouring, were arbitrarily oriented (Aberle and Nikora, 2006). Other 829 studies have also found flow parallel orientation of the long axes of grains, although a range 830 of causal mechanisms have been proposed, including low transport rates or the presence 831 of static armour layers (Hodge et al., 2013; Powell et al., 2016). 832

Our results cast doubt on the suitability of using the b-axis as a proxy for protrusion when estimating 2D entrainment thresholds, or in roughness calculations more generally. In these situations we recommend use of the c-axis instead of the b-axis since doing so reduces the discrepancy between τ_{cr}^* predicted by the 2D and 3D models by 62.8% as relative difference with respect to c-axis calculations (Figure 4f). Conversion to c-axis could be as simple as estimating a mean for c-to-b axis ratios obtained from a Wolman pebble count (Wolman, 1954) to use as a size multiplier on b-axis grain size distributions.

$_{ ext{\tiny B40}}$ Grain entrainment angles

The pivot (and other) angles that we calculate from our 3D XCT data and model are not limited to a defined range of values as the model does not rely on trigonometric functions, instead using a 3D vector space for calculations. In contrast, the Kirchner et al. (1990) model limits pivot angles to between 0° and 90° (e.g. see Table 1). Consequently we sometimes record pivot angles that are greater than 90°, for example for a grain where the contact points are above the grain's centre of mass. Such a grain will sit mostly

below the mean bed surface, but it is still subject to a component of the driving forces. 847 Although such grains are unlikely to be very mobile, the 3D model enables these grains 848 to be assigned a τ_{cr}^* . Therefore when calculating the range of τ_{cr}^* values over a gravel surface, we can incorporate all grains for which a proportion of their surface can be seen 850 from directly overhead. The inclusion of what might be considered outliers in a 2D grain 851 entrainment implementation account for some of the extreme values and scatter in the 852 data reported for the 3D results (Figures 4 and 5). Being able to calculate τ_{cr}^* for these grains shows that the 3D model can be applied beyond the realm of idealised situations. 854 To assess the impact of cohesion we have used a simple empirical formulation, which 855 shows the potential importance of the term and how easily it can be incorporated into

the 3D model. Cohesion effects will be a function of the matrix GSD, and so further work

is needed to develop and test the cohesion term including a test for field site dependence.

Fluvial response within the roughness layer

856

857

858

We used a logarithmic velocity profile assuming zero velocity at the local mean bed elevation to facilitate model comparison; however, this assumption ignores interfacial flows 861 that occur within the roughness layer. A more appropriate profile model to describe flow 862 over a coarse gravel bed assumes either a linear (Nikora et al., 2004; Mignot et al., 2009) or 863 a power law (Sarkar and Dey, 2010; Dey and Das, 2012) profile within the roughness layer 864 that increases from zero near the bed troughs coupled with a logarithmic profile upon en-865 tering the surface layer. By extracting roughness layer elevation boundaries from images 866 at a distance, say D₈₄, upstream of the grain, a profile may be constructed to accom-867 modate a roughness layer that is unique to each surface grain. These variations in local 868 hydraulic properties affecting the profile should generate more appropriate entrainment metrics across all surface grains in the sample. 870

The potential for grain mobilisation is spatially dependent on surrounding sediment; up-871 stream particles within close enough proximity are found to strongly affect grain entrain-872

ment (Measures and Tait, 2008). For our 3D model we developed an exposure factor that 873 is a physically-based weighted mean of exposed fraction of grain area. XCT scanned im-874 ages are rotated to provide multiple viewing angles of surface grains to mimic turbulence 875 fluctuation of flow velocity over the bed, where the horizontal component of the vector 876 is assumed to contribute to drag force during grain entrainment. Measures and Tait 877 (2008) used surface images of bed sediment to isolate individual grains, extract upstream 878 projected areas and their midpoint elevations, and then used these metrics for individual grain calculations. They considered two mechanisms of sheltering from upstream grains: 880 direct sheltering, which reduces exposed grain area, and remote sheltering, which modifies flow due to particle protrusion (Measures and Tait, 2008). Both mechanisms affect EF882 in our 3D entrainment model since a sheltering grain increases EF values with increased distance from the object grain.

885 Applications of 3D models

Although vector-based 3D entrainment models yield more realistic rolling motion physics than their 2D counterparts, their downside is the impracticality of collecting the 3D 887 scanned images that are necessary for model parameterisation. The time and costs involved in obtaining and processing sediment scans of riverbed samples is currently pro-889 hibitive for routine monitoring applications. The real value for developing vector-based 3D models to interpret scanned images of grains is their potential for improving existing 891 2D models (e.g. switching from b-axis to c-axis entrainment calculations). Furthermore, our finding that grain entrainment is primarily controlled by projection suggests that 893 high-resolution 2.5D measurement techniques such as Terrestrial Laser Scanning (TLS) offer exciting opportunities for in-situ, field based estimates of entrainment using exposure 895 as a proxy for projection.

Here we developed a 3D model that utilised simple constructs typically used in fluid mechanics of 2D entrainment models to facilitate comparison with a simple 2D model. Future development should focus on improving vector-based 3D modelling with respect to the fluid mechanics; three such areas of enhancement are computational fluid dynamics simulation over bed topography (Hardy, 2006; Hardy et al., 2005; Lane et al., 2004), coupling interfacial velocity profiles for the roughness layer with a logarithmic profiles for the surface layer (Nikora et al., 2004; Mignot et al., 2009; Sarkar and Dey, 2010; Dey and Das, 2012; Blois et al., 2014; Cooper et al., 2018), and entrainment effects due to turbulent pressure fluctuations (Amir et al., 2014; Cooper et al., 2018; Vollmer and Kleinhans, 2007). In doing so, a vector-based entrainment model representative of 3D rolling motion physics could serve as a benchmark from which enhancements to existing 2D models can be developed and tested.

