

1 **Title: In-situ shear modulus reduction with strain in stiff fissured**  
2 **clays and weathered mudstones**

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15 **Abstract**

16 The non-linear stress-strain behaviour of stiff clays and weak rocks at small and medium strains may  
17 be a critical consideration in the design of geotechnical structures. Empirical methods have been  
18 developed for estimating the maximum shear modulus and the normalised shear modulus reduction  
19 with strain of fine-grained soils. These are usually expressed as functions of the void ratio (or specific  
20 volume) and average effective (confining) stress, based on results from laboratory tests. However,  
21 the fidelity of these equations has not been widely evaluated in-situ.

22 This paper describes the use of in-situ measurements from an instrumented embankment to  
23 calculate the operational in-situ shear modulus of the underlying stiff clays and weathered  
24 mudstones at medium and large strains. It is shown that the shear modulus at very small strain of  
25 the weathered clays increased linearly with depth, consistent with empirical equations. The gradient  
26 of the normalised, non-linear stiffnesses of the clays were comparable with those measured in  
27 laboratory tests of fine-grained soils, at a range of strains. However, the values for the reference  
28 strain, where the maximum shear modulus reduces by 50%, were lower than was predicted by the  
29 empirical equations. Keywords: stiff clay, weathered mudstone, small-strain stiffness,  
30 instrumentation

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## 34 Introduction

35 The stress-strain behaviour of stiff clays and weak rocks is highly non-linear (Jardine et al. 1984;  
36 Atkinson 2000; Clayton 2011; O'Brien et al. 2023). Their stiffness reduces most rapidly with strain  
37 over the medium strain range of 0.001% to 0.1%. This corresponds to typical strain levels around  
38 geotechnical structures such as foundations, retaining walls and tunnels, which may vary from small  
39 (<0.001%) to large (up to 1%) prior to yield (Jardine et al. 1986; Mair 1993; Clayton 2011).

40 The reduction of in-situ ground stiffness at small, medium and large strains has been inferred from  
41 back-analyses of structural behaviour (Burland 1989; Ng et al. 1995; Ng et al, 1998; Clayton 2011)  
42 and is now an important design consideration for the serviceability of many geotechnical structures  
43 (BSI 2004; O'Brien et al. 2023). Figure 1 (adapted from Mair 1993; Ishihara 1996; Atkinson 2000 and  
44 Clayton 2011) shows the typical reduction in shear modulus ( $G$ ) from a maximum value ( $G_{max}$ ) at  
45 small strain (<0.001%) towards a lower modulus value at larger strains. The typical ranges of shear  
46 strain associated with common in-situ and laboratory testing methods, and applicable to  
47 geotechnical analyses, are also indicated.

48 Small strain stiffness can be measured using in-situ geophysical tests and in the laboratory using  
49 bender elements or resonant column apparatus (Clayton 2011). Stiffness at larger strains can be  
50 obtained from conventional and specialist triaxial testing of laboratory samples (Atkinson 2000;  
51 Clayton 2011) or from the back-analyses of structural behaviour at full-scale (Burland 1989; Menkiti  
52 et al. 2004; Kelly et al. 2018; Smith et al. 2018; Le et al. 2023). However, Hight et al. (2007) and  
53 O'Brien et al. (2023) describe a number of practical challenges related to the measurement of non-  
54 linear stiffness. For laboratory tests these include the potential for sample disturbance, the slow rate  
55 of testing and a limited number of samples or preferential sampling not representing the in-situ  
56 geological variation. For in-situ tests or back-analyses, challenges include the high cost, limited range  
57 of strain measurement and the relevance of the direction of measurement to that of the design  
58 loading. The maximum modulus ( $G_{max}$ ) from field measurements is often greater than that measured  
59 in the laboratory (Tatsuoka et al. 2003; O'Brien et al. 2023). There is also a 'data gap' between  
60 measurements of  $G_{max}$  at very small strain (<0.001%) and measurements of  $G$  in routine laboratory  
61 testing, which become less reliable below 0.01% strain. Additional complexities include stiffness  
62 anisotropy (Lings et al. 2000; Gasparre et al. 2007), and the dependence of stiffness on stress history  
63 and stress path (Atkinson et al, 1990; Hight & Higgins, 1995; Leroueil & Hight, 2003).

64 Atkinson (2000) advocated the use of simple analyses to assess in-situ ground stiffness for  
65 geotechnical design, where possible. This includes cases where movement is predominately one-  
66 directional, such as the settlement of a foundation or the horizontal movement at the top of a  
67 retaining wall. Empirical expressions for the secant shear modulus ( $G$ ) of clays at a range of strain  
68 values include those developed from the interpretation of a database of tests on fine-grained soils  
69 (Darendeli 2001; Vardanega & Bolton 2013) and the interpretation of laboratory and field data using  
70 easily-obtained parameters (Atkinson 2000; O'Brien et al. 2023).

71 The construction of the UK High Speed 2 (HS2) railway between London and Birmingham has  
72 provided an opportunity to obtain monitoring data from geotechnical structures including tunnels,  
73 cuttings and embankments built on or through a range of geological strata from the Cretaceous,  
74 Jurassic and Triassic periods. Among these was a fully-instrumented trial embankment constructed  
75 on weathered clays and mudstones of the Jurassic Charmouth Mudstone Formation (Lias Group) at a  
76 site near Banbury, Oxfordshire. This is a case of predominately vertical loading and ground  
77 deformation that is suited to the simple back-analysis approach advocated by Atkinson (2000).

78 This paper aims to assess values of operational in-situ shear moduli for stiff fissured clays and  
79 weathered mudstones, at a range of pre-yield strains (<1%) relevant for the serviceability of  
80 geotechnical a structure. This is achieved by measuring and analysing the surface loading, pore  
81 water pressures and ground deformations during the construction of an instrumented trial  
82 embankment on weathered clays and mudstones, interpreted using site investigation data and in-  
83 situ geophysical measurements.

## 84 Materials

85 An instrumented trial embankment was constructed on stiff fissured clays and weathered  
86 mudstones of the Charmouth Mudstone Formation. The settlement of the trial embankment was  
87 monitored to inform the design and construction of earthworks located on mudstone outcrops in  
88 central England for the High Speed 2 railway (Munro, 2021). Construction of the trial embankment  
89 began on 7 November 2020 and was completed on 9 December 2020, when the embankment had  
90 reached a height of 8.2 m (Menteth, 2024). The embankment was constructed using fill material  
91 excavated from a deep (15 m) cutting excavation located directly to the south.

## 92 The site

93 The trial embankment was located within an outcrop of the Charmouth Mudstone Formation  
94 approximately 14 km to the north of Banbury (52°11'17"N, 1°20'25"W), Oxfordshire (Figure 2). The  
95 Charmouth Mudstone of the Lias Group was formed approximately 183-199 Myr ago and was  
96 formerly known as the Lower Lias Clay (Cox et al. 1999). The Charmouth Mudstone Formation was  
97 deposited in shallow seas and subsequently exposed to overconsolidation and weathering during  
98 glacial and periglacial conditions in the last 0.2 M years. The lithology of the formation is principally  
99 mudstone with thin limestone and sandstone bands; with weathered clay and some superficial  
100 deposits at shallower depth. Based on downhole geophysical logs, Hobbs et al. (2012) described the  
101 Charmouth Mudstone Formation in this region (the East Midlands Shelf) as 100-150 m thick, with a  
102 remarkably uniform internal stratigraphy across the region. At the site there is a gradational  
103 weathering profile from the ground surface (Briggs et al. 2022), resulting from glacial, periglacial and  
104 contemporary weathering in this location (Quaternary Province 4: Foster et al. 1999).

105 Seven cable percussion (to 10 mbgl) and rotary cored (>10 mbgl) boreholes were drilled in the  
106 ground beneath the embankment (ground level c. 122 mAOD) and rotary cored samples were taken  
107 for laboratory testing, as part of the HS2 ground investigation. The weathering profile was recorded  
108 according to BS 5930: 2015+A1:2020 'Approach 4' for weak rocks. The borehole strata descriptions  
109 show weathered, firm to locally stiff fissured clay to 5 mbgl (117 mAOD) and weathered, stiff and  
110 very stiff fissured clay to 12 mbgl (110 mAOD). They show weathered, extremely weak fissured  
111 mudstone and unweathered extremely weak to very weak fissured mudstone below 12 mbgl (Figure  
112 3). A 2m thick band of calcareous siltstone (i.e. limestone) was observed at approximately 18 mbgl  
113 (104 mAOD). Both the transition from clay to mudstone (~12 mbgl) and the calcareous siltstone (~18  
114 mbgl) were visible in optical borehole images (not shown) obtained from beneath the centre of the  
115 embankment. Figure 3 shows the moisture content (%), plasticity index (%), specific volume, unit  
116 weight (kN/m<sup>3</sup>) and undrained shear strength (kPa) from HS2 ground investigation data obtained  
117 within 0.25 km of the trial embankment. The moisture content reduced from approximately 25%  
118 near the surface (<2.5 mbgl) to approximately 20% at greater depth. The plasticity index reduced  
119 from approximately 35% at the near surface to approximately 30% at greater depth. The bulk unit  
120 weight increased with depth from 20.5 kN/m<sup>3</sup> to 21.5 kN/m<sup>3</sup>. These are consistent with  
121 measurements in the Charmouth Mudstone Formation outcrop at this location (Briggs et al. 2022).

## 122 The trial embankment

123 Figure 2 shows a plan view of the trial embankment and the instrument locations. The embankment  
124 was constructed in stages between 7 November 2020 and 9 December 2020. It was approximately  
125 150 m long and 95 m wide (at the base), with a crest width of 55 m and a slope angle of  
126 approximately 23°. The height of the embankment was measured by aerial drone surveys during  
127 construction and reached a final value of 8.2 m. Surface water runoff ponds were located to the  
128 north and the east of the site.

129 Instruments were installed in two groups, beneath the centre and the eastern edge of the crest of  
130 the embankment, prior to construction (Figure 2). Each group comprised an RST Instruments  
131 LPTPC09-V-LP vibrating wire total earth pressure cell at the ground surface to measure the load from  
132 the embankment, three RST Instruments VW2100 vibrating wire piezometers to measure pore water  
133 pressure and an RST Instruments EXINLINE-1100 vibrating wire inline extensometer to measure  
134 vertical displacement through the ground profile (Table 1). A Campbell Scientific CS106 barometer  
135 was installed adjacent to the trial embankment to record barometric pressure (hPa) at hourly  
136 intervals. The instruments were installed between August and October 2020 and were logged at  
137 hourly intervals from 5 November 2020 until the end of construction on 9 December 2020. Data  
138 logging continued beyond December 2020 to measure the long term consolidation behaviour of the  
139 trial embankment, to inform the design and construction of HS2 (Menteth, 2024; Briggs et al. 2024).  
140 Data measured after December 2020 extended beyond the immediate, undrained response of the  
141 ground to construction of the trial embankment and were therefore not considered in the analyses  
142 presented in this paper.

## 143 Total pressure cells

144 Total pressure cells (PC1 and PC2) were installed in shallow pits at the ground surface prior to  
145 embankment construction, protected by a 300 mm thick layer of sand. They were calibrated during  
146 installation by the application of known weights and remained responsive to changes in barometric  
147 pressure throughout the monitoring period. Figure 4 shows that the pressure applied to the ground  
148 surface increased as the embankment construction progressed in a series of stages, with greater  
149 pressure beneath the centre of the embankment (PC1) than beneath the edge of the embankment  
150 crest (PC2). The pressures measured at the two locations diverged as construction progressed,  
151 owing to their different positions relative to the edge of the embankment crest.

152 The total pressure cell measurements in Figure 4 were initially used to estimate the unit weight of  
153 the embankment fill at three stages of construction for which drone survey data of the height were  
154 available. The cell pressure measurements and the back-calculated unit weight of the fill were used  
155 to determine the embankment height for other stages of construction, for which no drone survey  
156 data were available.

157 All calculations accounted for both the position of the cells beneath the embankment in relation to  
158 the edge of the embankment crest and for the error inherent in the measurements, owing to the  
159 difference in stiffness between the cells and the medium into which they are inserted, quantified by  
160 means of a cell action factor  $F_{cell}$  (Peattie & Sparrow 1954; Clayton & Bica 1993). Weiler & Kulhawy  
161 (1982) identified fifteen extraneous influences on pressure cell measurements in soil including the  
162 cell dimensions, lateral stress rotation and the relative stiffness of the pressure cell and the soil. A  
163 cell action factor ( $F_{cell}$ ) of 1.04 was adopted; that is, the measured pressure changes were assumed  
164 to be 4% greater than the true vertical stress changes beneath the embankment. This is consistent

165 with cell action factors of 1.04 given by Tory & Sparrow (1967) and  $1.04 \pm 0.03$  given by Talesnick  
166 (2013) for an infinitely stiff sensor.

