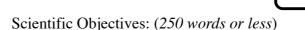
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Title:	Cenozoic Pacific Equatorial Age	Transect -	- Following the Palaeo	equator	
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Keywords: (5 or less)	Paleoceanography, Equatorial Pa	cific, Cend	ozoic	Area:	Eastern Equatorial Pacific
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Abstract: (400 words or less)

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As the largest ocean, the Pacific is intricately linked to major changes in the global climate system that took place during the Cenozoic. Throughout the Cenozoic the Pacific plate has had a northward component. Thus, the Pacific is unique, in that the thick sediment bulge of biogenic rich deposits from the currently narrowly focused zone of equatorial upwelling is slowly moving away from the equator. Hence, older sections are not deeply buried and can be recovered by drilling. Previous ODP Legs 138 and 199 were designed as transects across the paleo-equator in order to study the changing patterns of sediment deposition across equatorial regions, while this proposal aims to recover an orthogonal "age-transect" along the paleo-equator. Both previous legs were remarkably successful in giving us new insights into the workings of the climate and carbon system, productivity changes across the zone of divergence, time dependent calcium carbonate dissolution, bio- and magnetostratigraphy, the location of the ITCZ, and evolutionary patterns for times of climatic change and upheaval. Together with older DSDP drilling in the eastern equatorial Pacific, both Legs also helped to delineate the position of the paleo-equator and variations in sediment thickness from approximately 150°W to 110°W.

As we have gained more information about the past movement of plates, and where in time "critical" climate events are located, we now propose to drill an age-transect ("flow-line") along the position of the paleo-equator in the Pacific, targeting selected time-slices of interest where calcareous sediments have been preserved best. Leg 199 enhanced our understanding of extreme changes of the calcium carbonate compensation depth across major geological boundaries during the last 55 million years. A very shallow CCD during most of the Paleogene makes it difficult to obtain well preserved sediments, but we believe our siting strategy will allow us to drill the most promising sites and to obtain a unique sedimentary biogenic carbonate archive for time periods just after the Paleocene-Eocene boundary event, the Eocene cooling, the Eocene/Oligocene transition, the "one cold pole" Oligocene, the Oligocene-Miocene transition, and the Miocene, contributing to the objectives of the IODP Extreme Climates Initiative, and providing material that the previous legs were not able to recover.



We propose an ocean drilling cruise with the aim to achieve an age transect along the paleoequatorial Pacific spanning the early Eocene to Miocene (with earlier intervals being covered by previous ODP Legs). Drill sites target specific time-slices of interest, at locations that provide optimum preservation of calcareous sediments. Recovered cores will contribute towards (1) resolving questions of how and why paleo-productivity of the equatorial Pacific changed over time, (2) provide rare material to validate and extend the astronomical calibration of the geological time scale for the Cenozoic, (3) determine sea-surface and benthic temperature and nutrient profiles and gradients, (4) provide important information about the detailed nature of calcium carbonate dissolution and changes of the CCD, (5) enhance our understanding of bio- and magnetostratigraphic datums at the equator, as well as (6) provide information about rapid biological evolution and turn-over during times of climatic stress. (7) As our strategy also implies a paleo-depth transect, we hope to improve our knowledge about the reorganization of water masses as a function of depth and time. (8) Integrated with additional site-survey proposals, we intend to make use of the high level of correlation between tropical sediment sections and seismic stratigraphy to develop a more complete model of equatorial circulation and sedimentation. (9) Due to the northward component of the Pacific plate motion, our siting strategy also implies a limited N-S transect across the paleo-equator for some of the proposed time slices, providing additional information about N-S hydrographic gradients.

Please	describe	below	any	non-standard	measurements	technology	needed	to	achieve	the	proposed	scientific
objecti	ves.											

Proposed Sites:

		Water	Pe	netration (m)		
Site Name	Position	Depth (m)	Sed	Bsm	Total	Brief Site-specific Objectives	
PEAT-1B	141.8°W 12.3°N	4947	~200m	5m	~205m	Obtain calcareous sediments along an paleo-equatorial agetransect to decipher paleoceanography and paleoclimatology from the:	
PEAT-2B	140.8°W 11.5°N	4911	~200m	5m	~205m	middle Eocene	
PEAT-3B	138.3°W 10.6°N	4866	~200m	5m	~205m	middle/late Eocene	
PEAT-4B	131.7°W 7.5°N	4789	~250m	5m	~255m	E/O boundary	
PEAT-5B	128.2°W 7.6°N	4533	~320m	5m	~325m	Oligocene	
PEAT-6B	126.1°W 5.1°N	4399	~300m	5m	~305m	O/M boundary	
PEAT-7B	123.0°W 4.0°N	4511	~300m	5m	~305m	Miocene	
PEAT-8B	118.0°W 1.3°N	3965	~350m	5m	~355m	Miocene	

Following the Equator in Time and Space

A proposal for IODP drilling (16th Jan 2003, revised March 2004)

Introduction

The equatorial surface ocean circulation is inescapably linked to the trade wind system. The equatorial Pacific is the classic "world ocean" example of this linkage. It is dominated by wind-driven circulation and is largely unfettered by ocean boundaries. Here, the equator itself is characterized by a narrow zone of divergence that results from the change in the sign of the Coriolis effect and that gives rise to a band of high biologic productivity. The strength of the equatorial circulation and of this divergence is linked to the strength of the trade winds, which are in turn strongly tied to the global climate system. Variation in global climate, interhemispheric differences in temperature gradients, and marked changes in the ocean boundaries are all imprinted on the biogenicrich sediments that are accumulating in the equatorial zone. This proposal is designed to provide an understanding of equatorial Pacific circulation, and the carbonate production, deposition and dissolution, for the last 55 million years at a scale where orbital forcing can be measured. Combined with seismic reflection data following in the vein of Mitchell et al. (2003) and synthesized with earlier drilling (e.g. Moore et al., 2003; Moore et al., submitted) we can reconstruct equatorial Pacific history with high confidence and substantially improve upon work from the early stages of DSDP and recent ODP Legs.

Drilling for the history of the equatorial Pacific has been greatly simplified by a favorable motion for the Pacific Plate. Throughout the Cenozoic, the movement of the Pacific plate has had a northward component of around 0.25°/Myr. This northward movement transports the equatorial sediments gradually out from under the zone of highest sediment delivery, resulting in a broad mound of biogenic sediments (Fig. 1). This transport prevents the older equatorial sections from being buried deeply beneath the younger sections as the crust moves northward. The diminished overburden resulting from this transport also allows relatively good preservation of the biogenic debris. In addition it allows us to core nearly all sections using the advanced piston coring technique. The northward tectonic displacement, however, is not so large that a traverse of the equatorial zone (within 2° of the equator) was too rapid to record a reasonable period of equatorial ocean history. Typically drill sites remain within the equatorial zone for 10-20 Myr before passing beyond the northern edge of high biogenic sedimentation. Older equatorial sections are thus buried beneath a thin veneer of younger sediments as the crust moves northward. The diminished overburden resulting minimizes burial diagenesis of the biogenic debris. It also allows APC piston coring of much of this section with the right placement strategy for drill sites (Lyle et al., 2002).

In their summary of the Deep Sea Drilling results in the equatorial Pacific, van Andel et al. (1975) gave a general view of the development of the equatorial mound of sediments in the Pacific Ocean, based mostly upon three early drilling legs, DSDP legs 5, 8 and 16. They showed how both temporal and spatial variation in sediment accumulation rates resulted from plate movement, varying biologic productivity at the equatorial divergence, and varying carbonate preservation. The buildup of the Pacific equatorial mound of sediment has been more recently documented and discussed by Mitchell (1998, 2003; Fig. 1).

Drilling across the Pacific equatorial mound was addressed again some twenty years after the van Andel et al. (1975) compilation when ODP Leg 138 drilled an equatorial transect along 10 Ma crust, and then again 10 years later when ODP Leg 199 conducted a similar transect along 56 Ma crust. The newer drilling, coupled with major advances in geochronology, has documented the remarkable correlation of paleoceanographic events over thousands of km in the equatorial Pacific, caused by the large scale of Pacific

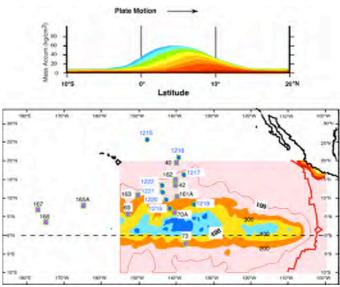


Figure 1: Top: Model cross section of equatorial sediment mound taking into account the northward drift of the Pacific plate. Bottom: Mapped thickness of the Pacific equatorial sediment mound. (Both modified from Mitchell, 1998; Mitchell et al., 2003).

equatorial circulation (Fig. 2). It is possible, with the addition of a relatively small number of new sites, to build detailed reconstructions of equatorial Pacific circulation through the Cenozoic.

The early drilling missed most of this detail because of the lack of important drilling technologies. APC coring, MST correlation, and core-log integration now allow the collection of relatively undisturbed sediments, the rebuilding of a continuous sediment column from individual cores, and the correlation to seismic reflection data.

That, and improved knowledge of the plate tectonic regime, will allow us to locate the areas of enhanced deposition rates associated with the paleo-equator. Combining multiple holes along the equator will result in a detailed record from the Pleistocene to the

Paleocene. These records will also be invaluable for the continued development of the Cenozoic time scale as well as for the paleoceanographic information they contain. ODP Legs 138 and 199 in the Pacific equatorial area have recovered excellent sections on which the detailed orbital tuning of the geologic time scale is being carried out. These sections give a much clearer picture of variation in sedimentation rates, isotopic evolution of the oceans, biologic

evolution and zoological provenance, variations in carbonate preservation, and variations in geochemical fluxes that result from paleoceanographic and paleoclimate changes. There are still parts of the Cenozoic time scale that require further refinement and verification of the proposed orbital tuning. The time scale older than the late Eocene has not been tuned at all, even though there is evidence of orbital frequencies in parts of the records recovered from this older interval (e.g., Norris and Röhl, 1999; Röhl

et al., 2001).