909 Conclusion

Kirchner et al. (1990) advocated a better understanding of the variability within gravel 910 bed surfaces, not simply between different locations. We have developed a vector-based 911 3D grain entrainment model that is parameterised using high-resolution XCT data of 912 individual grains within water-worked gravel beds. Our new approach explicitly accounts 913 for variability by making no simplifying assumptions about the grain geometry and loca-914 tion of the contact point(s) for pivoting, instead relying upon accurately measured grain 915 geometries and the spatial position all contact points. Our 3D entrainment model does 916 not rely on pivot angles to calculate critical entrainment shear stress; however, in addition 917 to pivot angles, bearing and tilt angles for the plane of rotation are easily calculated once 918 grain entrainment is established. 919

The 3D model presented here has made geometric entrainment formulations relevant and applicable to real-world data, removing the need for idealised situations and gross assumptions. In the event that one cannot fully resolve 3D subsurface grain geometry, we show that use of the grain c-axis, rather than the oft-used b-axis, for 2D grain entrainment calculations results in critical shear stress estimates that more closely resemble those calculated using the full 3D implementation. Furthermore, the new model ultimately

reveals that entrainment is predominantly controlled by (i) grain projection, (ii) plane of rotation bearing angle, and (iii) cohesion force due to grain contact with a matrix of fines. Whilst protrusion (and as we have shown here, projection) has long been recognised as a dominant control on critical shear stress, the impact of bearing angle and cohesion has not yet been accounted for. We demonstrate major sources of geometric error in 2D entrainment models to show where improves can be made in 2D model development; the magnitude of this error varies with changes to orientation for the plane of rotation.

Grain entrainment is a function of both the overall structure of the gravel bed, and the 933 overlying flow. Our model has significantly improved the representation of the former, but 934 still uses a simplified logarithmic flow profile to represent the latter. Lamb et al. (2008) 935 noted that local average velocity is not the only relevant velocity scale of interest and 936 that the role of turbulent fluctuations on grain entrainment should be considered, with 937 turbulent flow and pressure differentials aiding entrainment of grains from a bed (Vollmer 938 and Kleinhans, 2007). The structure of our vector-based 3D grain entrainment model is 939 such that it could be developed in the future to address the impact of flow properties on 940 grain entrainment. By improving flow properties to better replicate local hydraulics on vector-based 3D models, entrainment modelling of 3D scanned riverbed grains has the 942 potential for use as a benchmark for developing and testing enhancements in existing 2D models.

945 Appendices

A Equivalence of scalar triple products and POR cross products

Here the scalar triple product for axis of rotation (AOR) unit vector, λ , whose dot product with the $\mathbf{R} \times \mathbf{F}$ cross product is shown to be equivalent to projecting both vectors, \mathbf{R} and \mathbf{F} , onto the plane of rotation (POR) before calculating the scalar moment. The scalar triple product effectively produces a signed magnitude of the resulting moment of the

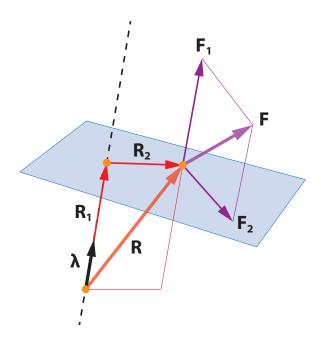


Figure A.1: Image of a plane of rotation (blue plane) and its orthogonal axis of rotation (dotted line). Arbitrary force and position vectors, \mathbf{F} and \mathbf{R} , that induce a moment $\mathbf{R} \times \mathbf{F}$ in 3D space are resolved into components parallel to the axis of rotation, \mathbf{F}_1 and \mathbf{R}_1 , and perpendicular to the axis of rotation, \mathbf{F}_2 and \mathbf{R}_2 , spanning the plane of rotation. The unit vector for the axis of rotation is $\boldsymbol{\lambda}$. Redrawn after Beer et al. (2013).

system. First, resolve each vector for resultant force, \mathbf{F} , and position of applied force, \mathbf{R} , into two components: one parallel to the AOR, \mathbf{F}_1 and \mathbf{R}_1 , and one orthogonal to the AOR, \mathbf{F}_2 and \mathbf{R}_2 , which lie within the POR (Figure A.1). Then the scalar triple product is given by

$$\lambda \cdot \mathbf{R} \times \mathbf{F} = \lambda \cdot (\mathbf{R}_1 + \mathbf{R}_2) \times (\mathbf{F}_1 + \mathbf{F}_2)$$

$$= \lambda \cdot (0 + \mathbf{R}_1 \times \mathbf{F}_2 + \mathbf{R}_2 \times \mathbf{F}_1 + \mathbf{R}_2 \times \mathbf{F}_2)$$

$$= 0 + 0 + 0 + \lambda \cdot \mathbf{R}_2 \times \mathbf{F}_2.$$

The first term of distributed cross products, $\mathbf{R}_1 \times \mathbf{F}_1$, in the second equality is zero because \mathbf{R}_1 and \mathbf{F}_1 are parallel vectors. The second and third terms, $\mathbf{R}_1 \times \mathbf{F}_2$ and $\mathbf{R}_2 \times \mathbf{F}_1$, in the third equality are zero because \mathbf{R}_1 and \mathbf{F}_1 are parallel to unit vector $\boldsymbol{\lambda}$; hence, the cross product of either of these two vectors with any other vector are obviously orthogonal to $\boldsymbol{\lambda}$. The remaining nonzero term is a scalar, equivalent to projecting force and position

vectors onto the POR, \mathbf{R}_2 and \mathbf{F}_2 , evaluating their vector cross product, $\mathbf{R}_2 \times \mathbf{F}_2$, and then calculating the signed magnitude of the result, $\operatorname{sgn}\{\mathbf{R}_2 \times \mathbf{F}_2\} \|\mathbf{R}_2 \times \mathbf{F}_2\|$. The latter, of which, is the consequence of taking the dot product of the unit vector $\boldsymbol{\lambda}$ with the resulting moment vector $\mathbf{R} \times \mathbf{F}$.