167 The elastic solution for pressures at the base of an embankment on elastic soil given by Perloff et al.  
168 (1967) was used to relate the embankment height and unit weight of the embankment fill to the  
169 pressure cell data (PC1 and PC2). The increase of embankment height ( $H$ ) was estimated from the  
170 measurements at PC1, below the centre of the embankment, using:

$$171 \quad H \approx \frac{\sigma_{PC1}}{\gamma_{emb} I_z F_{cell}}$$

172 Equation 1

173 where  $F_{cell}$  is the cell action factor,  $\sigma_{PC1}$  is the measured cell pressure at PC1 and  $\gamma_{emb}$  is the unit  
174 weight of the embankment fill. The influence factor ( $I_z$ ) was derived from the chart presented by  
175 Perloff et al. (1967). The unit weight of the embankment fill was calculated using Equation 1 using  
176 the known embankment heights measured by drone surveys at PC1 on 23 November, 2 December  
177 and 9 December 2020. The resulting value of  $22 \text{ kN/m}^3$  is similar to the bulk unit weight of the clay  
178 beneath the embankment (Figure 3).

179 The Perloff et al. (1967) influence factor beneath the edge of the crest of a long embankment with a  
180  $22.5^\circ$  slope angle was approximated using linear functions (Appendix A), and used together with the  
181 pressure cell measurements at PC2, to obtain a second estimate of the height of the embankment  
182 ( $H$ ):

$$183 \quad H \approx \begin{cases} \frac{\sigma_{PC2}}{\gamma_{emb} F_{cell} \sigma_{PC2} \sigma_{PC2,max}} & \text{when } \frac{\sigma_{PC2}}{\sigma_{PC2,max}} \leq 0.49 \\ \frac{\sigma_{PC2}}{\gamma_{emb} F_{cell} (1.134 \sigma_{PC2,max} - 0.275 \sigma_{PC2})} & \text{when } \frac{\sigma_{PC2}}{\sigma_{PC2,max}} \geq 0.49 \end{cases}$$

184 Equation 2

185 where  $F_{cell}$  is the cell factor,  $\sigma_{PC2}$  is the measured cell pressure at PC2,  $\sigma_{PC2,max}$  is the measured cell  
186 pressure at PC2 when the embankment is at maximum height and  $\gamma_{emb}$  is the unit weight of the  
187 embankment. Perloff et al. (1967) assumed a Poisson's ratio of 0.3 and did not consider any other  
188 values, but according to Poulos and Davis (1974) the effect of this parameter is small for  
189 embankments, like this one, that are relatively wide (width  $L$  / height  $H > 5$ ). Figure 5 shows the  
190 calculated increases in embankment height with time using the measurements at PC1 and PC2, the  
191 measurements of actual embankment height at PC1 at three loading stages on 23 November, 2  
192 December and 9 December 2020, and the eight loading stages selected for the shear stiffness  
193 analyses.

## 194 Piezometers

195 The vibrating wire piezometers (Table 1) installed beneath the centre of the embankment (PIEZO1)  
196 and the edge of the embankment crest (PIEZO2) were submerged in de-aired water before being  
197 lowered into the borehole (facing upwards). They were grouted in place using a water-cement-  
198 bentonite grout (2.0 : 1.0 : 0.3 by weight) to maintain hydraulic connectivity with the soil. The  
199 piezometers showed a hydrostatic pore water pressure profile below a water table approximately  
200 0.8-1 mbgl prior to embankment construction. The measurements from the shallower ( $\leq 20 \text{ mbgl}$ )  
201 piezometers beneath the centre of the embankment showed that pore water pressures rapidly  
202 increased in response to each loading stage (Briggs et al. 2024). This was followed by a slight  
203 reduction in pore water pressure (indicating a little drainage) beneath the edge of the embankment

204 crest between each loading stage, but this was small (1-2 kPa) relative to the applied changes in total  
205 stress (up to 190 kPa). The piezometers at greater depth (> 20 mbgl) showed a smaller response to  
206 the embankment loading, consistent with the attenuation of vertical stress with depth.

## 207 Extensometers

208 The vibrating wire in-line extensometers (Table 1) beneath the centre of the embankment (EXT1)  
209 and beneath the edge of the embankment crest (EXT2) were installed to depths of 60 mbgl and 50  
210 mbgl respectively. Each extensometer included six Borros hydraulic anchors installed at specified  
211 depths. These were connected to six displacement transducers in series, separated by stainless steel  
212 rods within a PVC sheath. During installation the anchors were hydraulically activated in ascending  
213 order from the base of the borehole, then grouted in place using a water-cement-bentonite grout  
214 (6.6 : 1.0 : 0.4 by weight). Table 2 shows the displacements of the extensometer anchors at various  
215 depths, relative to the deepest anchor at the base of the extensometer.

216 Table 2 shows that during embankment construction there was negligible displacement in EXT2  
217 between the anchor at 35 mbgl and the base anchor 50 mbgl. The displacement of the anchors at  
218 shallower depth (0, 2.5, 7.5, 15 & 25 mbgl) increased with each embankment loading stage. At EXT1,  
219 the measurements between the anchor at 40 mbgl and the base anchor at 60 mbgl showed some  
220 noise. They increased gradually to 1 mm at Stage 5 (27 November 2020), then reduced. This suggests  
221 that some dislocation of the base anchor may have occurred during loading stage 5; the potential for  
222 error was mitigated by using the displacement between adjacent extensometer anchors, rather than  
223 displacements relative to the base anchor, in calculations.

224 Figure 6 shows the relative displacement ( $\delta_{\text{Layer}}$ ) between adjacent extensometer anchors in each  
225 borehole during construction of the trial embankment. The relative displacements of the shallowest  
226 anchor pairs between 0 mbgl and 5 mbgl (EXT1) and between 0 mbgl and 2.5 mbgl (EXT2) were an  
227 order of magnitude greater than for the deeper anchors. They are therefore omitted from Figure 6  
228 for clarity, but are recorded in Table 3. The measurements show increasing relative displacements  
229 between the pairs of adjacent anchors down to at 40 mbgl at EXT1 (Figure 6(a)), and down to 35  
230 mbgl at EXT2 (Figure 6(b)). Note that the soil layers shown in Figure 6 are not of equal thickness.

## 231 In-situ testing

232 Downhole seismic tests were undertaken in four boreholes by a specialist contractor for the HS2  
233 ground investigation. Optical image logs were obtained from one borehole (DHGEO\_3) using a  
234 precision-machined prism and CCD camera. The boreholes were located approximately 250 m to the  
235 south of the trial embankment, at elevations between 134 and 136 mAOD. The borehole records  
236 showed weathered, stiff to very stiff fissured clay to 13 mbgl (113 mAOD), with mudstone below.  
237 The calcareous siltstone (i.e. limestone) was located within the mudstone at approximately 32 mbgl  
238 (102 mAOD).

239 P-wave and S-wave seismic velocities were measured at 1 m intervals of depth within plastic-lined  
240 boreholes, to 63 mbgl. The S-waves were generated by a sledgehammer striking the end of a timber  
241 sleeper at the ground surface. The P-waves were generated by vertically striking an acrylic plate at  
242 the ground surface with a sledgehammer. The seismic waves were detected by a BGK-7 multi  
243 element geophone having one vertical and six horizontal sensors, pneumatically clamped within the  
244 borehole at each successive test depth.

245 Figure 7 shows a linearly-increasing shear wave velocity ( $V_s$ ) with increasing depth within all four  
246 boreholes, to approximately 20 mbgl. The measurements in the mudstone at greater depth (> 20

247 mbgl) vary between individual boreholes. The borehole records showed no change in the visual  
 248 appearance of the mudstone weathering profile that might explain the increased variation in  
 249 geophysical measurements below 20 mbgl. Similarly, no change was visible in the optical images  
 250 from DHGEO\_3 (not shown). However, this depth is consistent with the transition between the  
 251 weathered (Class Ba) and the unweathered (Class A) material across the Charmouth Mudstone  
 252 Formation outcrop at the site location (Briggs et al. 2022). The measured compression wave  
 253 velocities ( $V_p$ ) shown in Figure 7 were less than for water (approx. 1500 m/s) at depths to 40 mbgl  
 254 and hence of limited use. This is typical of soft rocks (Clayton 2011; Poulos 2022).

255 The downhole geophysical measurements (Figure 7) and sample unit weight measurements (Figure  
 256 3) were used to produce a profile of shear modulus at very small strain ( $G_0$ ) for the weathered clay  
 257 and mudstone layers (up to 20 mbgl) using the relationship (Zisman 1933; Atkinson 2000; Poulos  
 258 2022):

$$259 \quad G_0 = \rho V_s^2 = \frac{\gamma_b}{g} V_s^2$$

260 Equation 3

261 where  $\rho$  is the bulk density ( $\text{kg/m}^3$ ),  $\gamma_b$  is the bulk unit weight ( $\text{kN/m}^3$ ),  $g$  is the acceleration of the  
 262 Earth's gravity ( $\text{m/s}^2$ ) and  $V_s$  is the shear wave velocity ( $\text{m/s}$ ). The shear modulus at very small strain  
 263 ( $G_0$ ) from the downhole geophysical measurements was considered as the maximum ( $G_{max}$ ).

264 For comparison, a profile of maximum shear modulus ( $G_{max}$ ) with depth was determined using the  
 265 unit weight and specific volume of the samples (Figure 3) as inputs for the Vardanega & Bolton  
 266 (2013) equation for fine-grained soils tested in laboratory conditions:

267

$$268 \quad \frac{G_{max}}{p'_r} = \frac{B}{(v)^{2.4}} \left( \frac{p'}{p'_r} \right)^{0.5}$$

269 Equation 4

270 where  $p'$  is the mean effective stress,  $p'_r$  is a reference stress (taken as 1 kPa) and  $v$  is the specific  
 271 volume of the triaxial samples obtained close to the trial embankment (Figure 3). A soil structure  
 272 coefficient,  $B$ , was selected for a typical fine-grained soil ( $B = 20,000$ ) and for an overconsolidated  
 273 aged clay ( $B = 50,000$ ), as described in Vardanega & Bolton (2013). Figure 8 shows an increasing  
 274 profile of  $G_{max}$  with depth at the site. This is in close agreement with the Vardanega & Bolton (2013)  
 275 profile for typical fine-grained soils ( $B = 20,000$ ) to 8 mbgl. There is greater scatter in the downhole  
 276 geophysical measurements below 8 mbgl. Therefore, separate linear (regression) fits for  $G_{max}$  were  
 277 derived for the weathered clay (0-8 mbgl) and the less-weathered clay and mudstone below (> 8  
 278 mbgl). The value of  $G_{max}$  at greater depth in the less-weathered clay and mudstone layers (> 11 mbgl)  
 279 lies closer to the Vardanega & Bolton (2013) profile for overconsolidated aged clay ( $B = 50,000$ ).

## 280 Methods

281 The essence of the approach was to use the embankment loading and extensometer data at known  
 282 construction stages to calculate the in-situ shear modulus profile of the Charmouth Mudstone  
 283 Formation beneath the embankment. Together with the downhole geophysical measurements,  
 284 these were used to determine the shear modulus reduction with strain curve. The method can be  
 285 summarised as:

286 1. Data from the pressure cells and aerial drone surveys of embankment height were used to obtain  
 287 the magnitude and distribution of loading at the ground surface at selected stages (loading stages)  
 288 during construction of the embankment.

289 2. Relative displacements measured between adjacent extensometer anchors were used to  
 290 determine the average vertical strains within selected layers below the embankment.

291 3. The vertical strains, together with the surface loading and elasticity equations, were used to  
 292 determine the representative stresses and strains below the embankment on a layer-by-layer basis  
 293 at selected loading stages.

294 4. Corresponding shear stresses and shear strains were used to calculate the operational secant  
 295 shear modulus and secant shear strain for each layer and loading stage.

296 6. A profile of maximum shear modulus against depth was obtained from down-hole seismic tests.

297 7. Finally, plots of normalised secant shear modulus against shear strain were obtained for each  
 298 layer and loading stage.

299 The shear modulus and shear modulus reduction curve were calculated for layers within the  
 300 weathered clay and weathered mudstone ground profile, to 20 mbgl. They were not calculated for  
 301 the unweathered mudstone at greater depth (>20 mbgl), because the measured displacements were  
 302 very small (<2 mm) below this depth (Table 2). The analyses assumed an immediate, undrained  
 303 ground response to the surface loading. The piezometers did show some drainage beneath the edge  
 304 of the embankment crest at shallower depth (up to 10 mbgl) between the loading stages (Briggs et  
 305 al. 2024), but the small (1-2 kPa) pore water pressure change relative to the applied loading (20-170  
 306 kPa) justifies the assumption of substantially undrained conditions.