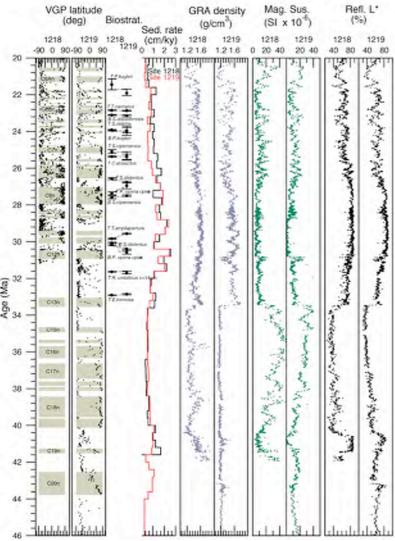


Figure 2 from Shipboard Scientific Party (2002): Coherence of sediment properties between widely separated drill sites in the equatorial Pacific is very high, allowing correlation of sediment properties over hundreds of kilometers. These two sites from ODP Leg 199 are more than 740 km apart.

Scientific objectives overview

The aim of this proposal is to recover the best possible carbonate sediment from overlapping drill sites at the paleo-equator of the Pacific, for time periods for which either no good-quality records exist yet, or that are otherwise important to answer specific scientific questions. Highlights include:

Interplay of carbonate productivity, dissolution, the carbon cycle, and basin-wide seismic stratigraphy:

Results from Leg 199 hinted that a very large change in the carbonate saturation state of the Pacific occurred at around the Eocene/Oligocene boundary, and a few million years earlier during the Eocene. Yet, due to the change in the CCD, only very limited carbonate material was recovered for the interval preceding this re-organization of the carbon cycle. We aim to recover a carbonate record across this important interval in order to answer whether it was changes in the surface water productivity that were at least partially responsible to initiate this change. Many of the climate proxies that provide information about the detailed carbon cycle processes are linked to wellpreserved carbonate material.

Furthermore, we would like to investigate the spatial variability of carbonate deposition for specific time slices. Analysis of older cores, combined with the sparse seismic lines available for the Pacific so far, indicate that in addition to temporal variations of carbonate deposition, there was also a spatial aspect, in that the Eocene and Oligocene pattern of carbonate deposition was more diffuse than the present-day sharply defined pattern of the equatorial upwelling zone. Sedimentary patterns are highly correlated on hundreds to thousands of kilometers in the Pacific, and by drilling along an agetransect at the position of the paleoequator in combination with seismic survey lines we can develop a highly detailed seismic stratigraphy for the Pacific basin.

Perfecting the astronomical calibration of bio- and magnetostratigraphic markers and the geological time-scale:

The determination of rates of geological processes are an important input for climate modelers, global stratigraphic correlation, and geological investigations that rely on the duration of events. Recent results have extended the material available for developing detailed time scales, but need more validation. The middle Eocene has proven extremely difficult to calibrate, mainly due to lack of high-quality carbonate records. New material would be invaluable to develop a highly detailed chrono-stratigraphy for the entire Cenozoic.

The Overall Strategy

To develop a detailed history of the Pacific equatorial current system, the strategy pursued in the most recent ODP legs (199 and 138) was to drill along a line of equal age of the ocean crust. This was done in order to obtain a ~N-S transect across the major E-W currents and to overcome the problem of relatively poor carbonate preservation in the Pacific, particularly in times when the CCD was very shallow.

Drilling at a paleo-ridgecrest allows the recovery of the shallowest sections available in the pelagic oceans, and thereby assures the best possible preservation of the carbonate sediments recovered. However, as the crust cools and sinks, the seafloor on which the sediments are deposited becomes deeper and deeper - and closer to the lysocline and CCD. Thus, the really well preserved part of the sections recovered in such "time-line" transects is restricted by the

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depth at which carbonate preservation is substantially diminished, as well as by the northward movement of sediment sections out of the region of high equatorial productivity. This limitation was exemplified by the results from ODP Leg 199, which recovered only limited amounts of carbonate prior to the Eocene/Oligocene boundary.

In this drilling effort we propose to overcome this limitation of the "timeline" strategy by pursuing an equatorial age transect, or "flow-line" strategy (Fig. 3), to collect well-preserved equatorial sections through the Cenozoic, while also making use of the Pacific plate motion to add an oblique transect across all time slices.

Figure 3: Age of ocean crust in the North Pacific (Renkin and Sclater, 1988) based on magnetic lineations compiled by Cande et al. (1986) and the tectonic histories of Hilde et al. (1977), Larson, (1976), and Rea and Dixon (1983). Contour interval is 5 M.y. (Age model from Berggren et al., 1985). Diagonally shaded area indicates regions where the age of the ocean floor is uncertain. Rectangles indicate estimated location of the paleoequator for the proposed intervals of study. Vertical dotted bars indicate approximate longitudinal positions of seismic transects conducted for ODP Legs 138 and 199.

We propose to drill a series of sites on the paleo-equator at key intervals in the evolution of the Cenozoic climate. These intervals span the very warm times of the Eocene, through the cooling of the late Eocene and Oligocene, the early Miocene time of relatively warm climates (or low ice volume), and into sections deposited during the development of the major southern and northern hemisphere ice sheets (Fig. 4). There are only very few previous sites that match our site selection criteria (Table 1). Each site is selected to be close to the geographic paleo-equator and on crust of an age slightly older than the intervals of particular interest.

In this way we track the paleoceanographic conditions at the paleo-equator, in the best preserved sediments obtainable. Integrated with additional site-survey proposals, we will also make use of the high level of correlation between tropical sediment sections and seismic stratigraphy to develop a more complete model of equatorial productivity and sedimentation.

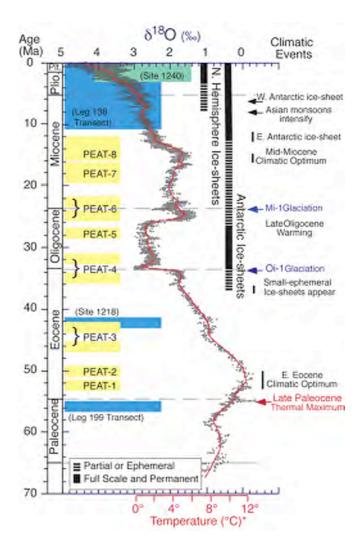


Figure 4: Time slices of specific climatic interest proposed for sampling. Sites with slightly older crustal ages on the Pacific paleoequator (after Zachos et al., 2001b).

Leg	Site	#Holes	Drilling	Carbonate finds (age in Ma)			
		(target)	tech.	age (Ma)	young	old	
DSDP 5	42	2	R	51-49	20	51	
DSDP 9	78	1	R	27-28	11	34	
DSDP 9	79	2	R	16-20	0	24.5	
ODP 199	1217	3	A/X	49-53	32	53.5, 34	
ODP 199	1218	3	A/X	49-53	11	42.5	
ODP 199	1219	1 of 2	A/X	49-53	19.5	34, 55.5	
ODP 199	1220	1 of 3	A/X	49-53	26, 53	34, 55.5	
ODP 199	1221	1 of 3	A/X	49-53	31, 54	55.5	
ODP 199	1222	1 of 2	A/X	49-53	30, 53	34, 55.5	

Table 1: Overview of older DSDP and ODP drill sites close to our proposed target area. Only sites close to target area are listed; only very few sites (bold) cover our time slice intervals, are in the right area, have the right crustal age to preserve calcium carbonate, or have employed modern drilling technology (APC-coring) to ensure recovery of undisturbed and continuous sediment sections.

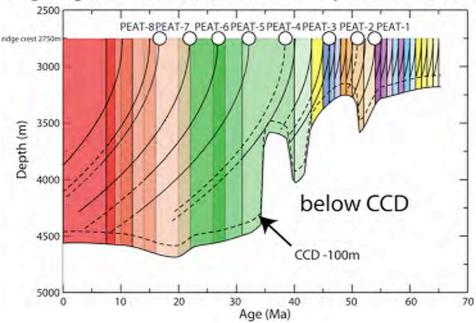
Selecting target ages

The times selected to be drilled in this proposal were chosen to pay particular attention to the overall climatic history of the Cenozoic, and to target particular times of marked changes in the climatic regime. The spacing of the sites will be determined by what we know of the paleo-lysocline and by results of previous

This separation of 5 Myr is a minimum for the shallow CCD of the Eocene; for good preservation of foraminifers an even closer spacing is needed.

The results of our paleo-equator reconstruction, and proposed drill site locations are shown in Fig. 6.

Targeting drill sites based on CCD history and subsidence



drilling. Where the CCD is particularly shallow, the spacing in time of agetransect sites needs to be closer than when the CCD is deep (Fig. 5). As a guide, Site 1218 was drilled on 42 Ma crust during times when the CCD was near 3.3 km. Nannofossil oozes were deposited at this location up to about 37 Ma before the crust at this site sank below the CCD.

Figure 5: Targeting Pacific equatorial age transect ("PEAT") drill sites for based on the CCD history (van Andel (1975), with new data from ODP Leg 199) to constrain the presence of calcium carbonate. Subsidence curves use a subsidence parameter k=0.35 (a conservative estimate with respect to whether curves trace above the CCD). Additional subsidence due to sediment loading was not modeled.

The Eocene (PEAT-1A to PEAT-4A)

The Eocene was a time of extremely warm climates that reached a maximum in temperatures near 52 Ma (Fig.4). From this maximum there was a gradual climatic cooling through to the Eocene-Oligocene Boundary. There appears to

have been a slight reversals in this trend within the middle Eocene, near 43 Ma, and in the late Eocene at 34-36 Ma, just prior to the marked drop in oxygen isotopes that marks the Eocene-Oligocene boundary and one of the most dramatic changes of the CCD (Fig. 5).

Proposed drilling targets based on revised seafloor age, CCD history and backtracking

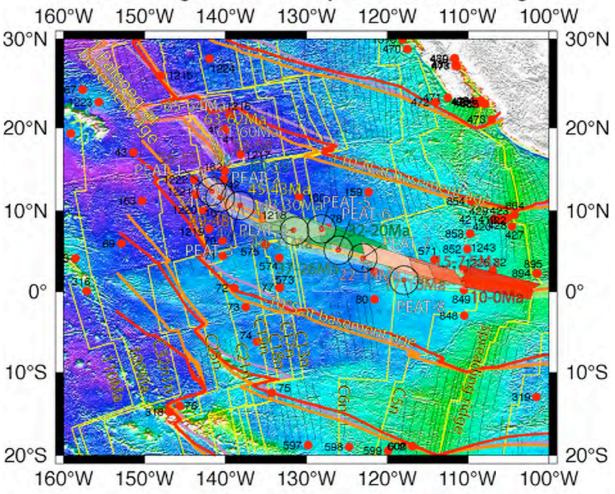


Figure 6: Proposed locations for site-survey work and drilling, corresponding to time slices targeted in this proposal. Shown are present-day bathymetry, revised magnetic anomaly isochrons (thin yellow lines, modified with new points from Petronotis 1991, 1994), and paleo-equator position at the crustal age obtained by backtracking, using fixed-hotspot stage poles from Koppers et al. (2001, red), Engebretson et al. (1985, orange), and palaeomagnetic poles from Sager & Pringle (1988, lilac). The shaded band lies within one degree north and south of the paleo-equator (averaged from fixed-hotspot rotation models). Colored areas correspond to those in Fig. 3 and represent time intervals of interest. These were obtained by intersecting the white paleo-equator area with the younger end of the time interval of interest, which was then backrotated to the older boundary of the time slice. This method requires correction if backtracking occurs across fracture zones.