B Exposure factor (EF) model and its adjusted value (EF_{adj})

B.1 Exposure factor, EF

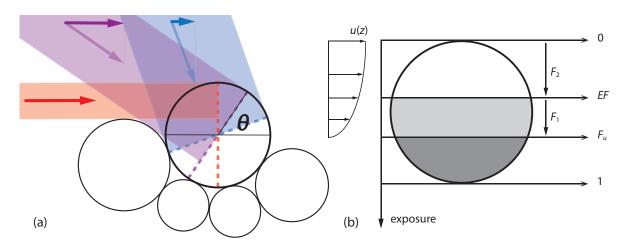


Figure B.1: (a) Development of the exposure factor, EF. For three different view angles, solid coloured bands show the areas of viewable stone whereas the dashed coloured lines indicate total area of the stone that could be viewed in the case of no upstream sheltering grains. Viewing angle range from 0° (directly overhead) to 90° (in the direction of flow). The exposure ratio, ER_{θ} , is the ratio of these two area values. (b) The adjusted exposure factor, EF_{adj} , is necessary because we are only interested in the relative exposure of the areas of the grain that are above the elevation at which the flow velocity is zero. To calculate EF_{adj} , we partition the face of the stone into two initial regions: above the mean bed elevation where the velocity profile is nonzero (white and light grey areas) and below mean bed elevation where flow velocity is zero (dark grey area). The upper region is further partitioned into the hidden fraction (light grey) and the exposed fraction (white) of the nonzero velocity profile.

The exposure factor, EF, for our 3D model is the functional equivalent of grain protrusion used in a 2D model, but is necessary since the elevation of sheltering effects from upstream grains are not laterally continuous across the face of the downstream grain. The exposure factor, $EF \in [0,1]$, is defined as the average proportion of the face of the grain that is exposed to the flow when the grain is viewed over multiple angles. We use the turbulence

of flow over the bed as a physical basis for developing a 3D exposure factor. One way 971 of picturing flow turbulence is to imagine a camera attached to the flow velocity vector, 972 V, that points in the same direction as the vector. Turbulence is a departure from the dominant downstream direction of flow, and so provides different viewing angles, θ , of 974 the bed surface (see colour bands in Figure B.1a). Viewed from overhead at $\theta = 0^{\circ}$, the 975 entire stone is seen (blue band) whereas looking downstream at $\theta = 90^{\circ}$, only a fraction 976 of the stone is exposed (red band). Under turbulent flow regimes, a sphere experiences viscous shear and pressure forces up to the moment of flow separation at about 115° 978 beyond the front of the sphere (Celik et al., 2014; Constantinescu and Squires, 2004) thus we can use overhead views of exposed areas in EF calculations. Let ER_{θ} be the exposure 980 ratio defined as the visible area of exposed stone (solid colour bands) divided by the total 981 area $(A_{\theta}, \text{ thick dashed colour lines})$ of the stone from the same viewing angle θ . Now 982 consider the magnitude of the horizontal component of the velocity vector, $\|\mathbf{V}\sin\theta\|$, as the speed of river current at an arbitrary point along the velocity profile, u, which is used 984 in calculating drag force. Observing that

$$(\mathbf{V}\sin\theta)^2 = \mathbf{V}\cdot\mathbf{V}\sin^2\theta = \|\mathbf{V}\|^2\sin^2\theta = \|\mathbf{V}\sin\theta\|^2,$$

the drag force on the stone is proportional to the product of dynamic pressure and stone area at the same viewing angle $\frac{1}{2}\rho(\mathbf{V}\sin\theta)^2A_{\theta}$ (equivalent to $\frac{1}{2}\rho u^2A$ from equation 9a). Hence, we can use the magnitude of the horizontal component of the velocity vector (illustrated by colour-matched vectors each from the same velocity \mathbf{V} , Figure B.1a) as a physically-based weighting mechanism in estimating a weighted mean of ER_{θ} , which is the exposure factor EF used in equation 9b. Thus EF is calculated by

$$EF = \frac{\sum_{\theta \in \Theta} ER_{\theta} \frac{1}{2} \rho(\mathbf{V} \sin \theta)^{2} A_{\theta}}{\sum_{\theta \in \Theta} \frac{1}{2} \rho(\mathbf{V} \sin \theta)^{2} A_{\theta}}$$

$$= \frac{\sum_{\theta \in \Theta} ER_{\theta} \sin^{2} \theta}{\sum_{\theta \in \Theta} \sin^{2} \theta} \quad \text{(spherical grains)}$$

$$\approx \frac{\sum_{\theta \in \Theta} ER_{\theta} \sin^{2} \theta}{\sum_{\theta \in \Theta} \sin^{2} \theta} \quad \text{(natural stones)}$$

Under idealised conditions of spherical grains, $\frac{1}{2}\rho \|\mathbf{V}\|^2 A_{\theta}$ is not dependent on viewing angle, so it factors out of both summations and cancels across the fraction. However, for natural stones, which are typically not spherical, this weighting mechanism provides a good approximation as an exposure factor. The summed denominator is necessary since the sum of all the weightings used in a weighted mean must be unity. To minimise edge effects when calculating EF on scanned images, we extended the edges of basket of grains by replicating copies of a binary image of sample particles in a circular pattern around the entire basket with a slight buffered overlap around the outer edge.

1000 B.2 Adjusted exposure factor, EF_{adj}

The EF is the proportion of the total face of the grain that is exposed to the flow, however 1001 for the entrainment calculation we are only interested in the proportion of the grain face 1002 that is exposed to nonzero flow velocities. Since the nonzero portion of the logarithmic 1003 velocity profile is defined as starting at the mean local bed elevation, which is generally 1004 above the base of a grain, it is necessary to rescale EF so that portions of the grain where 1005 the profile is zero are not included in the drag force calculations. Without this step, 1006 drag force calculations used in the moment balance will be underestimated resulting in increased critical shear values. The adjustment to EF depends upon whether the fraction 1008 of stone experiencing no flow (i.e. zero velocity profile) is greater or less than the fraction that is sheltered by upstream grains (as expressed by EF). Let F_u be the fraction of 1010 a stone above the mean local bed elevation where u > 0, and let $F_u = F_1 + F_2$ be the 1011

partition of this fraction such that $F_2 = EF$ is the exposure factor (see Figure B.1b). Then $F_1 = F_u - F_2 = F_u - EF$. Since EF represents the exposed fraction for the entire stone, the hidden fraction of F_u is F_1/F_u (light grey section) and the exposed fraction of F_u is the adjusted exposure factor $EF_{adj} = F_2/F_u$ (white section). EF_{adj} further reduces drag force on the stone due to the additional sheltering effects of EF, which are only partially accounted for by F_u (dark grey section). Then from the defined partition

$$EF_{adj} = \frac{F_2}{F_u} = \frac{F_u - F_1}{F_u} = 1 - \frac{F_u - EF}{F_u}, \quad (0 < EF \le F_u \le 1).$$