### 307 Calculation of stress increases associated with embankment construction

308 Construction of the trial embankment increased the total stresses in the underlying ground. The  
 309 increases in vertical, horizontal and shear stress ( $\sigma_z$ ,  $\sigma_x$  and  $\tau_{xz}$  respectively) were calculated for each  
 310 layer (between adjacent extensometer anchors) in each ground profile (beneath the centre and the  
 311 edge of the embankment crest) at each selected stage of construction (Figure 5). The changes in  
 312 stresses were calculated using the analytical equations for stress increments in an elastic half-space  
 313 under vertical loading and plane-strain conditions, derived by Gray (1936) and summarised in Poulos  
 314 & Davis (1974). These equations assume a linear elastic, homogeneous, isotropic material and can be  
 315 superimposed to derive solutions for geometrically more complicated loading scenarios, such as a  
 316 different location beneath an embankment. The equations are:

$$317 \quad \Delta\sigma_z = \frac{P}{\pi} \left[ \beta + \frac{x\alpha}{a} - \frac{z}{R_2^2} (x - b) \right]$$

318 Equation 5

$$319 \quad \Delta\sigma_x = \frac{P}{\pi} \left[ \beta + \frac{x\alpha}{a} - \frac{z}{R_2^2} (x - b) + \frac{2z}{a} \ln \frac{R_1}{R_0} \right]$$

320 Equation 6

$$321 \quad \Delta\tau_{xz} = -\frac{P}{\pi} \left[ \frac{z\alpha}{a} - \frac{z^2}{R_2^2} \right]$$

322 Equation 7



323 where  $\Delta\sigma_z$  is the change in vertical stress,  $\Delta\sigma_x$  is the change in horizontal stress (in the vertical cross-  
 324 sectional plane),  $\Delta\tau_{xz}$  is the change in shear stress,  $P$  is the surface load,  $x$  is the horizontal location,  $z$   
 325 is the vertical location (i.e. depth) and the geometry parameters are defined in Figure 9. The surface  
 326 load ( $P$ ) at each load stage was equal to the unit weight of the embankment fill ( $\gamma_{emb} = 22 \text{ kN/m}^3$ )  
 327 multiplied by the height of the embankment,  $H$  (as shown in Figure 5 and Table 2). The embankment  
 328 width geometry parameter,  $b$ , was half the embankment width. The slope width parameter,  $a$ ,  
 329 varied as the embankment height increased (for a slope angle of  $22.5^\circ$ ).

330 Changes in stress were calculated at the top ( $\Delta\sigma_{Top}$ ), midpoint ( $\Delta\sigma_{Mid}$ ) and base ( $\Delta\sigma_{Base}$ ) of the layers  
 331 beneath the centre of the embankment (EXT1) and beneath the edge of the embankment crest  
 332 (EXT2). These were used to derive the weighted average change in stress in each layer ( $\Delta\sigma_{LayerAve}$ )  
 333 using Simpson's rule (Atkinson 1989):

$$334 \quad \Delta\sigma_{LayerAve} = \frac{1}{6} [\Delta\sigma_{Top} + 4\Delta\sigma_{Mid} + \Delta\sigma_{Base}]$$

335 *Equation 8*

336 The weighted average changes in vertical stress ( $\Delta\sigma_{zLayerAve,Stage}$ ), horizontal stress ( $\Delta\sigma_{xLayerAve,Stage}$ ) and  
 337 shear stress ( $\Delta\tau_{xzLayerAve,Stage}$ ) were calculated for each layer, for each embankment loading stage.

### 338 Calculation of vertical strains

339 The average vertical strains were calculated for each stage of embankment construction and for  
 340 each layer ( $\varepsilon_{zLayer,Stage}$ ), from the relative displacement between adjacent pairs of extensometer  
 341 anchors,  $\delta_{zLayer,Stage}$  (Figure 6) and the initial layer thickness,  $Z_{0Layer}$  (that is, the initial extensometer  
 342 anchor spacing):

$$343 \quad \varepsilon_{zLayer,Stage} = \frac{\delta_{zLayer,Stage}}{Z_{0Layer}}$$

344 *Equation 9*

345 Vertical strains were calculated for three layers beneath the centre of the embankment at EXT1 (0 to  
 346 5 mbgl, 5 to 10 mbgl and 10 to 20 mbgl) and three layers beneath the edge of the embankment crest  
 347 at EXT2 (0 to 2.5 mbgl, 2.5 to 7.5 mbgl and 7.5 to 15 mbgl). These are the layers for which the  
 348 relative displacement between adjacent extensometer anchors was greater than 1.1 mm (Table 3).  
 349 The layers at greater depths (>20 mbgl), with relative displacements below this threshold, were  
 350 excluded from the analyses. The depth threshold of 20 mbgl corresponded with the transition from a  
 351 uniform to a more scattered shear wave velocity profile in the nearby downhole seismic tests (Figure  
 352 7), and with the transition from weathered (Class Ba) to unweathered (Class A) mudstone observed  
 353 across the Charmouth Mudstone Formation outcrop locally (Briggs et al. 2022).

### 354 Calculation of the in-situ shear modulus, shear strain and normalised shear 355 modulus

356 The operational secant shear modulus for each layer and embankment loading stage ( $G_{Layer,Stage}$ ) was  
 357 calculated using the stress-strain relationship (rearranged from Equation 1.36c in Poulos & Davis  
 358 1974):

$$359 \quad G_{Layer,Stage} = \frac{1}{2\varepsilon_{zLayer,Stage}} [(\Delta\sigma_{zLayerAve,Stage}(1 - \nu_u)) - (\nu_u\Delta\sigma_{xLayerAve,Stage})]$$

360

Equation 10

361 where  $\nu_u$  is the undrained Poisson's ratio (taken as 0.5). The plane-strain shear stress and shear  
 362 strain invariants were calculated for each layer and embankment loading stage to find the maximum  
 363 shear strains, for comparison with the Vardanega & Bolton (2013) laboratory test results. The  
 364 change in the plane-strain shear stress invariant, i.e. the radius of the Mohr circle in the cross-  
 365 sectional plane, was calculated using:

$$366 \quad \Delta\tau_{Layer Ave, Stage} = \sqrt{\frac{1}{4}(\Delta\sigma_{zLayer Ave, Stage} - \Delta\sigma_{xLayer Ave, Stage})^2 + \Delta\tau_{xzLayer Ave, Stage}^2}$$

367

Equation 11

368 The in-situ, plane-strain shear strain invariant ( $\gamma_{Layer, Stage}$ ) was calculated for each layer and for each  
 369 stage of embankment construction as:

$$370 \quad \gamma_{Layer, Stage} = \frac{\Delta\tau_{Layer Ave, Stage}}{G_{Layer, Stage}}$$

371

Equation 12

372 Finally, the secant shear modulus for each layer and embankment loading stage ( $G_{Layer, Stage}$ ) was  
 373 normalised by the maximum shear modulus ( $G_{max}$ ) at the midpoint of each layer. Values of  $G_{max}$  were  
 374 derived from linear regression fits (depth vs  $G_{max}$ ) to the downhole geophysical measurements for 0-  
 375 8 mbgl and 9-20 mbgl, as shown in Figure 8.

376 For comparison, a normalised secant shear modulus reduction with strain curve was also calculated  
 377 using the Vardanega & Bolton (2013) empirical relationship for fine-grained soils. This includes an  
 378 adjustment (referred to as a 'static adjustment') for shear strain rates in static laboratory tests:

$$379 \quad \frac{G}{G_{max}} = \frac{1}{1 + \left(\frac{\gamma}{\gamma_{ref}}\right)^\alpha}$$

380

Equation 13

381 and

$$382 \quad \gamma_{ref} = J\left(\frac{I_p}{1000}\right)$$

383

Equation 14

384 where  $\gamma$  is the shear strain,  $\gamma_{ref}$  is the reference shear strain at  $0.5G_{max}$ ,  $\alpha$  is a fitting parameter (set  
 385 equal to 0.736, as used in Vardanega & Bolton (2013)),  $I_p$  is the plasticity index (expressed as a  
 386 fraction rather than a percentage) and  $J$  is a regression coefficient relating  $I_p$  and  $\gamma_{ref}$  (where  $J = 2.2$   
 387 in Vardanega & Bolton (2013)). Curves were calculated for plasticity indices ( $I_p$ ) of 5% to 35%.

## 388 Results and discussion

389 Figure 10 shows the resulting graphs of normalised secant shear modulus against shear strain for six  
 390 layers beneath the embankment, and eight loading stages. For comparison, empirical curves defined  
 391 by the Vardanega & Bolton (2013) empirical equation (Equation 13) are shown for plasticity indices  
 392 ranging from 5% to 35%.

393 Figure 10 shows the expected behaviour of decreasing normalised secant shear modulus with  
 394 increasing shear strain within all the layers beneath the embankment. The measurements from the  
 395 shallowest layers (0-5 mbgl) are close to zero (less than 0.015) at  $\approx 1\%$  strain, and are located below  
 396 the Vardanega & Bolton (2013) empirical curves. The results from the soil layers between 2.5 mbgl  
 397 and 7.5 mbgl are close to the Vardanega & Bolton (2013) curve for a plasticity index ( $I_p$ ) of 15%.  
 398 These layers are the weathered, stiff and very stiff fissured clays with a plasticity index ( $I_p$ ) of 26 -  
 399 31% (Figure 3). The results from the stiffer (Figure 8), less-weathered clays and weathered  
 400 mudstones between 5 mbgl and 20 mbgl ( $I_p \approx 28\%$ ) are close to the Vardanega & Bolton (2013) curve  
 401 for a plasticity index ( $I_p$ ) of 5%. Therefore, the measurements from beneath the trial embankment  
 402 show decreasing values of reference strain ( $\gamma_{ref}$ ) with increasing depth (and a slight decrease in  
 403 plasticity; Figure 3). They fit the general reduction trend of the Vardanega & Bolton (2013) curves, as  
 404 determined by parameter  $\alpha$ , but show lower values of the reference strain ( $\gamma_{ref}$ ). The measurements  
 405 correspond with Vardanega & Bolton (2013) curves for lower values of the plasticity index compared  
 406 to index test data for the site (Figure 3). It should be noted that the data in Figure 10 assume  
 407 undrained conditions and that no volume change takes place (i.e.  $v_u = 0.5$ ). However, due to the high  
 408 stiffness of the materials, some undrained volume change may occur due to the compressibility of  
 409 the water or dissolved air (Briggs et al. 2024). The implications of this are explored in Appendix B.

410 Figure 11 shows the inferred vertical stress-strain plots for the six layers within the ground profile  
 411 beneath the trial embankment. Figure 11a shows that the vertical strains were approaching values  
 412 associated with yield ( $>1\%$ ) in the shallowest layers (up to 5 mbgl). While the secant modulus  
 413 decreased with vertical strain, the tangent modulus (given by the slope of the graph) increased. This  
 414 may be a result of the drainage and consolidation in these shallowest layers, particularly beneath the  
 415 edge of the embankment crest. Figure 11b shows that the vertical strains in the layers  $>5$  mbgl were  
 416 in the medium strain range (up to 0.08%). These layers showed decreasing secant and tangent  
 417 moduli with vertical strain. Figure 12 shows the operational secant shear modulus ( $G$ ) and shear  
 418 strain ( $\gamma$ ) of six layers within the ground profile beneath the centre and the edge of the crest of the  
 419 embankment. The deeper clay and mudstone layers (grey symbols) had the highest shear modulus  
 420 due to their greater in-situ stress and lower void ratio, in agreement with the geophysical  
 421 measurements (Figure 8). This reduced rapidly with shear strain, but reference to Figure 10 shows  
 422 that this was proportional to  $G_{max}$ . Figure 12 shows that the reduction of shear modulus with shear  
 423 strain in the shallower clay layers (black symbols) was more comparable to the mean curve for fine-  
 424 grained soils, as described in the Vardanega & Bolton (2013) database (where  $I_p = 39\%$ ,  $p' = 209$  kPa,  
 425  $G_{max} = 68$  MPa).

426 The shear modulus obtained from the downhole geophysical measurements (Figure 8) and the shear  
 427 stress-strain relationships obtained from the back-analyses (Figure 10 and Figure 12) show the  
 428 influence of weathering on the in-situ ground profile at the site. The weathered clay (0-8 mbgl)  
 429 exhibited a maximum shear modulus ( $G_{max}$ ) profile comparable with those measured in other fine-  
 430 grained materials, as demonstrated by the close fit to the Vardanega & Bolton (2013) equations.  
 431 However, the maximum shear modulus ( $G_{max}$ ) profile of the less-weathered clay and weathered  
 432 mudstone ( $>8$  mbgl) was larger and more variable than for the weathered clay. It was larger than  
 433 values derived from the Vardanega & Bolton (2013) equation for typical fine-grained soils (i.e.  $B =$   
 434  $20,000$ ), and was closer to those for overconsolidated aged clays (i.e.  $B = 50,000$ ). The shear modulus  
 435 ( $G$ ) of the deeper layers (7.5 to 15 mbgl and 10 to 20 mbgl) reduced more rapidly with shear strain  
 436 ( $\gamma$ ) than in the overlying layers. These less weathered, and hence more structured, clays and  
 437 mudstones were initially stiffer than the shallower, more weathered clays. However, the normalised  
 438 shear modulus ( $G/G_{max}$ ) in all layers reduced at a rate that was comparable to the Vardanega &  
 439 Bolton (2013) equation, over the range of medium strains relevant to geotechnical structures.

440 At intermediate depths (approximately 8-12 mbgl), the results showed a maximum shear modulus  
441 profile that was between that of the weathered clay and the mudstone. This transition compares to  
442 those in the gradational weathering profile in the wider Charmouth Mudstone Formation (Briggs et  
443 al. 2022), as shown by the results of visual inspection, soil classification tests and undrained  
444 unconsolidated (UU) triaxial compression tests.

