

# On-the-fly Full Hessian Kernel Calculations Based upon Seismic Wave Simulations

Yujiang Xie<sup>1</sup>, Catherine Rychert<sup>1</sup>, Nicholas Harmon<sup>1</sup>,

Qinya Liu<sup>2</sup>, and Dirk Gajewski<sup>3</sup>

<sup>1</sup> Ocean and Earth Science, University of Southampton, UK

<sup>2</sup> Department of Physics & Department of Earth Sciences, University of Toronto, Canada

<sup>3</sup> Institute of Geophysics, University of Hamburg, Germany

\*Corresponding author; E-mail:yujiang.xie@soton.ac.uk

## Abstract

Full waveform inversion (FWI) or adjoint tomography has routinely been performed to image the internal structure of the Earth at high resolution. This is typically done using the Fréchet kernels and the approximate Hessian or the approximate inverse Hessian because of the high computational cost of computing and storing the full Hessian. Alternatively, the full Hessian kernels can be used to improve inversion resolutions, convergence rates, and possibly mitigate inter-parameter tradeoffs. The storage requirements of the full Hessian kernel calculations can be reduced by compression methods, but often at a price of accuracy depending on the compression factor. Here we present

13 open-source codes to compute both Fréchet and full Hessian kernels on the fly (in the  
14 computer RAM) through simultaneously solving four wave equations, which we call  
15 QuadSEM. By recomputing two forward fields at the same time that two adjoint fields  
16 are calculated during the adjoint simulation, QuadSEM constructs the full Hessian  
17 kernels using the exact forward and adjoint fields. In addition, we also implement  
18 an alternative approach based on the wavefield storage method (WSM), which stores  
19 forward wavefields every  $k$ th ( $k \geq 1$ ) time step during the forward simulation and  
20 reads them back into memory during the adjoint simulation for kernel construction.  
21 Both Fréchet and full Hessian kernels can be computed simultaneously through the  
22 QuadSEM or the WSM code, only doubling the computational cost compared with the  
23 computation of Fréchet kernels alone. Compared with WSM, QuadSEM can reduce  
24 the disk space and I/O cost by three orders of magnitude in the presented examples  
25 while using 15,000 time steps. Numerical examples are presented to demonstrate the  
26 functionality of the methods, and the computer codes are provided with this contribu-  
27 tion.

## 28 **Introduction**

29 In the past thirty years, full waveform inversion (FWI), or sometimes interchangeably known  
30 as adjoint tomography in regional or global seismology, has become popular and widely  
31 used for imaging the Earth’s internal structures at multiple scales (e.g., Bamberger et al.,  
32 1982; Lailly, 1983; Tarantola, 1984, 1988; Gauthier et al., 1986; Igel et al., 1996; Pratt et al.,  
33 1998; Tape et al., 2009; Fichtner et al., 2009; Virieux and Operto, 2009; Liu and Gu, 2012;  
34 Zhu et al., 2012; French and Romanowicz, 2015; Bozdağ et al., 2016; Tromp, 2020). Sensi-  
35 tivity kernels which indicate the sensitivity of seismograms to model parameters are a key

36 component of full-waveform inversion or more generally the adjoint tomography. Typically,  
37 two types of sensitivity kernels are discussed in exploration, regional and global seismol-  
38 ogy context, the first-order derivatives of the seismological data functionals, Fréchet kernels  
39 (e.g., Dahlen et al., 2000; Tromp et al., 2005), and the second-order partial derivatives of  
40 functionals applied to an arbitrary model update, known as Hessian vector products or Hes-  
41 sian kernels (e.g., Pratt et al., 1998; Epanomeritakis et al., 2008; Fichtner and Trampert,  
42 2011; Métivier et al., 2013).

43 Fréchet kernels are widely used for waveform inversions or adjoint tomography via the  
44 scattering-integral methods (e.g., Chen et al., 2007a,b) or the adjoint methods (e.g., Lions,  
45 1968; Lailly, 1983; Tarantola, 1984, 1988; Tromp et al., 2005; Fichtner et al., 2006; Plessix,  
46 2006). The kernels can be computed via the correlations of incident forward wavefields  
47 with adjoint fields (e.g., Bamberger et al., 1982; Tromp et al., 2005). There are mainly two  
48 strategies to obtain the forward fields during the adjoint simulation for correlation. One is to  
49 write the forward fields onto a disk often with compression during the forward simulation and  
50 then during the adjoint simulation for the adjoint field, read the required forward fields back  
51 into the temporary memory. Compression schemes include the temporal-spatial compression  
52 (e.g., Fichtner et al., 2009) and the lossless or lossy compression (e.g., Hanzich et al., 2013;  
53 Lindstrom et al., 2016; Boehm et al., 2016), for instance, based on lossless or lossy compres-  
54 sion techniques (e.g., Unat et al., 2009; Weiser and Götschel, 2012; Götschel and Weiser,  
55 2015). The other approach is to only store the forward wavefield at selected time steps,  
56 called checkpoints, and during the adjoint simulation, re-solve the forward problem based  
57 on these selected time steps (e.g., Symes, 2007; Anderson et al., 2012; Komatitsch et al.,  
58 2016), for instance, based on the checkpointing algorithms (e.g., Griewank and Walther,  
59 2000; Charpentier, 2001; Walther and Griewank, 2004). For the elastic cases or the anelastic  
60 cases with the consideration of physical dispersion only, the forward fields can be completely  
61 reconstructed via the boundary values and the last snapshot of the forward fields during  
62 adjoint simulation (e.g., Gauthier et al., 1986; Tromp et al., 2005; Liu and Tromp, 2006).

63 The use of Hessian in FWI may increase convergence rates, improve model resolutions,  
64 and possibly mitigate inter-parameter tradeoffs. However, the exact computation of the  
65 full Hessian matrix is computationally prohibitive. Instead, the approximate Hessian or  
66 the approximate inverse Hessian has been computed based on the Gauss-Newton or the  
67 quasi-Newton approaches (e.g., Pratt et al., 1998; Virieux and Operto, 2009, and among  
68 others). In contrast to the computation of Fréchet kernels, which uses two fields, the forward  
69 and the adjoint fields, the computation of full Hessian kernels involves four fields (e.g.,  
70 Fichtner and Trampert, 2011), two forward fields and two adjoint fields. The utility the  
71 full or approximate Hessian kernels have been demonstrated in truncated-Newton FWI for  
72 exploration seismology (e.g., Métivier et al., 2013, 2014; Pan et al., 2017; Yang et al., 2018;  
73 Matharu and Sacchi, 2019), for instance, based on classical wavefield storage, checkpointing,  
74 and/or the finite-difference wave simulations. For large models, calculating the full Hessian  
75 kernels in the classical storage approach is challenging given the large space required to  
76 store the multiple wavefields and the associated I/O expense during simulations. Therefore,  
77 approximate Hessian kernels are used instead. For instance, for resolution analysis, they are  
78 estimated by a finite-difference approximation using the gradients from two nearby iterations  
79 (e.g., Zhu et al., 2012; Bozdağ et al., 2016). Luo et al. (2014) also derived the Hessian kernel  
80 formulas and then used the diagonal terms of the Hessian to construct four preconditioners  
81 for FWI and resolution analysis. Similar to those for Fréchet kernels, the formulas for these  
82 preconditioner operators involve the correlations of forward and adjoint fields and therefore  
83 can be computed based on the adjoint method (Tromp et al., 2005).

84 Compression methods have also been applied to the forward and adjoint wavefield stor-  
85 age for the computation of the Hessian kernels (Boehm and Ulbrich, 2015). In this case,  
86 as the decompressed wavefields also appear as a distributed source term in the two auxil-  
87 iary wave equations and errors are propagated to the two perturbed wavefields resulting in  
88 less accurate Hessian kernel construction (Boehm et al., 2016). For instance, using decom-  
89 pressed fields and the trust-region Newton PCG method only resulted in a small improve-

90 ment compared with LBFGS updates for FWI (Boehm and Ulbrich, 2015). This is mainly  
91 due to the trade-off between high compression factors and low storage requirements with  
92 the accuracy of conjugate-gradient (CG) update. For this reason, resolving small and/or  
93 weakly perturbed scatterers may still be challenging since the perturbed wavefields and the  
94 compressed/decompressed errors may be of the same order of magnitude for a high com-  
95 pression factor. Alternative Hessian kernels, for instance, the reduced Hessian kernels (e.g.,  
96 Epanomeritakis et al., 2008) have also computed and used in FWI for efficiency purposes.

97 In this paper, we present a computationally efficient method to construct the full Hessian  
98 kernels on the fly based on the second-order adjoint state methods (e.g., Fichtner and Trampert,  
99 2011) and the spectral element method (SEM, e.g., Seriani and Priolo, 1994; Faccioli et al.,  
100 1996, 1997; Komatitsch and Vilotte, 1998). We first review the theory on Fréchet and Hes-  
101 sian kernels. We then implement both the classical wavefield storage method (the full or  
102 adaptive time iteration) and the on-the-fly approach for the full Hessian kernel calculations.  
103 The latter approach, namely QuadSEM, is conducted by simultaneously solving four wave  
104 equations based on spectral-element simulations (utilizing open-source SPECFEM commu-  
105 nity codes). The on-the-fly approach is possible because only the last-state forward fields and  
106 the absorbing boundary fields need to be stored in the forward simulation for reconstructing  
107 the forward fields during the adjoint simulation for kernel reconstruction. We then present  
108 and discuss the results of the Fréchet and Hessian kernels for 2-D synthetic models. The  
109 related codes are published in the public domain for dissemination.

## 110 **Theory**

### 111 **Fréchet kernels**

112 Fréchet kernels, gradients or first-order derivatives of the seismic data functional,  $\chi$ , can  
113 be used to update the structural model from a chosen initial model via local optimization

114 rather than a costly global search. When the initial model is chosen sufficiently close to  
 115 the global minimum and when the source term is relatively accurate, the final model from  
 116 the local optimizations may be used to explain the observed data. By perturbing the data  
 117 functional as  $\delta\chi$  with respect to an isotropic model  $\mathbf{m}$ , we have (also see Tromp et al., 2005;  
 118 Fichtner and Trampert, 2011)

$$119 \quad \delta\chi = \int_V K_m \delta\mathbf{m} d^3\mathbf{x}, \quad (1)$$

120 where  $K_m$  denotes the Fréchet kernels and  $V$  denotes the model volume. Here we omit the  
 121 spatial and temporal dependencies of the kernels for simplicity unless stated otherwise. In  
 122 principle, the generic  $K_m$  can be expressed by different components depending on the choice  
 123 of model parameterization. For simplicity, we only show the case for a model parameteriza-  
 124 tion given by  $\mathbf{m} = (\rho, \alpha, \beta)$ , where  $\rho$  denotes the density and  $\alpha$  and  $\beta$  denote the P- and  
 125 S-wave speed, respectively. The kernels applied to the model perturbation in eq. (1) can be  
 126 further expressed as

$$127 \quad K_m \delta\mathbf{m} = \begin{pmatrix} K_\rho & K_\alpha & K_\beta \end{pmatrix} \begin{pmatrix} \delta\rho \\ \delta\alpha \\ \delta\beta \end{pmatrix}, \quad (2)$$

128 where  $\delta\mathbf{m} = (\delta\rho, \delta\alpha, \delta\beta)^T$ . As the computation of Fréchet kernels relies on the forward and  
 129 the adjoint fields computed from a given model, we rewrite the Fréchet kernels as a function  
 130 of the forward and adjoint fields

$$131 \quad \begin{pmatrix} K_\rho \\ K_\alpha \\ K_\beta \end{pmatrix} = \begin{pmatrix} K_\rho(\mathbf{s}^\dagger, \ddot{\mathbf{s}}, \mathbf{s}) \\ K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \\ K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \end{pmatrix}, \quad (3)$$

132 where  $\mathbf{s}$  denotes the forward displacement field and  $\mathbf{s}^\dagger$  denotes the adjoint field in this given  
 133 model. The  $\ddot{\mathbf{s}}$  is the second-order time derivative of  $\mathbf{s}$ , i.e., the forward acceleration field.