Now whenever $0 < F_u < EF \le 1$, EF already covers F_u in its entirety, so there is no need to reduce drag force any further than already accounted for by F_u . Hence, in this case $EF_{adj} = 1$. For computational efficiency, we combine both conditional domains for EF_{adj} into a single, unconditional domain, thus the EF_{adj} becomes

$$EF_{adj} = 1 - \max\left\{0, \frac{F_u - EF}{F_u}\right\}, \quad Fu, EF \in \left(0, 1\right].$$

$_{\scriptscriptstyle{1022}}$ C Empirical cohesive force model, \mathbf{F}_{C}

Cohesive force experiments were designed to determine the effect of burial depth, sand 1023 size, and clay fraction on the vertical tensile force required to extricate a 25 mm diameter 1024 spherical glass marble from a sand-clay mixture. The chosen marble size represents the 1025 lower end of the D_{50} size range from field data, (Hodge et al., 2013). The marble was 1026 also pulled from pure clay to establish a baseline from which treatment effects could be 1027 modelled relative to the effects of pure clay. Treatments used in the experiment were 1028 burial depth $(4 \,\mathrm{mm}, 8 \,\mathrm{mm} \text{ and } 12 \,\mathrm{mm})$, size range of sand $((0.1, 0.3] \,\mathrm{mm}, (0.5, 1.0] \,\mathrm{mm}$, 1029 and (1.0,2.0] mm), and clay fraction of the sand-clay mixture (0, 0.25, 0.50) and (0.75) as 1030 a $3 \times 3 \times 4$ main effects factorial design. 1031

For the experimental set up, three parts sand-clay mixture were thoroughly mixed with one part water, and then poured into a small tray that could accommodate fifteen buried marbles (three burial depths × five replicates) with sufficient spacing as to avoid surface deformation effects of neighbouring grains during force pulls. The sand-clay mixture was spread over the tray, and the surface was smoothed without compaction to assure consistency between pours. String was glued to each marble prior to vertical pulls. Force pulls were made of each free standing marble prior to burial to adjust their extraction force by the marble's weight. Marbles were then buried to their depth markings, and the surface was allowed to set undisturbed, yet still damp, until force measurements, which were obtained using a handheld force gauge.

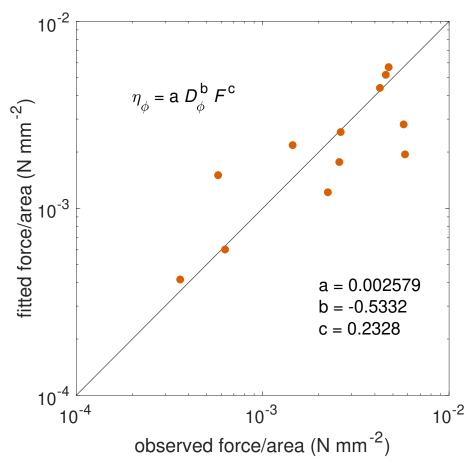


Figure C.1: Fitted versus observed cohesive force per surface area.

Force pulls for pure clay were found to be independent of burial depth as the simple linear regression was found to be insignificant (p = .526). Tensile force per buried surface area of the marble for all sand-clay mixtures had substantial nonlinear relationships with sand size and clay fraction. Therefore, a single power law model was developed where the response

was defined as the force per surface area for the mixtures divided by the overall mean of 1046 force per surface area for pure clay (i.e. the comparison baseline). Then the zero value clay 1047 fraction was adjusted to 0.001 before log-linearising the power law model for linear regress-1048 ing over mean sand sizes $(0.20 \,\mathrm{mm}, 0.75 \,\mathrm{mm})$ and $(1.50 \,\mathrm{mm})$ and clay fractions (0.001, 0.25, 0.25)1049 0.50 and 0.75). The significant regression of the linearised power law $(p < .001, R^2 = 0.54)$ 1050 was exponentiated and multiplied through by the baseline value for pure clay force per 1051 surface area yielding the final model (Figure C.1). Other empirical definitions could be de-1052 rived for use within our vector-based 3D model. Data and model development are located 1053 on our Github page at https://github.com/NERCPATCheS/VectorEntrainment3D. 1054

1055 D Algorithmic solution for 3D vector-based grain entrainment

An algorithmic solution is provided for calculating τ_{cr}^* so that more complexity can be 1056 added into the 3D framework (e.g. adding turbulence) without deriving an analytical 1057 solution. To arrive at a value of τ_{cr} requires finding the boundary shear stress, τ_b , used 1058 in Equation 8 that satisfies the static moment balance. Calculate entrainment for each 1059 stone in the sample from top to bottom as only surface stones experience shear stress. For 1060 each stone with entrainment potential, where initial force calculations are nonzero, get a 1061 list of N contact points of neighbouring stones that are within some specified distance. 1062 Then generate a list of $\binom{N}{2}$ contact pairs to construct vector pairs $(\mathbf{CV}_L, \mathbf{CV}_R)$, using 1063 the rule set to test all vector pairs for viability, keeping only the pairs that are viable 1064 for entrainment calculations. Next calculate a set of τ_{cr} values, where each value in the set $\{\tau_{cr}\}$ corresponds with each viable pair of contacts. The smallest τ_{cr} value in the set 1066 $\{\tau_{cr}\}$ and its corresponding pair of contacts $(\mathbf{CV}_L,\mathbf{CV}_R)$ are the entrainment solution 1067 for that particular stone. Calculate the corresponding τ_{cr}^* value, all entrainment angles 1068 using Equations 6, and any other useful information. What follows is the algorithm used 1069 to obtain a single τ_{cr} value for a given contact pair $(\mathbf{CV}_L, \mathbf{CV}_R)$ that will be added to 1070 the set $\{\tau_{cr}\}$ from which the entrainment solution for a given grain is $\tau_{cr} = \min\{\tau_{cr}\}$. 1071

Consider a small step length of boundary shear stress, $\Delta \tau_b$, to determine interval width,