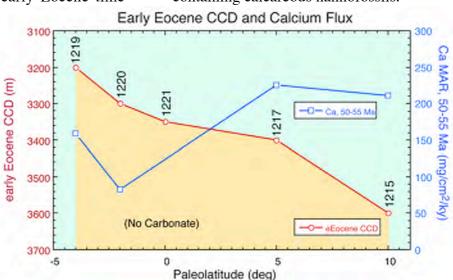
Throughout the Eocene, the CCD lay near 3.2 – 3.3 km depth. Thus, recovering well preserved carbonate sediments from the equatorial region is a substantial challenge, but not impossible if East Pacific Rise depth lay near the global average of 2.7 km.

Early and Middle Eocene (PEAT-1B, 2B)

Leg 199 drilled a N-S transect across the equatorial region at about 56 Ma. Sites on this transect had generally drifted below the CCD by 52-53 Ma. Thus we presently lack calcareous sediments from the region of the equatorial circulation system during this time of maximum Cenozoic warmth. Average (non-carbonate) accumulation rates at this time were moderate, showing only slight increases in some of the more northern sites on the Leg 199 transect (1215, 1216, 1220). What is particularly interesting in the records of Leg 199 is that the very shallow CCD of this early Eocene time

appears to deepen to the north, perhaps suggesting

Figure 7: Early
Eocene CCD and
C a mass
accumulation rates
as a function of
paleolatitude of
ODP Leg 199 Sites.



a northern source for the bottom waters (Fig. 7). Sites targeting this time interval would ideally give us sediments with sufficient carbonate material to better constrain the isotopic and biotic characteristics of the near surface equatorial waters.

Middle and Late Eocene (PEAT-3B)

Good paleomagnetic stratigraphy in Leg199 Sites allowed a much improved calibration of nannofossil and radiolarian biostratigraphic datums. Together, these stratigraphies allowed the development of a more detailed picture of temporal variations in sediment accumulation through the mid and upper Eocene of the tropical Pacific. These data show a factor of up to 2-3x increase in accumulation rates of siliceous ooze within the middle Eocene (41 to 45 Ma). At Site 1219 this interval of high biogenic flux is also marked by the preservation of intervals containing calcareous nannofossils.

High siliceous sedimentation occurs at a time of an apparent short reversal in the mid Eocene cooling trend (Fig. 4). It is difficult to interpret the cause of such a substantial change in silica flux during a very warm climatic regime. At the very least we need good carbonate recovery during this interval in order to apply the substantial array of carbonate-based proxies to this interval in order to evaluate the temperature and structure of the near-surface ocean.

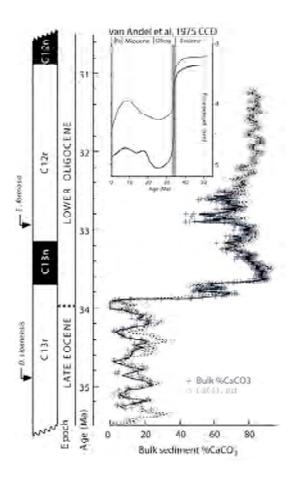
Eocene-Oligocene Boundary (PEAT-4B)

Results of ODP Leg 189, near the Tasman Rise (Exon, Kennett, Malone et al., 2001) confirmed that the Antarctic-Australian Passage opened to deep-water flow around the time of the Eocene-Oligocene boundary; however, there may have been substantial flow at shallower depths a few million years earlier. An increase in terrigenous flux measured at Site 690 in the Weddel Sea indicates that climatic cooling and the onset of ice rafting from Antarctica occurred at about 35 Ma (Joseph and Rea, 2002) - about a million years earlier than the Eocene

Figure 8: Bulk %CaCO₃ in ODP Leg 199 Site 1218 across the Oligocene – Eocene Boundary (P.A. Wilson et al., unpublished data). Note pronounced cyclical nature below and across the boundary, and very low %CaCO₃ in the uppermost Eocene and before.

Oligocene boundary. These results are borne out by drilling in Prydz Bay (O'Brien, Cooper, Richter et al., 2001) that show evidence of ice streams reaching down to the coast before the end of the Eocene.

Site 1218 contains the most complete Eocene-Oligocene boundary section recovered from the equatorial Pacific; however, it is far from pristine. Carbonate percentages drop markedly below the boundary and go to zero near 34 Ma (Fig. 8). At Site 1218, on the 40 Ma equator, there is a marked increase in the proportion of diatoms in the siliceous fraction of sediments in the uppermost Eocene.



This increased diatom abundance continues well into the lower Oligocene and coincides with the biggest turnover in the radiolarian fauna since the Cretaceous-Tertiary boundary. The large diatom abundance in the uppermost Eocene and lower Oligocene then drops off in the upper Oligocene. There is a comparable drop in measured terrigenous flux at Site 690 (Joseph and Rea, 2002).

This apparent latest Eocene cooling of the climate and increased primary productivity at low latitudes seems somewhat at odds with the apparent slight warming indicated by the oxygen isotopes (Fig. 2). These oddities, together with the major changes in planktonic assemblages, suggest an important restructuring of the upper mixed layer and thermocline waters in the Pacific that continues into the lower Oligocene. Well recovered and well preserved equatorial sections across this relatively short transition between warm and cold global climates will be very valuable in determining the impact of high latitude ocean boundary changes on climate, circulation, and productivity in the equatorial region.

The Oligocene (PEAT-5B)

This interval of time is noted for its markedly heavy oxygen isotopes (Fig. 4) and its relatively deep CCD (Fig. 5). It is generally thought that there was ice on

Antarctica during this interval, but not the large ice sheets to be found there later in the mid Miocene. There is no compelling evidence for ice sheets in the Northern Hemisphere during this time. Thus, there was apparently a relatively low global ice volume, relatively cold bottom waters, a relatively cold South Pole, and a relatively warm North Pole. This scenario of a "one cold pole" world has given rise to speculation on the impact of interhemispheric temperature imbalance on pole to equator temperature gradients and on the symmetry of the global wind systems. The extent to which such an imbalance may have affected the trade winds, the position of the Inter Tropical Convergence Zone, and the seasonal shifts in this zone, should be seen in the wind-driven currents of the equatorial region.

The older, low-resolution DSDP data indicate relatively high, but variable sediment accumulation rates during this interval and better carbonate preservation to the south of the equator (van Andel et al., 1975). In the Leg 199 equatorial transect, the highest accumulation rates encountered (>15 m/M.yr.) occurred in the lower part of the Oligocene, but these were in sites north of the Oligocene equator, or on relatively old (and therefore deep) crust. Thus we should expect a better preserved, thicker carbonate section

at the Oligocene equator.

Studies of Oligocene sections from Leg 199 and from other ODP sites (e.g., Paul et al., 2000, Zachos et al., 2001a) indicate the presence of strong eccentricity and obliquity cycles in carbonate preservation and suggest a strong (southern) high-latitude influence on the carbonate record. These cycles are leading to the development of an orbitally tuned time scale that reaches back to the base of the Oligocene. Such a time scale will make it possible to develop a very detailed picture of equatorial geochemical fluxes and of the degree of variability in the equatorial system of the Oligocene.

<u>The Latest Oligocene – Earliest Miocene</u> (PEAT-6B)

At the end of the Oligocene there is a dramatic rise in the oxygen isotope record (Fig. 4), whether through a decrease in global ice volume or increase in bottom water temperatures (or both) is uncertain. It is closely followed by a relatively short, sharp increase in oxygen isotope values that has been interpreted as a major glacial episode (Mi-1; Fig. 4) and correlated to a pronounced drop in sea level (Miller et al., 1991). This event is very close to the Oligocene-Miocene boundary. Although there are clear isotopic signals indicating major changes in ice volume, ocean temperatures and/or

ocean structure, this boundary has always been somewhat of an enigma. Unlike the major changes in the isotopic stratigraphy, the biostratigraphies of the planktonic microfossils show very little change at all across this boundary. In fact it is one of the most difficult epoch boundaries to pick using solely the microfossil biostratigraphies.

In Sites 1218 and 1219 of ODP Leg 199 this interval was well recovered; however, carbonate preservation still presented a problem for the classic foraminiferal stratigraphy. Both sites were deep and well within the lysocline. At the time Mio-Oligocene sediments were deposited, Site 1218 resided on 18 Myrold crust and was about 4.1 km deep. Site 1219 was on about 34 Myr-old crust and was about 4.5 km deep. There was a relative increase in the large diatoms near this boundary in the siliceous course suggesting fraction, increased productivity; however detailed, highresolution flux rates across this interval have yet to be determined. A well recovered section on the latest Oligocene equator, near the late Oligocene ridge crest, should provide both the resolution and the preservation required to better describe the changes in the equatorial ocean taking place at this time.

The Miocene (PEAT-7B)

The latest Oligocene through to the mid Miocene appears to have been a time of relative warmth comparable, at least in the oxygen isotope record, to the latest Eocene. However, the variability in the isotopic record of the early to mid Miocene seems a bit larger than that of the Eocene and may indicate more variability in climate and in global ice volume. The climatic "optimum" that marks the end of this interval, comes just before the major development of ice sheets on Antarctica and the marked increase in ice rafted debris in circum-Antarctic sediments.