134 **Hessian kernels**

135 **Components of Hessian kernels**

136 Similar to the first-order form of the Fréchet kernels as shown in eq. (1), the second-order  
 137 form or the Hessian operator can be written as (see Fichtner and Trampert, 2011)

$$138 \quad H(\delta\mathbf{m}_1, \delta\mathbf{m}_2) = \int_V K_m^1 \delta\mathbf{m}_2 d^3\mathbf{x} = \int_V (\mathbb{H}_a + \mathbb{H}_b + \mathbb{H}_c) \delta\mathbf{m}_2 d^3\mathbf{x}, \quad (4)$$

139 where  $K_m^1 = \mathbb{H}_a + \mathbb{H}_b + \mathbb{H}_c$  denotes the full Hessian kernels. Based upon the work of  
 140 Fichtner and Trampert (2011), we rewrite each part of the product as

$$141 \quad \mathbb{H}_a(\rho, \alpha, \beta) = \begin{pmatrix} K_\rho(\mathbf{s}^\dagger, \delta\ddot{\mathbf{s}}, \delta\mathbf{s}) \\ K_\alpha(\mathbf{s}^\dagger, \delta\mathbf{s}) \\ K_\beta(\mathbf{s}^\dagger, \delta\mathbf{s}) \end{pmatrix}, \quad (5)$$

$$142 \quad \mathbb{H}_b(\rho, \alpha, \beta) = \begin{pmatrix} K_\rho(\delta\mathbf{s}^\dagger, \ddot{\mathbf{s}}, \mathbf{s}) \\ K_\alpha(\delta\mathbf{s}^\dagger, \mathbf{s}) \\ K_\beta(\delta\mathbf{s}^\dagger, \mathbf{s}) \end{pmatrix}, \quad (6)$$

$$143 \quad \mathbb{H}_c(\rho, \alpha, \beta) = \begin{pmatrix} \rho^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\alpha + \rho^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\beta \\ \rho^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\rho + \alpha^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\alpha \\ \rho^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\rho + \beta^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\beta \end{pmatrix}, \quad (7)$$

146 where  $\delta\mathbf{s}$ ,  $\delta\ddot{\mathbf{s}}$ , and  $\delta\mathbf{s}^\dagger$  denote the perturbed wavefields based upon the model perturbation  
 147  $\delta\mathbf{m}_1 = \delta\mathbf{m} = (\delta\rho, \delta\alpha, \delta\beta)^\top$ . For simplicity, we use  $\delta\mathbf{m}$  as the model perturbation from this  
 148 point on wards. Eqs. (5)-(7) show a link between the Hessian kernels (e.g., Fichtner and Trampert,  
 149 2011) and the Fréchet kernels (e.g., Tromp et al., 2005). It implies that the implementation  
 150 framework for computing the Fréchet kernels can be used to compute the Hessian kernels by  
 151 replacing the regular field with its associated perturbed field. The  $\mathbb{H}_a$  can be computed with

152 the implementation of eq. (3) by replacing the forward fields with the perturbed forward  
 153 fields.  $H_b$  includes two contributions, i.e.,

$$154 \quad H_b = H_b^{(m)} + H_b^{(s)}, \quad (8)$$

155 where

$$156 \quad H_b^{(m)}(\rho, \alpha, \beta) = \begin{pmatrix} K_\rho(\delta \mathbf{s}_m^\dagger, \ddot{\mathbf{s}}, \mathbf{s}) \\ K_\alpha(\delta \mathbf{s}_m^\dagger, \mathbf{s}) \\ K_\beta(\delta \mathbf{s}_m^\dagger, \mathbf{s}) \end{pmatrix}, \quad (9)$$

$$157 \quad H_b^{(s)}(\rho, \alpha, \beta) = \begin{pmatrix} K_\rho(\delta \mathbf{s}_s^\dagger, \ddot{\mathbf{s}}, \mathbf{s}) \\ K_\alpha(\delta \mathbf{s}_s^\dagger, \mathbf{s}) \\ K_\beta(\delta \mathbf{s}_s^\dagger, \mathbf{s}) \end{pmatrix}. \quad (10)$$

159 The former is due to the perturbation of the model, and the latter is due to the per-  
 160 turbation of the adjoint source which is defined as the approximate Hessian kernels in  
 161 Fichtner and Trampert (2011). Both the  $H_b^{(m)}$  and  $H_b^{(s)}$  can be computed with the implemen-  
 162 tation of eq. (3) by replacing the adjoint fields with the associated perturbed adjoint fields.  
 163 The  $H_b^{(m)}$  and  $H_b^{(s)}$  are computed to determine the  $H_b$  component for the full Hessian kernel  
 164 calculations, i.e., considering the model perturbation and the adjoint source perturbation  
 165 as well in one or two adjoint simulations. The  $H_b^{(s)}$  can be computed separately to obtain  
 166 the approximate Hessian kernels, i.e., only accounting for the adjoint source perturbation.  
 167 The construction of  $H_c$  is straightforward based upon the Fréchet kernels  $K_m$  and the model  
 168 perturbation  $\delta \mathbf{m}$ .

## 169 Perturbed wavefields

170 Eqs. (5)-(10) show that the Hessian kernels can be computed with the same implementation  
 171 framework as that for the Fréchet kernel calculations, e.g., by eq. (3) using the adjoint  
 172 methods. Any other packages for wave simulations and Fréchet kernel computation can be



173 redesigned and adapted to compute the Hessian kernels with additional effort to compute  
 174 the perturbed forward fields  $\delta\mathbf{s}$  and  $\delta\ddot{\mathbf{s}}$ , and the perturbed adjoint field  $\delta\mathbf{s}^\dagger$  due to the model  
 175 perturbation  $\delta\mathbf{m}$  and the perturbation of the adjoint source.

176 The  $H_a$  component accounts for the perturbation of the forward fields (Fichtner and Trampert,  
 177 2011), e.g.,

$$178 \quad \delta\mathbf{s} = \lim_{v \rightarrow 0} \frac{1}{v} [\mathbf{s}(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, t) - \mathbf{s}(\mathbf{m}_r; \mathbf{x}, t)], \quad (11)$$

179 where  $\mathbf{m}_r$  denotes the reference model, and  $r = 0, 1, 2, \dots, N$  represents the iteration number.  
 180 The initial model is set to  $\mathbf{m}_0$ . The same consideration applies to the perturbed acceleration  
 181 field  $\delta\ddot{\mathbf{s}}$  for density kernel calculations.

182 The  $H_b$  component consists of two contributions:  $H_b^{(s)}$  and  $H_b^{(m)}$ , where the former  $H_b^{(s)}$   
 183 relies on the approximate perturbed adjoint field

$$184 \quad \delta\mathbf{s}_s^\dagger = \mathbf{s}_s^\dagger(\mathbf{m}_r; \mathbf{x}, T - t) - \mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t). \quad (12)$$

185 In the equation, the  $\mathbf{s}_s^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)$  field is generated by the adjoint source  $\mathbf{f}^\dagger(\mathbf{m}_r +$   
 186  $v\delta\mathbf{m}; \mathbf{x}, T - t)$ , and  $\mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)$  is generated by the adjoint source  $\mathbf{f}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)$ .  
 187 The adjoint sources could be the traveltime adjoint source, the waveform adjoint source, or  
 188 any other adjoint source based on the choices of seismic data functional  $\chi$  as discussed in  
 189 Tromp et al. (2005). The  $H_b^{(s)}$  and  $H_b^{(m)}$  need to be computed for the  $H_b$  determination in  
 190 order to compute the full Hessian kernels:  $H_a + H_b^{(s)} + H_b^{(m)} + H_c$ . The perturbed adjoint  
 191 field for the  $H_b^{(m)}$  may be given by

$$192 \quad \delta\mathbf{s}_m^\dagger = \lim_{v \rightarrow 0} \frac{1}{v} [\mathbf{s}_m^\dagger(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, T - t) - \mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)], \quad (13)$$

193 where the two adjoint fields  $\mathbf{s}_m^\dagger(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, T - t)$  and  $\mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)$  are generated  
 194 through the perturbed and unperturbed model from the same adjoint source  $\mathbf{f}^\dagger(\mathbf{m}_r; \mathbf{x}, T - t)$ .  
 195 Thereafter, the total perturbed adjoint field is

$$196 \quad \delta\mathbf{s}^\dagger = \delta\mathbf{s}_s^\dagger + \delta\mathbf{s}_m^\dagger. \quad (14)$$

197 The  $\delta \mathbf{s}_s^\dagger$  and  $\delta \mathbf{s}_m^\dagger$  may be computed in one adjoint simulation with the perturbation account-  
198 ing for both the model and the adjoint source simultaneously, or computed separately in two  
199 adjoint simulations considering the two perturbations individually.

200 The computation of  $H_c$  relies on the Fréchet kernels and model perturbation, see eq. (7).  
201 It has also been shown that  $H_c$  is non-zero when the model is parametrized as  $\rho$ ,  $\alpha$ , and  $\beta$  but  
202 zero when the model is given in terms of density and elastic moduli (see Fichtner and Trampert,  
203 2011).

## 204 Implementation

205 The computation of the full Hessian kernels relies on the regular and perturbed fields as men-  
206 tioned above. Its implementation is straightforward based on the wavefield storage method  
207 (WSM), which saves the forward fields at full or adaptive time steps and reads each saved  
208 time step of the forward fields back into temporary memory during the adjoint simulation  
209 for kernel constructions (See Appendix B for the WSM method based on the full or adaptive  
210 time integration scheme). Or only two adjoint wave equations need to be solved simulta-  
211 neously during the adjoint simulation. However, in light of the large storage requirement  
212 by the WSM, here we focus on showing how the full Hessian kernels are computed by the  
213 on-the-fly approach implemented in QuadSEM, which involves simultaneously solving four  
214 wave equations during the adjoint simulation. For the following examples, we only consider  
215 cases with purely elastic models.

## 216 Forward simulation

217 Figure 1 shows the comparison between the classical SEM and the QuadSEM during the  
218 forward simulations. In comparison to one model used by the classical SEM, QuadSEM  
219 carries wavefield simulations for two models simultaneously, e.g.,  $\mathbf{m}_1$  and  $\mathbf{m}_2$ , where  $\mathbf{m}_2 =$   
220  $\mathbf{m}_1 + v\delta\mathbf{m} = \mathbf{m}_1 + \Delta\mathbf{m}$ . In this case, the wavefields, including displacement  $\mathbf{s}$ , velocity  $\mathbf{v}$ ,

221 acceleration  $\ddot{\mathbf{s}}$ , and the boundary contribution  $\mathbf{b}$  (a general symbol used to refer to velocity  
222 or traction fields on the coupled boundary with an external model) are computed for the two  
223 models at each time step. The displacement seismograms  $\mathbf{s}(\mathbf{x}_r, t)$  are computed by spatial  
224 interpolation of fields at the receiver  $\mathbf{x}_r$  at each time step. The grid-point locations and mesh  
225 topology database files are shared by the two models used simultaneously in the forward  
226 simulation with QuadSEM, and only arrays/files related to model material properties such  
227 as  $\rho$ ,  $\alpha$ , and  $\beta$  need to be defined separately for the two models. The CPU and memory  
228 requirements for QuadSEM are about twice the cost of the classical SEM simulation. The  
229 forward simulations for either the classical SEM or the QuadSEM are designed to provide  
230 the absorbing boundary fields, the last state of the forward field, and the seismograms at  
231 receivers, for subsequent simulations.