 $[\tau_1, \tau_2] = [n-1, n] \Delta \tau_b$, iterated over $n = 1, 2, \ldots$, defining a partitioned range of τ_b values 1073 (for example, $[\tau_1, \tau_2] = [0, \Delta \tau_b]$ for n = 1, and $[\tau_1, \tau_2] = [\Delta \tau_b, 2\Delta \tau_b]$ for n = 2, and so on). 1074 Using either moment expression in Equation 7 or Equation 15a (i.e. not set to zero) to 1075 define a moment function $m_i = f(\tau_i)$ for i = 1, 2, the moment interval corresponding to 1076 $[\tau_1, \tau_2]$ is $[m_1, m_2]$. Our goal is to iterate $[\tau_1, \tau_2]$ until $[m_1, m_2]$ contains zero, which occurs 1077 whenever $m_1 \times m_2 < 0$. Since $[m_1, m_2]$ is usually far from containing zero at n = 1, it 1078 will take $k = \lfloor m_1/(m_1 - m_2) \rfloor > 0$ iterations (where $\lfloor \bullet \rfloor$ is the floor function) for interval 1079 $[m_1, m_2]$ to be relatively close to zero (note that if k < 0, then $\tau_{cr} < 0$; set $\tau_{cr} = \infty$). 1080 Therefore, start by iterating $[\tau_1, \tau_2] = [n-1, n] \Delta \tau_b$ over $n = k, k+1, \ldots$ until $m_1 \times m_2 < 0$. 1081 Once we reach an interval $[\tau_1, \tau_2]$ that generates $m_1 \times m_2 < 0$, we then use a binary search 1082 algorithm as follows: (i) split the interval $[\tau_1, \tau_2]$ in half at the mean, $\bar{\tau} = (\tau_1 + \tau_2)/2$, 1083 and generate moment intervals $[m_1, m_{\bar{\tau}}]$ and $[m_{\bar{\tau}}, m_2]$ corresponding to $[\tau_1, \bar{\tau}]$ and $[\bar{\tau}, \tau_2]$, 1084 respectively, (ii) test each half interval $[m_1, m_{\bar{\tau}}]$ and $[m_{\bar{\tau}}, m_2]$ to find which half contains 1085 zero, (iii) choose a new $[\tau_1, \tau_2]$ that corresponds with the moment interval that contains 1086 zero going back to step (i) and repeat until max $\{|m_1|, |m_2|\} \approx 0$, and (iv) calculate $\bar{\tau}$ 1087 from the final $[\tau_1, \tau_2]$ saving its value to the set $\{\tau_{cr}\}$ for the current $(\mathbf{CV}_L, \mathbf{CV}_R)$ pair. 1088 Get the next $(\mathbf{CV}_L, \mathbf{CV}_R)$ pair for the current grain, setting $[\tau_1, \tau_2] = [0, \Delta \tau_b]$ at n = 1, 1089 and repeat the algorithm for all valid $(\mathbf{CV}_L, \mathbf{CV}_R)$ pairs of the current stone. Repeat 1090 the algorithm for all remaining stones in the sample to obtain a τ_{cr}^* value plus angles and 1091 other metrics for each stone with entrainment potential. A flowchart detailing the steps 1092 taken in the algorithmic solution is provided (Figure D.1.) 1093 Binary search algorithms are computationally efficient as they have a geometric rate of 1094 convergence in narrowing $[\tau_1, \tau_2]$ intervals to obtain a τ_{cr} value. By initialising the search 1095 for zero in $[m_1, m_2]$ over n = k, k + 1, ..., the $[\tau_1, \tau_2]$ interval used in the binary search is 1096 obtained in just a few iterations. Combining both of the these enhancements with parallel 1097 processing the for-loops over contact pairs $(\mathbf{CV}_L, \mathbf{CV}_R)$, this algorithm is highly efficient. 1098 The Λ values in Equation 15c indicated the potential for $\tau_{cr} < 0$, which is possible due 1099 to the relaxed $\beta < |90^{\circ}|$ assumption of the current rule set for viable $(\mathbf{CV}_L, \mathbf{CV}_R)$ pairs. 1100

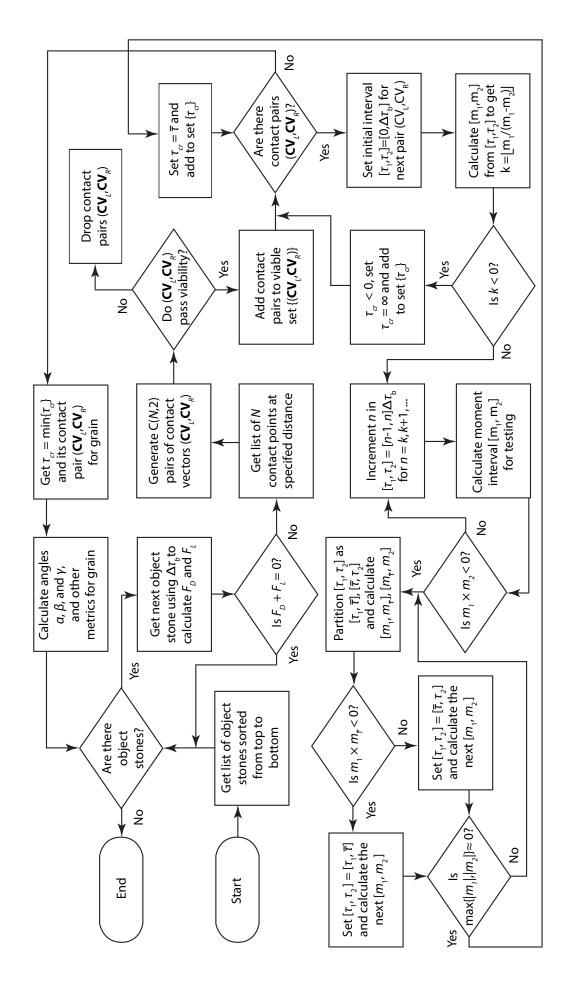


Figure D.1: Flowchart of algorithmic solution for 3D vector-based grain entrainment.

For our analysis, we chose to ignore any negative τ_{cr} values as extraneous information.