Although the spectral character of the carbonate records and the variability of lower to mid Miocene and early Oligocene isotopic records appear comparable (Fig. 4), the mean value of the lower to mid Miocene record is about 1‰ lighter than that of the lower Differences between Oligocene. planktonic biota of these two intervals are also quite striking. In the Oligocene, evolutionary change appears to be minor and biostratigraphic zones span long intervals of time; while in the early to mid Miocene there is a burst of evolutionary vigor. What was it that caused this profound change in the nature of the ocean records? The only major ocean boundary change proposed for the time

near the Oligocene-Miocene boundary was the opening of the Drake passage to deep flow; however, there is some debate as to the exact timing of this event (Barker, 2001; Pagani et al., 1999), and its direct impact on the tropical ocean is uncertain. It may be that, as in the Eocene – Oligocene boundary section, the link lies in the shallow intermediate waters that provide nutrients to lower latitude upwelling regions.

For the equatorial region, an even more pertinent question is, what changes were occurring in the Miocene tropical ocean that led to this burst of Miocene evolution? Well recovered sections from these intervals near the ridge crest should give us the best preserved, most strategically located samples to answer questions concerning the changing environment of the near-surface equatorial ocean.

The Mid Miocene (PEAT-8B)

In principle, the age-transect strategy of this proposal would not be complete without data from the Plio-/Pleistocene. However, in addition to logistical reasons in terms of cruise length, near paleo-equatorial records have already been targeted by ODP Legs 138 (Pisias, Mayer, Janecek, et al., 1995) and 202, which provide information about the development of Northern Hemisphere

glaciation. Our last proposed site focuses instead on the interesting events following a Mid-Miocene maximum in deposition (van Andel et al., 1975). There is a wide latitude range of CaCO₃ deposition during the earliest Neogene, with a relatively sharp transition to a narrower CaCO₃ belt after 20 Ma (Lyle, 2003). CaCO₃ mass accumulation rates in the central equatorial Pacific recovered from the 18-19 Ma 'famine', and in the period between 14 and 16 Ma reached a second maximum in carbonate deposition, which is also evident in the seismic stratigraphy of the equatorial sediment bulge (Knappenberger, 2000). We designed PEAT-8B to recover a ridge-crest record at the early-mid Miocene sedimentation maximum on 16 Ma old crust. We plan to position the Miocene sites somewhat to the south of the estimated paleoequatorial position in order to maximize the time these sites remain within the equatorial zone, to allow for some error in positions (evidence suggests a southward bias of the equatorial sediment mound relative to the hotspot frame of reference; Knappenberger, 2000), and to place the interval of maximum interest above the basal hydrothermal sediments.

Tuning the Time Scale

The perfect geologic timescale has always been a sought-after goal. Determination of the rates at which earth processes take place and how these rates change are key not only to developing an earth history but also to understanding the nature of the processes themselves. Radiometrically constrained timescales were a great leap forward in this quest for a chronostratigraphy; however, there are a multitude of stratigraphic, geochemical and radiometric difficulties that constrain the accuracy and the applicability of radiometric techniques - not the least of which is the general increase in uncertainty of a radiometric date with increasing age. The paleomagnetic timescale has provided an important set of time markers for the Cenozoic; however, this approach is also far from perfect. Individual reversals are "dated" ultimately by radiometric techniques and by extrapolations that assume constant seafloor spreading rates. In addition, reversals are often widely spaced in time, and thus require some form of time interpolation between the reversals.

The discovery of orbitally driven variations in the earth's climate, and their preservation in the marine sedimentary record, has been the latest advance in our search for the perfect chronostratigraphy. The application of this technique is not

without assumptions of its own, including the predictability of the exact beat of the climatic metronome. However, if constrained by a paleomagnetic reversal stratigraphy, and if the same ages are obtained for geologic boundaries in different regions with different sensitivities to each of the different orbital forcings, we can develop a strong confidence in the time scale. A complete orbitally tuned Cenozoic timescale can provide estimates of process rates that have comparable precision throughout the Cenozoic.

We are working towards astronomically tuned Cenozoic time scale. As of now, there is an anchored astronomical time scale back to 30 Ma (mostly from ODP Leg 154), apart from an interval from 16-20 Ma that requires additional confirmation. Recent results from Leg 199 have resulted in an extension of the tuning across the Eocene/Oligocene boundary. Several other ODP Legs provide suitable data for the time interval 35-42 Ma (ODP Legs 171B and 177), 55-57 (ODP Leg 171B), 62 to 65 Ma (ODP Legs 165 and 171B). From this perspective, the current proposal would be useful to narrow existing gaps in our tuning efforts, particularly in the Eocene slice targeted by PEAT-3B. Additional drilling will be required for the time interval from 52 to

62 Ma, which is difficult to achieve in the Pacific due to the very shallow CCD, and will most likely have to be provided by data from the Atlantic.

Siting Strategy

In pursuing the history of the equatorial Pacific Ocean through both time-line and flow-line transects, we have two major advantages over the efforts that took place in the earlier days of scientific ocean drilling. Table 1 demonstrates that, although previous drill-sites have targeted the general area, they mostly do not fulfill all of our criteria in terms of (1) a sufficient number of holes to obtain a continuous record, (2) modern coring technology to obtain undisturbed sediments, (3) being located inside the paleo-equatorial zone, or (4) on the right crustal age, to ensure the presence of calcium carbonate at the targeted time slice.

We have developed the technical capability of recovering complete, relatively undisturbed sections, and we have established a seismic stratigraphy for the equatorial Pacific that allows us to map discrete intervals of time throughout the equatorial mound of biogenic sediments (Mayer et al, 1985; Knappenberger 2000; Moore et al, 2002, Mitchell et al., 2003). This later capability is key to two aspects of the present proposal. First, it facilitates the location of

the thickest part of the mound for each of the time intervals. This temporal mound "crest" we take to be the "geographic" equator for each time interval traced in the seismic records (Fig. 9).

Depth-dependent dissolution potentially complicates the interpretation of mound thickness as the "geographic" equator, and needs to be resolved through modeling as well as new data obtained through this proposal.

We commonly use locations of plate rotation poles, determined with paleomagnetic measurements, reconstruct site positions through most of the Cenozoic; however, there is always some error involved in this procedure. With the equatorial divergence zone being only a few degrees of latitude wide, the unquestioning acceptance of an equator determined by plate model rotations could lead to the placement of a site substantially off the geographic equator (Fig. 9). The divergence of surface waters at the equator and the resulting narrow zone of high productivity (at least since the beginning of the Oligocene) give an unambiguous locus of maximum sediment flux and characteristic sediment facies.

Van Andel et al. (1975) used this approach to establish a rotation pole for the Pacific plate that has compared well to subsequent refinements using other

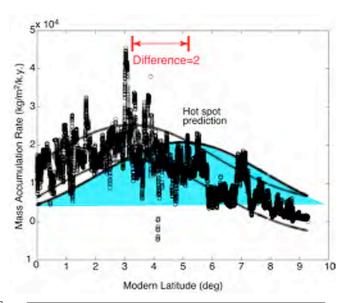


Figure 9: Estimated bulk sediment accumulation rates for the time interval 16.3 –20 Ma based on a seismic stratigraphic interpretation of seismic line recorded in an Pacific equatorial transect along 40 Ma crust, combined with estimates of sediment bulk density from drilling results (Knappenberger, 2000). Note the difference in latitudinal position of the maximum in flux rates and the estimated position of the equator based on a fixed hotspot plate rotation model.

techniques. However, van Andel et al. (1975) depended solely on drilled sections to establish the center of the sediment mound for different times in the past. We can use seismic records, tied to recovered sedimentary sections, to map out sediment thickness variation in both time and space, and to pinpoint the most desirable site location for each of the intervals we wish to sample.

Such an approach for locating our sites has the added advantage that it allows us to develop a map view of sediment thickness for each broad time interval selected, overcoming the problems associated with spatially varied deposition caused by bottom boundary layer processes. It also allows us to tie in previously recovered sections in the region and to study sediment fluxes in the tropical Pacific in three dimensions. The mapping approach will require additional survey data, as older data are mostly of poor quality, or not well located for this problem. However, this will be a prime objective of site survey cruises – existing Ventura, Ewing and other seismic data demonstrate that this mapping will be straightforward.

With the ability to map the time evolution of the sediment mound, all that is remaining in order to develop a agetransect is to cross check the seismic mapping results with independent estimates of the age of the ocean crust. We can establish the age of the crust both

through mapped crustal magnetic lineations (Petronotis, 1991; Petronotis et al., 1994; Müller et al., 1997) and through interpolation of basement ages between previously drilled sites (see figs. 1 and 3 in Mitchell, 1998). The mapping

approach should work well back to about 40 Ma. In the older Eocene sections, fixed hotspot plate rotation models are

questionable, and need be supplemented with coarser polar wander paths (Sager and Pringle, 1988). Furthermore, general circulation models of the near-surface tropical Eocene ocean (e.g., Huber, 2002: Fig. 10) leave us with a much more complicated view of the possible spatial pattern of tropical divergence and sedimentation than that seen in the younger Cenozoic. However, recent results from ODP Leg 199, when combined with results from older DSDP sites, give a comparatively simple mid Eocene patterns of siliceous sediment accumulation, even if the reconstructed latitudinal positions of the sites are open to question (Fig. 11). The Eocene sedimentation pattern does not look like one derived from the Neogene, however, and carbonate deposition appears to be displaced away from the equatorial

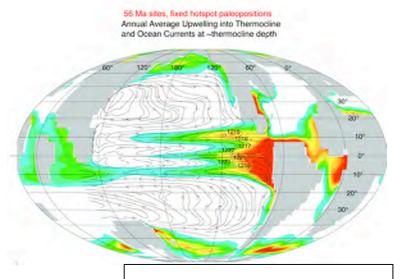
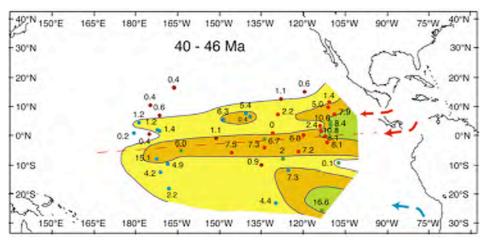


Figure 10: Modeled patterns of divergence in the 56 Ma oceans (Huber, 2002). Paleo-positions of Leg 199 sites shown for reference.



carbonate
deposition and
the interplay
between
carbonate
dissolution,
surface water
productivity,
and export of

Figure 11: from Moore et al. (2003, 2004 submitted). Average sediment accumulation for the interval 40-46 Ma, given in m/My. Site positions are backtracked to estimated position at 43 Ma. Site location filled circles are coloured according to sediment type for the time interval: blue = carbonate, green = siliceous carbonate, red = siliceous; brown = clay. A red dashed line approximate center of the marks the equatorial divergence (geographic equator). Contours are at 1, 5, and 10 m/My. The red dashed line indicates the approximate geographic paleo-equator based on the sediment archive - with a notable difference compared to the fixed hotspot based rotation.

region. Additional seismic coverage is needed to the west and east of the Leg 199 transect to better define the older Eocene sections and to more confidently locate regions of maximum thickness.