445 The Poulos & Davis (1974) elasticity equations enabled back-analyses for the simple case of  
446 predominately vertical loading and ground deformation at the trial embankment. It includes  
447 assumptions of undrained loading, linear elasticity and isotropic, homogenous ground stiffness. The  
448 assumption of undrained behaviour in the clay and mudstone is justified by the relatively short  
449 duration of the embankment trial construction (32 days) and the short (1-7 day) intervals between  
450 the embankment loading stages.

451 The elastic half space model assumes a constant shear modulus throughout the ground profile, but  
452 the results showed that the shear modulus increased linearly with depth. However, it is well-known  
453 that vertical stress changes beneath loaded areas are insensitive to nonlinear stress-strain  
454 behaviour, stiffness anisotropy and increasing stiffness with depth (Burland et al. 1977). This was  
455 confirmed by supplementary finite element analyses in Sigma/w (GEO-SLOPE International Ltd,  
456 2013) assuming a linear elastic material (not reported in this paper). This showed that the calculated  
457 changes in vertical total stress were not sensitive to the use of a constant or a linearly increasing  
458 shear modulus profile. Further, these analyses showed that, for the geometry of the trial  
459 embankment, the calculated changes in horizontal total stress were also insensitive (<1% difference)  
460 to a stiffness increasing with depth, and to a stiffness anisotropy in the range indicated by laboratory  
461 and in-situ measurements in clays and mudstones of 1.5 to 2 horizontal:vertical (Mitchell & Soga  
462 2005; Clayton 2011). Burland (2012) has also shown that stiffness anisotropy has a limited influence  
463 on the change in vertical stresses beneath a uniform surface load, such as an embankment.

## 464 Conclusions

465 Instrumentation installed beneath a trial embankment was used to measure the settlement of the  
466 underlying foundation of weathered clays and weathered mudstones, in response to the staged  
467 construction of an 8.2 m high, clay fill embankment. The measurements showed the vertical  
468 deformation of the foundation in response to the applied surface load. Complementary, in-situ  
469 measurements of shear modulus using downhole geophysical methods showed that the foundation  
470 maximum shear modulus increased with depth up to 20 mbgl. The calculated distributions of stress  
471 increase and measured strains were used to determine the secant shear modulus of the foundation  
472 strata at a range of depths and shear strains. This led to the following conclusions.

473 1. The maximum shear modulus ( $G_{max}$ ) of the Charmouth Mudstone Formation increases with depth  
474 and is influenced by the in-situ weathering profile. Measurements within the weathered, stiff and  
475 very stiff fissured clays compare well with the Vardanega & Bolton (2013) empirical correlation for  
476 typical fine-grained soils (i.e.  $B = 20,000$ ), up to a depth of 8 mbgl. Below this depth, the maximum  
477 shear modulus ( $G_{max}$ ) is 50-100 MPa greater than for typical fine-grained soils. At depth (> 11 mbgl)  
478 the Vardanega & Bolton (2013) equation for overconsolidated aged clay (i.e.  $B = 50,000$ ) is more  
479 comparable to the maximum shear modulus ( $G_{max}$ ). The measured shear modulus profile aligns with  
480 the transitions from weathered clay (<13 mbgl) to weathered mudstone ( $\approx$ 13-20 mbgl) and  
481 unweathered mudstone (>20 mbgl) shown in the corresponding borehole records. These compare  
482 with the gradational weathering profile in the wider Charmouth Mudstone Formation outcrop at the  
483 site location (Briggs et al. 2022).

484 2. The normalised secant shear modulus ( $G/G_{max}$ ) of weathered clays and mudstones were  
485 determined using extensometers, a known surface load and complementary geophysical  
486 measurements of the maximum shear modulus ( $G_{max}$ ). Extensometer anchors installed at multiple  
487 depths beneath an increasing surface load, such as an embankment under construction, allow the  
488 shear modulus of the ground to be calculated for multiple stress increments and for a range of strain  
489 values. Critically, it is possible to obtain in-situ shear modulus measurements at a range of strains  
490 that are relevant for the serviceability design of geotechnical structures (<1% strain). These are not  
491 routinely measured in laboratory triaxial tests or in materials that are difficult to sample, such as stiff  
492 clays and weak rocks.

493 3. The Vardanega & Bolton (2013) empirical equation for normalised secant shear modulus  
494 reduction with strain compared with in-situ measurements from the weathered clays and  
495 mudstones beneath the trial embankment (0 mbgl to 20 mbgl). The in-situ measurements from the  
496 shallower, more plastic clay layers showed larger values of reference strain ( $\gamma_{ref}$ ) than the deeper  
497 layers of less-weathered, more structured and less plastic clay and mudstone. This is in agreement  
498 with the Vardanega & Bolton (2013) empirical correlations for fine-grained soils of varying plasticity  
499 index. However, the plasticity indices of the Vardanega & Bolton (2013) curves that fit the in-situ  
500 measurements ( $I_p$  of 5% - 15%) are much lower than the measured plasticity indices of the clays and  
501 mudstones beneath the trial embankment ( $I_p$  of 26% - 31%). Therefore, the values for the reference  
502 strain ( $\gamma_{ref}$ ) that compare to the in-situ measurements are lower than would be predicted by the  
503 empirical equations.

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## 514 Competing interests

515 Competing interests: The authors declare there are no competing interests.

## 516 Data availability

517 The data presented in this paper are available online via the University of Bath institutional  
518 repository (Briggs, 2024a) and may be accessed at <https://doi.org/10.15125/BATH-01353>.

519

520 

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- 642
- 643

## 644 Figures and Tables

- 645 *Figure 1: A shear stiffness reduction curve showing the typical strain range for geotechnical structures, analysis types and*  
 646 *the approximate range of different measurement methods. Redrawn from Mair (1993), Ishihara (1996), Atkinson (2000),*  
 647 *Clayton (2011) and O'Brien et al. (2023).*
- 648 *Figure 2: The location of the trial embankment showing (a) a plan of the embankment and the location of instrumentation,*  
 649 *(b) the site location (52°11'17"N, 1°20'25"W) within the outcrop of the Charmouth Mudstone Formation in central England.*
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 654 *Oxfordshire (Briggs et al. 2022) are shown for comparison.*
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- 657 *Figure 5: The change in embankment height with time, back-calculated from cell pressure measurements (PC1 & PC2) and*  
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- 659 *Figure 6: The relative displacement between extensometer anchors installed at various depths during construction of the*  
 660 *trial embankment at (a) EXT1 and (b) EXT2. Note that the extensometer anchors were not equally spaced.*
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- 667 *Figure 9: The geometry parameters for the distributed vertical embankment loading equations described in Poulos & Davis*  
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- 669 *Figure 10: A normalised secant shear modulus reduction curve with strain for layers beneath the trial embankment, derived*  
 670 *from monitoring data. These are compared to results from the Vardanega & Bolton (2013) model for fine-grained soils with*  
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- 675 *Figure 12: The secant shear modulus,  $G$  (MPa) vs shear strain,  $\gamma$  (%), for layers within the ground profile beneath the*  
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 677 *for fine-grained soils (where  $I_p = 39\%$ ,  $p' = 209$  kPa,  $G_{max} = 68$  MPa).*
- 678 *Figure A 1: Influence factors (points) from the Perloff et al. (1967) chart and fitted linear regressions (solid lines). The*  
 679 *dashed line shows the fit for a 22.5° slope angle.*
- 680 *Figure B 1: A normalised secant shear modulus reduction curve with strain for layers beneath the trial embankment, derived*  
 681 *from monitoring data and assuming  $v_u$  equal to 0.458. These are compared to results from the Vardanega & Bolton (2013)*  
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- 684 *Table 2: The extensometer anchor displacements at various depths within EXT 1 and EXT2, relative to the base anchor*  
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 686 *stages.*
- 687 *Table 3: The relative displacement between extensometer anchors installed at various depths within EXT1 and EXT2 (mm).*  
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 689 *that the extensometer anchors were not equally spaced.*

## Tables

Table 1: Instrumentation installed beneath the trial embankment (see Figure 2).

Measurement type	Instrument type (& model)	Instrument location & depth (mbgl)	Measuring range/resolution
Total pressure (kPa)	Vibrating wire total earth pressure cell* (LPTPC09-V-LP)	PC1 at 0.3 PC2 at 0.3	A 31.7 cm diameter cell calibrated to measure pressure between 0 and 175 kPa, logged at 0.1 kPa resolution
Pore water pressure (kPa)	Vibrating wire piezometers* (VW2100)	PIEZO1 at 10, 20, 34 PIEZO2 at 7.5, 15, 25 (note: not shown in analyses)	Pore pressure between 0 and 350 kPa (at 10, 20 m) Pore pressure between 0 and 700 kPa (All others) Measurements logged at 0.1 kPa resolution
Vertical ground displacement (mm)	Vibrating wire inline extensometers (EXINLINE-1100)	EXT1 at 0, 5, 10, 20, 30, 40 & 60 EXT2 at 0, 2.5, 7.5, 15, 25, 35 & 50	Tape measurement at 0.02 mm resolution

\* Calibrated by the manufacturer in compliance with BS EN ISO/IEC 17025:2017 (British Standards Institution, 2017)

Table 2: The extensometer anchor displacements at various depths within EXT 1 and EXT2, relative to the base anchor (mm). The measurements are shown to 1 decimal place for dates corresponding to eight known embankment loading stages.

Stage	Date & time	Emb. Height (m)	Anchor location	Anchor displacement relative to the base anchor (mm to 1dp)													
				EXT1	EXT1	EXT1	EXT1	EXT1	EXT1	EXT1	EXT1	EXT2	EXT2	EXT2	EXT2	EXT2	EXT2
			Anchor depth (mbgl)	60	40	30	20	10	5	0	50	35	25	15	7.5	2.5	0
1	07/11/2020 (18:00)	1.16		0.0	-0.4	-0.4	-0.4	-0.4	-0.5	-4.3	0.0	0.0	0.0	-0.1	-0.3	-0.5	-6.0
2	20/11/2020 (18:00)	3.69		0.0	-0.2	-0.4	-0.7	-1.1	-1.5	-16.6	0.0	0.0	-0.1	-0.5	-1.4	-2.1	-16.1
3	23/11/2020 (18:00)	4.33		0.0	-0.4	-0.7	-1.1	-1.8	-2.4	-21.5	0.0	0.0	-0.1	-0.7	-1.9	-3.0	-22.7
4	26/11/2020 (18:00)	4.82		0.0	-0.9	-1.3	-1.8	-2.8	-3.5	-27.5	0.0	0.0	-0.1	-1.1	-2.7	-4.0	-32.3
5	27/11/2020 (18:00)	5.45		0.0	-1.0	-1.4	-1.9	-3.1	-3.9	-29.3	0.0	0.0	-0.2	-1.2	-3.0	-4.6	-35.4
6	30/11/2020 (18:00)	5.86		0.0	-0.3	-0.7	-1.4	-2.8	-3.8	-31.8	0.0	0.0	-0.2	-1.4	-3.8	-5.7	-42.3
7	02/12/2020 (18:00)	6.59		0.0	-0.2	-0.7	-1.5	-3.2	-4.4	-34.8	0.0	0.0	-0.2	-1.6	-4.7	-7.0	-49.7
8	09/12/2020 (18:00)	8.23		0.0	-0.3	-0.9	-2.0	-4.3	-6.1	-43.8	0.0	0.0	-0.3	-2.3	-5.7	-9.6	-57.9

Table 3: The relative displacement between extensometer anchors installed at various depths within EXT1 and EXT2 (mm). The measurements are shown to 1 decimal place for dates corresponding to eight known embankment loading stages. Note that the extensometer anchors were not equally spaced.