Acknowledgements: This research was enabled by NERC grants NE/K013386/1 to University of Southampton and NE/K012304/1 to Durham University. We would like to extend our sincere gratitude to the University of Southampton postgraduate and undergraduate students and the Geography technical team for flume and laboratory assistance. We would also like to acknowledge use of the University of Southampton μ -VIS X-Ray Imaging Centre. Finally, we would like to thank the editor and both reviewers for their constructive comments, which improved the quality of this paper.

m_{109} References

- J. Aberle and V. Nikora. Statistical properties of armored gravel bed surfaces. Water

 Resources Research, 42(11), 2006. doi: 10.1029/2005wr004674.
- J. R. Agudo, S. Dasilva, and A. Wierschem. How do neighbors affect incipient particle motion in laminar shear flow? *Physics of Fluids*, 26(5), 2014. doi: 10.1063/1.4874604.
- Sharif Ahmed, John Harkness, Louis Le Pen, William Powrie, and Antonis Zervos.
- Numerical modelling of railway ballast at the particle scale. *International Journal*
- for Numerical and Analytical Methods in Geomechanics, 40(5):713–737, 2016a. doi:
- 1117 10.1002/nag.2424.
- Sharif Ahmed, Trudy Naugler Klassen, Samuel Keyes, Michael Daly, David L. Jones,
- Mark Mavrogordato, Ian Sinclair, and Tiina Roose. Imaging the interaction of roots
- and phosphate fertiliser granules using 4d x-ray tomography. Plant and Soil, 401(1-2):
- 125–134, 2016b. doi: 10.1007/s11104-015-2425-5.
- Sk Zeeshan Ali and Subhasish Dey. Hydrodynamics of sediment threshold. *Physics of Fluids*, 28(7), 2016. doi: 10.1063/1.4955103.
- Mohammad Amir, Vladimir I. Nikora, and Mark T. Stewart. Pressure forces on sedi-

- ment particles in turbulent open-channel flow: a laboratory study. *Journal of Fluid Mechanics*, 757:458–497, 2014. doi: 10.1017/jfm.2014.498.
- Christophe Ancey, Tobias Boehm, Magali Jodeau, and Philippe Frey. Statistical description of sediment transport experiments. *Physical Review E*, 74(1), 2006. doi: 10.1103/PhysRevE.74.011302.
- A. Armanini and C. Gregoretti. Incipient sediment motion at high slopes in uniform flow condition. Water Resources Research, 41(12), 2005. doi: 10.1029/2005wr004001.
- Ronel Barzilai, Jonathan B. Laronne, and Ian Reid. Effect of changes in fine-grained matrix on bedload sediment transport in a gravel-bed river. Earth Surface Processes and Landforms, 38(5):441–448, 2013. doi: 10.1002/esp.3288.
- F.P. Beer, E.R. Johnston Jr., D.F. Mazurek, P.J. Cornwell, and B.P. Self. *Vector Mechanics for Engineers*. McGraw Hill, New York, NY, 10 edition, 2013.
- G. Blois, J. L. Best, G. H. S. Smith, and R. J. Hardy. Effect of bed permeability and hyporheic flow on turbulent flow over bed forms. *Geophysical Research Letters*, 41(18): 6435–6442, 2014. doi: 10.1002/2014gl060906.
- J. S. Bridge and S. J. Bennett. A model for the entrainment and transport of sediment grains of mixed sizes, shapes, and densities. *Water Resources Research*, 28(2):337–363, 1992. doi: 10.1029/91wr02570.
- John M. Buffington and David R. Montgomery. A systematic analysis of eight decades of incipient motion studies, with special reference to gravel-bedded rivers. *Water Resources Research*, 33(8):1993–2029, 1997. doi: 10.1029/96WR03190.
- John M. Buffington, William E. Dietrich, and James W. Kirchner. Friction angle measurements on a naturally formed gravel streambed: Implications for critical boundary shear
 stress. Water Resources Research, 28(2):411–425, 1992. doi: 10.1029/91WR02529.
- Ben Callow, Ismael Falcon-Suarez, Sharif Ahmed, and Juerg Matter. Assessing the carbon

- sequestration potential of basalt using x-ray micro-ct and rock mechanics. *International Journal of Greenhouse Gas Control*, 70:146–156, 2018. doi: 10.1016/j.ijggc.2017.12.008.
- George Casella and Roger L. Berger. *Statistical Inference*. Duxbury Press, Pacific Grove, CA, 2nd edition edition, 2001.
- Ahmet O. Celik, P. Diplas, and C. L. Dancey. Instantaneous pressure measurements on a spherical grain under threshold flow conditions. *Journal of Fluid Mechanics*, 741:60–97, 2014. doi: 10.1017/jfm.2013.632.
- C. Chen, A. I. Packman, and J. F. Gaillard. Using x-ray micro-tomography and porescale modeling to quantify sediment mixing and fluid flow in a developing streambed. *Geophysical Research Letters*, 36, 2009. doi: 10.1029/2009gl037157.
- C. Chen, A. I. Packman, D. X. Zhang, and J. F. Gaillard. A multi-scale investigation of interfacial transport, pore fluid flow, and fine particle deposition in a sediment bed.

 Water Resources Research, 46, 2010. doi: 10.1029/2009wr009018.
- V. Clausnitzer and J. W. Hopmans. Pore-scale measurements of solute breakthrough using
 microfocus x-ray computed tomography. Water Resources Research, 36(8):2067–2079,
 2000. doi: 10.1029/2000wr900076.
- G. Constantinescu and K. Squires. Numerical investigations of flow over a sphere in the subcritical and supercritical regimes. *Physics of Fluids*, 16(5):1449–1466, 2004. doi: 10.1063/1.1688325.
- James R. Cooper, Annie Ockleford, Stephen P. Rice, and D. Mark Powell. Does the permeability of gravel river beds affect near-bed hydrodynamics? *Earth Surface Processes* and *Landforms*, 43(5):943–955, 2018. doi: 10.1002/esp.4260.
- S. Dey. Sediment threshold. *Applied Mathematical Modelling*, 23(5):399–417, 1999. doi: 10.1016/s0307-904x(98)10081-1.
- S. Dey and R. Das. Gravel-bed hydrodynamics: Double-averaging approach. Journal of

- Hydraulic Engineering-Asce, 138(8):707–725, 2012. doi: 10.1061/(asce)hy.1943-7900. 1175 0000554.
- Digisens. Digi xct 3.2, 2014. URL www.digisens3d.com. 1177