A strategy for understanding the interplay between the CCD, CaCO₃ dissolution, and productivity

The Pacific, specifically the equatorial upwelling zone, is the largest oceanic source of CO_2 to the atmosphere, and controls atmospheric CO_2 levels (Dore et al., 2003). The release and uptake of CO_2 is the direct consequence of calcium

biogenic carbonate from the surface the sediment waters Distinguishing between the effects of carbonate dissolution and productivity has been a field of intense study in the past. An important objective of this proposal is to address the detailed workings of depth dependent carbonate dissolution, which is intricately linked to the climate system and paleoceanography. In the classical model for carbonate dissolution, accumulation rates locally decrease linearly from a lysocline down to a carbonate compensation depth (CCD), reflecting linearly increasing rate of dissolution. The depth of both of these mapable surfaces is variable spatially and in time, and interacts with climatic and physical processes. The equatorial Pacific is one of the classical areas where the lysocline-CCD model was first developed, but there has been little subsequent effort to test it, a necessary step considering that the functional form of dissolution is now known to depend in

a more complex way on organic carbon burial and water mass properties. The age-transect will provide the necessary additional data with which to test the carbonate paradigm, and to recover previously unavailable carbonate material from important Paleogene time slices in the Pacific.

Specifically, the recovery of shallowly buried carbonate sediments from near the paleo-equatorial upwelling zone would contribute significantly towards separating the various processes that affect carbonate deposition preservation, and reduce some of the processes that affect climatic proxy records, such as diagenetic recrystallisation (Pearson et al., 2001). The Neogene productivity has been strongly oriented parallel to the equator, so differences in carbonate thicknesses at a common latitude but differing depth will permit the effect of dissolution to be isolated, following Lyle, 2003 and Mitchell, 2003. In addition, the strategy adopted in this proposal will provide new data throughout the Cenozoic with which it will be possible to map the spatial evolution of the equatorial CCD with time. This is so because the northward component of the Pacific plate movement results in the multiple recovery of the same time slice in different sites, but with a slightly different paleolatitude.

Recovering more detailed records from the best possible material will also allow a better understanding of physical processes that might affect or hinder our interpretation of carbonate proxy records, such as the "carbonate ion effect" – an observed and modeled influence of the carbonate ion concentration on stable isotope fractionation in carbonate (Spiro et al., 1997; Zeebe and Wolf-Gladrow, 2001).

Preliminary work with Ewing seismic data (Mitchell et al., 2003) has revealed a surprising lack of dissolution with depth in the westerly region of this study area. Our aim is to develop a more extensive 3D model for the stratigraphy of the equatorial Pacific deposits that links all existing core data using a grid of highresolution seismic reflection profiles. The numerical stratigraphic model will then be used to assess carbonate dissolution, and in particular the spatial pattern of sharp changes in dissolution such as the extremely abrupt change in the CCD at the Oligocene/Eocene boundary, which has been linked to a possible abrupt onset of continental weathering. The sediment archive recovered in this proposal will allow the application the substantial array of carbonate-based proxies with which the wider regional seismic study can be ground-truthed.

Reconstructing paleoceanographic properties and SST

A large number of paleoceanographic interpretations rely on obtaining proxy data (stable isotope measurements, elemental ratios such as Mg/Ca, seasurface temperature estimates from faunal distributions and isotope data, Alkenone proxies, geochemical productivity and burial indicators, etc.). In turn, a very large number of these measurements rely on the presence of biogenic calcium carbonate. For the Pacific, the drilling strategy we propose here is conceptually the best approach to recover this important material with the best possible preservation, and the least amount of diagenetic effects for a long intervals throughout the Cenozoic, and will thus contribute to the objectives of the IODP Extreme Climates Initiative.

Spatial range considerations

In order to recover the best preserved and most complete carbonate record from the paleo-equatorial Pacific, the "agetransect" siting strategy necessarily implies a more restricted N-S transect, even though the northward movement of the Pacific Plate does allow us to recover identical time slices multiple times at different paleo-latitudes (separated by several degrees). However, we note that the regional seismic study to be developed as part of our site-survey work would give

us the opportunity to begin integrating data from older drill sites with the new drilling, both the IODP-626 proposed drilling and that from ODP Legs 199 and 138. Addition of new data from existing drill sites will flesh out the age transect so that a more complete model of the evolution of the equatorial Pacific can be developed.

We are proposing to measure changes in biogenic sedimentation over a critical interval, to understand an impressive but poorly documented change in C-org deposition and a turnover in the biogenic silica dynamics in the Pacific. These changes coincide with a minimum in the deposition of carbonate that needs better documentation (Lyle, 2003).

Paleomagnetic objectives

One important aspect of this proposal is recovery the of high-quality paleomagnetic data. Results from ODP Leg 199 demonstrate that high-quality data can be recovered from nearequatorial carbonate. Leg 199 succeeded in recovering almost all magnetic reversals from the Paleogene through to the present. However, Leg 199 did not recover biogenic carbonate sediments through most of the Eocene, nor for ages younger than the lower Miocene. Thus, while the paleomagnetic record during these times was of high quality, global

correlation is hindered by the lower mass accumulation rate, the absence of a detailed isotope stratigraphy, and sparser biostratigraphic control. In order to facilitate the development of an integrated magneto- and biostratigraphic framework with a stable isotope stratigraphy (necessary to enable global correlation), a recovery of magnetic reversals within carbonate sediment is desirable. This pre-requisite contributed to the strategy described in this proposal. In addition, further detailed magnetostratigrapic data, particularly from the Eocene, will help to resolve the suggestion that the geographic equator, as determined from the biogenic sediment bulge, might not coincide with the paleoequator position backtracked with a fixedhotspot reference frame (Moore et al., 2003, 2004 (submitted); Tarduno, 2003).

Ancillary benefits (MORB, basement)

Our proposed drilling aims to recover basement samples at all sites. A transect of MORB samples from a fixed location in the absolute mantle reference frame would be a unique sample suite and, while not one of the primary objectives of this proposal, should be of strong potential interest to mantle geochemists. In addition, a transect of basalt samples along the flow-line that have been erupted in similar formation-water environments

should be of interest for low-temperature alteration studies (see, e.g., Elderfield & Schultz, 1996).

<u>Possible problems that could be</u> encountered: cherts, dolomite

While not generally a problem for Neogene sites in the Pacific, drilling and recovery difficulties related to chert layers are a necessary consideration when drilling Paleogene sediments. While early DSDP coring was generally defeated by chert, better cutting shoes and active heave compensation have helped in the management of chert layers during drilling.

Previous legs have shown the general occurrence of cherts in certain parts of the Eocene Pacific. Prior to ODP Legs 198 and 199, the general belief was that sediment that has never been deeply buried, like the sites proposed here, are likely to contain only oozes. However, Leg 199 drilling results indicate that cherts are present even with shallow burial, particularly at the boundary between the lower and middle Eocene. These sediments were still unlithified, as radiolarian ooze was recovered right up to the upper boundary of chert zones, with radiolarian ooze and carbonates below.

Understanding where chert zones lie will help to minimize sediment loss. Data from Leg 199 suggest that chertified

intervals were deposited slowly. While it is unlikely that the sites in this proposal will avoid chert altogether, the specific planning of this proposal aims for calcium carbonate (with higher sedimentation rates than chert-bearing radiolarian ooze), as well as a lower abundance of silica rich formations. Thus, we believe that the chert problem can be overcome.

For this proposal, we suggest the technical preparation of half-length APC facilities, as this would overcome some of the sediment loss problems encountered during Leg 199 and other equatorial Pacific cruises, in case chert is encountered. The sites proposed here that could be affected by chert horizons include PEAT-1B through 4B.

Another possible problem worthy of consideration is that several of the Leg 199 sites showed abundance of dolomitization near the basement of recovered holes (early Paleogene age). To avoid such sediment in our time slices, we have chosen basement ages that are slightly older than the target age. This procedure necessarily involves a trade-off between maximum recovery of target carbonate, and avoiding dolomitized sediments and compensating for potential age uncertainties (Fig. 5).

Planning the Site Survey

The exact determination of drill site locations should be accomplished through seismic survey during a site survey cruise. We can employ our current knowledge of sea-floor age, plate rotation models, and the history of the CCD to plan this survey. We note that proposals have now been submitted to accomplish site-survey work: a proposal led by Lyle to the US NSF, and a proposal led by Mitchell to the UK IODP. The proposed track for the site survey is shown in Fig. 12.

The determination of possible areas for cruise surveys depends on three types of information: 1) a map of sea-floor age as determined from magnetic lineations and previous drilling results, 2) plate rotation models that allow us to predict which locations once were underneath the paleoequator, and 3) a detailed history of the CCD. The reconstruction of the CCD was pioneered by van Andel (1975), and can now be supplemented with additional results from recent drilling.

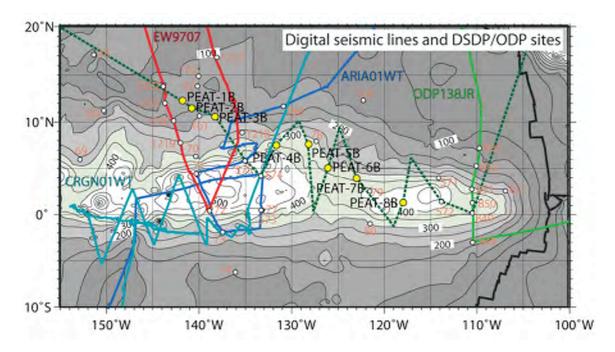
Fig. 5 illustrates the resulting CCD history for the Pacific. Importantly, the early Cenozoic time interval coincides with a very shallow CCD, which will make the detailed planning of future drill-sites a more crucial and important task. In addition, a shallow CCD implies that one can obtain calcareous rich sediments over much shorter time intervals.

The reconstructed history of the CCD can then be used as a guide as to how time intervals of particular interest (Fig. 4) can be targeted. This approach is shown in figure 5, where we have plotted "ideal" crustal ages to drill each of the 12 time intervals of interest while remaining above the CCD. Colored intervals correspond to those in figure 6.