## 232 **Simultaneous backward and adjoint simulations**

233 The strategy of using simultaneous backward and adjoint simulations was adopted for in-  
234 stance in the SPECFEM2D (<https://geodynamics.org/cig/software/specfem2d/>) and  
235 the SPECFEM3D (<https://geodynamics.org/cig/software/specfem3d/>) packages to com-  
236 pute the Fréchet kernels on the fly. A workflow for computing the Fréchet kernels using the  
237 classical SEM method is shown in Figure A1. For purely elastic models, the backward simu-  
238 lation is a time-reversed reconstruction of the forward field using the last state of the forward  
239 field as a starting point. The absorbing boundary contributions saved in the forward simu-  
240 lation are re-injected into the backward simulation when the forward field is reconstructed  
241 backward in time. The simulations for the backward reconstruction and the adjoint wavefield  
242 are performed simultaneously so that the corresponding time slices of the forward and adjoint  
243 wavefields can be accessed both in memory to construct the Fréchet kernels. A similar idea is  
244 adopted in the QuadSEM, and as shown in Figure 2 and Figure A2, where both the regular  
245 and perturbed forward wavefields, as well as the regular and perturbed adjoint wavefields  
246 for the two models are simultaneously computed every time step, so that the Fréchet and

247 full Hessian kernels can be constructed on the fly as wavefield products are computed and  
248 integrated over time steps. As indicated by Figure 2, the calculations of Fréchet kernels (by  
249 Arrows 1 and 2) and the full Hessian kernels (by Arrows 1, 2, 4, and 5, or Arrows 1, 2, 3, 4,  
250 and 5 depending on the adjoint source of 5) are simultaneously performed on the fly since all  
251 required wavefields are computed for each time step. Alternatively, the approximate Hessian  
252 kernels can be computed by the solutions indicated by Arrows 1, 2, 3, and 4. The QuadSEM  
253 degenerates to classical SEM when solutions indicated by 4 and 5 or 3 are not used.

254 Although the forward and adjoint wavefields for  $\mathbf{m}_1$  and  $\mathbf{m}_2$  are combined to compute  
255 kernels in the QuadSEM, the same mesher database is used except for those variables or  
256 matrices that define  $\mathbf{m}_1$  and  $\mathbf{m}_2$ . The memory cost is small since only one time step of the  
257 various fields and the integrated kernels are kept in memory. The Fréchet kernels need 3 (1 in  
258 the forward and 2 in the adjoint) simulations, while the QuadSEM needs 6 (2 in the forward  
259 and 4 in the adjoint) simulations for the simultaneous computation of both Fréchet and  
260 full Hessian kernels. A simultaneous computation of both Fréchet and approximate Hessian  
261 kernels also requires 6 (2 in the forward and 4 in the adjoint) simulations. Therefore, roughly  
262 QuadSEM doubles the memory and CPU time required for the simultaneous computation  
263 for both Fréchet and full Hessian kernels compared to the requirement for Fréchet kernels  
264 alone.

## 265 Numerical examples

### 266 Models

267 To test the numerical implementation of QuadSEM, three models are considered in this study.  
268 First, a homogeneous 2D model (*Model 1*) that is 800 km in the horizontal direction and 360  
269 km in the vertical direction and with density  $\rho=2900 \text{ kg/m}^3$ , P-wave speed  $\alpha=8000 \text{ m/s}$ ,  
270 and S-wave speed  $\beta=4800 \text{ m/s}$ , is used as a starting background model to generate initial

271 wavefields and waveforms. The second and the third models are perturbed versions of the  
272 homogeneous model. The second model (*Model 2*) has an additional +10% perturbation in  
273  $\alpha$  and  $\beta$  over a  $10 \text{ km} \times 10 \text{ km}$  area centered at the horizontal location of 335 km and depth  
274 of 135 km. The third model (*Model 3*) includes three anomalies that are  $8 \text{ km} \times 10 \text{ km}$ ,  
275 centered at a depth of 115 km at horizontal locations of 120 km, 180 km, and 240 km,  
276 respectively, with +10% perturbations in  $\alpha$  and  $\beta$ . No density perturbation is considered for  
277 the second and third model.

278 For all three models, we use 400 elements in the horizontal direction and 360 elements  
279 in the depth direction with  $5 \times 5$  Gauss-Lobatto-Legendre (GLL) points for each element,  
280 which leads to about 500  $m$  and 250  $m$  grid point spacing in the horizontal and vertical  
281 direction, respectively. The mesher and related databases for these models are built by the  
282 internal mesher tool of the SPECFEM2D package with slightly changing for the two models.  
283 Figure 3 shows the locations of these model perturbations (Part II, blue boxes in the last  
284 column) and the source-receiver geometry, together with the kernel images discussed in the  
285 section of Single source-receiver combination. These models are chosen to illustrate the  
286 differences in the calculation of Hessian kernels between the single source-receiver pair and  
287 single-source multiple-receiver case.

## 288 **Single source-receiver combination**

289 We first examine the kernel calculation for a single source-receiver combination based on  
290 *Model 1* and *Model 2*. We place a point source at  $(x, z)=(100 \text{ km}, -260 \text{ km})$ , and a standard  
291 Ricker wavelet with the dominant frequency of 0.5 Hz is used. A single receiver is placed  
292 on the top surface of the model at  $(x, z)=(600 \text{ km}, 0 \text{ km})$ . The simulations use  $dt = 0.01 \text{ s}$   
293 and run for a total of 7,000 time steps. Adjoint sources for cross-correlation traveltimes  
294 (Tromp et al., 2005) are first calculated based on the first P-wave arrival recorded by the  
295 two-component seismograms. The P-wave is primarily sensitive to P-wave speed,  $\alpha$ , so only  
296  $\alpha$  kernels are shown in the examples below. As discussed in the section of Simultaneous

297 backward and adjoint simulations, QuadSEM computes the Fréchet kernels using the same  
298 solutions of the forward and adjoint equations as the classical SEM (see Figure 2).

299 First, we examine the Fréchet kernel for *Model 2* (i.e.,  $\mathbf{m}_1 + v\delta\mathbf{m}$ ), and the full Hessian  
300 kernel for the model perturbation from *Model 1* (i.e.,  $\mathbf{m}_1$ ) to *Model 2*. The approximate  
301 Hessian kernels are computed for *Model 1* but using the adjoint source computed from  
302 the seismograms/measurements of *Model 2*. We show the Fréchet kernel, the approximate  
303 Hessian kernel, and the full Hessian kernel for P wavespeed  $\alpha$  in the first row of Figure 3  
304 (Part I) with a zoomed-in version around the perturbations given in the first row of Figure 3  
305 (Part II). The  $H_c$  is restricted to the perturbation indicated by the black box (see Figure 3c  
306 and its expression of eq. (7)). Note that the black box here is the  $H_c$  Hessian kernel with  
307 a negative value of  $10^{-9}$ , not the model perturbation although they are located in the same  
308 position. The  $H_b^{(s)}$  kernel is mostly invisible in Figure 3c except those around the black box  
309 due to its relatively small amplitude. The  $H_a$  and  $H_b^{(m)}$  are separated by the black box.  
310 Substantial differences are observed between the approximate and full Hessian kernels.

311 It takes the QuadSEM about a total of 37.57 mins with a maximum memory usage of  
312  $\sim 2.48$  GB to simultaneously compute both the Fréchet and full Hessian kernels using 4  
313 cores on a standard laptop (with 2.3 GHz Dual-Core Intel Core i5 processor and 8GB 2133  
314 MHz LPDDR3 memory). The computation of Fréchet kernels alone by the classical adjoint  
315 method takes about 18.76 mins with a maximum memory usage of 1.5 GB using the same  
316 computer. Therefore, in this case, a simultaneous computation of both the Fréchet and full  
317 Hessian kernels via QuadSEM roughly takes about 2 times the CPU time and  $\sim 1.53$  times  
318 the memory when compared to the computation of Fréchet kernels alone. The memory  
319 cost is slightly less than 2 because we use the same variables and matrices and these are  
320 independent of the wavefields generated from the two input models. The storage required  
321 for the QuadSEM is small due to the on-the-fly nature of the calculations, which takes 751  
322 MB disk space to store the absorbing boundary fields, the last-state forward fields as well  
323 as the seismograms. The wavefield storage method (WSM) stores the fields at all time steps

324 and requires about 513 GB disk space to store these fields.

## 325 **One source and three receivers**

326 We also show an example with one source and three receivers for the calculation of Hessian  
327 kernels, where *Model 1* is used as the background model (i.e.,  $\mathbf{m}_1$ ) and Hessian kernels are  
328 computed with respect to the perturbation in *Model 3* (i.e.,  $\mathbf{m}_1 + v\delta\mathbf{m}$ , and the  $v\delta\mathbf{m}$  here  
329 indicates a new perturbation for the three scatterers). The source is placed at  $(x, z) =$   
330  $(150 \text{ km}, -260 \text{ km})$  with the same source time function as in the section of Single source-  
331 receiver combination. Three receivers are placed on the top surface of the model located at  
332 horizontal locations of 100 km, 200 km, and 300 km, respectively. The total number of time  
333 steps and time intervals are the same as the example in the section of Single source-receiver  
334 combination.

335 The second row of Figure 3 (Part I) shows the Fréchet kernel, the approximate Hessian  
336 kernel, and the full Hessian kernel computed for  $\alpha$ , again for a traveltime adjoint source  
337 measured from the first-arrival of the P-wave. A zoomed-in version around the perturbations  
338 is given in the second row of Figure 3 (Part II). More detailed descriptions about the Fréchet  
339 and Hessian kernels are given in the figure caption. The computational cost for this example  
340 is almost the same as for that in the section of Single source-receiver combination since the  
341 forward and adjoint simulation time is almost independent of the number of receivers. There  
342 is one additional step in the window picking and computation of the adjoint source, which  
343 the cost is negligible compared to the field calculations. A few selected time steps of the  
344 regular wavefields and their perturbations computed by the QuadSEM are shown in Figure 4  
345 and Figure 5. For the on-the-fly implementation, QuadSEM, the key output files used in the  
346 forward simulation and in the simultaneous backward and adjoint simulation are presented  
347 in Figure 6.

## 348 **Kernel comparisons**

349 For both the one- and multi-receiver cases shown in Figure 3, we found substantial differences  
350 between the approximate Hessian kernels and the full Hessian kernels, in agreement with  
351 previous work (Fichtner and Trampert, 2011). Most notably, the amplitudes of the full  
352 Hessian kernels can be up to 100% stronger than those of the approximate Hessian kernels  
353 within the first and second Fresnel zones. These areas are covered by  $H_a$ ,  $H_b^{(m)}$ , and  $H_c$  in  
354 the full Hessian kernels and usually omitted in the calculation of the approximate Hessian  
355 kernels. The greater positive values of the Hessian kernels in the vicinity of the perturbation  
356 suggest that the inversion using the full Hessian kernels will result in better illumination in  
357 the region of the model perturbation. In comparison, when only the approximate Hessian  
358 kernel is used, the model updates tend to be distributed along the entire kernel.