Stage	Date & time	Emb. Height (m)	Anchor location	Relative displacement between the extensometer anchors (mm to 1dp)												
				EXT1	EXT1	EXT1	EXT1	EXT1	EXT1	EXT2	EXT2	EXT2	EXT2	EXT2	EXT2	
			Anchor depths (mbgl)	40-60	30-40	20-30	10-20	5-10	0-5		35-50	25-35	15-25	7.5-15	2.5-7.5	0-2.5
1	07/11/2020 (18:00)	1.16		-0.4	0.0	0.0	0.0	0.0	-3.8		0.0	0.0	0.0	0.0	0.0	-5.5
2	20/11/2020 (18:00)	3.69		0.0	0.0	-0.3	-0.5	-0.4	-15.1		0.0	0.0	-0.4	-0.9	-0.8	-14.0
3	23/11/2020 (18:00)	4.33		-0.4	-0.3	-0.4	-0.7	-0.6	-19.2		0.0	0.0	-0.6	-1.2	-1.1	-19.7
4	26/11/2020 (18:00)	4.82		-0.9	-0.4	-0.5	-1.0	-0.7	-24.0		0.0	0.0	-0.9	-1.6	-1.4	-28.3
5	27/11/2020 (18:00)	5.45		-1.0	-0.4	-0.6	-1.2	-0.8	-25.4		0.0	0.0	-1.0	-1.8	-1.6	-30.8
6	30/11/2020 (18:00)	5.86		-0.3	-0.5	-0.7	-1.4	-1.0	-28.0		0.0	0.0	-1.2	-2.4	-1.9	-36.6
7	02/12/2020 (18:00)	6.59		0.0	-0.5	-0.8	-1.6	-1.2	-30.4		0.0	0.0	-1.4	-3.1	-2.3	-42.7
8	09/12/2020 (18:00)	8.23		-0.3	-0.6	-1.1	-2.3	-1.8	-37.7		0.0	-0.3	-2.0	-3.5	-3.9	-48.4

## Figures

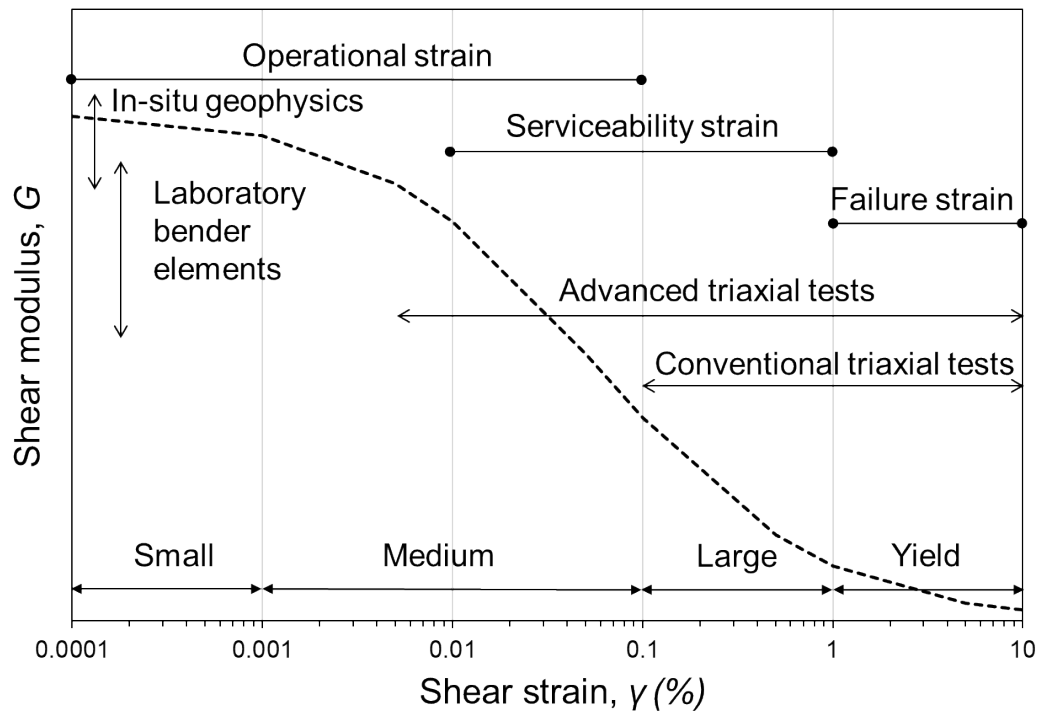


Figure 1: A shear stiffness reduction curve showing the typical strain range for geotechnical structures, analysis types and the approximate range of different measurement methods. Redrawn from Mair (1993), Ishihara (1996), Atkinson (2000), Clayton (2011) and O'Brien et al. (2023).

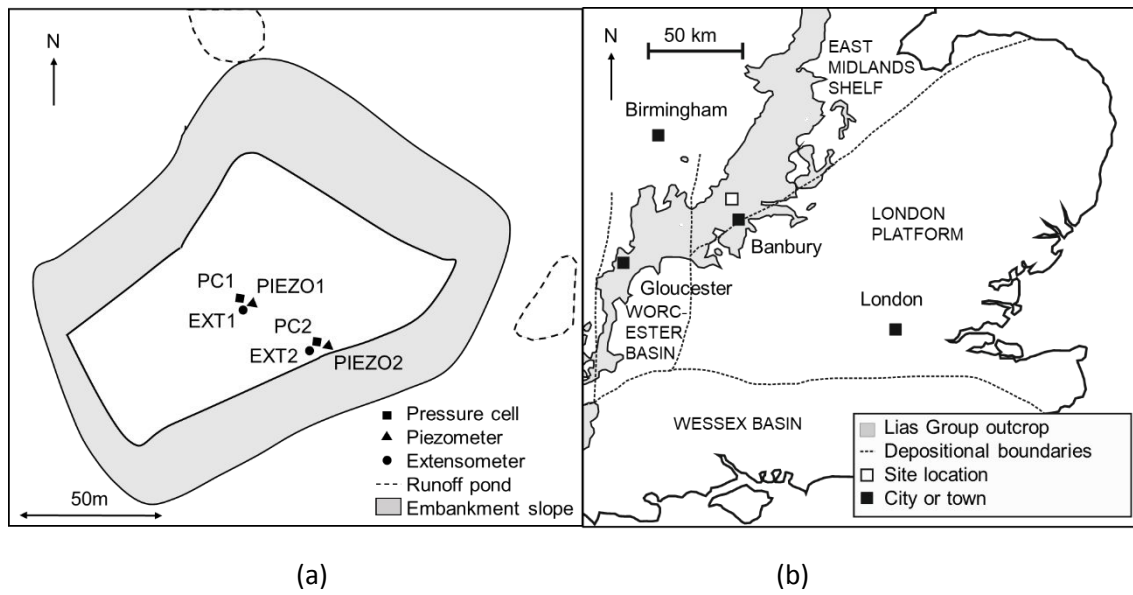


Figure 2: The location of the trial embankment showing (a) a plan of the embankment and the location of instrumentation, (b) the site location ( $52^{\circ}11'17''\text{N}$ ,  $1^{\circ}20'25''\text{W}$ ) within the outcrop of the Charmouth Mudstone Formation in central England.

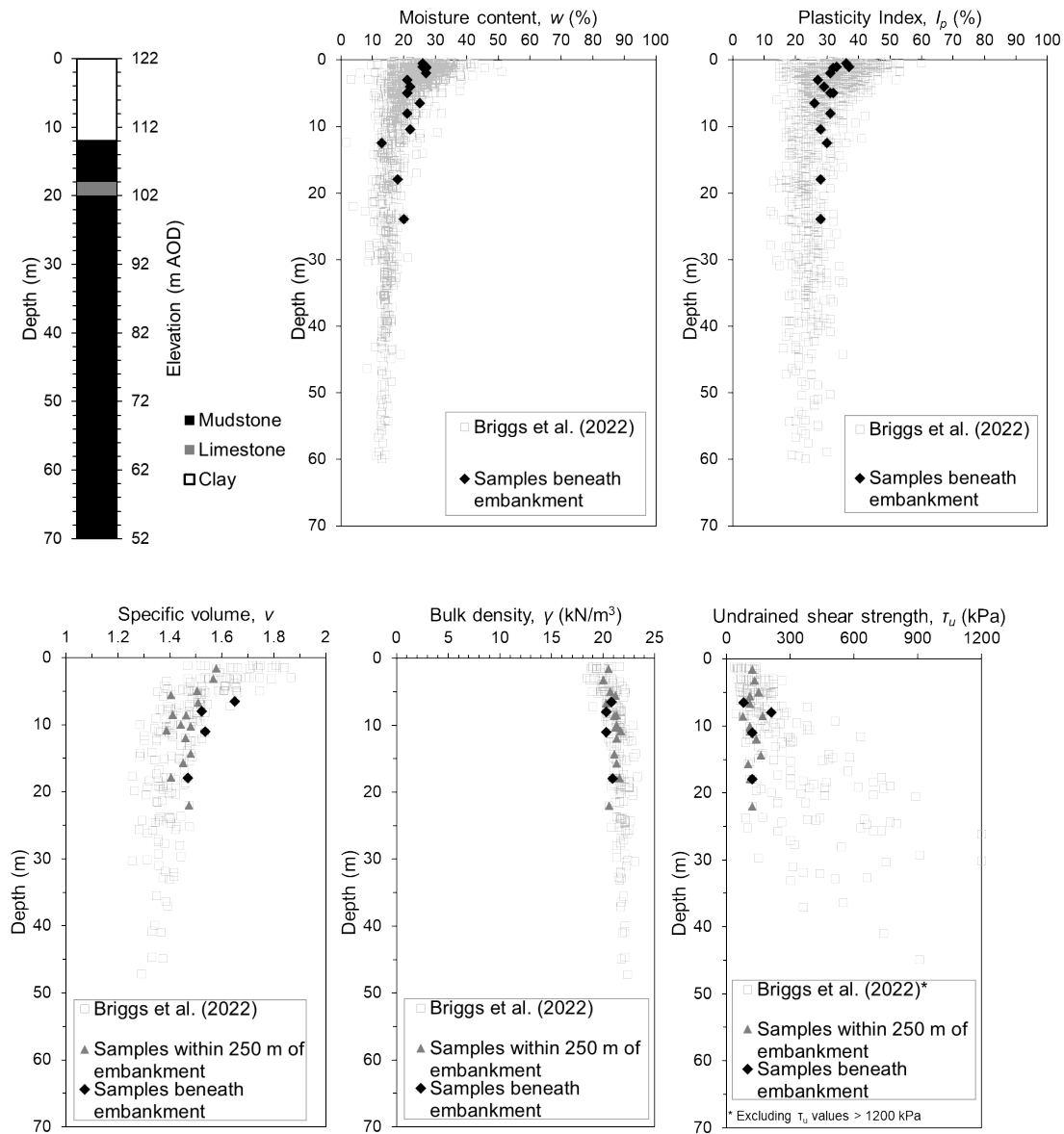


Figure 3: The ground profile derived from HS2 ground investigation data obtained beneath or near the trial embankment, showing (a) the geological profile shown in borehole strata descriptions, (b) the moisture content (%) profile, (c) the plasticity index (%) profile, (d) the specific volume profile, (e) the bulk density ( $\text{kN/m}^3$ ) profile and (f) the undrained shear strength (kPa) profile from unconsolidated undrained triaxial tests. Data from the Charmouth Mudstone Formation in Oxfordshire (Briggs et al. 2022) are shown for comparison.



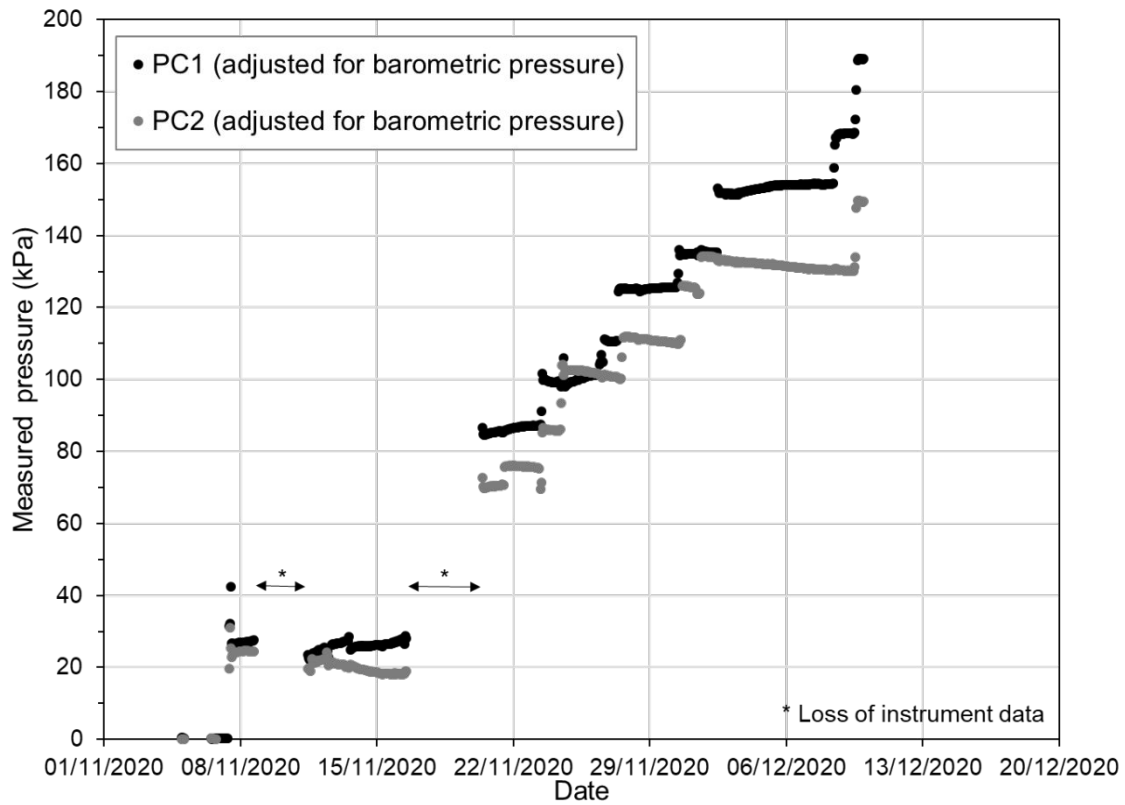


Figure 4: Total pressure (kPa) measured beneath the centre of the embankment (PC1) and beneath the edge of the embankment crest (PC2) using total earth pressure cells during embankment construction (November to December 2020).