In order to implement our age transect approach, the crustal ages shown in Fig. 5 have to be translated to specific locations on today's ocean floor. In particular, for each crustal age shown in figure 5, we attempt to locate those sites that were positioned at the paleo-equator during the time interval of interest.

Our methodology is to use the digital agegrid of sea-floor age from Müller et al., 1997, modified and improved with additional magnetic anomaly picks from Petronotis (1991, 1994) and DSDP/ODP basement ages. For this grid, each point is then backrotated in time to zero age, using the fixed-hotspot stage-poles from Koppers et al. (2001) and Engebretson et al. (1985), and the paleo-pole data from Sager and Pringle (1988). From the backtracked latitudes for each grid point we then obtained the paleo-equator at the crustal age by contouring.

Figure 12: Proposed site survey trackline (dashed), with drillsite locations and available digital seismic reflection data from previous cruises. The new trackline is designed to locate drillsites, and to fill in gaps in the coverage of the older surveys. Light green: Venture-1; blue: Ariadne-1; blue-green: Crossgrain-1; red: Ewing EW9707. Only proposed Site PEAT-4B is in the vicinity of an existing trackline (CROSSGRAIN), but due to technical problems no records were collected for ~10 hours as the Site location was approached.



The results are shown in Fig. 6, together with proposed locations for site-survey work for this proposal, which take into account the plate-rotation with respect to the paleo-equator, moved slightly towards the south to accommodate a potential error in the fixed hotspot rotation model (Moore et al., 2003; 2004 submitted). The model presented in Fig. 6 is partly corroborated by comparison with previous drilling results. For example, DSDP Site 78, near proposed survey area PEAT-5, has a basement age as predicted.

Presently we have two excellent seismic transects across the Pacific equatorial mound, plus a few other older seismic lines that provide useful ties (Fig. 12). To accurately locate the sites that sample all the intervals proposed here, further seismic survey is needed. However, based on the simple plate rotation models and on the patterns of sedimentation derived from seismic transects (Fig. 6) and from drilling results (Fig. 8) we can focus on the survey areas for the Oligocene and Miocene intervals by interpolation between Venture and EW9709 track lines (Fig. 12). For the early Pliocene interval existing seismic lines may give us information to tentatively identify a site location for further detailed survey. For the Eocene intervals a comprehensive interpretation of existing seismic lines and drilling data will guide us in a more complete survey of the 34 –54 Ma crust. One of the proposed sites (PEAT-5B) is close to previous drilling sites (DSDP-78).

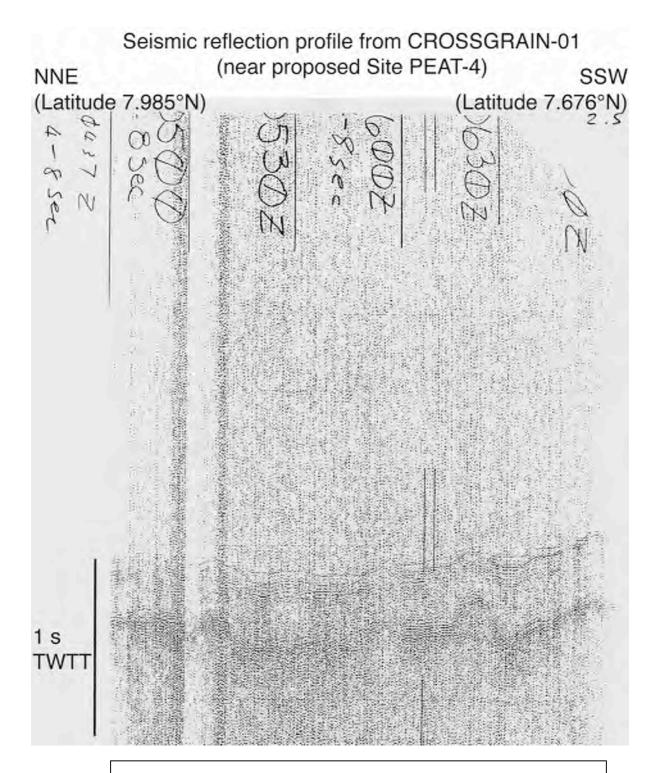


Figure 13: Example of seismic reflection profile from CROSSGRAIN-1 survey (NGDC #15040172, T. Washington, Scripps, 1987) very close to PEAT-4 (Figure 12). Unfortunately the seismic data were collected at transit speed, resulting in a low-quality survey. However, this section establishes that the sediment cover's thickness is near our prediction (~0.3-0.4 s TWT, ~200-300m). We cannot yet distinguish any features about the sediment section, nor can we optimize the location based upon existing data. Exact positions of proposed sites will have to await additional proposed site survey work.

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Site Description Forms:

Site description forms for the 8 proposed drillsites are attached here. Because of the poor quality of most of the existing seismic records, we have a good indication of sediment thickness at only five of the sites, based on previous DSDP and ODP drilling. The nature of the existing records is demonstrated by a seismic line near Site PEAT-4B (Figure 13).

In the absence of data, we have estimated the sediment thickness on the basis of our understanding of the regional depositional patterns. Water depths for the sites are range from ~4 to 5 km. At each site we would like to plan on 3 APC/XCB cores through the entire section, and logging as appropriate, with particular emphasis on obtaining FMS logs for stratigraphic correlation.

Because there is no modern site survey data, we have not submitted forms 2, 3, 4, or 5 for each site. Form 2 would be blank in the absence of useful data; we anticipate a standard logging program which will be outlined in detail with the logging contractor after site surveys are completed. There has never been a safety problem at open-ocean pelagic sites, and the lithologic summary information must also await a site survey and analysis of the accompanying piston cores.

Clearly a site-survey cruise will be required before this can become a fully mature drilling proposal, and in that eventuality we will provide a complete and accurate description of each proposed drill site. We note that proposals have now been submitted to accomplish sitesurvey work: a proposal led by Lyle to the US NSF, and a proposal led by Mitchell to the UK IODP.

iSAS/IODP Site Summary Forms:

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

[<u>New</u>]

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form Submitted:	24th Mar. 2004
Site Specific Objectives with	Obtain calcareous sediments from the region of the equatorial circulation system during maximum Cenozoic warmth (52-54 Ma). Constrain the isotopic and biotic characteristics of the near surface equatorial waters. Gain information about the nature of very shallow CCD at this time.
objectives in proposal)	
List Previous Drilling in Area:	ODP 1221, 1222; DSDP 42, 162

Section B: General Site Information

Site Name: (e.g. SWPAC-01A)	PEAT-1B	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #	Area or Location:	Eastern equatorial Pacific
Latitude:	Deg: 12°N	Min: 18	Jurisdiction:	International
Longitude:	Deg: 141°W	Min: 48	Distance to Land:	1621 km (Hawaii)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4800-5000 m

Section C: Operational Information

	Sediments B							Bas	asement		
Proposed Penetration:	200m	200m 5m									
(m)	What is the total	sed. thi	ickness? ~200m	n	1						
								Penetrati	on: 205	m	
General Lithologies:	Pelagic cl ooze, chalk	•	siliceous oo chert.	ze,	calc	areous	basalt				
Coring Plan: (Specify or check)	1-2-3-APC, XCB a	as need	ed								
Wireline Logging Plan:	Standard 7	Tools			Spec	ial Tool	s		LWD		
r iaii.	Neutron-Porosity		Borehole Tele	eviewei	· 🔲 📗	Formatio	n Fluid Samp	ling 🗌	Density-Neutron		
	Litho-Density		Nuclear Magne Resonance	etic		Borehole 7	Temperature		Resistivity-Gamma R	ау 🗆	
	Gamma Ray		Geochemical			Borehole S	Seismic		Acoustic		
	Resistivity		Side-Wall Core Sampling	÷							
	Acoustic										
	Formation Image					Others ()		Others ()		
Max.Borehole Temp.:	Expected va	lue (For Riser Dri	lling))						
Mud Logging:	Cuttings Sampling Intervals										
(Riser Holes Only)	fro	m	m	to			m,		m intervals		
	fro	m	m	to			m,		m intervals		
		Ва	sic Sampling	Inte	rval.	s: 5m					
Estimated days:	Drilling/Coring:		Logging				Tot	tal On-S	ite:5		
Future Plan:	Longterm	Bore	ehole Observ	atio	n Pl	an/Re-	entry Pla	an			
Hazards/	Please che	ck fo	llowing List	of P	oter	ntial Ha	azards		What is your Weat		
Weather:	Shallow Gas		omplicated Seabed Co.			rothermal A			window? (Prefera period with the reas		
	Hydrocarbon	□ s	oft Seabed		Lands	lide and Turl	bidity Current			,	
	Shallow Water Flow		Currents		Metha	nne Hydrate					
	Abnormal Pressure	□ F	ractured Zone		Diapir	and Mud Vo	olcano				
	Man-made Objects	□ F	ault		High 7	Temperature					
	H ₂ S	□ H	ligh Dip Angle		Ice Co	onditions					
	CO ₂										

iSAS/IODP Site Summary Forms:

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

<u>New</u>

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Submitted:	24th Mar. 2004 Obtain calcareous sediments from the region of the equatorial circulation system during maximum Cenozoic warmth (50-52 Ma). Constrain the isotopic and biotic
Objectives with	characteristics of the near surface equatorial waters. Gain information about the nature of very shallow CCD at this time.
List Previous Drilling in Area:	DSDP 42, DSDP 162, ODP 1221

Section B: General Site Information

onerar one miro	munon		
PEAT-2B	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #	Area or Location:	Eastern equatorial Pacific
Deg: 11°N	Min: 30	Jurisdiction:	International
Deg: 140°W	Min: 48	Distance to Land:	1761 km (Hawaii)
WGS 84			
Primary: X	Alt:	Water Depth:	4800-5000 m
	PEAT-2B Deg: 11°N Deg: 140°W WGS 84	PEAT-2B If site is a reoccupation of an old DSDP/ODP Site, Please include former Site # Deg: 11°N Min: 30 Deg: 140°W Min: 48 WGS 84	of an old DSDP/ODP Site, Please include former Site # Deg: 11°N Min: 30 Jurisdiction: Deg: 140°W Min: 48 Distance to Land:

Section C: Operational Information

	Sediments						Basement				
Proposed	200m 5m										
Penetration: (m)	What is the total sed. thickness? ~200m m										
· /							Total I	Penetrati	on: 205	m	
General Lithologies:	Pelagic cl	ay, si	iliceous oo	ze, o	calca	areous	basalt				
	ooze, chalk	and o	chert.								
Coring Plan:	1-2-3-APC, XCB	as neede	4								
(Specify or check)	1 2 3 m e, neb	as neede									
Wireline Logging	Standard 7	Гools		;	Speci	ial Tools	S		LV	WD	
Plan:	Neutron-Porosity	, <u> </u>	Borehole Tele	eviewer		Formation	n Fluid Samp	oling 🔲	Density-Neu	tron	
	Litho-Density		Nuclear Magne Resonance	etic	ш	Borehole T & Pressure	Temperature		Resistivity-Ga	ımma Ray 🗆	
	Gamma Ray		Geochemical			Borehole S	Seismic		Acoustic		
	Resistivity		Side-Wall Core Sampling	e							
	Acoustic		Jumpung								
	Formation Image				(Others ()		Others ()	
Max.Borehole Temp.:	Expected va	lue (F	or Riser Dri	lling)							
Mud Logging:	Cuttings Sampling Intervals										
(Riser Holes Only)	fro	m	m	to			m,		m inte	ervals	
	fro	m	m	to			m,		m inte	ervals	
		Bas	ic Sampling	Inter	rvals	:: 5m					
Estimated days:	Drilling/Coring:		Logging				To	tal On-S	ite:5		
Future Plan:	Longterm	Rorel	hole Observ	vatio	n Pla	an/Re-	entry Pl	an			
	Longierm	Боге		alioi			citery is				
TT 1/											
Hazards/ Weather:			lowing List						What is you window? (F		
,, 54,11511	Shallow Gas	Cor	nplicated Seabed Co	ndition	Hydro	othermal Ac	ctivity		period with the		
	Hydrocarbon	So	ft Seabed		Landsli	ide and Turt	oidity Current				
	Shallow Water Flow	☐ Cu	rrents		Methan	ne Hydrate					
	Abnormal Pressure	☐ Fra	actured Zone		Diapir a	and Mud Vo	olcano				
	Man-made Objects	☐ Fa	ult		High Te	emperature					
	H ₂ S	Hi	gh Dip Angle		Ice Con	nditions					
	CO ₂										

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

<u>New</u>

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Submitted:	24th Mar. 2004 Obtain calcareous sediments from the middle and late Eocene as part of the
Objectives with	equatorial age transect to investigate reason for 2-3x increase in accumulation rates of siliceous ooze within the middle Eocene (41 to 45 Ma). Evaluate the temperature and structure of the near-surface ocean. Investigate the nature of a strong and rapid change of CCD, identified during ODP Leg 199.
List Previous Drilling in Area:	DSDP 161, DSDP 42, ODP 1221, ODP 1220

-				
Site Name: (e.g. SWPAC-01A)	PEAT-3B	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #	Area or Location:	Eastern equatorial Pacific
Latitude:	Deg: 10°N	Min: 36	Jurisdiction:	International
Longitude:	Deg: 138°W	Min: 18.0	Distance to Land:	2046 km (Hawaii)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4800-5000 m

			Sediments						Basement				
Proposed Penetration: (m)	200m						5n	n					
	What is the total	sed. t	hickness?	~200m	n	1							
						_			tal Penetra	tion: 205		m	
General Lithologies:	Pelagic cl ooze, chalk	-			ze,	calc	areous	basa	lt				
Coring Plan: (Specify or check)	1-2-3-APC, XCB	as nee	eded										
Wireline Logging Plan:	Standard 7	Γool	s			Spec	cial Tool	s		L	WD		
Pian:	Neutron-Porosity		Bo	rehole Tele	eviewei	r 🔲	Formatio	n Fluid S	ampling [Density-Ne	utron		
	Litho-Density			lear Magnonance	etic		Borehole 7		ure	Resistivity-C	Samma R	ау 🗆	
	Gamma Ray		Geo	chemical			Borehole S	Seismic		Acoustic			
	Resistivity		ı	-Wall Core	e								
	Acoustic		_										
	Formation Image						Others ()		Others ()		
Max.Borehole Temp.:	Expected va	lue	(For Ri	ser Dri	lling))							
Mud Logging:	Cuttings Sampling Intervals												
(Riser Holes Only)	fro	m		m	to			m,		m int	tervals		
	fro	m		m	to			m,		m int	tervals		
			asic Sa				<u>s: 5m</u>	,					
Estimated days:	Drilling/Coring:	5		Logging	:1				Total On-	Site:6			
Future Plan:	Longterm	Boı	rehole (Observ	vatio	n Pl	lan/Re-	entry	Plan				
Hazards/	Please che	ck f	ollowir	ng List	of P	otei	ntial Ha	azards	s	What is yo	ur Weati	her	
Weather:	Shallow Gas	_	Complicated	_			lrothermal A		_	window? (
		Ш			Ш					period with	the reas	ons)	
	Hydrocarbon		Soft Seabed	i		Lands	slide and Tur	bidity Curr	rent 🗌				
	Shallow Water Flow		Currents			Metha	ane Hydrate						
	Abnormal Pressure		Fractured Z	Cone		Diapi	r and Mud Vo	olcano					
	Man-made Objects		Fault			High '	Temperature						
	H ₂ S		High Dip A	ingle		Ice Co	onditions						
	CO ₂												

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

<u>New</u>

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form Submitted:	24th Mar. 2004
Site Specific Objectives with	Eocene/Oligocene Boundary: obtain well recovered and well preserved equatorial sections across this relatively short transition between warm and cold global climates to determine the impact of high latitude ocean boundary changes on climate, circulation, and productivity in the equatorial region.
List Previous Drilling in Area:	DSDP 160, DSDP 575, ODP 1218

beenon b. o.	onerar proc miro	illiation		
Site Name: (e.g. SWPAC-01A)	РЕАТ-4В	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #		Eastern equatorial Pacific
Latitude:	Deg: 7°N	Min: 31.4	Jurisdiction:	International
Longitude:	Deg: 131°W	Min: 39.9	Distance to Land:	2616 km (Clipperton Island)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4800-5000 m

		Se	ediments				Basement				
Proposed Penetration:	300m					5n	n				
(m)	What is the total	sed. thick	kness? ~300m	n	1						
							Total F	enetrati	on: 305	m	
General Lithologies:	Pelagic cl ooze, chalk	•	liceous oo chert.	ze,	calca	reous	basalt				
Coring Plan: (Specify or check)	1-2-3-APC, XCB a	ns needed	I								
Wireline Logging Plan:	Standard 7	Tools			Specia	al Tool	S				
i ian.	Neutron-Porosity		Borehole Tele	eviewei		Formatio	n Fluid Samp	ling 🔲	Density-Neutron		
	Litho-Density		Nuclear Magne Resonance	etic		orehole T Pressure	Temperature		Resistivity-Gamma	Ray 🗆	
	Gamma Ray		Geochemical		\Box	orehole S	Seismic		Acoustic		
	Resistivity		Side-Wall Core Sampling	e							
	Acoustic										
	Formation Image				C	Others ()		Others ()		
Max.Borehole Temp. :	Expected va	lue (F	or Riser Dri °C	lling)							
Mud Logging:	Cuttings Sampling Intervals										
(Riser Holes Only)	froi	n	m	to			m,		m interval	S	
	froi	n	m	to			m,		m interval	s	
	1101		ic Sampling		rvals	· 5m	111,		111 111002 7 442	2	
Estimated days:	Drilling/Coring:		Logging				Tot	al On-S	Site:6		
Future Plan:	Longterm	Borel	nole Observ	vatio	n Pla	n/Re-	entry Pla	an			
Hazards/	Please chee	ck fol	lowing List	of P	otent	tial Ha	azards		What is your We		
Weather:	Shallow Gas		nplicated Seabed Co			thermal Ac			window? (Preference period with the re		
	Hydrocarbon	Sof	t Seabed		Landslie	de and Turt	bidity Current				
	Shallow Water Flow	Cui	rrents		Methane	e Hydrate					
	Abnormal Pressure	Fra	ctured Zone		Diapir a	nd Mud Vo	olcano				
	Man-made Objects	☐ Fau	ılt		High Te	mperature					
	H ₂ S	Hig	gh Dip Angle		Ice Con	ditions					
	CO ₂										

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

[<u>New</u>]

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form Submitted:	24th Mar. 2004
Site Specific Objectives with Priority (Must include general objectives in proposal)	Oligocene: obtain well recovered and well preserved equatorial sections during a time of relatively deep CCD to extend and supplement existing astronomical time calibrations. Investigate how a "one cold pole" world might impact the interhemispheric temperature imbalance on pole to equator temperature gradients and on the symmetry of the global wind systems. Gain new information on the position of trade winds, the position of the Inter Tropical Convergence Zone, and the seasonal shifts in this zone, as reflected by the wind-driven currents of the equatorial region.
List Previous Drilling in Area:	DSDP 78, DSDP 159, DSDP 160, DSDP574

Site Name: (e.g. SWPAC-01A)	PEAT-5B	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #	Area or Location:	Eastern equatorial Pacific
Latitude:	Deg: 7°N	Min: 38.4	Jurisdiction:	International
Longitude:	Deg: 128°W	Min: 11.1	Distance to Land:	2235 km (Clipperton Island)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4500-4600 m

		S	ediments				Basement					
Proposed	320m					5n	n					
Penetration: (m)	What is the total sed. thickness? ~320m m											
, ,							Total I	Penetrati	on: 325		m	
General Lithologies:	Pelagic cl	ay, s	iliceous oc	oze,	calc	areous	basalt					
	ooze, chalk	and	chert.									
Carina Dlan	1-2-3-APC, XCB	as neede	·d									
Coring Plan: (Specify or check)	1 2 3 Ai C, ACD	as neede	, u									
Wireline Logging	Standard 7	Гools			Spec	ial Tool	S					
Plan:	Neutron-Porosity	,	Borehole Te	leviewe		Formatio	n Fluid Samp	oling 🔲	Density-Ne	utron		
	Litho-Density		Nuclear Magr Resonance	netic	ш	Borehole 7	Temperature		Resistivity-C	Samma Ra	у 🗆	
	Gamma Ray		Geochemical			Borehole S	Seismic		Acoustic			
	Resistivity		Side-Wall Con Sampling	re								
	Acoustic											
	Formation Image					Others ()		Others ()		
Max.Borehole Temp.:	Expected va	lue (I	For Riser Dr	illing)								
Mud Logging:	Cuttings Sampling Intervals											
(Riser Holes Only)	fro	m	m	to			m,		m in	tervals		
	fro	m	m	to			m,		m in	tervals		
		Bas	sic Sampling	z Inte	rvals	s: 5m						
Estimated days:	Drilling/Coring:	5.5	Logging	g:0.5			То	tal On-S	ite:6.0			
Future Plan:	Longterm	Bore	hole Obser	vatio	n Pl	an/Re-	entry Pl	an				
	2.01.8.01.11					,	<i>y</i>					
II/												
Hazards/ Weather:			llowing Lis						What is yo window? (
,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	Shallow Gas		mplicated Seabed C	ondition	Hydi	rothermal A	ctivity		period with			
	Hydrocarbon	Sc	oft Seabed		Landsl	lide and Tur	bidity Current					
	Shallow Water Flow	☐ Cu	urrents		Metha	ne Hydrate						
	Abnormal Pressure	☐ Fr	ractured Zone		Diapir	and Mud Vo	olcano					
	Man-made Objects	☐ Fa	nult		High T	Temperature						
	H ₂ S	☐ Hi	igh Dip Angle		Ice Co	onditions						
	CO ₂											