359 In the multi-receiver case, we observe similarly higher amplitude in the Hessian kernels  
360 near the three model perturbations (Figure 3f) (Part I and II); whereas, for the approximate  
361 Hessian kernels, the sensitivity has high amplitudes around the middle anomaly only. This  
362 again suggests that using the full Hessian kernels in the inversion will focus model perturba-  
363 tions closer to the actual anomalies and hence provide better resolution for small anomalies  
364 within the model.

## 365 **Discussions**

366 The Hessian kernels are typically used with the Fréchet kernels for computing the model  
367 update or search direction based upon truncated Newton optimization (e.g., Nash, 1985;  
368 Grippo et al., 1989; Nash, 2000). This may potentially generate more accurate results and  
369 quicker convergence compared with L-BFGS based optimization for multi-parameter full-  
370 waveform inversion (FWI) (e.g., Métivier et al., 2013, 2014; Pan et al., 2017; Yang et al.,  
371 2018; Matharu and Sacchi, 2019).

372 The QuadSEM implementation can compute both the Fréchet and full Hessian kernels



373 simultaneously, requiring only double the computational cost of Fréchet kernels alone. To fur-  
374 ther reduce the computational costs, the source encoding techniques (e.g., Tromp and Bachmann,  
375 2019) or a reverse propagation of a superposition of forward and residual wavefields (e.g.,  
376 Robertsson et al., 2021) may be considered as well.

377 An important question remains as to whether the additional costs of the simultaneous  
378 computation of the Fréchet and full Hessian kernels at twice the computational cost can  
379 be offset by the more rapid convergence of the non-linear inversion. As high-performance  
380 computing becomes more accessible and efficient, this may become much less of a concern.

381 In addition to the model expressed in terms of density, P-wave and S-wave velocities,  
382 the approximate Hessian kernels and the full Hessian kernels can be expressed by different  
383 model components with similar derivations. In this study, we only show kernel examples for  
384 elastic models to demonstrate the on-the-fly approach. For anelastic models, a checkpointing  
385 method (e.g., Komatitsch et al., 2016) may be needed, and we are in the process of devel-  
386 oping such method to compute the anelastic full Hessian kernels by applying similar idea  
387 as QuadSEM but storing the two forward fields at some checkpoints in order to reconstruct  
388 them accurately during the adjoint simulation. The idea of QuadSEM is not limited to the  
389 SEM, and other numerical solvers that simulate seismic wave propagation can be adapted  
390 to compute the Fréchet kernels as well the full Hessian kernels.

## 391 Conclusions

392 We present QuadSEM, a package for on-the-fly full Hessian kernel calculation through si-  
393 multaneously solving four wave equations, which is designed to simultaneously compute the  
394 Fréchet and full Hessian kernels on the fly with only about double of the computational  
395 cost for the calculation of Fréchet kernels alone. In the QuadSEM, the WSM is also imple-  
396 mented as it is complementary to the on-the-fly approach. While the examples presented  
397 in this paper are rather specific to three elastic models, the underlying idea is very gen-

398 eral. The QuadSEM trades off the computational cost with storage and I/O, and improves  
399 the accuracy of full Hessian kernel calculations by combining the exact forward and adjoint  
400 wavefields on the fly in temporary memory. This makes it possible to use the accurate full  
401 Hessian information for multi-parameter FWI based upon the spectral-element and adjoint  
402 methods. It potentially provides a step forward in improving FWI to better image and un-  
403 derstand Earth structure, particularly in regions characterized by weak and/or small scale  
404 heterogeneities.

## 405 **Data and Resources**

406 No field data were used in this work. Models and the QuadSEM codes can be freely down-  
407 loaded via <https://github.com/yujiangxie/QuadSEM> or requested from the authors.

## 408 **Acknowledgments**

409 We acknowledge funding from the Natural Environment Research Council (NE/M003507/1)  
410 and the European Research Council (GA 638665) for supporting this work. Q. Liu is sup-  
411 ported by the Natural Sciences and Engineering Research Council of Canada (NSERC)  
412 Discovery Grant No. 487237. Some funding is also provided by the Wave Inversion Tech-  
413 nology (WIT) consortium. We thank all the developers of the SPECFEM2D/3D packages  
414 for their continued community support. We thank Fichtner Andreas for his helpful com-  
415 ments which significantly improved our methods after revision. We would also like to thank  
416 the Editor-in-Chief, Allison Bent, the Associate Editor and one anonymous referee for their  
417 valuable suggestions that greatly helped to improve the paper. Y. Xie would like to thank  
418 Daniel Peter for continued discussions on the Hessian kernels.

419 **Author contributions:** methodology and software, Y. Xie; validation, Y. Xie, C. Rychert,  
420 N. Harmon, Q. Liu and D. Gajewski; writing-original draft preparation, Y. Xie, C. Rychert,  
421 N. Harmon; writing-review and editing, Q. Liu and D. Gajewski; supervision, C. Rychert,  
422 N. Harmon, Q. Liu and D. Gajewski. All authors have read and agreed to the published  
423 version of the manuscript and the authors declare no conflict of interest.

## References

- Anderson, J. E., L. Tan, and D. Wang, 2012. Time-reversal checkpointing methods for RTM and FWI, *Geophysics*, **77**, S93–S103.
- Bamberger, A., G. Chavent, C. Hemony, and P. Lailly, 1982. Inversion of normal incidence seismograms, *Geophysics*, **47**, 757–770.
- Boehm, C., M. Hanzich, J. de la Puente, and A. Fichtner, 2016. Wavefield compression for adjoint methods in full-waveform inversion, *Geophysics*, **81(6)**, R385–R397.
- Boehm, C., and M. Ulbrich, 2015. A semismooth Newton-CG method for constrained parameter identification in seismic tomography, *SIAM Journal on Scientific Computing*, **37**, S334–S364.
- Bozdağ, E., D. Peter, M. Lefebvre, D. Komatitsch, J. Tromp, J. Hill, N. Podhorszki, and D. Pugmire, 2016. Global adjoint tomography: first-generation model, *Geophys. J. Int.*, **207**, 1739–1766.
- Charpentier, I., 2001. Checkpointing schemes for adjoint codes: application to the meteorological model Meso-NH, *SIAM J. Sci. Comp.*, **22**, 2135–2151.
- Chen, P., T. H. Jordan, and L. Zhao, 2007a. Full three-dimensional tomography: a comparison between the scattering-integral and adjoint-wavefield methods, *Geophys. J. Int.*, **170**, 175–181.
- Chen, P., L. Zhao, and T. H. Jordan, 2007b. Full 3D Tomography for the crustal structure of the Los Angeles region, *Bulletin of the Seismological Society of America*, **97**, 1094–1120.
- Dahlen, F., S.-H. Hung, and G. Nolet, 2000. Fréchet kernels for finite-frequency travel-times—I. Theory, *Geophys. J. Int.*, **141**, 157–174.
- Epanomeritakis, I., V. Akçelik, O. Ghattas, and J. Bielak, 2008. A Newton-CG method for large-scale three-dimensional elastic full-waveform seismic inversion, *Inverse Problems*, **24**, 034015.
- Faccioli, E., F. Maggio, R. Paolucci, and A. Quarteroni, 1997. 2D and 3D elastic wave propagation by a pseudospectral domain decomposition method, *J. Seismology*, **1**, 237–

451 251.

452 Faccioli, E., F. Maggio, A. Quarteroni, and A. Tagliani, 1996. Spectral-domain decomposition  
453 methods for the solution of acoustic and elastic wave equations, *Geophysics*, **61**, 1160–  
454 1174.

455 Fichtner, A., H. Bunge, and H. Igel, 2006a. The adjoint method in seismology — I. Theory,  
456 *Physics of the Earth and Planetary Interiors*, **157**, 86–104.

457 Fichtner, A., B. L. N. Kennett, H. Igel, and H. P. Bunge, 2009. Full waveform tomography  
458 for upper-mantle structure in the Australasian region using adjoint methods, *Geophys. J.*  
459 *Int.*, **179**, 1703–1725.

460 Fichtner, A., and J. Trampert, 2011. Hessian kernels of seismic data functionals based upon  
461 adjoint techniques, *Geophys. J. Int.*, **185**, 775–798.

462 French, S. W., and B. Romanowicz, 2015. Broad plumes rooted at the base of the Earth’s  
463 mantle beneath major hotspots, *Nature*, **525**, 95–99.

464 Gauthier, O., J. Virieux, and A. Tarantola, 1986. Two-dimensional nonlinear inversion of  
465 seismic waveforms: numerical results, *Geophysics*, **51**, 1387–1403.

466 Götschel, S., and M. Weiser, 2015. Lossy compression for PDE-constrained optimization:  
467 Adaptive error control, *Computational Optimization and Applications*, **62**, 131–155.

468 Griewank, A., and A. Walther, 2000. Algorithm 799: revolve: an implementation of check-  
469 pointing for the reverse or adjoint mode of computational differentiation, *ACM Trans.*  
470 *Math. Softw.*, **26**, 19–45.

471 Grippo, L., F. Lampariello, and S. Lucidi, 1989. A Truncated Newton Method with Non-  
472 monotone Line Search for Unconstrained Optimization, *J. Optim. Theory Appl*, **60**, 401–  
473 419.

474 Hanzich, M., F. Rubio, G. Aguilar, and N. Gutierrez, 2013. Efficient lossy compression for  
475 seismic processing, 75th EAGE Conference and Exhibition, EAGE, Expanded Abstracts,  
476 Th–16–08.

477 Igel, H., H. Djikpesse, and A. Tarantola, 1996. Waveform inversion of marine reflection

478 seismograms for P impedance and Poisson's ratio, *Geophys. J. Int.*, **124**, 363–371.

479 Komatitsch, D., and J. P. Vilotte, 1998. The spectral element method: an effective tool to  
480 simulate the seismic response of 2D and 3D geological structures, *Bull. Seism. Soc. Am.*,  
481 **88**, 368–392.

482 Komatitsch, D., Z. Xie, E. Bozdağ, E. S. d. Andrade, D. Peter, Q. Liu, and J. Tromp, 2016.  
483 Anelastic sensitivity kernels with parsimonious storage for adjoint tomography and full  
484 waveform inversion, *Geophys. J. Int.*, **206**, 1467–1478.

485 Lailly, P., 1983. The seismic inverse problem as a sequence of before stack migrations: Con-  
486 ference on Inverse Scattering, Theory and Application, Society for Industrial and Applied  
487 Mathematics, Expanded Abstracts, 206–220.

488 Lindstrom, P., P. Chen, and E. Lee, 2016. Reducing disk storage of full-3D seismic waveform  
489 tomography (F3DT) through lossy online compression, *Comput. Geosci.*, **93**, 45–54.

490 Lions, J., 1968. Contrôle optimal de systèmes gouvernés par des équations aux dérivées  
491 partielles, Dunod Gauthier-Villars.

492 Liu, Q., and J. Y. Gu, 2012. Seismic imaging: From classical to adjoint tomography, *Tectono-*  
493 *physics*, **566**, 31–66.

494 Liu, Q., and J. Tromp, 2006. Finite-Frequency Kernels Based on Adjoint Methods, *Bulletin*  
495 *of the Seismological Society of America*, **96**, 2383–2397.

496 Luo, Y., R. Modrak, and J. Tromp, 2014. Strategies in Adjoint Tomography. In: *Freedon*  
497 *W., Nashed M., Sonar T. (eds), Handbook of Geomathematics*, Springer.