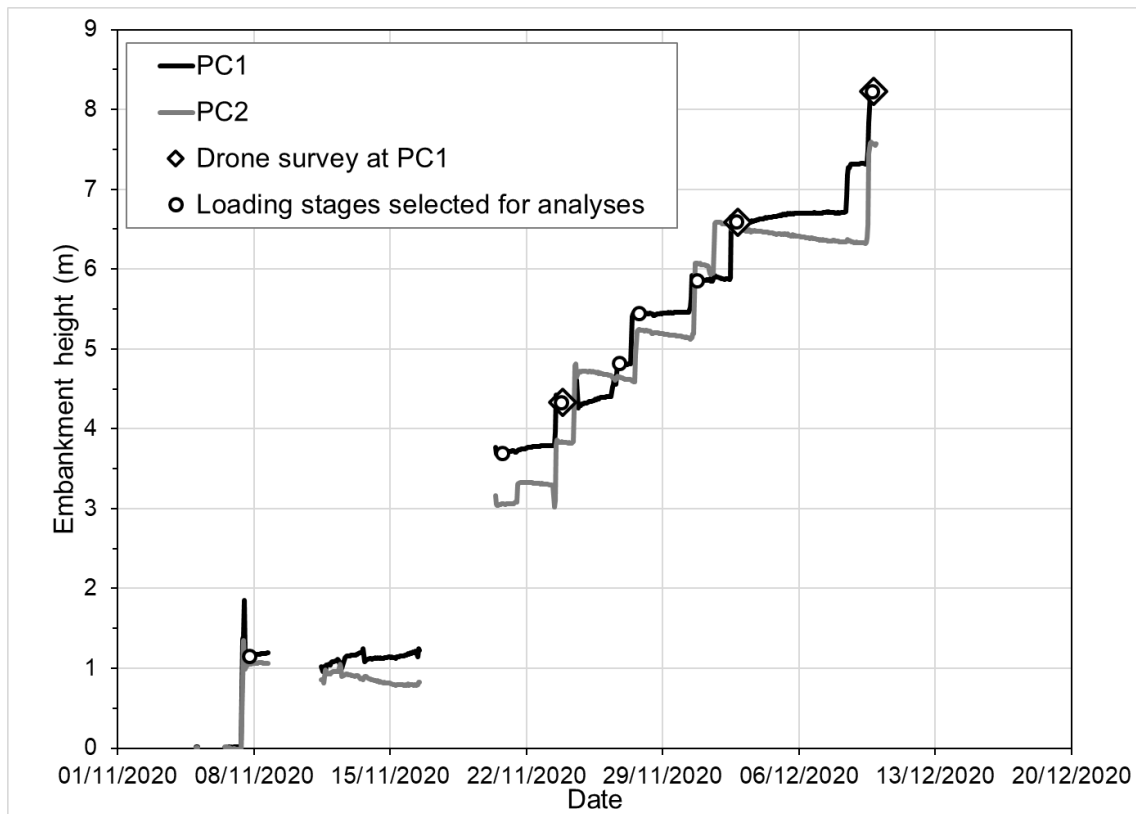
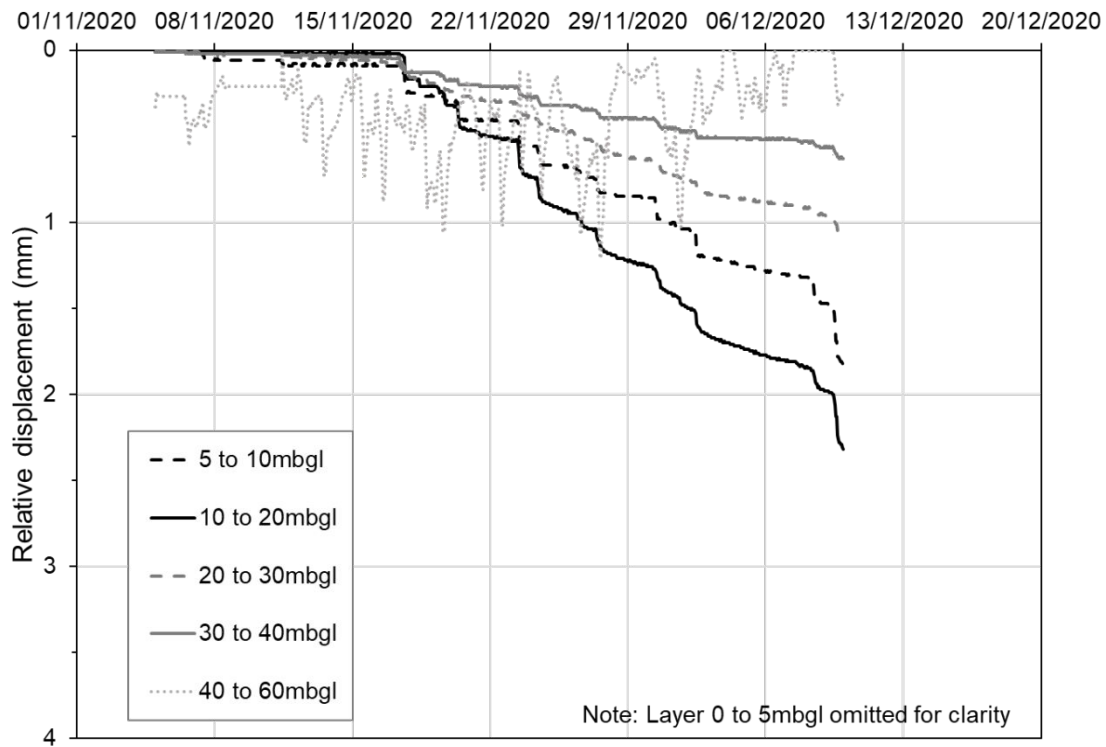
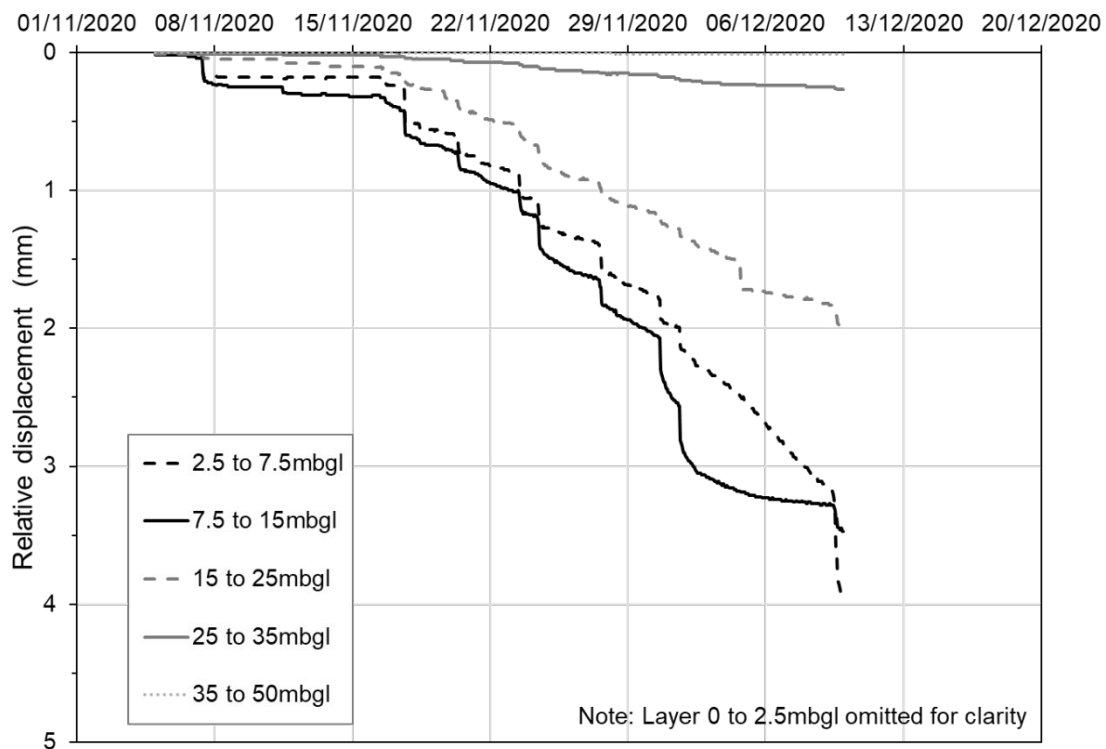


Figure 5: The change in embankment height with time, back-calculated from cell pressure measurements (PC1 & PC2) and the known embankment height from drone survey measurements at PC1.

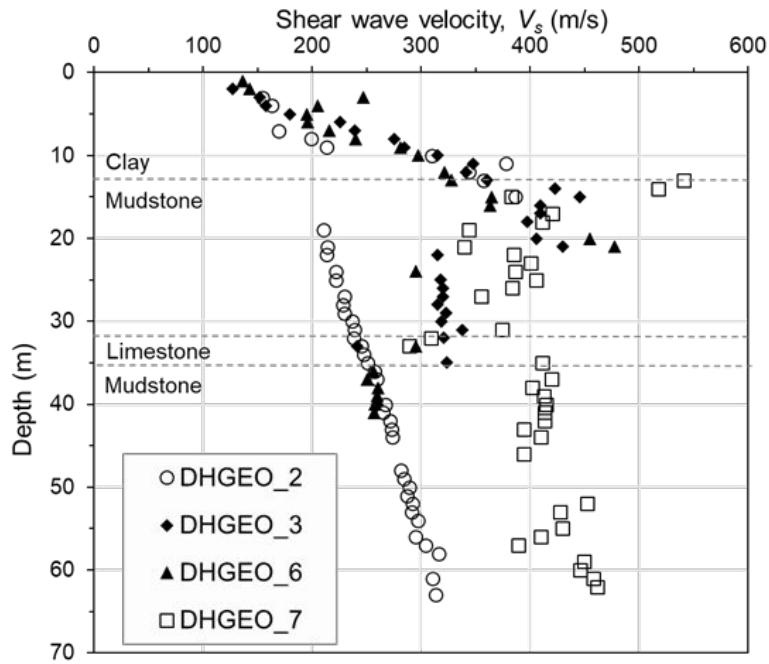


(a)

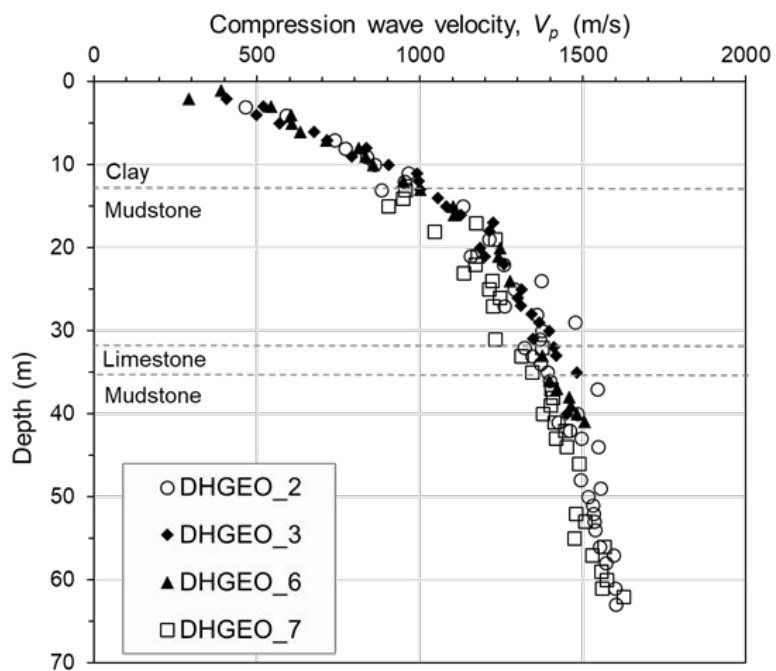


(b)

Figure 6: The relative displacement between extensometer anchors installed at various depths during construction of the trial embankment at (a) EXT1 and (b) EXT2. Note that the extensometer anchors were not equally spaced.



(a)



(b)

Figure 7: Downhole geophysical measurements of (a) shear wave velocity (m/s) and (b) compression wave velocity (m/s), at four boreholes (DHGEO\_2, DHGEO\_3, DHGEO\_6 & DHGEO\_7) located to the south of the trial embankment.