1176

- Norman R. Draper and Harry Smith. Applied Regression Analysis. Wiley-Interscience, 1178 New York, third edition edition, 1998. ISBN 978-0-471-17082-2. 1179
- J. D. Fenton and J. E. Abbott. Initial movement of grains on a stream bed effect of relative protrusion. Proceedings of the Royal Society of London Series a-Mathematical 1181 Physical and Engineering Sciences, 352(1671):523–537, 1977. doi: 10.1098/rspa.1977. 1182 0014. 1183
- Z. Y. Ge, B. Sahiner, H. P. Chan, L. M. Hadjiiski, P. N. Cascade, N. Bogot, E. A. 1184 Kazerooni, J. Wei, and C. A. Zhou. Computer-aided, detection of lung nodules: False 1185 positive reduction using a 3d gradient field method and 3d ellipsoid fitting. Medical 1186 Physics, 32(8):2443–2454, 2005. doi: 10.1118/1.1944667. 1187
- Volume Graphics. Vgstudio max 2.1, 2011. URL www.volumegraphics.com. 1188
- R. J. Hardy. Modeling granular sediment transport over water worked gravels within a 1189 CFD framework. River Flow 2006, Vols 1 and 2. Taylor & Francis Group, 2006. ISBN 0-1190 415-40815-6. URL <GotoISI>://WOS:000241916500096. Times Cited: 0 International 1191 Conference on Fluvial Hydraulics Sep 06-08, 2006 Lisbon, PORTUGAL. 1192
- R. J. Hardy, S. N. Lane, M. R. Lawless, J. L. Best, L. Elliott, and D. B. Ingham. Develop-1193 ment and testing of a numerical code for treatment of complex river channel topography 1194 in three-dimensional cfd models with structured grids. Journal of Hydraulic Research, 1195 43(5):468–480, 2005. doi: 10.1080/00221680509500145. 1196
- Heather Haynes, Elisa Vignaga, and William M. Holmes. Using magnetic resonance 1197 imaging for experimental analysis of fine-sediment infiltration into gravel beds. Sedi-1198 mentology, 56(7):1961–1975, 2009. doi: 10.1111/j.1365-3091.2009.01064.x. 1199

- Rebecca A. Hodge, David A. Sear, and Julian Leyland. Spatial variations in surface sed-
- iment structure in riffle-pool sequences: a preliminary test of the differential sediment
- entrainment hypothesis (dseh). Earth Surface Processes and Landforms, 38(5):449–465,
- 2013. doi: 10.1002/esp.3290.
- 1204 C. E. Johnston, E. D. Andrews, and J. Pitlick. In situ determination of particle friction
- angles of fluvial gravels. Water Resources Research, 34(8):2017–2030, 1998. doi: 10.
- 1029/98wr00312.
- James W. Kirchner, William E. Dietrich, Fujiko Iseya, and Hiroshi Ikeda. The variability
- of critical shear stress, friction angle, and grain protrusion in water-worked sediments.
- Sedimentology, 37(4):647–672, 1990. doi: 10.1111/j.1365-3091.1990.tb00627.x.
- Maarten G. Kleinhans, Cécile R. L. P. N. Jeukens, Chris J. G. Bakker, and Roy M. Frings.
- Magnetic resonance imaging of coarse sediment. Sedimentary Geology, 208(3):69–78,
- 2008. doi: http://dx.doi.org/10.1016/j.sedgeo.2008.07.002.
- P. D. Komar and P. A. Carling. Grain sorting in gravel-bed streams and the choice of
- particle sizes for flow-competence evaluations. Sedimentology, 38(3):489–502, 1991.
- 1215 P. D. Komar and Z. Li. Pivoting analyses of the selective entrainment of sediments by
- shape and size with application to gravel threshold. Sedimentology, 33(3):425–436, 1986.
- doi: 10.1111/j.1365-3091.1986.tb00546.x.
- Michael P. Lamb, William E. Dietrich, and Jeremy G. Venditti. Is the critical shields stress
- for incipient sediment motion dependent on channel-bed slope? Journal of Geophysical
- 1220 Research: Earth Surface, 113(F2):n/a-n/a, 2008. doi: 10.1029/2007JF000831.
- S. N. Lane, R. J. Hardy, L. Elliott, and D. B. Ingham. Numerical modeling of flow pro-
- cesses over gravelly surfaces using structured grids and a numerical porosity treatment.
- Water Resources Research, 40(1), 2004. doi: 10.1029/2002wr001934.
- Eckhard Limpert, Werner A. Stahel, and Markus Abbt. Log-normal distributions across
- the sciences: Keys and clues. BioScience, 51(5):341–352, 2001. doi: 10.1641/
- 0006-3568(2001)051[0341:LNDATS]2.0.CO;2.

- C. H. Ling. Criteria for incipient motion of spherical sediment particles. *Journal of Hydraulic Engineering-Asce*, 121(6):472–478, 1995. doi: 10.1061/(asce)0733-9429(1995) 121:6(472).
- 1230 MathWorks. Matlab 9.4, 2018. URL www.mathworks.com.
- Donald McAlister. The law of the geometric mean. *Proceedings of the Royal Society of London*, 29, 1879.
- R. Measures and S. Tait. Quantifying the role of bed surface topography in controlling sediment stability in water-worked gravel deposits. *Water Resources Research*, 44(4), 2008. doi: 10.1029/2006wr005794.
- E. Mignot, D. Hurther, and E. Barthelemy. On the structure of shear stress and turbulent kinetic energy flux across the roughness layer of a gravel-bed channel flow. *Journal of Fluid Mechanics*, 638:423–452, 2009. doi: 10.1017/s0022112009990772.
- M. C. Miller, I. N. McCave, and P. D. Komar. Threshold of sediment motion under
 unidirectional currents. *Sedimentology*, 24(4):507–527, 1977. doi: 10.1111/j.1365-3091.
 1977.tb00136.x.
- J. J. More and D. C. Sorensen. Computing a trust region step. Siam Journal on Scientific and Statistical Computing, 4(3):553–572, 1983. doi: 10.1137/0904038.
- M. Nabi, H. J. de Vriend, E. Mosselman, C. J. Sloff, and Y. Shimizu. Detailed simulation
 of morphodynamics: 2. sediment pickup, transport, and deposition. Water Resources
 Research, 49(8):4775–4791, 2013. doi: 10.1002/wrcr.20303.
- P. Naden. An erosion criterion for gravel-bed rivers. Earth Surface Processes and Landforms, 12(1):83–93, 1987. doi: 10.1002/esp.3290120110.
- V. Nikora, K. Koll, I. McEwan, S. McLean, and A. Dittrich. Velocity distribution in the roughness layer of rough-bed flows. *Journal of Hydraulic Engineering-Asce*, 130(10): 1036–1042, 2004. doi: 10.1061/(asce)0733-9429(2004)130:10(1036).