498 Matharu, G., and M. Sacchi, 2019. A subsampled truncated-Newton method for multipa-  
499 rameter full-waveform inversion, *Geophysics*, **84(3)**, R333–R340.

500 Métivier, L., F. Bretaudeau, R. Brossier, S. Operto, and J. Virieux, 2014. Full waveform  
501 inversion and the truncated Newton method: quantitative imaging of complex subsurface  
502 structures, *Geophysical Prospecting*, **62**, 1353–1375.

503 Métivier, L., R. Brossier, J. Virieux, and S. Operto, 2013. Full Waveform Inversion and the  
504 Truncated Newton Method, *SIAM Journal on Scientific Computing*, Society for Industrial

505 and Applied Mathematics, **35**, B401–B437.

506 Nash, S. G., 1985. Preconditioning of truncated-Newton methods, SIAM J. Sci. Stat. Com-  
507 put., **6**, 599–616.

508 Nash, S. G., 2000. A survey of truncated-Newton methods, Journal of Computational and  
509 Applied Mathematics, **124**, 45–59.

510 Pan, W., K. A. Innanen, and W. Liao, 2017. Hessian-free Gauss-Newton full-waveform inver-  
511 sion via l-BFGS preconditioned conjugate-gradient algorithm, Geophysics, **82(2)**, R49–  
512 R64.

513 Plessix, R.-É., 2006. A review of the adjoint-state method for computing the gradient of a  
514 functional with geophysical applications, Geophys. J. Int., **167**, 495–503.

515 Pratt, R., C. Shin, and G. J. Hick, 1998. Gauss-Newton and full Newton methods in  
516 frequency-space seismic waveform inversion, Geophys. J. Int., **133**, 341–362.

517 Robertsson, J. O., F. Andersson, and R.-É. Plessix, 2021. Efficient snapshot-free reverse  
518 time migration and computation of multiparameter gradients in full waveform inversion,  
519 Geophysics, **0**, 1–79.

520 Seriani, G., and E. Priolo, 1994. Spectral element method for acoustic wave simulation in  
521 heterogeneous media, Finite Elements in Analysis and Design, **16**, 337–348.

522 Symes, W. W., 2007. Reverse time migration with optimal checkpointing, Geophysics, **72**,  
523 SM213–SM221.

524 Tape, C., Q. Liu, A. Maggi, and J. Tromp, 2009. Adjoint tomography of the southern  
525 California crust, Science, **325**, 988–992.

526 Tarantola, A., 1984. Inversion of seismic reflection data in the acoustic approximation, Geo-  
527 physics, **49**, 1259–1266.

528 Tarantola, A., 1988. Theoretical background for the inversion of seismic waveforms, including  
529 elasticity and attenuation, Pure appl. Geophys., **128**, 365–399.

530 Tromp, J., 2020. Seismic wavefield imaging of Earth’s interior across scales, Nature Reviews  
531 Earth & Environment, **1**, 40–53.

- 532 Tromp, J., and E. Bachmann, 2019. Source encoding for adjoint tomography, *Geophys. J.*  
533 *Int.*, **218**, 2019–2044.
- 534 Tromp, J., C. Tape, and Q. Liu, 2005. Seismic tomography, adjoint methods, time reversal  
535 and banana-doughnut kernels, *Geophys. J. Int.*, **160**, 195–216.
- 536 Unat, D., T. Hromadka, and S. Baden, 2009. An adaptive sub-sampling method for in-  
537 memory compression of scientific data, *Data Compression Conference, IEEE, DCC'09*,  
538 262–271.
- 539 Virieux, J., and S. Operto, 2009. An overview of full waveform inversion in exploration  
540 geophysics, *Geophysics*, **74**, WCC1–WCC26.
- 541 Walther, A., and A. Griewank, 2004. Advantages of Binomial Checkpointing for Memory-  
542 reduced Adjoint Calculations. In: Feistauer M., DolejšV., Knobloch P., Najzar K. (eds)  
543 *Numerical Mathematics and Advanced Applications*, Springer.
- 544 Weiser, M., and S. Götschel, 2012. State trajectory compression for optimal control with  
545 parabolic PDEs, *SIAM J. Sci. Comput.*, **34**, A161–A184.
- 546 Yang, P., R. Brossier, L. Métivier, J. Virieux, and W. Zhou, 2018. A Time-Domain Precon-  
547 ditioned Truncated Newton Approach to Visco-acoustic Multiparameter Full Waveform  
548 Inversion, *SIAM Journal on Scientific Computing*, **40**, B1101–B1130.
- 549 Zhu, H., E. Bozdağ, D. Peter, and J. Tromp, 2012. Structure of the European upper mantle  
550 revealed by adjoint tomography, *Nat. Geosci.*, **5**, 493–498.



551 **Full mailing address for each author**

552 **Dr. Yujiang Xie**

553 Ocean and Earth Science, National Oceanography Centre Southampton (NOCS)

554 University of Southampton Waterfront Campus

555 European Way, Southampton SO14 3ZH, UK

556

557 **Prof. Catherine Rychert**

558 Ocean and Earth Science, National Oceanography Centre Southampton (NOCS)

559 University of Southampton Waterfront Campus

560 European Way, Southampton SO14 3ZH, UK

561

562 **Prof. Nicholas Harmon**

563 Ocean and Earth Science, National Oceanography Centre Southampton (NOCS)

564 University of Southampton Waterfront Campus

565 European Way, Southampton SO14 3ZH, UK

566

567 **Prof. Qinya Liu**

568 Department of Physics & Department of Earth Sciences, University of Toronto

569 22 Ursula Franklin Street, Toronto, ON, M5S 3B1, Canada

570 **Prof. Dirk Gajewski**

571 Department of Earth Sciences, Universität Hamburg

572 Room: 1432, Geomatikum, Bundesstrasse 55

573 20146 Hamburg, Germany

574 **List of Figure Captions**

575

576 **Figure 1.** Sketch illustrating the workflow of forward simulation for classical SEM vs.  
577 QuadSEM.

578 **Figure 2.** Sketch illustrating the workflows for the simultaneous backward and adjoint  
579 simulations for classical SEM vs. QuadSEM.

580 **Figure 3.** Fréchet and Hessian kernels computed for the investigated models.

581 **Figure 4.** Four selected time steps of the five wavefields computed by the QuadSEM.

582 **Figure 5.** A few time steps of selected perturbed fields computed on the fly using the  
583 first-order finite-difference approximation.

584 **Figure 6.** Key files output from the forward simulation and the simultaneous backward and  
585 adjoint simulation in the QuadSEM.

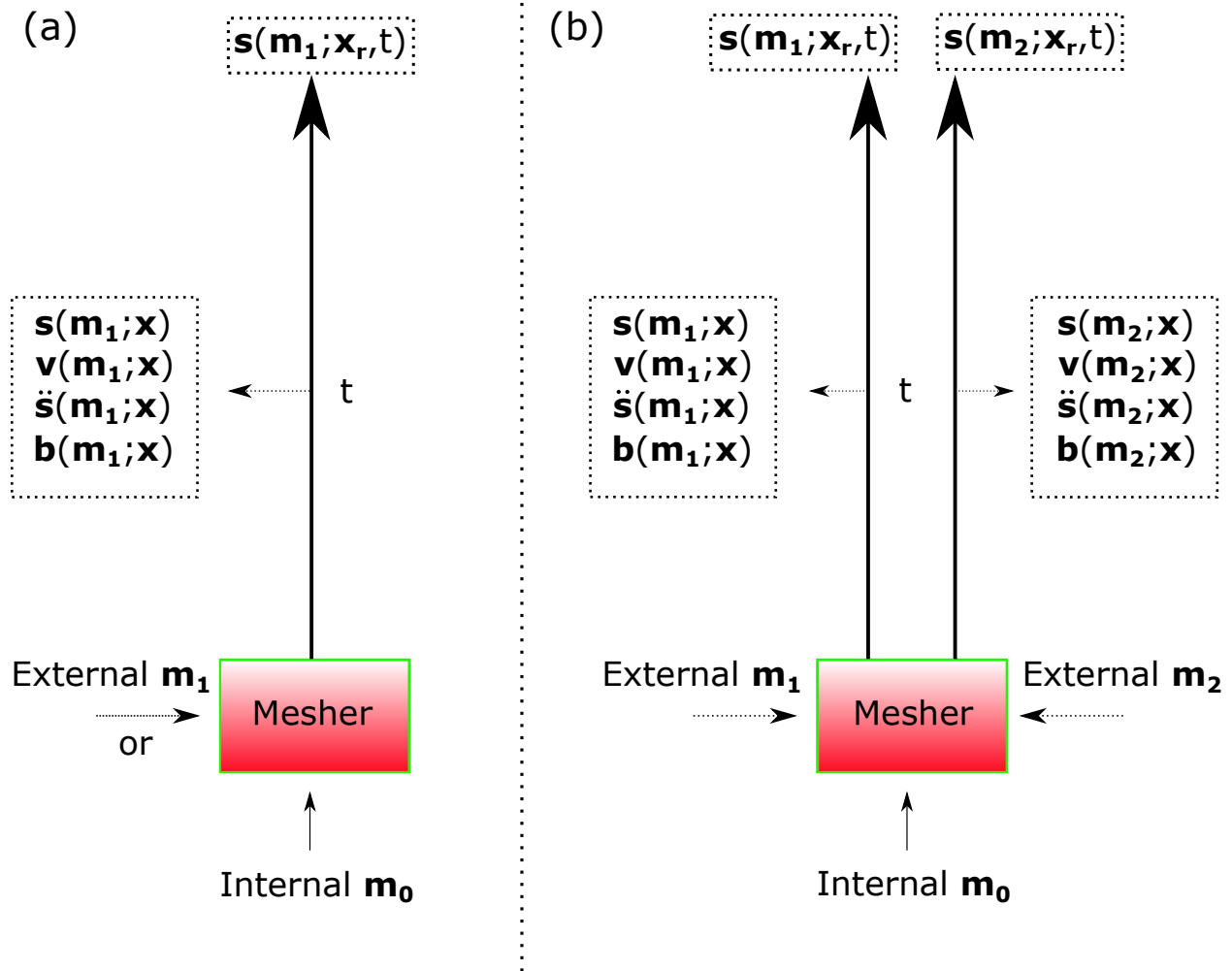


Figure 1: Sketch illustrating the workflow of forward simulation for classical SEM vs. QuadSEM. (a) In classical SEM forward simulation, a single model is used and it is set either by the internal mesher (e.g.,  $\mathbf{m}_0$ ) or importing from external file ( $\mathbf{m}_1$ ) after the mesher is set up. (b) In the QuadSEM forward simulation, two models ( $\mathbf{m}_1$  and  $\mathbf{m}_2$ ) are imported into the internal mesher, where  $\mathbf{m}_2 = \mathbf{m}_1 + v\delta\mathbf{m} = \mathbf{m}_1 + \Delta\mathbf{m}$ , and  $\mathbf{m}_0$  will be omitted when external models are loaded.