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002



Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form Submitted:	24th Mar. 2004
Site Specific Objectives with Priority	Latest Oligocene: recover sections on the latest Oligocene equator to investigate the major changes in the equatorial ocean taking place around the Oligocene/Miocene boundary, including a possible major glaciation and/or major changes in ice volume. Provide the least possible dissolved material to improve bio- and magnetostratigraphy of this critical time interval.
List Previous Drilling in Area:	DSDP 78, DSDP 574, DSDP 79

Site Name: (e.g. SWPAC-01A)	PEAT-6B	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #		Eastern equatorial Pacific
Latitude:	Deg: 5°N	Min: 5.2	Jurisdiction:	International
Longitude:	Deg: 126°W	Min: 5.8	Distance to Land:	2068 km (Clipperton Island)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4300-4500 m

		Se	ediments				Basement				
Proposed Penetration:	300m					5m	1				
(m)	What is the total	sed. thic	kness? ~300m	n	1						
								Penetrat	ion: 305	m	
General Lithologies:	Pelagic clay, siliceous ooze, calcareous bas ooze, chalk and chert.										
Coring Plan: (Specify or check)	1-2-3-APC, XCB	as neede	d								
Wireline Logging Plan:	Standard 7	Tools			Special	Tools	3				
I laii.	Neutron-Porosity	_	Borehole Tel	eviewe	· 🗌 Fo	ormation	n Fluid Samp	oling 🔲	Density-Neutron		
	Litho-Density		Nuclear Magn Resonance	etic		rehole T Pressure	emperature		Resistivity-Gamma	Ray 🗆	
	Gamma Ray		Geochemical		Boı	rehole S	eismic		Acoustic		
	Resistivity		Side-Wall Cor Sampling	re							
	Acoustic										
	Formation Image				Oth	ners ()		Others ()		
Max.Borehole Temp.:	Expected va	lue (F	or Riser Dri	illing)	•						
Mud Logging:	Cuttings San	npling	Intervals								
(Riser Holes Only)	fro	m	m	to			m,		m interval	S	
	fro	m	m	to			m,		m interval	s	
		Bas	ric Sampling	Inte	rvals: .	5m					
Estimated days:	Drilling/Coring:		Logging				То	tal On-S	Site:6.5		
Future Plan:	Longterm	Bore	hole Obser	vatio	n Plan	/Re-	entry Pl	an			
Hazards/	Please che	ck fol	lowing List	t of P	otenti	al Ha	zards		What is your Wee		
Weather:	Shallow Gas		mplicated Seabed Co		Hydrothe				window? (Prefer period with the re-		
	Hydrocarbon		ft Seabed		Landslide	and Turb	idity Current		1	,	
	Shallow Water Flow	Cu	ırrents		Methane F	Hydrate					
	Abnormal Pressure	Fra	actured Zone		Diapir and	Mud Vo	lcano				
	Man-made Objects	☐ Fa	ult		High Temp	perature					
	H ₂ S	Hi	gh Dip Angle		Ice Condit	ions					
	CO ₂										

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

<u>New</u>

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form Submitted:	24th Mar. 2004
Site Specific	Miocene: obtain well preserved sections to investigate what led to the large burst of evolutionary change during this time in equatorial regions. Obtain material to further study the larger climatic variability compared to Eocene.
List Previous Drilling in Area:	DSDP 78, DSDP 79

beenon b. o.	onerar proc mire.			
Site Name: (e.g. SWPAC-01A)	РЕАТ-7В	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #		Eastern equatorial Pacific
Latitude:	Deg: 3°N	Min: 59.1	Jurisdiction:	International
Longitude:	Deg: 123°W	Min: 1.3	Distance to Land:	1787 km (Clipperton Island)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	4400-4600 m

	Sediments						Basement					
Proposed	300m					51	n					
Penetration: (m)	What is the total sed. thickness? ~300m m											
, ,	TY MAK 15 the count seek allements in 2 2 2 2 2						Total	Penetrati	ion: 305		m	
General Lithologies:	Pelagic cl	ay, s	siliceous oc	ze,	calc	areous	basalt					
	ooze, chalk	ζ.										
Coring Plan:	1-2-3-APC, XCB	as neede	ed.									
(Specify or check)	1 2 3 m e, neb	as neede										
Wireline Logging	Standard 7	Tools			Spec	ial Too	ls					
Plan:	Neutron-Porosity		Borehole Tel	eviewei	. 🗆	Formatio	on Fluid Sam	oling 🔲	Density-Ne	utron		
	Litho-Density		Nuclear Magn Resonance	etic	ш	Borehole & Pressur	Temperature e		Resistivity-C	amma Ra	у 🗆	
	Gamma Ray		Geochemical			Borehole	Seismic		Acoustic			
	Resistivity		Side-Wall Cor Sampling	e								
	Acoustic		2									
	Formation Image					Others ()		Others ()		
Max.Borehole Temp.:	Expected va	lue (l	For Riser Dri	lling)								
Mud Logging:	Cuttings Sampling Intervals											
(Riser Holes Only)	fro	m	m	to			m,		m int	ervals		
	fro	m	m	to			m,		m int	ervals		
	Basic Sampling Intervals: 5m											
Estimated days:	Drilling/Coring:		Logging				To	tal On-S	Site:6.5			
Future Plan:	Longterm Borehole Observation Plan/Re-entry Plan											
	Longierm	2010				, 110	circi y 1 i					
II/								1				
Hazards/ Weather:			llowing List						What is yo window? (
,, 54,11511	Shallow Gas		omplicated Seabed Co	ondition	Hyd	rothermal A	ctivity		period with			
	Hydrocarbon	□ So	oft Seabed		Lands	lide and Tu	rbidity Current					
	Shallow Water Flow	□ C	urrents		Metha	nne Hydrate						
	Abnormal Pressure	☐ Fi	ractured Zone		Diapir	r and Mud V	/olcano					
	Man-made Objects	☐ Fa	ault		High 7	Temperature	:					
	H ₂ S	П Н	igh Dip Angle		Ice Co	onditions						
	CO ₂											
	2											

Form 1 - General Site Information

Please fill out information in all gray boxes Revised 7 March 2002

<u>New</u>

Section A: Proposal Information

Title of Proposal:	Cenozoic Pacific Equatorial Age Transect
Date Form	24th Mar. 2004
Submitted:	
	Middle Miocene: To recover a ridgecrest calcium carbonate record of the early-
Site Specific	mid Miocene sedimentation maximum on 16 Ma old crust, spanning 16 to 12 Ma.
Objectives with	
Priority	
(Must include general	
objectives in proposal)	
	DSDP 79, 571
List Previous	
Drilling in Area:	

beenon b. o.	onerar proc miro	illiation		
Site Name: (e.g. SWPAC-01A)	РЕАТ-8В	If site is a reoccupation of an old DSDP/ODP Site, Please include former Site #	Area or Location:	Eastern equatorial Pacific
Latitude:	Deg: 1°N	Min: 19	Jurisdiction:	International
Longitude:	Deg: 117°W	Min: 58.5	Distance to Land:	1462 km (Clipperton Island)
Coordinates System:	WGS 84			
Priority of Site:	Primary: X	Alt:	Water Depth:	3900-4100 m

	Sediments						Basement					
Proposed Penetration:	350m 5m											
(m)	What is the total sed. thickness? ~350m m											
								enetrati	on: 355		m	
General Lithologies:	Pelagic cl ooze, chalk	•	siliceous oo	ze,	calc	areous	basalt					
Coring Plan: (Specify or check)	1-2-3-APC, XCB	as need	led									
Wireline Logging Plan:	Standard 7	Tools			Spec	ial Tool	S					
i ian.	Neutron-Porosity		Borehole Tele	eviewe		Formation	n Fluid Samp	ling 🔲	Density-Neutron			
	Litho-Density		Nuclear Magne Resonance	etic		Borehole T & Pressure	emperature		Resistivity-Ga	mma Ra	ау 🗆	
	Gamma Ray		Geochemical			Borehole S	Seismic		Acoustic			
	Resistivity		Side-Wall Core Sampling	e								
	Acoustic											
	Formation Image					Others ()		Others ()		
Max.Borehole Temp.:	Expected va	lue (For Riser Dri	lling)							
Mud Logging:	Cuttings Sampling Intervals											
(Riser Holes Only)	fro	m	m	to			m,		m inte	rvals		
	fro	m	m	to			m,		m inte	rvals		
		Вс	isic Sampling	Inte	rvals	s: 5m						
Estimated days:	Drilling/Coring:		Logging				Total On-Site:7					
Future Plan:	Longterm	Bor	ehole Observ	vatio	n Pl	an/Re-	entry Pla	an				
Hazards/	Please che	ck fo	ollowing List	of P	oten	ntial Ha	nzards		What is you			
Weather:	Shallow Gas	_ (Complicated Seabed Co			rothermal Ac			window? (P period with the			
	Hydrocarbon		Soft Seabed		Landsl	lide and Turt	oidity Current		period will in	ic rease	,,,,	
	Shallow Water Flow		Currents		Metha	ne Hydrate						
	Abnormal Pressure		Fractured Zone		Diapir	and Mud Vo	olcano					
	Man-made Objects		Fault		High 7	Геmperature						
	H ₂ S		High Dip Angle		Ice Co	onditions						
	CO ₂											