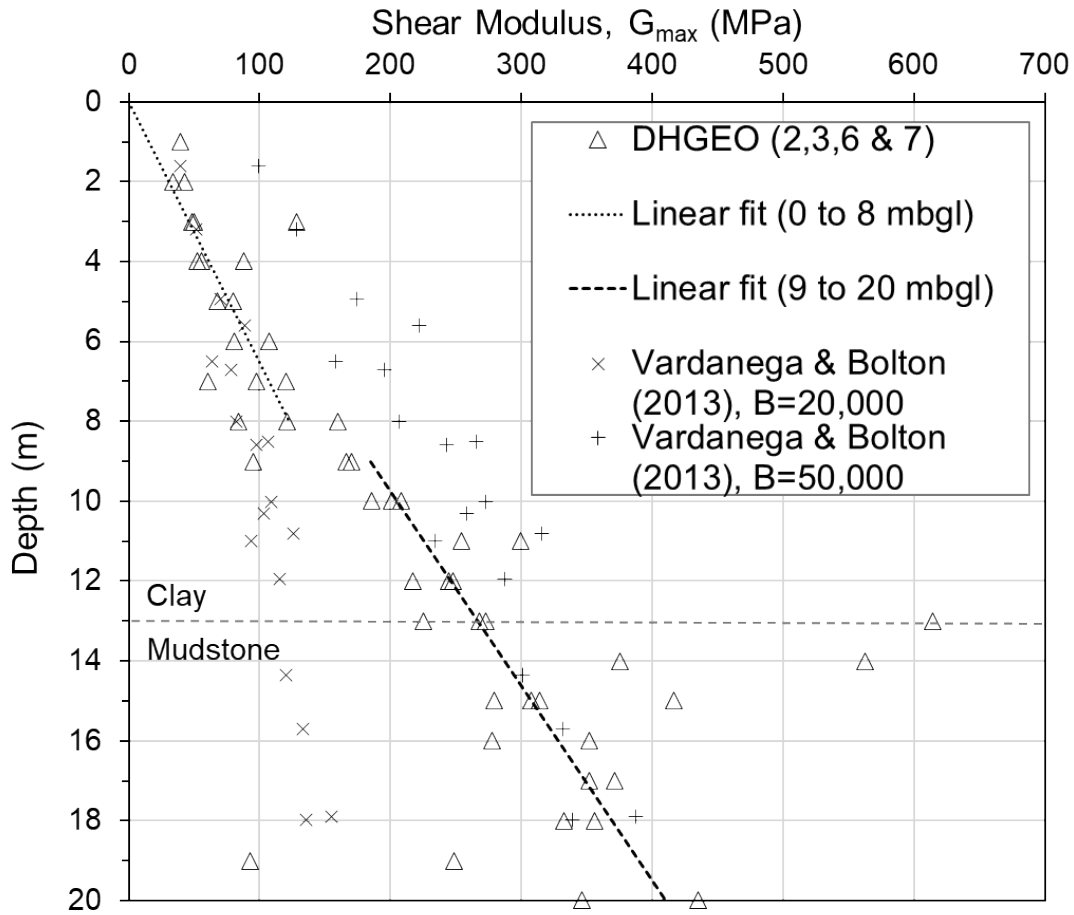


Figure 8: The maximum shear modulus profile ( $G_{max}$ ) derived from the downhole geophysical measurements. The Vardanega & Bolton (2013) model for fine-grained soils, plotted using the specific volume of the triaxial data (Figure 3), is shown for comparison. Linear regressions for  $G_{max}$  are shown for the weathered clay (0 to 8mbgl) and the transition to less-weathered clay and mudstone below (9-20 mbgl).

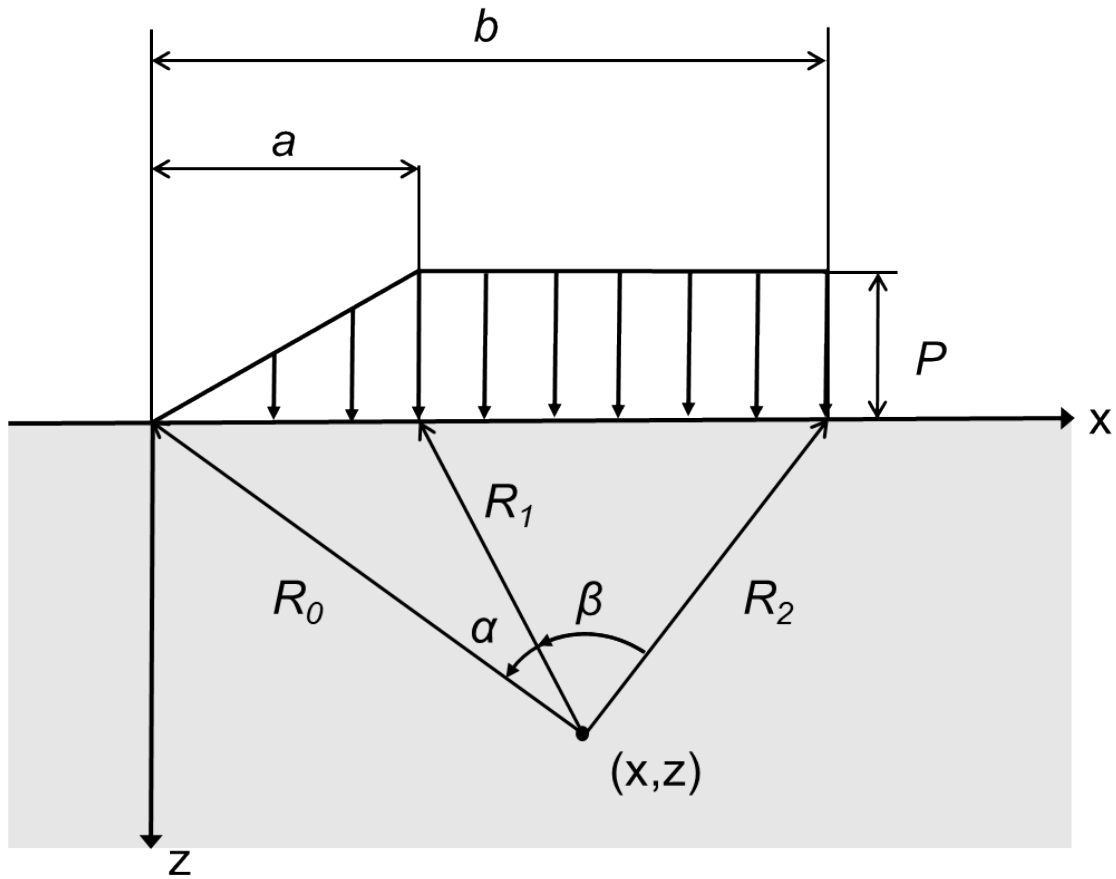


Figure 9: The geometry parameters for the distributed vertical embankment loading equations described in Poulos & Davis (1974).

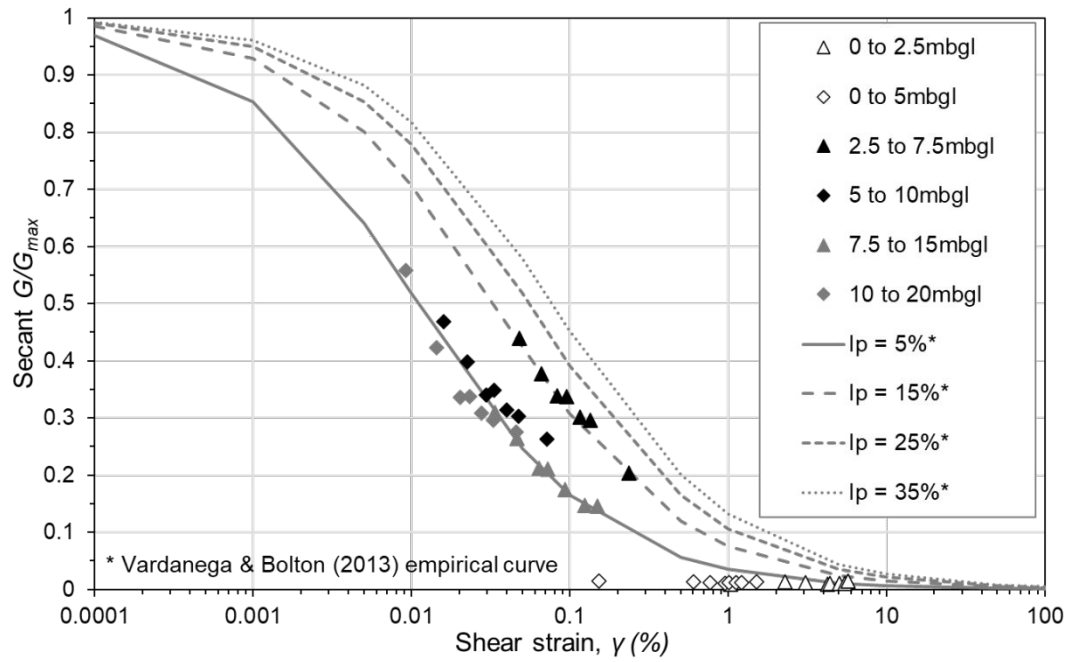


Figure 10: A normalised secant shear modulus reduction curve with strain for layers beneath the trial embankment, derived from monitoring data. These are compared to results from the Vardanega & Bolton (2013) model for fine-grained soils with plasticity indices ( $I_p$ ) of 5% to 35%.

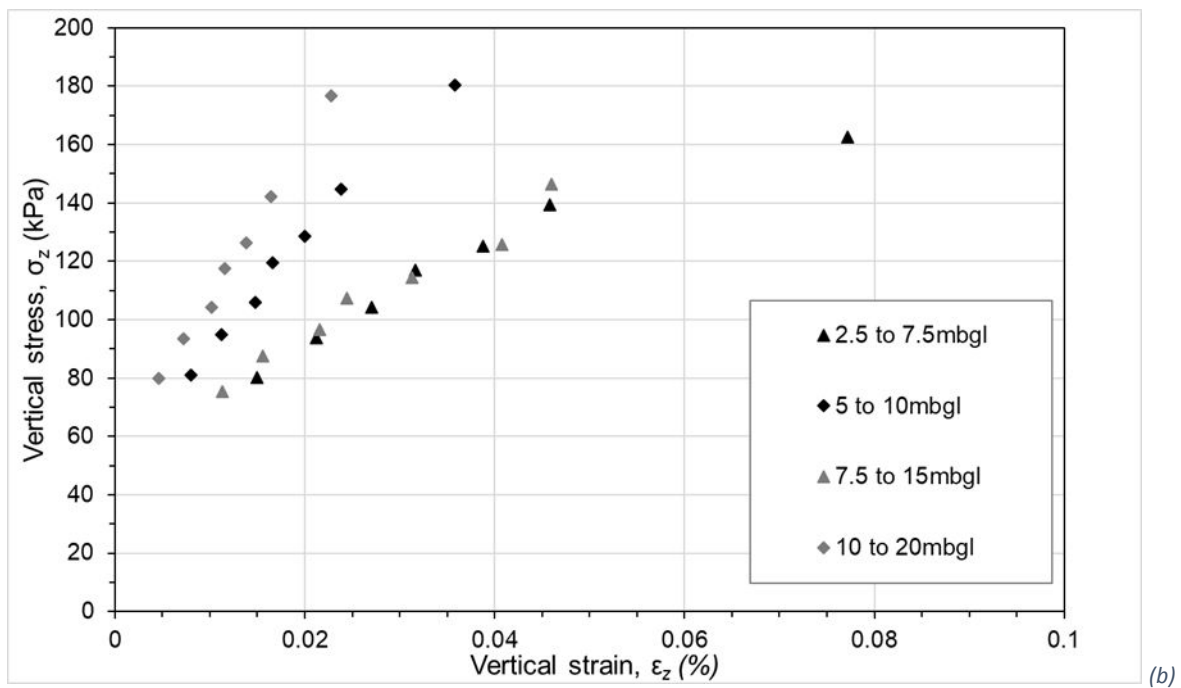
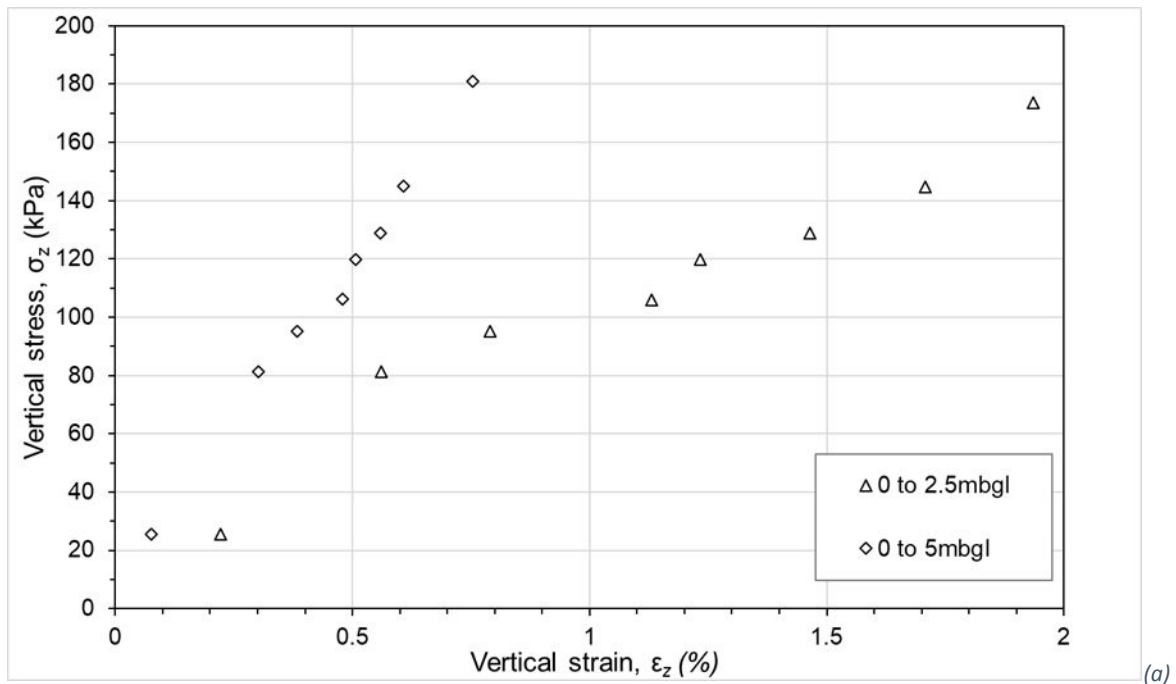


Figure 11: The average vertical stress ( $\sigma_z$ ) vs strain ( $\epsilon_z$ ) within the soil layers beneath the trial embankment, for eight loading stages, shown for (a) near surface layers (0 to 5 mbgl) at strains approaching yield and (b) deeper layers (2.5 to 20 mbgl) at medium strains.