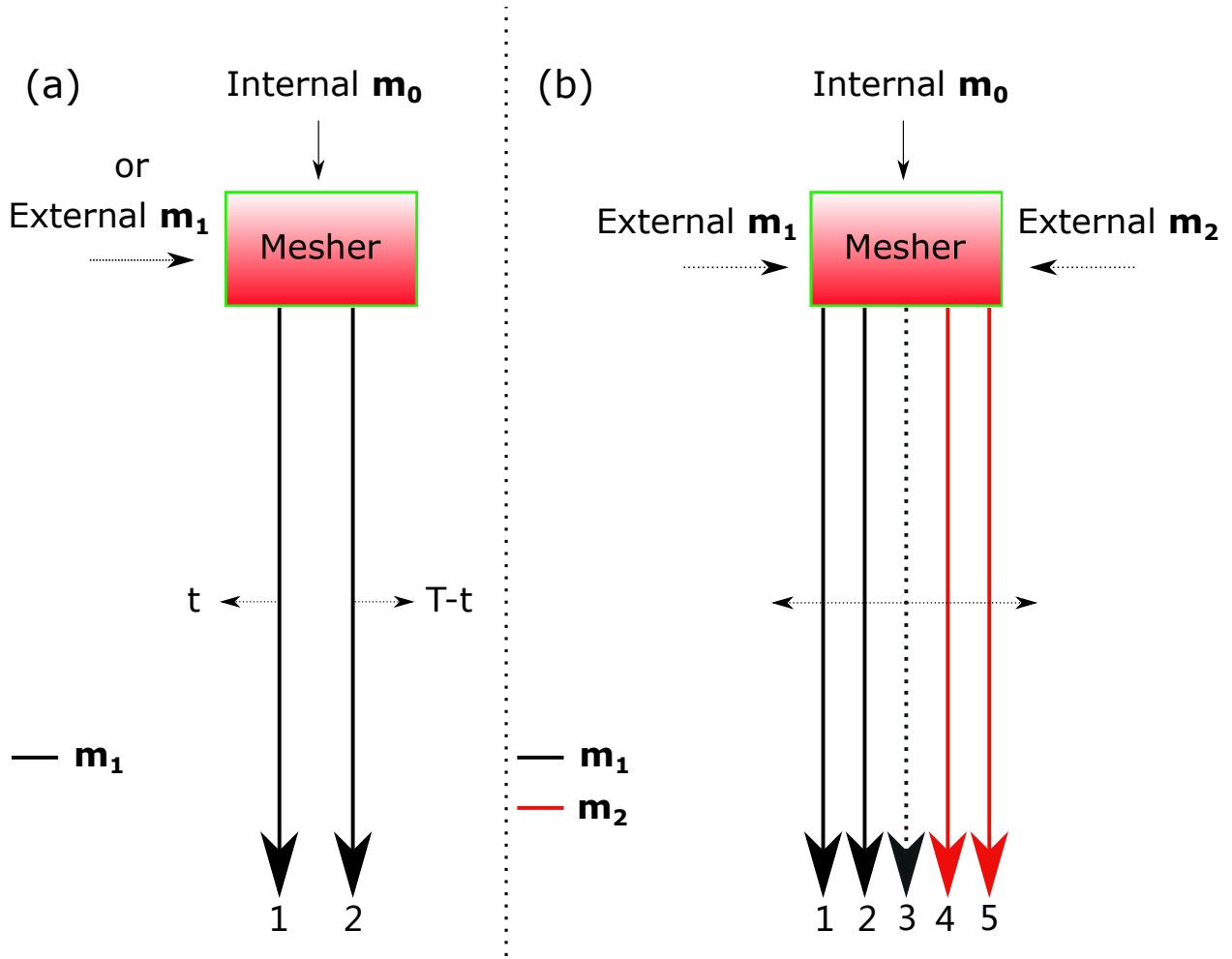


Figure 2: Sketch illustrating the workflows for the simultaneous backward and adjoint simulations for classical SEM vs. QuadSEM. (a) In the simultaneous backward and adjoint simulation of the classical SEM. Each arrow represents the solution for one wave equation with Arrow 1 indicating the backward simulation (i.e. the reconstruction of the forward field) and Arrow 2 indicating the adjoint simulation. (b) In the simultaneous backward and adjoint simulation of the QuadSEM, Arrows 1, 2, and 3 indicate the solutions of the wave equations for model  $\mathbf{m}_1$ , and Arrows 1 and 2 perform the same as in (a), and Arrow 3 performs the same as Arrow 2 but its adjoint source is computed by using the measurements of  $\mathbf{m}_2$  to account for the perturbation of the adjoint source. The red Arrows 4 and 5 indicate the computation of the backward and adjoint fields for model  $\mathbf{m}_2$ .

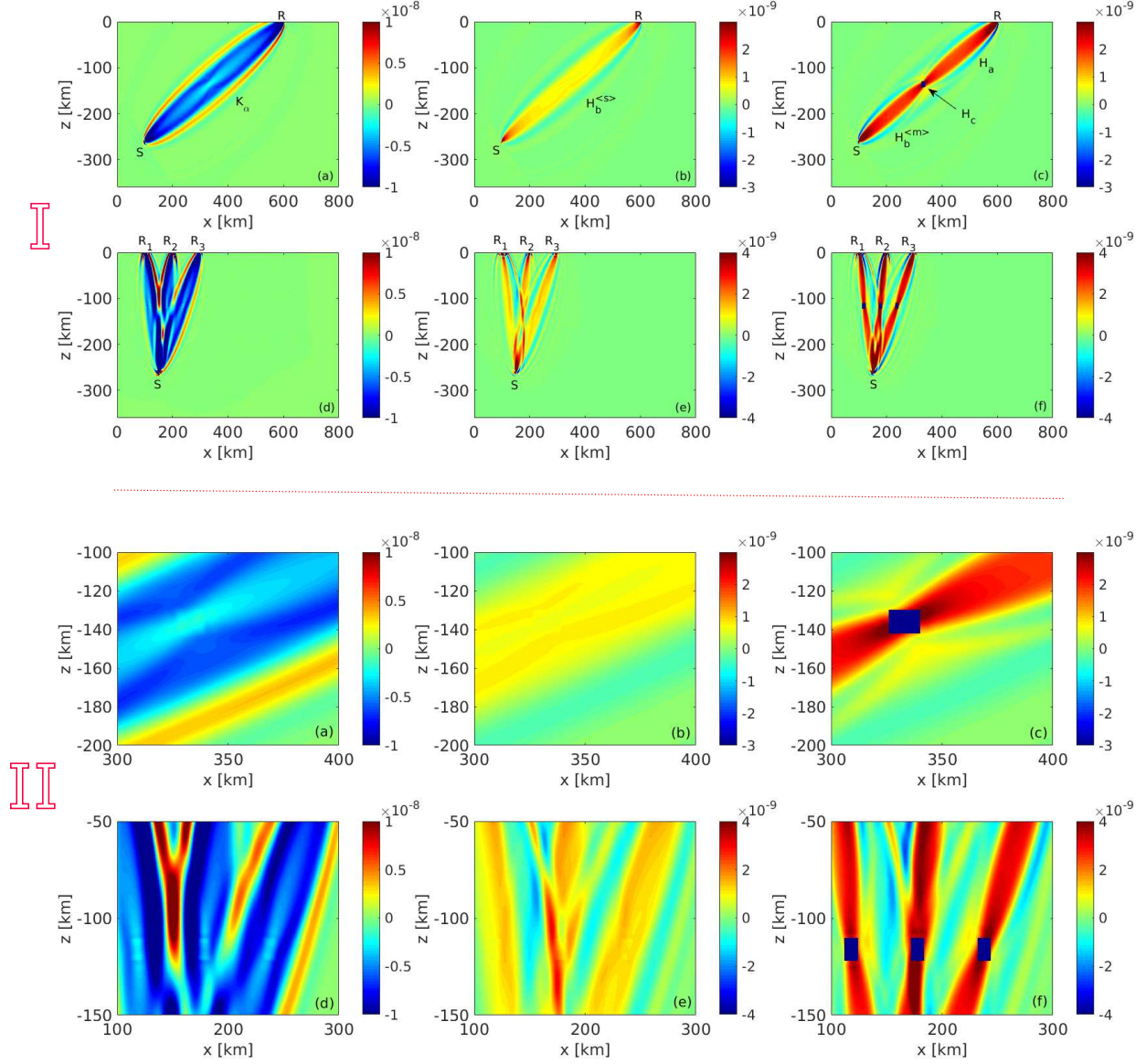


Figure 3: Part I: Fréchet and Hessian kernels computed for *Model 2* (top row) and *Model 3* (bottom row) as discussed in section of Numerical examples. In the top row we show (a) the Fréchet kernel  $K_\alpha$ , (b) the approximate Hessian kernel  $H_b^{(s)}$ , and (c) the full Hessian kernel for  $\alpha$  for the single source single station case with a single scattering object. Similarly, Panels (d), (e), and (f) in the bottom row show the various kernels for the case of a single source and three stations with three scattering objects. The kernel unit for all sub-figures is  $[s m^{-2}]$ . A zoomed view of the perturbations within Part I is shown in Part II.

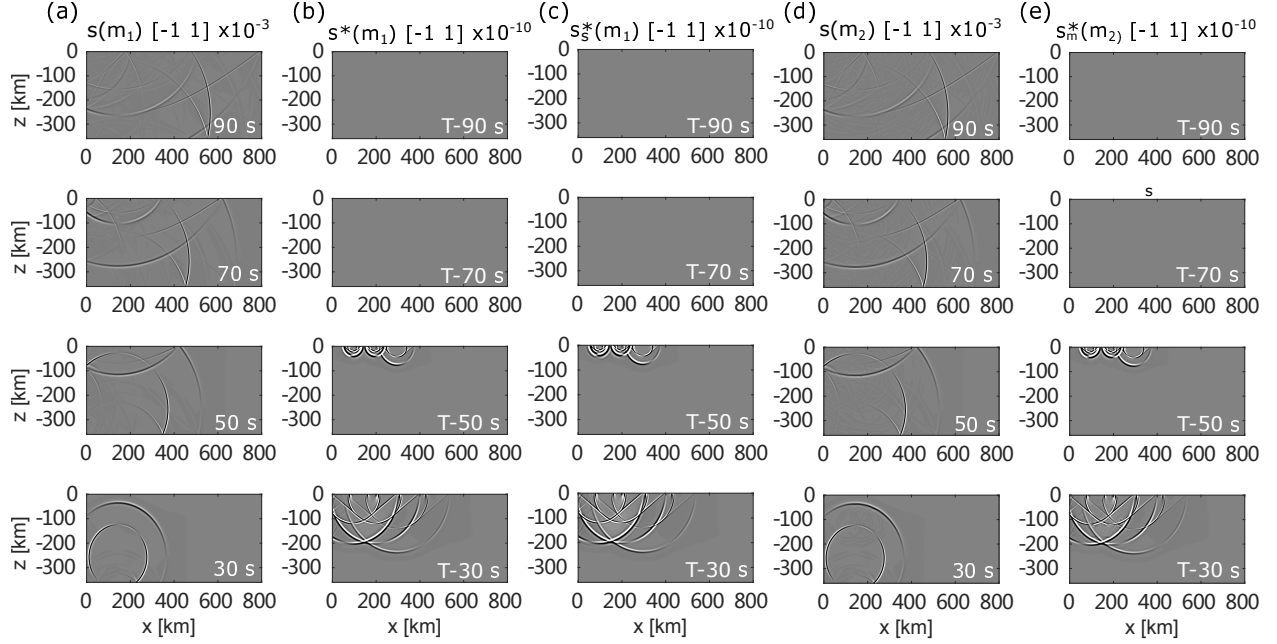


Figure 4: Four selected time steps of the five wavefields computed by the QuadSEM using the on-the-fly approach. (a) The forward fields recorded at times 30 s, 50 s, 70 s, and 90 s for model  $\mathbf{m}_1$ . (b) The adjoint fields for the same model but recorded at reversed times of T-90 s, T-70 s, T-50 s, and T-30 s, where  $T = 100$  s in this test. (c) The adjoint fields generated by the adjoint source computed from the measurements for  $\mathbf{m}_2$ . (d) and (e) show the similar simulation as (a) and (b) but for model  $\mathbf{m}_2$ , instead of  $\mathbf{m}_1$ . (b) and (e) look similar due to the use of the same adjoint source, but they are different after the adjoint fields travel through the scatterers.