- Jorge Nocedal and Stephen Wright. Numerical Optimization. Springer-Verlag, New York, 1252 1999. ISBN 0-387-98793-2. 1253
- Lyman Ott and Michael Longnecker. An Introduction to Statistical Methods and Data 1254 Analysis. Duxbury Press, Pacific Grove, CA, 5th ed edition, 2001. 1255
- A. I. Packman, A. Marion, M. Zaramella, C. Chen, J. F. Gaillard, and D.T. Keane.
- Development of layered sediment structure and its effects on pore water transport and 1257
- hyporheic exchange. Water, Air, & Soil Pollution: Focus, 6(5-6):433-442, 2006. doi: 1258
- 10.1007/s11267-006-9057-y.1259

1263

- A. N. Papanicolaou, P. Diplas, N. Evaggelopoulos, and S. Fotopoulos. Stochastic in-1260 cipient motion criterion for spheres under various bed packing conditions. Journal of 1261 $Hydraulic\ Engineering-Asce,\ 128(4):369-380,\ 2002.\ doi:\ 10.1061/(asce)0733-9429(2002)$ 1262 128:4(369).
- A. Pierret, Y. Capowiez, L. Belzunces, and C. J. Moran. 3d reconstruction and quantifi-1264 cation of macropores using x-ray computed tomography and image analysis. Geoderma, 1265 106(3-4):247-271, 2002. doi: 10.1016/s0016-7061(01)00127-6. 1266
- D. Mark Powell, Annie Ockelford, Stephen P. Rice, John K. Hillier, Thao Nguyen, Ian 1267 Reid, Nicholas J. Tate, and David Ackerley. Structural properties of mobile armors 1268 formed at different flow strengths in gravel-bed rivers. Journal of Geophysical Research: 1269 Earth Surface, 121(8):1494–1515, 2016. doi: 10.1002/2015JF003794. 1270
- Philip J. Pritchard and John W. Mitchell. Fox and McDonald's Introduction to Fluid 1271 Mechanics. John Wiley & Sons, Inc., Hoboken, NJ, 9 edition, 2015. 1272
- John C. Russ. The Image Processing Handbook. CRC Press, Boco Raton, Fl, 6th edition, 1273 2011. 1274
- S. Sarkar and S. Dey. Double-averaging turbulence characteristics in flows over a gravel 1275 bed. Journal of Hydraulic Research, 48(6):801–809, 2010. doi: 10.1080/00221686.2010. 1276 526764. 1277

- Johannes Schindelin, Ignacio Arganda-Carreras, Erwin Frise, Verena Kaynig, Mark Lon-
- gair, Tobias Pietzsch, Stephan Preibisch, Curtis Rueden, Stephan Saalfeld, Benjamin
- Schmid, Jean-Yves Tinevez, Daniel James White, Volker Hartenstein, Kevin Eliceiri,
- Pavel Tomancak, and Albert Cardona. Fiji: an open-source platform for biological-
- image analysis. Nature Methods, 9(7):676–682, 2012. doi: 10.1038/nmeth.2019.
- Mark W. Schmeeckle, Jonathan M. Nelson, and Ronald L. Shreve. Forces on stationary
- particles in near-bed turbulent flows. Journal of Geophysical Research, 112(F2), 2007.
- doi: 10.1029/2006jf000536.
- A. Shields. Application of similarity principles and turbulence research to bed-load
- movement. translated from "anwendung der aehnlichkeitsmechanik und der turbulen-
- zforschung auf die geschiebebewegung" by w. i. ott and j. c. van uchelen. Mitteil.
- 1289 Preuss. Versu-chanst. Wasser, Erd, Schiffsbau, 26, 1936.
- 1290 Stefan Vollmer and Maarten G. Kleinhans. Predicting incipient motion, including the
- effect of turbulent pressure fluctuations in the bed. Water Resources Research, 43(5):
- n/a-n/a, 2007. doi: 10.1029/2006WR004919.
- Patricia L. Wiberg and J. Dungan Smith. Calculations of the critical shear stress for
- motion of uniform and heterogeneous sediments. Water Resources Research, 23(8):
- 1471–1480, 1987. doi: 10.1029/WR023i008p01471.
- D. Wildenschild and A. P. Sheppard. X-ray imaging and analysis techniques for quantify-
- ing pore-scale structure and processes in subsurface porous medium systems. Advances
- in Water Resources, 51:217–246, 2013. doi: 10.1016/j.advwatres.2012.07.018.
- D. Wildenschild, J. W. Hopmans, C. M. P. Vaz, M. L. Rivers, D. Rikard, and B. S. B.
- 1300 Christensen. Using x-ray computed tomography in hydrology: systems, resolutions, and
- limitations. Journal of Hydrology, 267(3-4):285–297, 2002. doi: 10.1016/s0022-1694(02)
- 1302 00157-9.
- 1303 M. G. Wolman. A method of sampling coarse river-bed material. Transactions, American
- 1304 Geophysical Union, 35(6):951-, 1954.

- 1305 F. C. Wu and Y. J. Chou. Rolling and lifting probabilities for sediment entrainment.
- $_{1306}$ Journal of Hydraulic Engineering-Asce, 129(2):110-119, 2003. doi: 10.1061/(asce)
- 0733-9429(2003)129:2(110).
- 1308 F. C. Wu and K. H. Yang. Entrainment probabilities of mixed-size sediment incorporating
- near-bed coherent flow structures. Journal of Hydraulic Engineering-Asce, 130(12):
- 1187-1197, 2004. doi: 10.1061/(asce)0733-9429(2004)130:12(1187).
- ¹³¹¹ Fu-Chun Wu and Men-Rong Jiang. Numerical investigation of the role of turbulent
- bursting in sediment entrainment. Journal of Hydraulic Engineering-Asce, 133(3):329–
- 334, 2007. doi: 10.1061/(asce)0733-9429(2007)133:3(329).
- 1314 E. M. Yager, J. W. Kirchner, and W. E. Dietrich. Calculating bed load transport in
- steep boulder bed channels. Water Resources Research, 43(7):n/a-n/a, 2007. doi:
- 10.1029/2006WR005432.