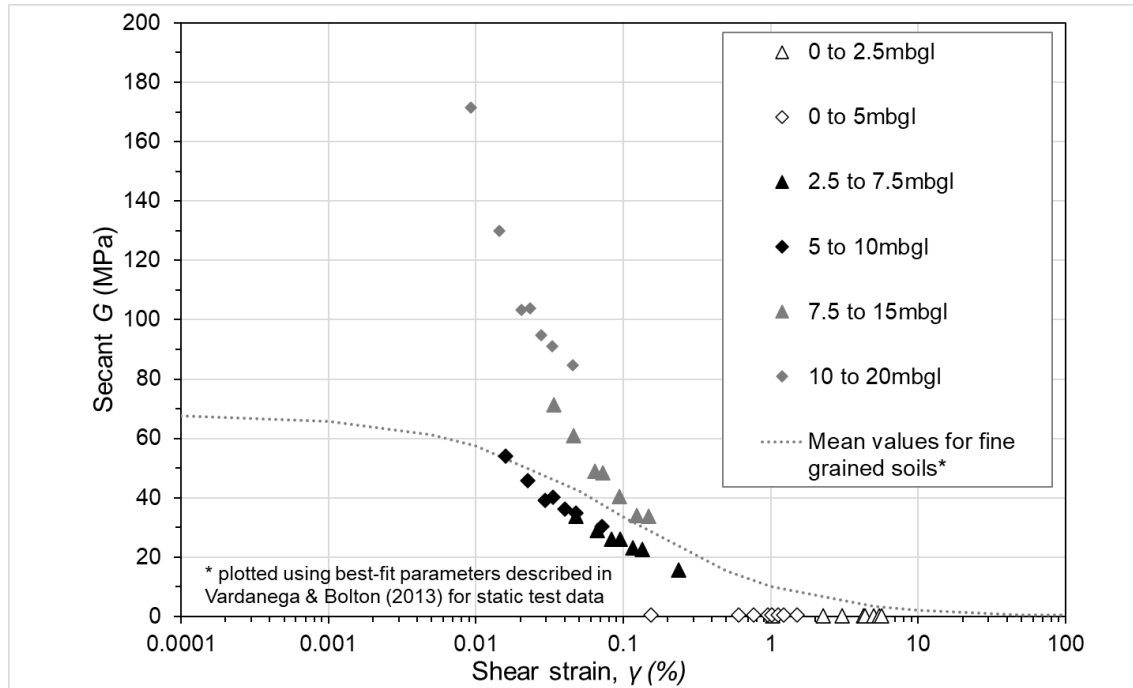


Figure 12: The secant shear modulus,  $G$  (MPa) vs shear strain,  $\gamma$  (%), for layers within the ground profile beneath the embankment. For comparison, a best-fit curve from the Vardanega & Bolton (2013) database is shown for the mean values for fine-grained soils (where  $I_p = 39\%$ ,  $p' = 209$  kPa,  $G_{max} = 68$  MPa).

## 1 Appendix A

2 Pressure cell PC2 was installed beneath the edge of the embankment crest and was influenced by its  
 3 proximity to the embankment slope. Elastic solutions for stresses at the base of an embankment, as  
 4 presented by Perloff et al. (1967) were used to quantify the influence of the embankment slope on  
 5 the measured stresses. These solutions link the embankment height at each loading stage ( $H$ ),  
 6 embankment crest width  $L$  and distance to the embankment centre line ( $x$ ) to the predicted  
 7 influence factor ( $I_z = \sigma_z / (H\gamma_{emb})$ ). Perloff et al. (1967) presented solutions for  $\nu = 0.3$  and various  
 8 slope angles ( $\theta = 15, 30, 45$  and  $60$  degrees) and crest widths ( $L/H = 0, 0.5, 1, 3$  and  $5$ ) in chart  
 9 form. The  $x$ -location of PC2 in Perloff's coordinate system is given by:

$$10 \quad \frac{x}{H} = \frac{L}{H} + \frac{1}{\tan \theta} \left[ \frac{H_{final}}{H} - 1 \right]$$

11 *Equation A 1*

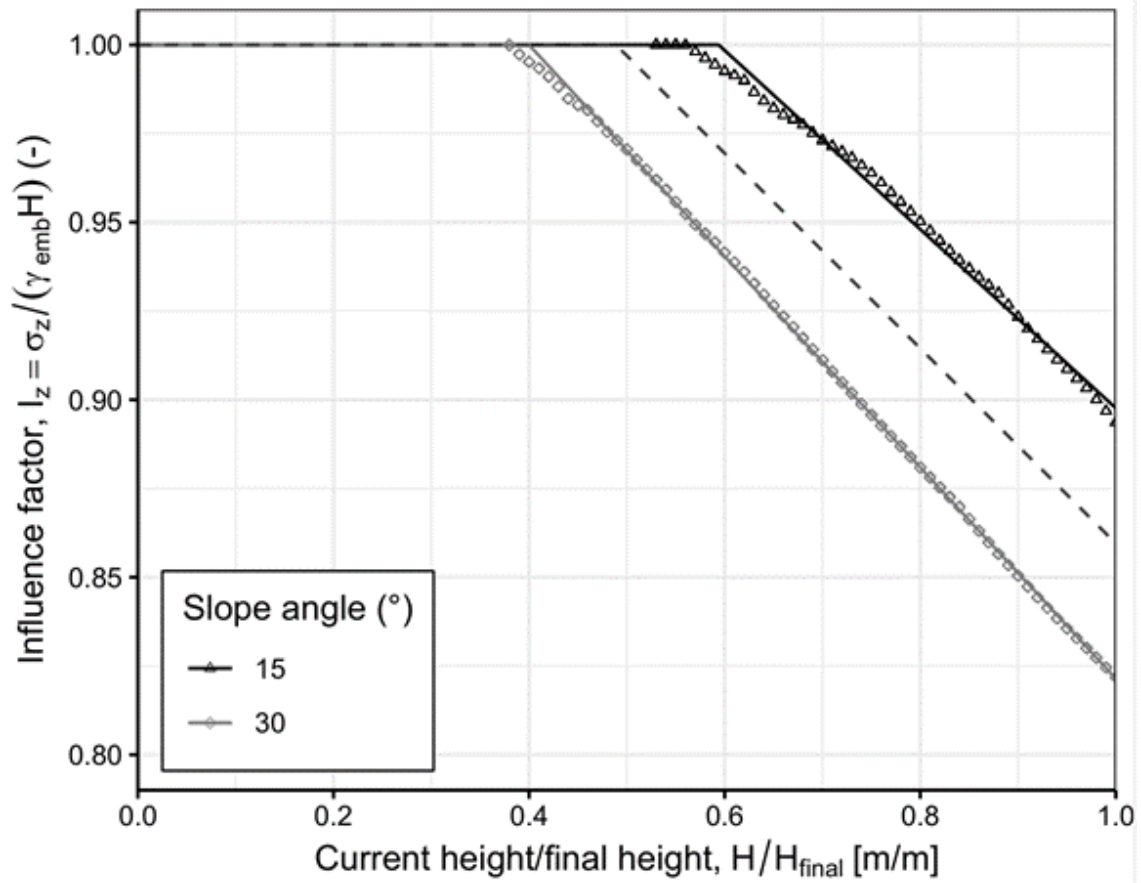
12 Where  $H_{final}$  is the final embankment height, i.e. 8.2 m

13 The Perloff et al. (1967) charts were digitised for  $L/H = 5$  and for  $\theta = 15$  and  $\theta = 30$  degrees. The  
 14 influence factors for PC2 were subsequently determined. The results (Figure A 1) show that when  
 15 the embankment height is still relatively low, measured stresses in PC2 will not be affected by the  
 16 embankment slope. However, the influence factors reduced in almost linear fashion beyond a  
 17 certain embankment height threshold. As the embankment had a slope angle close to 22.5 degrees  
 18 (halfway between the 15- and 30-degree cases), the intercepts and slopes of the fits for the 15 and  
 19 30 degree data were averaged. This resulted in:

$$20 \quad I_z = \frac{\sigma_z}{\gamma_{emb}H} = \min \left\{ \begin{array}{l} 1 \\ 1.134 - 0.275 \frac{H}{H_{final}} \end{array} \right.$$

21 *Equation A 2*

22 This is shown in Figure A 1.



23

24 Figure A 1: Influence factors (points) from the Perloff et al. (1967) chart and fitted linear regressions (solid lines). The  
 25 dashed line shows the fit for a 22.5° slope angle.

26

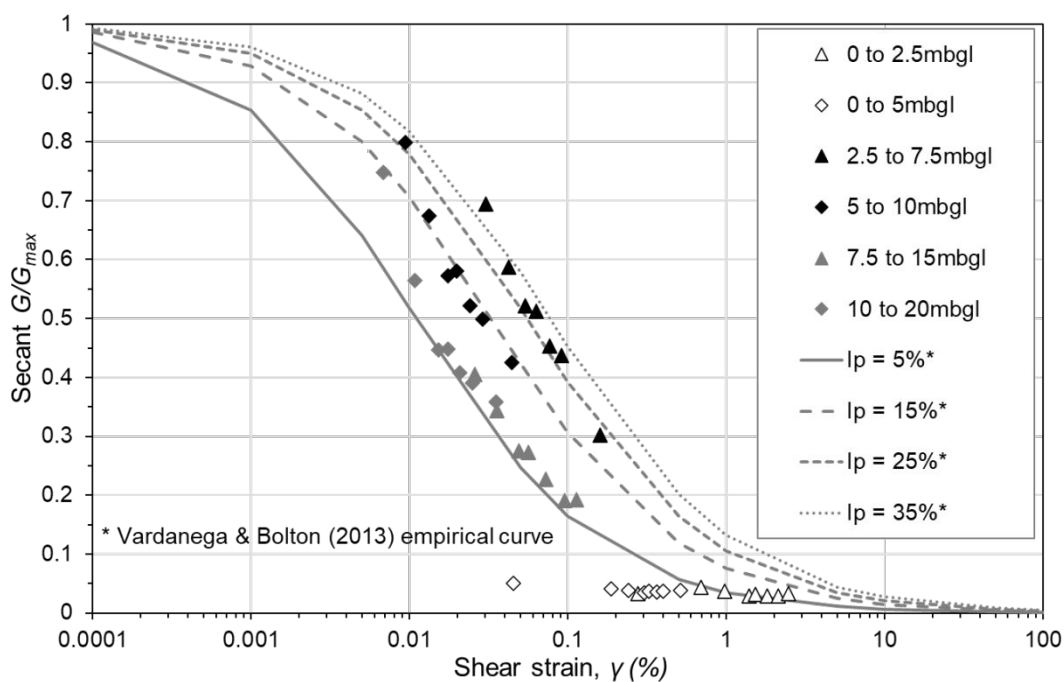
## 27 Appendix B

28 The interpretation of the measurements beneath the trial embankment assumed fully undrained  
 29 conditions, with no volume change (i.e.  $\nu_u = 0.5$ ). However, Briggs et al. (2024) showed that  
 30 Skempton's (1954)  $B$  value can reduce below unity in genuinely undrained conditions in stiff clays  
 31 and mudstones such as those of the Charmouth Mudstone Formation. This is because their high  
 32 stiffness relative to that of water makes the  $B$  values sensitive to very small reductions of the in-situ  
 33 saturation ratio. For example, a small reduction in saturation ratio to 0.995 (i.e. 99.5%) can reduce  
 34 the  $B$  value from unity to 0.2. The influence of the  $B$  value on the undrained Poisson's Ratio ( $\nu_u$ ) can  
 35 be calculated using:

$$36 \quad \nu_u = \frac{3\nu + B(1 - 2\nu)}{3 - B(1 - 2\nu)}$$

37 Equation B 1

38 Briggs et al. (2024) showed that  $B$  values beneath the trial embankment were approximately 0.6  
 39 during construction. A drained Poisson's Ratio ( $\nu$ ) was estimated as 0.4. Figure B1 shows the data  
 40 replotted with the assumptions that  $B = 0.6$ ,  $\nu = 0.4$  and therefore  $\nu_u = 0.458$ . Figure B 1 shows that  
 41 the data from the shallower layers (black symbols) move up (i.e. higher  $G/G_{max}$ ) and to the left (i.e.  
 42 lower shear strain) relative to their positions when  $\nu_u = 0.5$  (Figure 10). The deeper layers (grey  
 43 symbols) are less affected and remain close to the Vardanega & Bolton (2013) curve fit for  $I_p = 5\%$ .



44

45

46 *Figure B 1: A normalised secant shear modulus reduction curve with strain for layers beneath the trial embankment, derived*  
 47 *from monitoring data and assuming  $\nu_u$  equal to 0.458. These are compared to results from the Vardanega & Bolton (2013)*  
 48 *model for fine-grained soils with a plasticity index ( $I_p$ ) of 5% to 35%.*

49