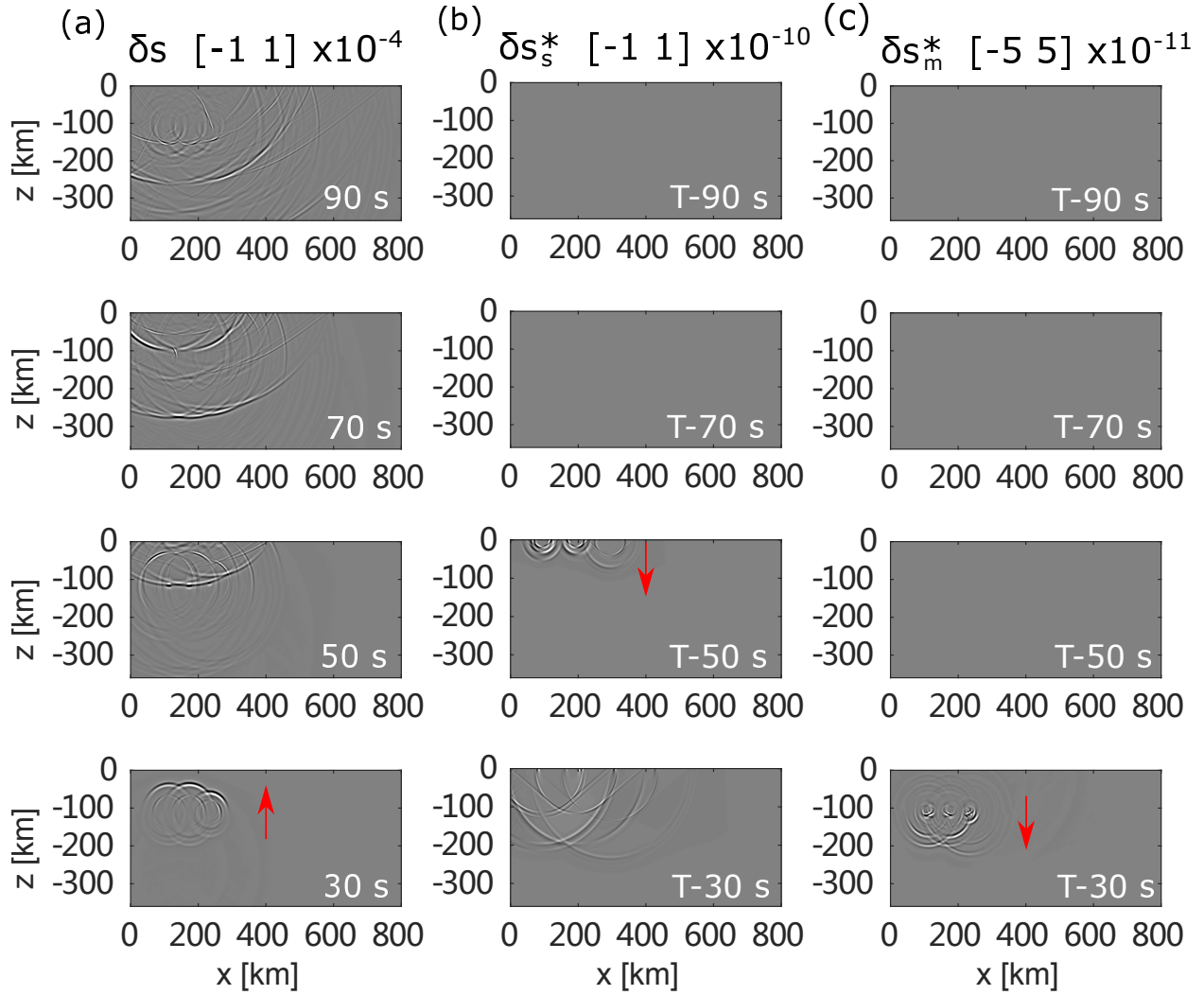


Figure 5: A few time steps of selected perturbed fields computed on the fly using the first-order finite-difference approximation. (a) shows the perturbed forward fields. (b) shows the perturbed adjoint fields due to the perturbation of the adjoint source. (c) shows the perturbed adjoint fields due to the perturbation of the model. The perturbed fields, e.g., generated around the red arrows, are due to the perturbations either from the model or from the adjoint source.

Forward output	Backward and adjoint output
Database*****.bin	
absorb_elastic_bottom*****.bin absorb_elastic_left*****.bin absorb_elastic_right*****.bin absorb_elastic_bottom_m2_*****.bin absorb_elastic_left_m2_*****.bin absorb_elastic_right_m2*****.bin	proc*****_rho_kappa_mu_kernel.dat proc*****_rho_kappa_mu_kernel_Ha.dat proc*****_rho_kappa_mu_kernel_Hbm.dat proc*****_rho_kappa_mu_kernel_Hbs.dat proc*****_rho_kappa_mu_kernel_Hc.dat proc*****_rho_kappa_mu_kernel_Habc.dat
AA.S****.BXX.semd AA.S****.BZX.semd AA.S****.BXX.semd_m2 AA.S****.BZX.semd_m2	proc*****_rhop_alpha_beta_kernel.dat proc*****_rhop_alpha_beta_kernel_Ha.dat proc*****_rhop_alpha_beta_kernel_Hbm.dat proc*****_rhop_alpha_beta_kernel_Hbs.dat proc*****_rhop_alpha_beta_kernel_Hc.dat proc*****_rhop_alpha_beta_kernel_Habc.dat
lastframe_elastic*****.bin lastframe_elastic_m2_*****.bin	

Figure 6: Key files output from the forward simulation and the simultaneous backward and adjoint simulation in the QuadSEM. The left column shows the files output from the forward simulation. The first row shows the meshing database which includes the internal model to be replaced by the two external models before the main time loop in the simultaneous backward and adjoint simulation. The second row shows the absorbing boundary fields, where the shadow part indicates files output for model  $\mathbf{m}_2$ . The third and fourth rows show the seismograms registered at the receivers and the last state of the forward field. These files output in the forward simulation will be used in the simultaneous backward and adjoint simulation. The right column shows the key files output in the simultaneous backward and adjoint simulation, including the Fréchet kernels, the approximate Hessian kernels ('Hbs'), and the full Hessian kernels ('Habc'), etc. In the right column, the top part shows kernels for the  $\rho$ ,  $\kappa$ , and  $\mu$  parameter set and the bottom part shows kernels for the  $\rho$ ,  $\alpha$ , and  $\beta$  parameter set.



## 586 **Simultaneous backward and adjoint simulations for the** 587 **computation of Fréchet and Hessian kernels on the fly**

588 Figure A1 shows the simultaneous backward and adjoint simulations for the computation  
589 of Fréchet kernels on the fly, where the backward simulation is designed to reconstruct the  
590 forward fields backward in time during the adjoint simulation. In this way, the Fréchet  
591 kernels can be constructed on the fly since the forward fields for time  $t$  and the adjoint  
592 fields for time  $(T - t)$  or vice versa can be simultaneously accessed. The  $T$  indicates the  
593 total simulation time. This on-the-fly strategy can be extended to compute the Hessian  
594 kernels but the solutions of two forward and two adjoint equations are combined. Figure A2  
595 shows the simultaneous computation of several forward and adjoint fields for constructing  
596 the Hessian kernels on the fly.

## 597 **Computation of Hessian kernels by wavefield storage** 598 **method**

599 Hessian kernels can be computed when the required wavefields are available. To get the  
600 required fields, we use one forward simulation and two adjoint simulations (see Figure A3).  
601 The forward simulation is to compute and save four forward fields, that is  $\mathbf{s}(\mathbf{m}_1)$ ,  $\mathbf{s}(\mathbf{m}_2)$ ,  
602  $\ddot{\mathbf{s}}(\mathbf{m}_1)$ ,  $\ddot{\mathbf{s}}(\mathbf{m}_2)$ , where  $\mathbf{m}_2 = \mathbf{m}_1 + v\delta\mathbf{m}$ . The first adjoint simulation (Adjoint simulation I)  
603 is designed to compute and save the adjoint field  $\mathbf{s}^\dagger(\mathbf{m}_2)$ . The second adjoint simulation  
604 (Adjoint simulation II) is a simultaneous adjoint simulation and the Hessian calculation.  
605 Figure A3 shows the computation of full Hessian kernels. It can be similarly changed for the  
606 computation of approximate Hessian kernels, where one needs to store  $\mathbf{s}_s^\dagger(\mathbf{m}_1)$  in the Adjoint  
607 simulation I. Figure A4 shows each component of the Hessian kernels with respect to the  
608 three model parameters. For this example, the *Model 2* and the single source-receiver pair

609 (as the section of Single source-receiver combination) are used. Only the first P-wave arrival  
610 is used for the calculation of the traveltimes adjoint sources.

611 **List of Figure Captions for Appendix**

612 **Figure A1.** Forward simulation and the simultaneous backward and adjoint simulation for  
613 computing the Fréchet kernels.

614 **Figure A2.** Simultaneous backward and adjoint simulation in the QuadSEM for the com-  
615 putation of Hessian kernels on the fly.

616 **Figure A3.** A workflow illuminating the computation of full Hessian kernels by the wave-  
617 field storage method (WSM).

618 **Figure A4.** Four components of the full Hessian kernels with respect to the model given in  
619  $\rho$ ,  $\alpha$ , and  $\beta$ .

620

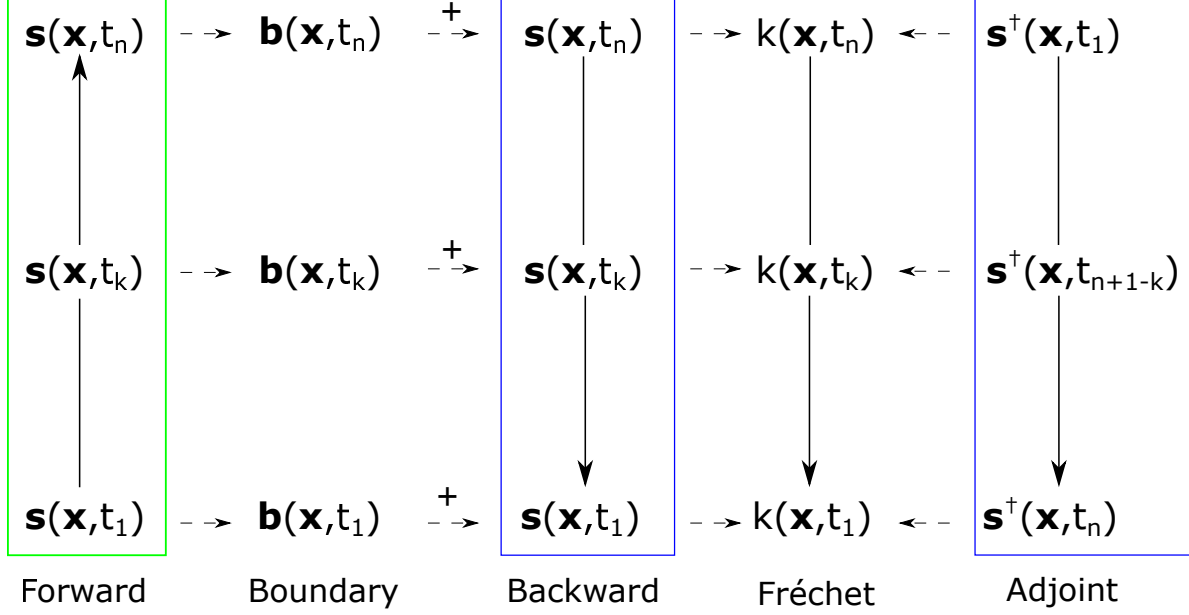


Figure A1: Forward simulation (green rectangle) and the simultaneous backward and adjoint simulation (two blue rectangles) for computing the Fréchet kernels. The forward simulation is started from the first time step  $t_1$  and ended at the last time step  $t_n$ . The absorbing boundary field  $\mathbf{b}(\mathbf{x}, t_k)$  of each time step  $t_k$  and the last state field  $\mathbf{s}(\mathbf{x}, t_n)$  are stored in the forward simulation. The backward simulation takes the last state field as a start point and reconstructs the forward field backward in time. In each time step, the absorbing boundary field  $\mathbf{b}(\mathbf{x}, t_k)$  is re-injected into the backward simulation to reconstruct the forward fields (called backward fields here). The adjoint simulation is started from the time-reversed adjoint source from the receivers. The Fréchet kernels at each time step or at a sub-sampled time step are constructed on the fly based upon the backward and adjoint fields. If each time step is used, the kernels are summed at each time step until the final step as  $K_m = \sum_{k=1}^n K(\mathbf{x}, t_k)\delta t$ , where  $\delta t$  is time interval in the simulation.

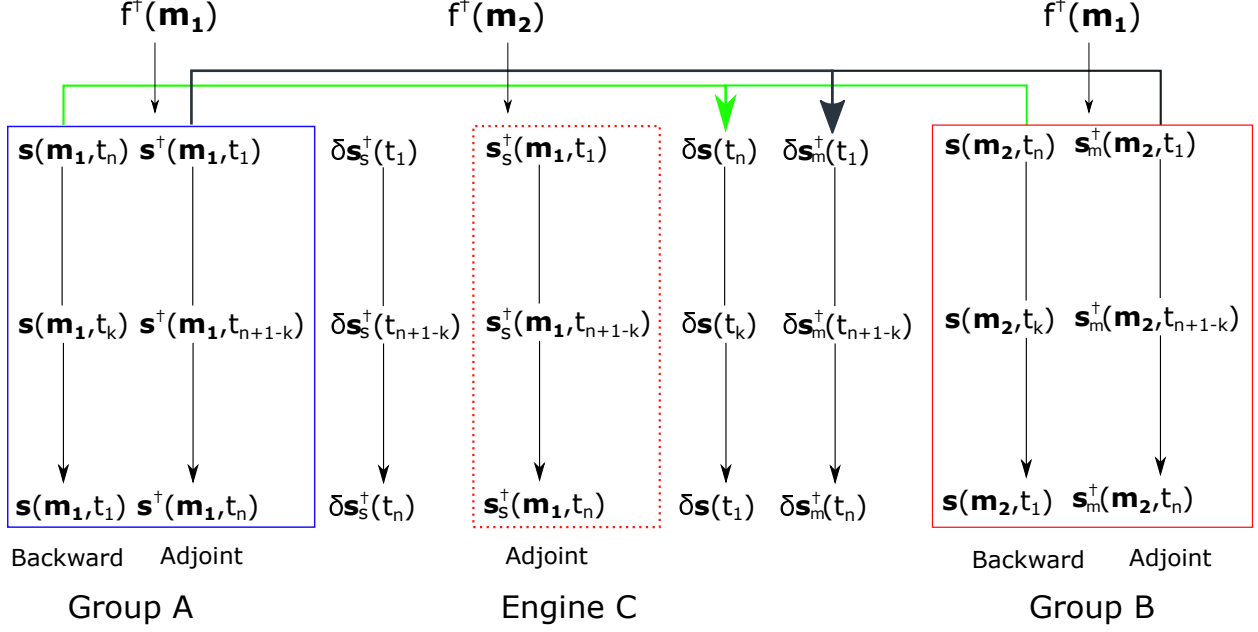


Figure A2: Simultaneous backward and adjoint simulation in the QuadSEM for the computation of the Hessian kernels on the fly. Group A: the solutions of forward and adjoint equations are combined and used for model  $\mathbf{m}_1$ , where one solution is solved for the backward simulation and the other is solved for the adjoint simulation. Group A is designed to compute the backward and adjoint fields for model  $\mathbf{m}_1$ . On the right side, Group B combines the solutions of the forward and adjoint equations for model  $\mathbf{m}_2$ . Engine C represents one solution of the adjoint equation designed to compute the adjoint field due to the perturbation of the adjoint source  $f^\dagger(\mathbf{m}_2)$ . The simulation in Engine C is the same as the adjoint simulation in Group A except the source term is different. We design the workflow to show the computation of each required wavefield. In the full Hessian kernel calculations, we use Group A and Group B, just changing the adjoint source of Group B from  $f^\dagger(\mathbf{m}_1)$  to  $f^\dagger(\mathbf{m}_2)$ . Since all the fields are computed on the fly for each designed time step (each time step or adaptive time step), the perturbed fields to be used in the calculation of Hessian kernels can be computed also on the fly.

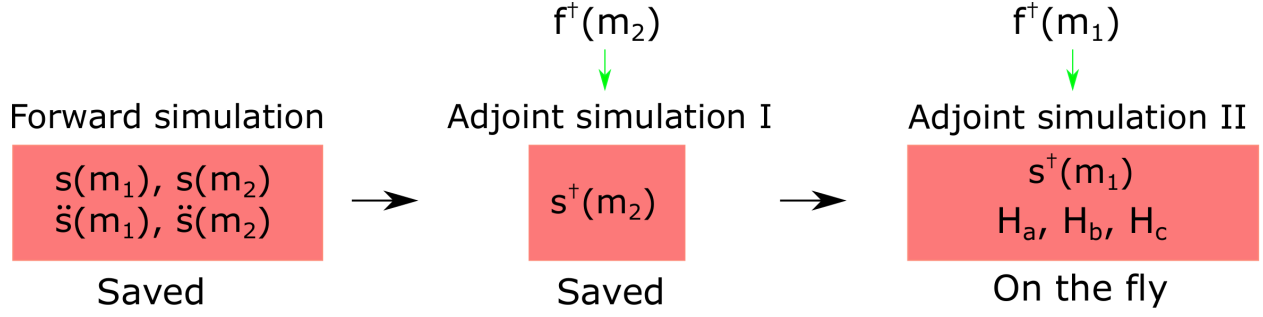


Figure A3: A workflow illuminating the computation of full Hessian kernels by the wavefield storage method (WSM) using the adaptive time integration or storing the fields at all time steps. The first step (Forward simulation) is designed to compute and save the forward fields and the second step (Adjoint simulation I) is to compute and save one adjoint field due to the perturbations of model and adjoint source. The last step (Adjoint simulation II) is to compute one adjoint field  $\mathbf{s}^+(\mathbf{m}_1)$  on the fly, and read one time step of the saved five fields into the temporary memory for the computation of the full Hessian kernels. The  $\mathbf{f}^\dagger(\mathbf{m}_1)$  and  $\mathbf{f}^\dagger(\mathbf{m}_2)$  denote the two adjoint sources computed from the measurements of the two models.

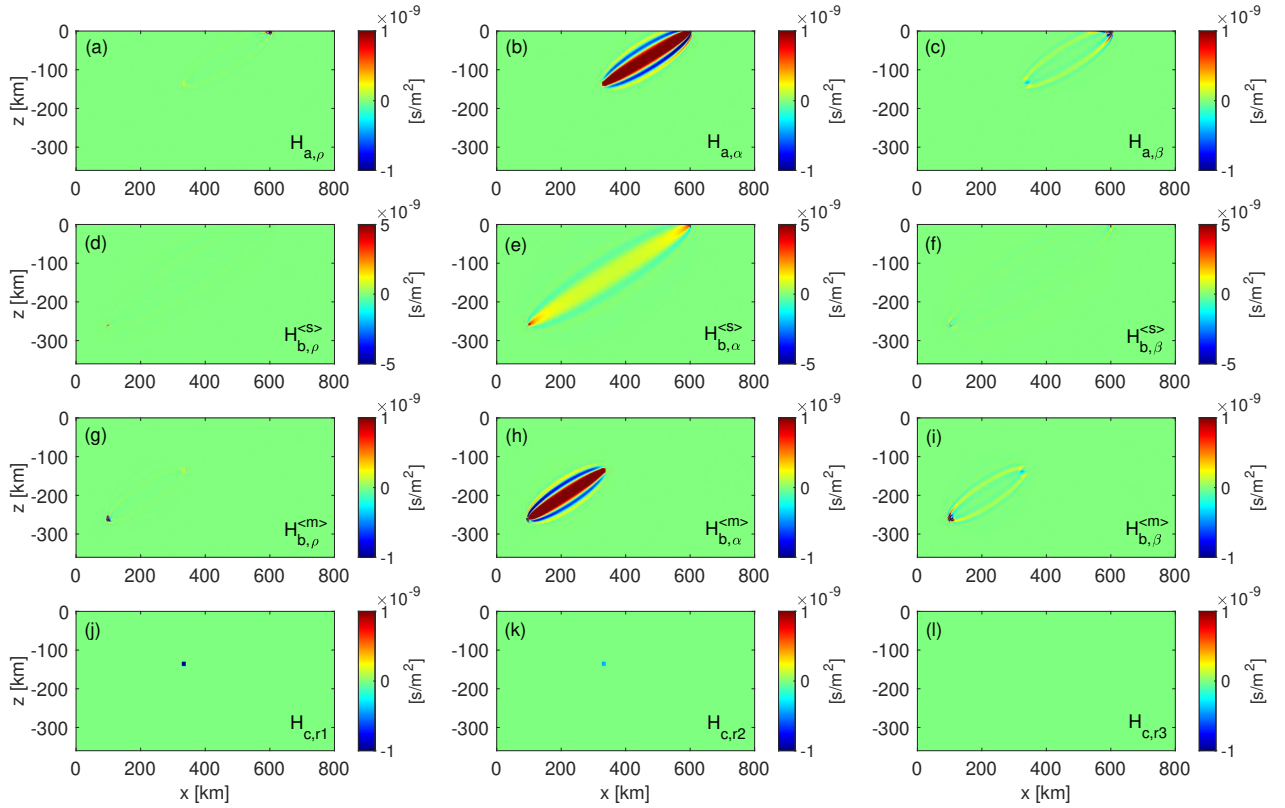


Figure A4: Four components of the full Hessian kernels with respect to the model given in  $\rho$ ,  $\alpha$ , and  $\beta$ . The top first row shows the  $H_a$  component with respect to the three model parameters. Only the  $H_{a,\alpha}$  is well observed since only the first P-wave arrival is used for the adjoint source calculation. The second row shows the  $H_b^{(s)}$  component, which is approximate Hessian kernels due to the perturbation of the adjoint source to the adjoint field. The third row shows the  $H_b^{(m)}$  component which is due to the perturbation of the model for the adjoint field. The bottom row shows the  $H_c$  component. Only the kernels for  $H_{c,r1}$  and  $H_{c,r2}$  are observed since the  $K_\beta$  is very small due to the use of the first P-wave arrival only. The  $ri$  (where  $i = 1, 2, 3$ ) indicates the three rows in the  $H_c$  expression. The full Hessian kernels are obtained by summing the four components together.